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Pressure-temperature conditions and significance of Upper Devonian eclogite and amphibolite facies metamorphisms in southern French Massif central

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Abstract – The southwestern French Massif central in western Rouergue displays an inverted metamorphic sequence with eclogite and amphibolite facies units forming the top of the nappe stack. They are often grouped into the leptyno-amphibolite complex included, in this area, at the base of the Upper Gneiss Unit. We sampled garnet micaschists and amphibolites to investigate their metamorphic history with isochemical phase diagrams, thermobarometry and U-Pb zircon dating. Our results demonstrate that two different tectono-metamorphic units can be distinguished. The Najac unit consists of biotite-poor phengite-garnet micaschists, a basic-ultrabasic intrusion containing retrogressed eclogites and phengite orthogneisses. Pressure and temperature estimates on micaschists with syn-kinematic garnets yield a prograde with garnet growth starting at 380°C/6–7 kbar, peak pressure at 16 kbar for 570°C, followed by retrogression in the greenschist facies. The age of high pressure metamorphism has been constrained in a recent publication between ca. 383 and 369 Ma. The Laguépie unit comprises garnet-free and garnet-bearing amphibolites with isolated lenses, veins or dykes of leucotonalitic gneiss. Thermobarometry and phase diagram calculation on a garnet amphibolite yield suprasolidus peak P-T conditions at 710°C, 10 kbar followed by retrogression and deformation under greenschist and amphibolite facies conditions. New U-Pb analyses obtained on igneous zircon rims from a leucotonalitic gneiss yield an age of 363 ± 3 Ma, interpreted as the timing of zircon crystallization after incipient partial melting of the host amphibolite. The eclogitic Najac unit records the subduction of a continental margin during Upper Devonian. It is tentatively correlated to a Middle Allochthon, sandwiched between the Lower Gneiss Unit and the Upper Gneiss Unit. Such an intermediate unit is still poorly defined in the French Massif central but it can be a lateral equivalent of the Groix blueschists in the south Armorican massif. The Uppermost Devonian, amphibolite facies Laguépie unit correlates in terms of P-T evolution to the Upper Gneiss Unit in the Western French Massif central. This Late Devonian metamorphism is contemporaneous with active margin magmatism and confirms that the French Massif central belonged to the continental upper plate of an ocean-continent subduction system just before the stacking of Mississippian nappes.

Keywords: Variscan / Rouergue / HP metamorphism / partial melting


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1 Introduction

The evaluation of pressure-temperature-time paths in orogenic belts is fundamental to reconstruct the geotectonic evolution before, during and after the orogenesis. Specifically, pre-collisional eclogite and blueschist high-pressure (HP) units formed in the subducting slab (Ernst, 2001) and amphibolite to granulite high-temperature (HT) units belonging to the upper plate (Hydmann et al., 2005) may contain crucial information on the dynamics of pre-orogenic subductions (Brown, 2006). The stacking of these subduction-related HP and HT metamorphic rocks onto greenschist facies units during exhumation and/or collision can lead to the formation of inverted metamorphic sequences (Peacock, 1987) where metamorphic conditions increase upward in the tectonic stack and where the boundary between units is marked by localized deformation and sharp contrasts in metamorphic conditions.

Pioneer studies on the Variscan French Massif central (FMC) showed that pre-Visean metamorphic units have a complex tectonic distribution (see Burg et al., 1984, 1989; Ledru et al., 1989, 1994; Santallier et al., 1994; Faure et al., 2005, 2009; Lardeaux, 2014; Lardeaux et al., 2014), with greenschist facies units forming the lowermost and uppermost parts of the tectonic pile (the para-autochthon and the so-called epizonal units, respectively). Amphibolite facies units locally preserving relic of eclogites and HP granulites rocks are classically grouped into the Lower Gneiss Unit (LGU) and the Upper Gneiss Unit (UGU) in the French Massif central (see Bellot and Roig, 2007) but investigations on the neighboring Armorican massif (Bosse et al., 2000, 2005; Ballèvre et al., 2009; Pitra et al., 2010; Ballèvre et al., 2014) pointed out the existence of a Late Devonian ophiolitic and eclogite/blueschist-bearing Middle Allochthon sandwiched between the upper and lower units. Berger et al. (2010a, 2010b) proposed to introduce a Middle Allochthon in the FMC located between LGU and UGU and composed of HP to UHP eclogites with host orthogneiss and micaschists but also of ophiolites marked by a post-oceanic equilibration in the granulite facies (Santallier et al., 1994; Berger et al., 2010b). This subdivision, still immature for the French Massif central, group together rocks with contrasted metamorphic evolutions. More detailed metamorphic petrology studies are thus needed to propose a better scheme of subdivision and coherent geodynamic interpretations. Furthermore, this metamorphic organization is not integrated in the most recent compilations on the metamorphic evolution of the French Massif central (Faure et al., 2009; Lardeaux, 2014) in which most HP metamorphic rocks are included within the Upper Gneiss Unit encompassing the enigmatic leptyno-amphibolite complex (Santallier et al., 1988). The latter is defined as a bimodal association of felsic and mafic gneisses frequently containing relics of eclogite and granulite facies rocks and often located at the base of UGU, sometimes in LGU. These apparent conflicting views are, on the one hand due to the poor quality of exposures in the French Massif central, and on the other hand to the lack of detailed quantitative metamorphic studies coupled to petrochronological investigations that are to propose robust pressure-temperature-time (P-T-t) paths.

Recent dating performed on eclogites from Southern Massif central yielded Upper Devonian to Lowermost Carboniferous ages (Lotout et al., 2018, 2020) while the HP metamorphism and the closure of oceanic domains in the French Massif central are typically considered to be around 410–390 Ma (Eo-varican cycle or D0 event of Faure et al., 2009; see also Lardeaux, 2014). Ages of HP metamorphism published between 1980 and 2000 have to be re-evaluated.
and, more generally, petrochronological data on metamorphic units are required to better constrain the Devonian evolution of the French Massif central. These data are crucial to determine if the West European Variscan belt is polycyclic and marked by two phases of convergence and ocean closure (Faure et al., 2005, 2008, 2009; Lardeaux, 2014) or monocyclic with a single event of continental collision (Ballèvre et al., 2009) or somewhere between the two endmembers models.

This paper addresses the question of the subdivision of amphibolite and eclogite facies metamorphic units and their geological meaning in southern French Massif central near Najac in Western Rouergue (Burg et al., 1984, 1989). According to most authors, the metamorphic units in this area are thought to belong to the leptyno-amphibolite complex and are often integrated into the UGU (Lardeaux, 2014; Lotout et al., 2018). Our study shows that their constitutive rocks display contrasted P-T-t-d histories. Therefore, we provide new P-T estimations on phengite-garnet micaschists from Najac and on migmatic amphibolites from Laguépie. These results are coupled to U-Pb zircon dating in the latter unit and interpreted considering the new and robust multi-method ages provided by Lotout et al. (2018) on an eclogite from Najac.

2 Geological setting

The Variscan French Massif central consists of several tectono-metamorphic units that were stacked before the Visean (see Faure et al., 2009; Lardeaux, 2014). Despite the poor quality of their exposures, the metamorphic units are well defined on geological maps and in the literature in Limousin, Haut-Allier and Rouergue domains (Dubuisson et al., 1988; Ledru et al., 1989; Ledru et al., 1994; Burg et al., 1989a). This stack is characterized by an apparent inverse metamorphic gradient, with amphibolite facies and eclogite facies units (Lower Gneiss and Upper Gneiss Units; LGU and UGU, respectively) commonly lying above greenschist units (Para-Autochthon; see Burg et al., 1984; Lardeaux, 2014). At the scale of the French Massif central, the pre-Carboniferous tectono-metamorphic evolution is subdivided into three phases following (Faure et al., 2005, 2008, 2009, 2017): D0 represent the Late Silurian/Early Devonian (420–400 Ma) HP metamorphic phase; D1 relates to amphibolite facies migmatisation in LGU and UGU associated with a top-to-the-southeast kinematics, it occurred during Middle and Upper Devonian (385–375 Ma); D2 (360–350 Ma) is marked by top-to-the-northwest kinematics under amphibolite facies conditions. However, recent petrochronological investigations performed by Lotout et al. (2018, 2020) yielded Upper Devonian to Early Carboniferous ages for HP metamorphism in southern French Massif central (Najac and Lévézou). These new findings suggest that the D1 and D2 phases as defined by Faure et al. (2009) are also contemporaneous with formation and exhumation of eclogites in southern French Massif central.

Outcrops of the Viaur and Aveyron valleys near Najac and Laguépie were described in details by Collomb (1970), Bodinier and Burg (1981) and Burg et al. (1984). The para-autochthon consists of greenschist facies metapelites (chlorite-muscovite ± biotite bearing) with rare intercalation of felsic, often meta-volcanic, orthogneiss (Fig. 1). The Lower Gneiss Unit is separated from the para-autochthon by a low angle mylonitic shear zone. It is made of orthogneiss representing former felsic plutons with rare lenses of amphibolite facies biotite-muscovite ± garnet metapelites. The contact with the Najac-Carmaux unit is not well exposed but it is interpreted as a thrust contact considering that the latter contains eclogite facies rocks that were not found in the Lower Gneiss Unit (see Burg et al., 1984). The Najac-Carmaux klippe as defined by Burg et al. (1984) consists of an association of phengite-rich micaschists (often garnet-bearing, with minor amounts of biotite and locally accompanied by chloritoid and staurolite; Delor et al., 1987), phengite-bearing porphyroclastic orthogneisses and basic-ultrabasic complexes known as the Najac and Laguépie massifs (Delor et al., 1986; Bodinier et al., 1986; Lotout et al., 2018). Bodinier and Burg (1981) and Bodinier et al. (1986) interpret the eclogitic layered intrusion south of Najac as a former laccolith that intruded the precursor of micaschists and orthogneiss prior to high-pressure metamorphism. We observed small (max 10 cm width) xenoliths of phengite orthogneiss included into eclogite facies metabasites south of Najac, close to Mergieux. This supports the conclusion considering the Najac basic-ultrabasic massif as a former intrusion within the orthogneiss. Bodinier and Burg (1981) and Bodinier et al. (1986) also noted that the Najac eclogitic basic-ultrabasic massif and the Laguépie amphibolite recorded distinct metamorphic histories but so far, no metamorphic studies have been performed on the amphibolites.

Metamorphic foliations in the Viaur and Aveyron valleys are dipping to the east (20–70° dip) except in the eastern part of the map (Fig. 1). They are affected by small scale overturned N-S folds with an axial plane facing west. Stretching lineations are generally trending N350 with low plunges (<30°). Asymmetric criterion for sense of shear are well expressed along the mylonitic band separating the para-autochthon and the Lower Gneiss Unit is top-to-the-north (D2 phase of Faure et al., 2009). The eclogitic orthogneisses and micaschists of the Najac unit also carry a N280–300-trending stretching lineation with top-to-the-southeast kinematics (deduced from C/S structures and asymmetric feldspar porphyroclasts in orthogneisses).

Northeast of Najac, the whole metamorphic massif is intruded by the Carboniferous Villefranche de Rouergue
Fig. 1. Geological setting of the Najac and Laguépie units: (a) geological map modified after Burg et al. (1989a). The red arrows indicate the kinematics deduced from C’/C/S structures; (b) interpretative W-E cross section with indication of lithology and metamorphic facies for each unit.
granodioritic pluton (Burg et al., 1989b). It is bounded by the Villefranche fault to the west, which separates the metamorphic units from the large Permian to Mesozoic Quercy basin and it is unconformably overlain by Permian and Cenozoic deposits to the south.

3 Samples description and mineral chemistry

In order to compare the metamorphic evolution of Najac and Laguépie units, we focus our study on several samples of phengite-garnet micaschists surrounding the basic-ultrabasic complex and on amphibolites and associated leucotonalitic gneisses at Laguépie. The orthogneisses around Najac are deformed but their mineral assemblage did not fully reequilibrate during metamorphism (except for some late kinematic Si-rich white micas); these samples were consequently disregarded in this study.

3.1 The Laguépie unit

Samples were collected in an abandoned quarry along the Aveyron valley, west of Saint Martin de Laguépie. The amphibolites are homogeneous amphibole-plagioclase rocks alternating with discontinuous leucotonalitic orthogneisses forming isolated lenses (several millimetres to 20 cm in width) (Fig. 2a). In the eastern side of the quarry, one outcrop exposes discontinuous and rounded levels of melanocratic garnet amphibolite wrapped into large veins of leucogness. Contacts between both lithologies can be either sharp or lobed, occasionally resembling cauliflower structures in migmatites (Burg and Vanderhaeghe, 1993) (Fig. 2b–d). The leucogness locally cuts across the melanocratic levels and it can also forms deformed lenses within the amphibolites (Fig. 2c). These observations suggest that the unconnected felsic lenses were formed in situ in the amphibolites while large felsic veins or dykes containing lobed fragments of garnet amphibolites are the
results of melt drainage and segregation. Accordingly, these field observations suggest that the precursor leucotonalitic magma formed after incipient partial melting of the amphibolite.

The most common lithology of the Laguépie unit is an amphibolite made of amphibole, plagioclase and quartz. Plagioclase is commonly cloudy due to advanced sericitization and frequently contains epidote neoblasts (Fig. 3a, b). Amphibole porphyroclasts are pleochroic from blue-green to brownish and small needles of light colored green amphibole are found at the contact with plagioclase (Fig. 3a, b). Quartz is common, but never exceed 10 vol.%. The texture is porphyroclastic with oblique localized cataclastic shear bands consisting of crushed chlorite, actinolite and albite grains (Fig. 3a). Secondary calcite and pyrite-chlorite veins cut shear bands indicating a top-to-the-north sense of shear.

Garnet amphibolite levels consist of a mineral assemblage of plagioclase-amphibole–garnet ± titanite ± quartz ± k-feldspar with large (up to 3 cm width), subrounded, pre-kinematic porphyroclastic garnet rich in quartz inclusions (Fig. 3c). Elongated porphyroclasts of amphibole and plagioclase alternating with rare quartz ribbons define a foliation wrapping the garnet (Fig. 3c). Titanite, apatite, zircon and pyrite are the most common accessory phases. Retrogression is evidenced by the cloudy aspect of plagioclase, the development of light-colored amphibole needles rimming large brownish porphyroclastic grains, growth of chlorite within fractures cutting across garnet and thin cataclastic shear bands made of chlorite, albite and epidote.

Leucotonalitic gneisses consist of an assemblage of quartz–albite-chloritized biotite + K-feldspar ± garnet and display an augen texture (Fig. 3d) with large porphyroclastic feldspars. Garnet are rounded and small (≤1 cm) compared to those observed the amphibolite layer. Their modal abundance usually decreases from the contact with the amphibolite to the core of the leucocromatic orthogneiss vein. Chlorite, probably replacing a former biotite, forms elongated domains aligned along the foliation or C−C’ shear bands indicating a top-to-the-north sense of shear.

One amphibolite sample (LAG3b) and one garnet-amphibolite sample (AJAH08), collected close to a leucotonalitic vein, were analysed for their mineral composition. The garnet-free amphibolite LAG3b is made of oligoclase (Xan: 0.13–0.17) with albigitic rims (Xan: down to 0.07). Amphibole appears as large brown-green crystals characterized by low-Ti (0.94–1.25 wt% TiO2) magnesiohöyburne (Mg#: 54–55, Si: 6.72–6.78 a.p.f.u., Fig. 4a) core surrounding Si-rich and Ti-poor thin rim (6.84–7.09 Si a.p.f.u.; 0.33–0.69 wt % TiO2 Fig. 4a). Unaltered plagioclase grains in the garnet and thin cataclastic shear bands made of chlorite, albite and epidote.

3.2 Micaschists of the Najac unit

Micaschists of the Najac units are homogeneous light gray-blue rocks displaying a penetrative planar fabric with poorly defined mineral lineation. Granoblastic and elongated quartz and plagioclase form millimeter-scale layers alternating with phengite-rich bands (Fig. 3a). Biotite is present in low volumetric proportions but is preferentially concentrated at the contact between garnet and white mica (Fig. 3e). In most samples, garnet forms minute (<300 μm in diameter) subhedral grains within the phengitic layers and show pressure shadows filled with biotite and phengite (Fig. 3e). Garnet is rich in quartz and rutile inclusions sometimes delineating a foliation that can be parallel or oblique to the main planar fabric. NJC08 contains centimetric syn-kinematic fibrous snowball garnet porphyroblasts with quartz inclusions and enclosed into phengite stacks (Fig. 3f), substantiating that garnet growth occurred during deformation. Accessory minerals consist of rutile (partly transformed into ilmenite), apatite and zircon.

Garnet chemical zoning is obvious in sample NJC11 where compositions range from Alm₄₄Sp₄₄Pr₅₄Grs₂₅ in the core to Alm₆₆Sp₄₄Pr₅₄Grs₂₅ in the rim (Fig. 5b), while its Fe₉# decreases from core to rim (96 to 92). Garnet in NJC12a shows a more homogenous composition with core at Alm₆₈Sp₄₄Pr₅₄Grs₂₂ and a rim at Alm₇₂Sp₃₅Pr₅₄Grs₁₉ at nearly constant Fe₉# (91–93). White micas are phengitic with Si (per formula unit, on the basis of 11 oxygens) ranging from 3.30 to 3.40 in NJC08, 3.18 to 3.36 in NJC12a and 3.17–3.35 in NJC11; and Fe+Mg (a.p.f.u.) between 0.20 and 0.45. Phengite crystals with the lowest Si content are found close to biotite and in the rim of large grains. Biotite has generally low Ti (0.09–0.11 a.p.f.u.) with Mg# varying from 0.45 and 0.50 and low Si (2.70–2.80 a.p.f.u.) in the three analyzed samples. Plagioclase is an oligoclase with Xan comprised between 0.14–0.16 in NJC11 and NJC12a.

4 Pressure-temperature estimations

4.1 Amphibolites from the Laguépie unit

Garnet amphibolite AJAH08 collected in the outcrop shown in Figure 2c, d was selected for detailed thermobarometric investigations because it contains garnet, a good tracer to recover P-T conditions. The sample’s bulk composition has been calculated using modal proportions (estimated by pixel counting on thin section images) and core composition of unaltered rock forming minerals (garnet, amphibole, plagioclase, titanite and quartz). Two different estimations were made and the results were differing by less than 5% (relative) on the different oxides proportions. Phase diagrams were calculated with PerpleX (Connolly, 2009) in the system Na₂O-CaO-FeO-MnO–MgO-TiO₂–Al₂O₃–SiO₂–H₂O. Water content has been estimated on the basis of the modal proportion of amphiboles, considering they contain approximately 2 wt% H₂O. The calculation uses the thermodynamic database of Holland and Powell (2011, update ds62) and solid solution models adapted for amphibolite to granulite facies mafic rocks (see Green et al., 2016). Ferric iron was not considered because its content in minerals (estimated by stoichiometry) is very low and no oxides containing ferric iron...
Fig. 3. Photomicrographs of representative samples from the Najac and Laguépie units. Laguépie samples: (a) amphibole-plagioclase-ilmenite amphibolite LAG3b representing the dominant lithology of the Laguépie amphibolite (1,9558°E, 44,1485°N); (b) chlorite-epidote-albite shear band oblique to the main foliation and indicating a top to the north sense of shear (sample LAG3b); (c) garnet amphibolite AJAH08 with foliation molded around the garnet porphyroclast; (d) leucogneiss from sample AJAH08 showing a porphyroclastic augen texture. Note the C/S structure marked by chlorite formed after solidification and cooling of the leucotonalitic orthogneiss; (e) Representative texture of the Najac unit garnet-phengite micaschists (sample NJC12a; 1,9754°E, 44,2159°N). Note the pressure shadows filled with biotite and phengite around garnet indicating that the high pressure assemblage is syn-kinematic. Greenish chlorite partly replaces garnet and biotite in all samples; (f) large porphyroclastic snowball and skeletal garnet from sample NJC11 (1,9693°E; 44,2196°N), again showing that the high-pressure mineral assemblage developed during deformation.
Fig. 4. Minerals compositions in the amphibolites from the Laguépie unit: (a) amphibole composition in the classification diagram by Leake et al. (1997); (b) garnet compositions in AJAH08.

Fig. 5. Minerals composition in the micaschists from the Najac unit: (a) Si versus Fe+Mg diagram for white micas showing the phengitic substitution; (b) garnet zoning profile in NJC11; (c) garnet composition for three representative samples (NJC8, NJC11 and NJC12a).
were observed. The mineral assemblage observed in the garnet amphibolite AJAH08 (plagioclase-amphibole-garnet ± titanite ± quartz) cannot be reproduced in the isochemical phase diagram because the calculation systematically predicts the presence of small amounts (<4 vol.%) of clinopyroxene in fields where amphibole, garnet and plagioclase are stable. This can be a consequence of retrograde metamorphism that has transformed clinopyroxene into amphibole or inaccuracies of the calculated model. Garnet, plagioclase, amphibole, quartz and titanite without orthopyroxene, epidote, ilmenite or rutile coexist between 8 and 14 kbar and 600 and 740 °C (Fig. 6). XAn and Mg# isopleths for plagioclase and amphibole, respectively, cross between 8 and 10 kbar at 600–700 °C fields where low amounts of clinopyroxene are predicted (<4 vol.%) (Fig. 6). The overlapping zone of various garnet isopleths (Fe#, grossular, spessartine contents) yields P-T conditions of 710 ± 15 °C and 10 ± 1 kbar (Fig. 6). As the garnet cores display higher Fe# (up to 84) and grossular content (up to 26 mol%) compared to the rim (Fe# down to 80 and 26 grossular mol%), the slight zoning observed in garnet could trace a small increase in temperature from 690 to 730 °C at nearly constant pressure (Fig. 6). This P-T domain locates close to the ilmenite-titanite and ilmenite-rutile phase transitions but also clinopyroxene (<4 vol.%) and melt present (<2 vol.%) fields (Fig. 6); consequently, it lies above the vapor-present solidus calculated for the amphibole-plagioclase-ilmenite amphibolite LAG3b. It is worth to note that the calculated phase diagram closely reproduces the observed modal proportion of garnet (7 vol.%) around 700 °C and 10 kbar.

Application of garnet-amphibole-plagioclase-quartz barometry (Kohn and Spear, 1990) with hornblende-plagioclase thermometry (Holland and Blundy, 1994) yields 640–700 °C, 8–9 kbar. LAG3b plagioclase amphibole pairs yield similar albeit slightly lower temperature conditions of 640–680 °C at a fixed pressure of 9 kbar. The magnesi hornblende and albitic rims in AJAH08 equilibrated at lower temperature (480–560 °C).

4.2 Micaschists from the Najac unit

Sample NJC11 and NJC12a were selected for isochemical phase diagram calculations. They were calculated in the system Na₂O-K₂O-FeO-MnO-MgO-TiO₂-Al₂O₃-SiO₂-H₂O using the Holland and Powell (1998) thermodynamic database revised in 2002, with solid solution models from Tajcmanova et al. (2009) for biotite, Newton et al. (1980) for plagioclase, Holland and Powell (1998) for garnet, Coggon and Holland (2002) for white mica, Holland and Powell (1998) for chlorite. We ignored ferric iron because calculated Fe³⁺ in silicates (by charge balance or stoichiometry) is systematically very low and Fe³⁺ bearing oxides are absent. Before drawing the isochemical section using the bulk rock composition measured by XRF spectrometry (CAF service, Stellenbosch University, South Africa), we tested different bulk water content using T-X sections because calculated pyrope content in garnet is

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**Table 1**

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<th>Component</th>
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**Fig. 6.** Isochemical phase diagram built for the bulk composition of the garnet amphibolite in sample AJAH08 of the Laguépie unit: (a) phase diagram also showing the isopleth for XAn in plagioclase and Mg# of amphibole (corresponding to composition determined with the electron microprobe). The small triangles represent the results of garnet-amphibole plagioclase-quartz thermobarometry; (b) Fe#, grossular and spessartine isopleths in garnet corresponding to composition measured by electron microprobe. Also shown is the vapor-saturated solidus for the amphibolite LAG3b.
extremely sensitive to bulk rock H2O content. The best results, i.e. convergence of garnet isopleths and reproduction of the observed mineral assemblages, were obtained with the H2O content calculated using the micas modal proportion and considering a H2O content of \( \sim 5 \) wt\% in white mica.

Garnet in NJC11 preserved a well-developed bell shaped Mn profile. Core composition (spss: 27mol\% and Fe#: 96) are in equilibrium in greenschist facies conditions (360–400 °C, 6–7 kbar) slightly above the garnet-in phase boundary in a field where muscovite, biotite, plagioclase, garnet, quartz and rutile are in equilibrium (Fig. 7a). Intermediate compositions between core and rim draw a linear prograd path passing at 420 °C/8 kbar and 460 °C/12 kbar. Garnet rims equilibrated at 510 ± 20 °C and 14 ± 1 kbar in a field where phengitic mica, biotite, plagioclase, garnet, quartz and rutile are stable with minor amounts of sodic amphibole and clinopyroxene (Fig. 7a). Computed modal proportion for garnet and biotite are low (<6 vol.\% and <10 vol.\%, respectively). Isopleths of Si content in phengite (3.20–3.35 p.f.u.) and of Fe# in garnet (93–94) cross at 510 ± 30 °C and 15 ± 2 kbar.

**Fig. 7.** Isochemical phase diagrams computed for compositions corresponding to garnet micaschists from the Najac unit: (a) sample NJC11 used to determine the prograde P-T path; (b) sample NJC12a used to determine the peak P-T conditions and the retrograde path. Garnet in/out and chlorite in/out curves are indicated in bold to better localize the retrogression in greenschist facies conditions.
The field corresponding to the high-pressure assemblage observed in NJC12a (phengite-quartz-biotite-garnet-plagioclase-rutile) lies within 500–570 °C and 13–16 kbar under vapor under-saturated conditions (Fig. 7b). Free H₂O coexist with the observed assemblage at higher temperatures. At higher pressure conditions, the formation of sodic amphibole and pyroxene appears as a result of plagioclase and biotite breakdown while titanite is stable below 13–14 kbar and 520–560 °C (Fig. 7b). Calculated Fe# in garnet is temperature sensitive. Measured values (90–92) correspond to temperatures ranging between 520 and 585 °C for 10 and 20 kbar of pressure (Fig. 7b). These isopleths cross those representing temperatures ranging between 520 and 585 °C.

Metamict zircons (Corfu et al., 2003) appear as a result of plagioclase and biotite breakdown. At higher temperatures, the prograde path at 385 °C in the Najac eclogite has been recently determined by multi-pressure (Fig. 7b). These isopleths cross those representing temperatures ranging between 520 and 585 °C. The Najac unit encompasses the eclogitic basic-ultrabasic massif, the phengite-bearing orthogneiss and the phengite-garnet micaschists (Fig. 1). Micaschists record early garnet retrogression of these phase or to inaccuracies of the thermodynamic calculations. Calculated grossular content in this P-T domain (0.20–0.28, increasing with decreasing pressure) is slightly above those determined with electron-microprobe (0.18–0.22). The pressure conditions at which measured and calculated grossular content are overlapping (0.20–0.22) are around 17 kbar. Low-Si (2.72–2.78 a.p.f.u.) biotite with Mg# of 45–48 and Ti between 0.10 and 0.11 p.f.u. is in equilibrium between 460–520 °C and 2–7 kbar (Fig. 7b); this is consistent with the presence of chlorite in the retrograde assemblage and the higher modal amounts of biotite calculated in greenschist facies conditions (around 16 vol.%). Biotite stable in the eclogite facies has much higher calculated Si content (>2.90 a.p.f.u.), compositions that were not observed in our samples.

### 5 Zircon U-Pb geochronology

The age of high-pressure, low temperature metamorphism in the Najac eclogite has been recently determined by multimeter geochronology (prograde path at 385–383 Ma, HP peak at 377 ± 3 Ma, exhumation around 369 ± 13 Ma, Sm-Nd and Lu-Hf isochron plus U-Pb zircon and apatite dating by Lotout et al., 2018) but the age of amphibolite facies metamorphism in the Laguépie has not been constrained. A large leucocratic biotite lens (AJAH09, 20 × 30 cm) included in the amphibolite was sampled, from which about 40 zircon grains were extracted using heavy liquids and magnetic separation techniques at the Geosciences Environment Toulouse laboratory.

Zircon images were acquired using a cold-cathode optical cathodoluminescence (CL) at the University of Mons using a Cambridge Image Technology model 8200 Mk5 system. They are generally small (length <200 μm), with minute quartz and apatite inclusions. These zircons show two different textures (Fig. 8a):

- cores of subhedral elongated grains or crystal fragments with dark shade (CL or transmitted light) displaying chaotic or no zoning and being often cracked with fractures extending into the rims. These features are typical of metamict zircons (Corfu et al., 2003);
- luminescent part, often light brown and translucent, riming the dark cores and showing oscillatory growth zoning. These rims are usually 15 to 50 μm large and are relatively rare compared to dark zircon cores. These luminescent rims were also found as isolated fragments.

U-Pb analyses were performed at Geosciences Montpellier (University of Montpellier, France) using a Teledyne G2 excimer laser probe coupled to a ThermoFinnigan Element XR high-resolution ICP-MS (AE-TSE ISO regional facility of the OSU OREME). The probe was set at 25 μm diameter, with an ablation frequency of 4 Hz and a fluency of 6 J/cm². More details on analytical setup can be found in Bosch et al. (2011). Unknowns were calibrated against zircon 91500 standard reference material and GJ-1 was used as secondary standard (both reference materials were analyzed before and after each 5 unknowns). U-Pb ages and common Pb corrections were calculated using the IsoplotR code by Vermeesch (2018). Twelve analyses of GJ-1 zircon yield an upper intercept of 605 ± 14 Ma with 606 ± 7 Ma and 598 ± 5 Ma for the average 207Pb/235U and 206Pb/238U ages, respectively. This is in good agreement with the recommended values, i.e. an upper intercept of 608.5 ± 1.5 Ma calculated from slightly discordant ID-TIMS analyses (Jackson et al., 2004).

Dark metamict cores are U-rich (often >1000 ppm and up to 4650 ppm) and yield strongly discordant U-Pb ellipses (Fig. 8b; Tab. 1). The ablation spectra were moreover irregular and consequently, none of the analyzed zircon cores were used to calculate an age. About 15 spots were shot in the luminescent zoned rims but only 9 were not contaminated by the U-rich metamict cores or by apatite inclusions in depth (evidenced by irregular ablation spectra and/or a continuous increase in U or Pb count rates). These analyses show regular ablation spectra for the 206Pb/238U ratios (i.e. comparable to those of reference zircons 91500 and GJ-1) but more irregular ones for those involving 206Pb. U contents range between 420 and 1044 ppm and Th/U ratios (Tab. 1) are on average slightly higher (0.02–0.04) compared to cores (Th/U: 0.01–0.02 with one analysis at 0.04). The data are discordant (1 to 14% of discordance, 4 ellipses are overlapping Concordia) with 206Pb/235U ages ranging between 356 ± 9 Ma and 374 ± 9 Ma and 206Pb/238U ages between 370 ± 12 Ma and 419 ± 14 Ma. Age calculation using a discordia regression line for rim analyzes yields a lower intercept at 363 ± 3 Ma (MSWD = 2) (Fig. 8c). The weighted average of 206Pb/238U common Pb corrected age is slightly older at 365 ± 3 Ma (Fig. 8d; Tab. 1).

### 6 Discussion

#### 6.1 Two distinct tectono-metamorphic units with contrasted P-T paths

The Najac and Laguépie units are grouped into a single tectono-metamorphic unit by different authors (Burg et al., 1984, 1989; Lardeaux, 2014; Faure et al., 2017; Lotout et al., 2018) but we show hereafter that they display different metamorphic histories and ages.

The Najac unit encompasses the eclogitic basic-ultrabasic massif, the phengite-bearing orthogneiss and the phengite-garnet micaschists (Fig. 1). Micaschists record early garnet...
nucleation around 360–400 °C, 6–7 kbar (Fig. 7a) followed by syn-kinematic growth (Fig. 7a) and equilibration of garnet rims and phengite around 510–570 °C and 14–17 kbar (∼340 °C/GPa) in the eclogite facies (Fig. 7a, b). Late- to post-kinematic biotite and chlorite developed preferentially around garnet (Fig. 7b) and are indicative of retrogression in the greenschist facies at 460–520 °C and 2–7 kbar (Fig. 7b). Lotout et al. (2018) obtained slightly higher peak conditions for an eclogite from the basic-ultrabasic Najac massif (600 °C, 18 kbar; Fig. 9) but similar thermal gradient at peak pressure (330 °C/GPa) (Fig. 9). Multi-methods dating performed by Lotout et al. (2018) on an eclogite from Najac constrained the age of prograde evolution around 383–385 ± 3 Ma and the pressure peak at 377 ± 3 Ma (Fig. 9). The P-T path determined for the Najac unit (Fig. 9) is different than those determined for the paragneisses and micaschists of the Lower Gneiss and Upper Gneiss units in Limousin and French Massif central (Bellot and Roig, 2007; Faure et al., 2009; Lardeaux, 2014; Do Couto et al., 2016), both showing pervasive recrystallization in the amphibolite facies (600–700 °C, 5–10 kbar) during Devonian D1 and D2 events (Faure et al., 2009).
The Laguépie unit comprises garnet-bearing amphibolite, garnet-free amphibolite and a leucotonalitic orthogneiss forming isolated lenses and, sporadically, veins and dykes within amphibolites (Fig. 2). The isolated, unconnected leucotonalitic lenses, the lobed contact between the garnet amphibolite and the leucoamphibolite magma as well as the sodic nature of the felsic material support that it formed after partial melting of the amphibolite. Garnet in the amphibolite sample AJAH08 preserves core to rim zoning attesting for isobaric melting of the amphibolite. Garnet in the amphibolite and the leucotonalite magma as well as the sodic nature of the felsic material support that it formed after partial melting of the amphibolite. Garnet in the amphibolite sample AJAH09 preserves high temperature and low pressure conditions corresponding to the amphibolite and greenschist facies after the partial melting event. U-Pb ages of luminescent zoned zircon rims in a leucotonalitic lens from Laguépie unit (sample AJAH09) yield an age of 363 ± 3 Ma (Fig. 8; Tab. 1) which is interpreted as the timing of felsic magma crystallization following partial melting of the amphibolite.

These conditions are just above the vapor-present solidus calculated for the bulk composition of garnet-bearing and garnet-free amphibolites from Laguépie. Retrogression is evidenced by the development of low-Al magnesiohornblende rims around tschermakite in equilibrium with oligoclase and epidote (TTbl-P around 510°C) heating from 690 to 730°C (Fig. 6). The main plagioclase-garnet-amphibole-quartz-ilmenite-melt assemblage equilibrated around 700−730°C at 10±1 kbar. (~720°C/GPa; ~20°C/km) on a garnet-amphibolite residue. These conditions are just above the vapor-present solidus calculated for the bulk composition of garnet-bearing and garnet-free amphibolites from Laguépie. Retrogression is evidenced by the development of low-Al magnesiohornblende rims around tschermakite in equilibrium with oligoclase and epidote (TTbl-P around 510−580°C in the epidote-amphibolite and/or greenschist facies. The potassic texture of garnet, amphibole and plagioclase and the crystallization of albite, chlorite and actinolite within cataclastic shear zones were acquired in the amphibolite and greenschist facies, respectively. Solid state deformation characterized by a top-to-the-north sense of shear (the D2 event of Faure et al., 2009) affected the Laguépie unit during retrogression in P-T conditions corresponding to the amphibolite and greenschist facies after the partial melting event. U-Pb ages of luminescent and zoned zircon rims in a leucotonalitic lens from Laguépie unit (sample AJAH09) yield an age of 363 ± 3 Ma (Fig. 8) which is interpreted as the timing of felsic magma crystallization following partial melting of the amphibolite.

Our petrological and chronological study demonstrates that these two units have contrasted pressure-temperature-time-deformation histories (Fig. 9) and cannot be assigned to a single tectono-metamorphic unit. The implications of this result are discussed below.

6.2 Correlations with tectono-metamorphic units of the Western European variscan belt

The metamorphic units of the FMC are characterized by contrasted P-T evolutions. The metapelites of the parautochthon have a metamorphic evolution within the greenschist facies locally reaching amphibolite facies near
large Carboniferous plutons (Bellot and Roig, 2007). Both the Lower Gneiss (Lower Allochthon by Santallier et al., 1994) and Upper Gneiss Units (Bellot and Roig, 2007; Lardeaux, 2014; called Middle Allochthon by Santallier et al., 1994) display evidences for pervasive recrystallization in the amphibolite facies, around 600–750°C, 5–10 kbar (see Faure et al., 2009; Do Couto et al., 2016) after a phase of high pressure metamorphism, mostly recorded in the UGU. The uppermost epizonal units (called Upper Allochthon by Santallier et al., 1994) are poorly characterized, they are mostly considered as greenschist facies meta-sedimentary and meta-igneous rocks (Roig et al., 1996; Lardeaux, 2014). The nappe stack in the southern Armorican massif is either described similarly to the FMC by Faure et al. (2005, 2008) or subdivided into various allochthons by Ballèvre et al. (2009, 2014). In both the FMC and AM, a Middle Allochthon has been defined on the basis of its structural position within the nappe stack and the metamorphic conditions recorded by its constitutive rocks. Berger et al. (2010a) described zoisite and kyanite eclogites (~660°C, 29 kbar) exhumed at temperatures below 650°C. They are located close to the contact between LGU and UGU and were included into a Middle Allochthon. Ballèvre et al. (2014) proposed to assign the Groix and Bois de Céné blueschists (<570°C, up to 17 kbar and also devoid of recrystallization in the amphibolite facies, Ballèvre et al., 2003) to a Middle Allochthon but its structural position is uncertain and debated (see Ballèvre et al., 2009 and also Faure et al., 2005). A Middle Allochthon would thus define as a unit comprising cold HP rocks devoid of Late Devonian amphibolite facies recrystallization and located between LGU and UGU.

The Najac unit is often included within the Upper Gneiss Unit on maps published by several authors (Lardeaux, 2014; Faure et al., 2017; Lotout et al., 2018). However, the UGU, containing most of retrogressed eclogite lenses in the FMC (formed during the 420–400 Ma D0 event of Faure et al., 2009), also include migmatitic paragneisses having experienced temperatures up to 700–750°C (Middle Devonian D1 event of Faure et al., 2008, 2009) followed by retrogression in the amphibolite facies (Upper Devonian D2 event; Faure et al., 2009). The P-T path of the Najac unit is unusual in the French Massif central (Fig. 9; see Bellot and Roig, 2007; Faure et al., 2009 and Lardeaux, 2014 for reviews on the metamorphic structure and evolution of the French Massif central). Rocks forming the Najac unit have not recorded the D1 and D2 amphibolite facies metamorphic phases preserved in LGU and UGU and instead preserve high-pressure, low temperature (340°C/GPa at peak pressure) event contemporaneous with a top-to-the-southeast sense of shear. High-pressure rocks included in the Middle Allochthon of the FMC and the AM display metamorphic evolutions similar to the one of the Najac unit, i.e. a HP stage followed by an isothermal exhumation path without recrystallization in the amphibolite facies (Berger et al., 2010a; Ballèvre et al., 2003, 2014). Delor et al. (1986) described glaucophane relics in Najac eclogites and already proposed that they belong to the same tectono-metamorphic unit than the Groix and Bois de Céné units in the Armorican massif. High-pressure rocks from these Middle Allochthon occurrences in Brittany (Guiraud et al., 1987; Ballèvre et al., 2003) are characterized by a cold high pressure metamorphism (<500°C, up to 18 kbar; 7–8°C/km) followed by exhumation at low temperature (<500°C) that compares to Najac micaschists.

Fig. 9. Proposed P-T path (orange arrows) for the Najac and the Laguépie units. Grey P-T paths are from the literature: Groix blueschists from Ballèvre et al. (2003), Cellier éclogites from Ballèvre et al. (2014); Najac éclogite from Lotout et al. (2018); UGU eclogite from Bellot and Roig (2007); UGU paragneiss from Bellot and Roig (2007) and Schulz et al. (2009), Champtoceaux migmatites after Pitra et al. (2010). Age of the HP metamorphism in the Groix Island from Bosse et al. (2005), age of pressure peak in the Najac eclogite from Lotout et al. (2018).
(Fig. 9a). The base of the Lower Allochthon (Lower Gneiss Unit of Faure et al., 2005) unit in the Armorican massif also contains eclogitic lenses and HP orthogneiss (in the Cellier unit) having quite similar peak conditions (600 °C, 2 GPa; 10°C/km; Fig. 9a) (Ballèvre et al., 2014) when compared to the Najac unit. However, they cannot be considered as the lateral equivalent of the Najac unit as the latter clearly thrusts over the Lower Gneiss Unit (Fig. 1).

Published ages for HP metamorphism in the Middle Allochthon are not strictly equivalent in different locations. Lotout et al. (2018) performed multi-method dating (U–Pb zircon and apatite; Sm–Nd and Lu–Hf isochrons) on a retrogressed eclogite from Najac and obtained Upper Devonian ages (383–369 Ma) interpreted as the age of high-pressure metamorphism. Sm–Nd, Rb–Sr and Ar–Ar geochronology by Bosse et al. (2005) also yielded 370–355 Ma for high pressure metamorphism and exhumation of the Groix blueschists. Berger et al. (2010a) proposed that the age of HP metamorphism in Middle Allochthon of Limousin is Lower Devonian but this is based on questionable and sparse ages obtained on thin uraniferous zircon rims surrounding uranium-rich Ordovician cores. Upper Devonian (~380 Ma) ages were also obtained on these zircon rims but they were tentatively attributed to D1 Late-Devonian high temperature metamorphism recorded in Limousin (cf. Faure et al., 2008; Melleton et al., 2009) while the dated eclogite is almost not affected by retrogression. New careful analyses done in the zircon rims from the same sample investigated by Berger et al. (2010a) yielded younger Upper Devonian ages (work in progress).

Considering the metamorphic evolution, the age and the structural position (thrusting over the Lower Gneiss Unit) of the high-pressure Najac unit, we thus propose that it represents the lateral equivalent of the Groix and Bois de Cénét unit in the Armorican massif and the Limousin eclogites of Berger et al. (2010a) and assign this unit to the Middle Allochthon. Further exploration of phengitic, biotite-poor garnet micaschists and associated eclogites is needed in the French Massif central to better characterize a potential Middle Allochthon and to correlate it with well-known occurrences of the Armorican massif.

The Laguèpie unit records amphibolite facies metamorphism peaking at 710–730 °C, 10 kbar, corresponding to metamorphic conditions obtained in the Upper Gneiss Unit elsewhere in the French Massif central (Burg et al., 1984; Bellot and Roig, 2007; Schulz, 2009; Lardeaux, 2014; Do Couto et al., 2016; Fig. 9b). The UGU hosts retrogressed eclogites enclosed into migmatic paragneiss, formed during the D0 and D1 events, respectively (Faure et al., 2009; Melleton et al., 2009; Do Couto et al., 2016). High pressure metamorphism and partial melting were followed by retrogression in the amphibolite facies around 600–700 °C, 5–10 kbar during the Upper Devonian/Lower Carboniferous D2 event (Faure et al., 2009). Association of mafic and felsic rocks recrystallized in the amphibolite facies is common in the Leptyno-Amphibolite Complex (LAC), usually attributed to the UGU (see Ledru et al., 1989, 1994 and Lardeaux, 2014) and more rarely to the LGU (Bellot and Roig, 2007). Mafic rocks from the LAC often display relic assemblages attesting for eclogite and HP granulite facies metamorphism (Santallier et al., 1988) and their igneous precursors are thought to have been emplaced in an ocean-continent transition zone (Lardeaux, 2014). If the Laguèpie unit belongs to the LAC, it must be envisaged that the some of the mafic-felsic associations can also originate after partial melting of a mafic precursor. The age of partial melting and felsic melt crystallization has been constrained at ca. 363 ± 3 Ma in the Laguèpie unit. While Faure et al. (2009) usually ascribe the partial melting event to the 385–375 Ma Devonian D1 phase, U-Pb dating indicates that amphibolite facies (up to 730 °C) partial melting at Laguèpie occurred during Late Devonian. It was followed by top-to-the-north shearing under amphibolite to greenschist facies conditions probably during Lowermost Carboniferous corresponding to the kinematic proposed for the D2 event (Faure et al., 2017).

In the Armorican massif, robust P–T estimates by Pitra et al. (2010) yielded peak temperature above 650 °C at pressure conditions ranging from 8 to 11 kbar for migmatic paragneisses of the eclogite and HP orthogneiss bearing Champtoeaux unit. However the structural position of this units is debated: Faure et al. (2005, 2008) include it into the UGU by correlation with FMC occurrences; Ballèvre et al. (2009) and Pitra et al. (2010) place it at the base of an Upper Allochthon while the upper parts of this allochthon involve a Cadomian basement with low grade metasediments and ophiolites (Faure et al., 2005, 2008; Ballèvre et al., 2009; Ducassou et al., 2011). It is however included within the Lower Allochthon by Ballèvre et al. (2014) by correlation within Iberian massifs. The Laguèpie and Champtoeaux units both yield peak temperature conditions around 700 °C and 10 kbar (Fig. 9). Their metamorphic evolutions are thus more comparable to the Upper Gneiss Unit in FMC, as suggested by Faure et al. (2005, 2008). Considering the conditions and age of amphibolite facies metamorphism (363 Ma) along with the deformation post-dating the anatectic stage, the Laguèpie units correlates with the LAC in the FMC and the Champtoeaux Unit in the AM, that we tentatively ascribe to the Upper Gneiss Unit as defined in the literature (see Lardeaux, 2014; Fig. 9).

6.3 Geodynamic significance

The results presented in this study and those published recently by Lotout et al. (2018, 2020) are not compatible with the subdivision in tectono–metamorphic phases proposed by Faure et al. (2005, 2008, 2009). In these latter papers, the D0 phase correspond to the formation of HP units between 420 and 400 Ma during the closure of an oceanic domain separating Gondwana and Armorica. However, Lotout et al. (2018) provided Upper Devonian ages for the Najac eclogites and Lotout et al. (2020) obtained Upper Devonian to Early Carboniferous ages for HP metamorphism in Lévezou. These data suggest that HP metamorphism is younger in southern French Massif central compared to other occurrences (Pin and Lancelot, 1982; Ducrot et al., 1983; Paquette et al., 1995). Furthermore, partial melting in UGU as recorded in the Laguèpie unit is usually ascribed to the 385–375 Ma D1 event by Faure et al. (2009) but it is dated at 363 Ma at Laguèpie. It is however affected by the Late Devonian/Lower Carboniferous D2 top-to-the-north shearing (Faure et al., 2017). These discrepancies suggest that either (i) new and old ages for HP
metamorphism should be critically evaluated because dating HP metamorphism is a difficult and challenging task or (ii) metamorphic events are about 20 Myr younger in southern French Massif central compared to northern parts or (iii) that final collision between Armorica and Gondwana occurred during Upper Devonian/Early Carboniferous and not Lower Devonian, as suggested by Lardeaux (2014) and Ballèvre et al. (2009, 2014).

If we consider that Upper Devonian HP/LT metamorphism (Lotout et al., 2018, 2020) is coeval with partial melting in UGU (Mellet et al., 2009; this study) and active margin magmatism as represented by tonalites and diorites intruding the UGU (375–355 Ma; Bernard-Griffiths et al., 1985; Shaw et al., 1993; Bertrand et al., 2001; Pin and Paquette, 2002); these data are compatible with an active subduction system forming a paired metamorphic belt during Late Devonian (Miyashiro, 1967; Brown, 2006). The Najac unit, comprising a basic/ultrabasic intrusion within orthogneisses and micaschists all metamorphosed under eclogite facies conditions, either represents fragments of a continental lower plate or the accretionary prism in the subduction system (proposed for the Groix blueschists, Ballèvre et al., 2014). The UGU, and accordingly, the Laguépie unit, are, in this proposition, part of the upper plate of the subduction system, similarly to what has been proposed by Lardeaux (2014) and Lardeaux et al. (2014). Nappe stacking that caused superimposition of the Najac and Laguépie units occurred after 363 Ma, probably during the Early Carboniferous 345–335 Ma D3 phase of Faure et al. (2009) because the Najac micaschists were not affected by the Late Devonian 700 °C/10 kbar metamorphism observed at Laguépie.

7 Conclusion

Petrological investigations on metamorphic samples from the Najac and Laguépie units in the southwestern FMC reveal two distinct P-T evolutions. The Najac unit is characterized by syn-kinematic eclogitic metamorphism in micaschists and basic rocks implying subduction of a continental margin at the end of the Devonian (around 377 Ma) followed by cold exhumation in the greenschist facies. The Laguépie unit records amphibolite facies metamorphism (710 °C, 10 kbar) dated at 363 Ma and later affected by deformation and retrogression under amphibolite to greenschist facies conditions. These two units were formerly grouped into the leptyno-amphibolite complex belonging to the Upper Gneiss Unit. Comparison with tectono-metamorphic units in the French Massif central and the Armorican massifs suggest that the Najac unit can be integrated into the HP/LT Middle Allochthon while the Laguépie unit belongs to the Upper Gneiss Unit. The French Massif central consequently preserves the remnants of an Upper Devonian paired metamorphic belt, where the Middle Allochthon represent subducted material attached to the slab or within an accretionary prism while the Upper Gneiss Unit was located within the upper plate of the subduction system. Lower Carboniferous continental collision further led to the stacking of metamorphic units with contrasted pre-Carboniferous histories. The stack, in its current configuration, is showing an apparent inverted metamorphic gradient, well expressed in the Rouergue domain.

The structure of the tectono-metamorphic units near Najac can be considered as a reference for further studies on the structure and evolution of metamorphic units preserved in the French Massif central. Future investigations should focus on obtaining precise P-T paths on pre-Carboniferous metamorphic units belonging to the FMC and exploration of phengite-bearing biotite-poor garnet micaschists (locally enclosing eclogites) possibly marking a still poorly defined Middle Allochthon. Altogether, these results should be used to propose an integrated tectono-metamorphic framework that can be useful to test the different geodynamic scenarios proposed for the Variscan orogeny in Western Europe.

Supplementary material

Table S1. Tables of mineral chemistry for samples of Najac micaschists and Laguépie amphibolites.

The Supplementary material is available at https://www.bsgf.fr/10.1051/bsgf/2020033/olm.

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