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## Elusive ice clouds in the upper mesosphere of Venus

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### Introduction

H<sub>2</sub>O ice clouds in the mesospheres of terrestrial planets have been observed for Earth as Polar Mesospheric Clouds (PMCs) and are thought to be condensed on meteoric smoke [1, 2]. In contrast, on Mars both H<sub>2</sub>O and CO<sub>2</sub> ice clouds have been observed [3, 4]. These clouds play an important role in redistributing condensable material and the chemistry of the mesospheres of Mars and Earth. The possibility of sub-visible H<sub>2</sub>O ice clouds on Venus was speculated in 1983 by Turco, et al. [5] based on a limited set of observations (four Pioneer entry probes), however, no high-altitude ice clouds have ever been observed on Venus. More recent satellite observations have shed new light on the extreme low temperature conditions of the upper atmosphere [6]. Through satellite observation, experimental studies and modelling, our understanding of cloud formation and composition within terrestrial mesospheric clouds on Earth and Mars has improved significantly in recent decades. Here we apply this improved understanding of mesospheric clouds and the breadth of atmospheric observations to the possibility of ice clouds on Venus and importantly whether they could be detected in satellite observations.

Figure 1 shows an example polar temperature profile for Venus and the calculated saturation ratio with respect to both H<sub>2</sub>O and CO<sub>2</sub> ice. The temperature drops below 80 K between 120 – 130 km, causing significant supersaturations with respect to the solid phases of CO<sub>2</sub> and particularly H<sub>2</sub>O, conditions which are conducive to ice cloud formation.

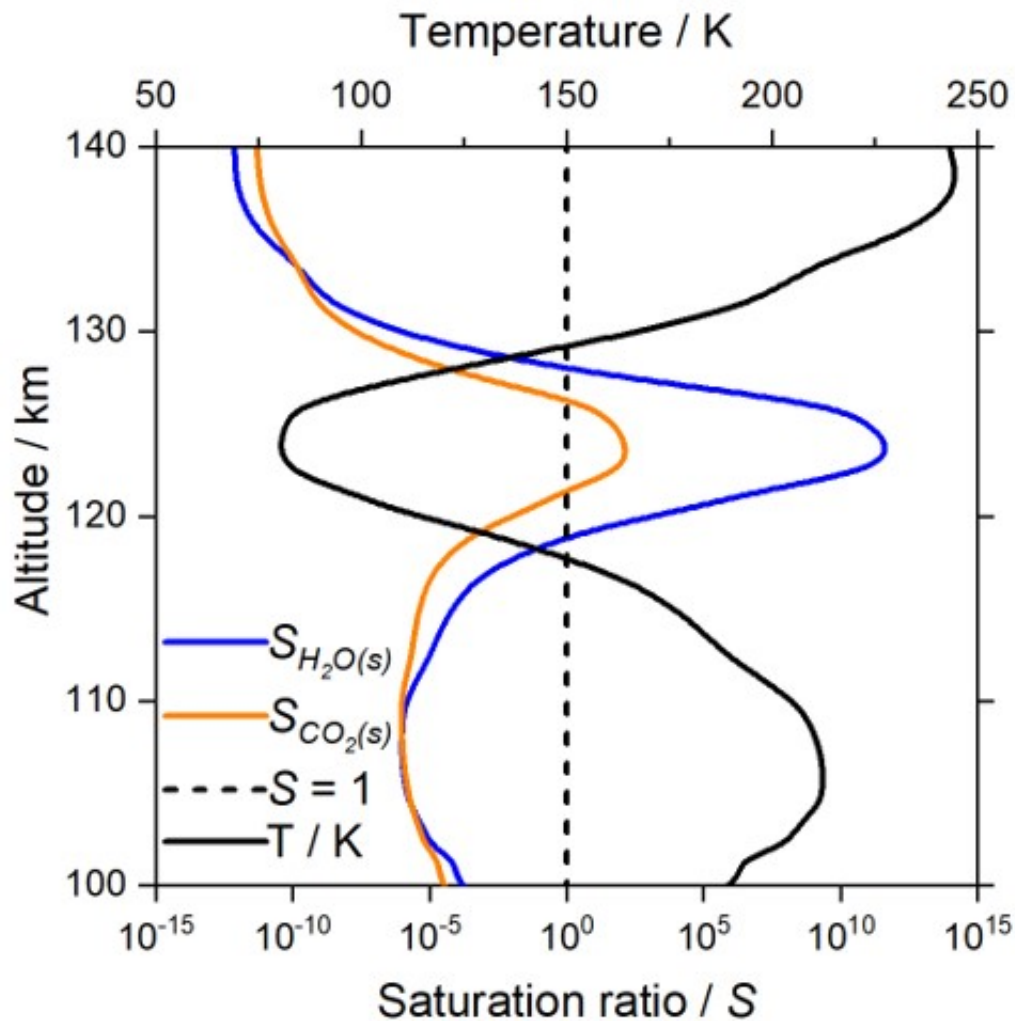


Figure 1: Saturation ratios ( $S$ ) for crystalline  $H_2O$  and  $CO_2$  ice derived from Venus Express observations. Here we use an average temperature profile at  $70\text{--}80^\circ$  Lat with an assumed constant  $H_2O$  volume mixing ratio of 2 ppm [6, 7].

### Findings

The first step in ice cloud formation is nucleation, which we evaluate homogeneously and heterogeneously for both  $H_2O$  and  $CO_2$  ice, nucleated by vapour deposition using Classical Nucleation Theory (CNT), an approach applied previously to Martian mesospheric clouds [8, 9]. Several nucleating particles are discussed and considered in the heterogeneous calculations, including meteoric smoke and in the case of  $CO_2$  nucleation,  $H_2O$  ice particles [8, 10, 11]. This theory is

applied to a planet-wide set of temperature observations (160+ occultations) from the SOIR instrument onboard Venus Express to help determine possible cloud locations and frequency. Preliminary results indicate that in the presence of suitable nanometre-sized nucleating particles, conditions exist on Venus for both H<sub>2</sub>O and CO<sub>2</sub> to nucleate, especially where the coldest temperatures are reached in the polar regions above 120 km. If visible clouds are possible in Venus mesosphere condensation of the primary constituent (CO<sub>2</sub>, significantly more abundant than H<sub>2</sub>O) onto the initial ice seeds is the likely pathway.

The growth of CO<sub>2</sub> ice particles are then predicted under favourable nucleation conditions, using a 1D model which accounts for growth, sublimation, and sedimentation of the particles. The model shows that the clouds form in a region of the atmosphere where the pressure is very low ( $< 1 \times 10^{-2}$  Pa). The ice particles therefore sediment rapidly into a warmer layer of the atmosphere and sublimate, so that they are very short-lived (typically  $< 5$  mins). Once peak particle sizes and lifetimes are determined, the extinction coefficient can be calculated for a realistic particle population using Mie theory, and hence the probability of the clouds being observed in occultation by a satellite spectrometer can be evaluated.

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