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The MALINA oceanographic expedition: how do changes in ice cover, permafrost and UV radiation impact biodiversity and biogeochemical fluxes in the Arctic Ocean?

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Abstract. The MALINA oceanographic campaign was conducted during summer 2009 to investigate the carbon stocks and the processes controlling the carbon fluxes in the Mackenzie River estuary and the Beaufort Sea. During the campaign, an extensive suite of physical, chemical and biological variables were measured across seven shelf–basin transects (south–north) to capture the meridional gradient between the estuary and the open ocean. Key variables such as temperature, absolute salinity, radiance, irradiance, nutrient concentrations, chlorophyll *a* concentration, bacteria, phytoplankton and zooplankton abundance and taxonomy, and carbon stocks and fluxes were routinely measured onboard the Canadian research icebreaker CCGS *Amundsen* and from a barge in shallow coastal areas or for sampling within broken ice fields. Here, we present the results of a joint effort to compile and standardize the collected data sets that will facilitate their reuse in further studies of the changing Arctic Ocean. The data set is available at https://doi.org/10.17882/75345 (Massicotte et al., 2020).

1 Introduction

The Mackenzie River is the largest source of terrestrial particles entering the Arctic Ocean (see Doxaran et al., 2015, and references therein). During the past decades, temperature rise, permafrost thawing, coastal erosion and increasing river runoff have contributed to intensifying the export of terrestrial carbon by the Mackenzie River to the Arctic Ocean (e.g., Tank et al., 2016). Furthermore, the environmental changes currently happening in the Arctic may have profound impacts on the biogeochemical cycling of this exported carbon. On one hand, reduction in sea ice extent and thickness exposes a larger fraction of the ocean surface to higher solar radiation and increases the mineralization of this carbon into atmospheric CO_2 through photo-degradation (Miller and Zepp, 1995; Bélanger et al., 2006b). On the other hand, the possible increase in nutrients brought by Arctic rivers may contribute to higher autotrophic production and sequestration of organic carbon (Tremblay et al., 2014).

Given that these production and removal processes are operating simultaneously, the fate of Arctic river carbon transiting toward the Arctic Ocean is not entirely clear. Hence, detailed studies about these processes are needed to determine if the Arctic Ocean will become a biological source or a sink of atmospheric CO_2. With regard to this question, the MALINA oceanographic expedition was designed to document and gain insight into the stocks and the processes controlling carbon fluxes in the Mackenzie River and the Beaufort Sea. Specifically, the main objective of the MALINA oceanographic expedition was to determine how (1) primary production, (2) bacterial activity and (3) organic-matter photo-oxidation influence carbon fluxes and cycling in the Canadian Beaufort Sea. In this article, we present an overview of an extensive and comprehensive data set acquired from a coordinated international sampling effort conducted in the Mackenzie River and in the Beaufort Sea in August 2009.

2 Study area, environmental conditions and sampling strategy

2.1 Study area and environmental conditions

The MALINA oceanographic expedition was conducted between 30 July and 25 August 2009 in the Mackenzie River and the Beaufort Sea systems (Fig. 1). Figure 2 shows an overview of the sea ice conditions that prevailed during the expedition. In Fig. 2a, a true-color image from MODIS Terra reveals how the sea ice pack was fragmented toward the end of the expedition, specially near the 200 m isobath (identified by the continuous red line). On the shelf, the sea ice concentration was higher at the beginning of the expedition. During the 4-week cruise, the ice concentration gradually decreased toward the north (Fig. 2b).

The Mackenzie River basin is the largest in northern Canada and covers an area of approximately 1 805 000 km^2, which represents around 20 % of the total land area of

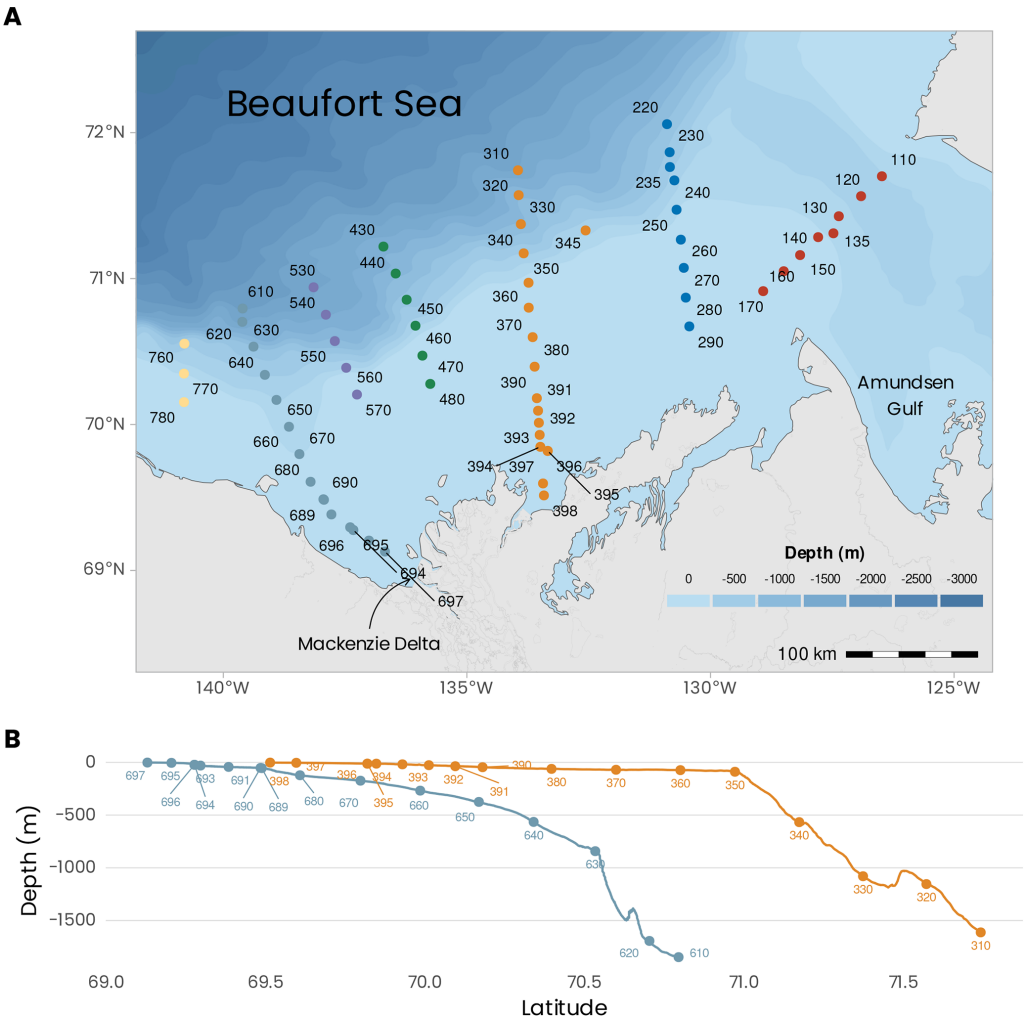

Figure 1. **(a)** Localizations of the sampling sites visited during the MALINA 2009 campaign. The colors of the dots represent the seven transects visited during the mission. **(b)** Bathymetric profiles for transects 600 and 300. Bathymetric data from General Bathymetric Chart of the Oceans (GEBCO) (https://download.gebco.net/).

Canada (Abdul Aziz and Burn, 2006). Between 1972 and 2016, the average monthly discharge (recorded at the Arctic Red River station) varied between 3296 and 23 241 $m^3 s^{-1}$ (shaded area in Fig. 3a). The period of maximum discharge usually occurs at the end of May, with decreasing discharge until December, whereas the period of low and stable discharge extends between December and May. During the MALINA oceanographic cruise, the daily discharge varied between 12 600 and 15 100 $m^3 s^{-1}$ (red segment in Fig. 3a; see also Ehn et al., 2019). Draining a vast watershed, the Mackenzie River annually delivers on average 2100 and 1400 Gg C yr^{-1} of particulate organic carbon (POC) and dissolved organic carbon (DOC), respectively, into the Arctic Ocean (Stein and Macdonald, 2004; Raymond et al., 2007). During the expedition conducted onboard the CCGS *Amundsen*, the air temperature recorded by the foredeck meteorological tower varied between −2 and 11 °C (Fig. 3b). The average air temperature was 3 °C and usually remained above 0 °C.

2.2 General sampling strategy

The sampling was conducted over a network of sampling stations organized into seven transects identified with three digits: 100, 200, 300, 400, 500, 600 and 700 (Fig. 1a). Stations were sampled across these seven shelf–basin transects (south–north) to capture the meridional gradient between the estuary and the open ocean (except for transect 100 across the mouth of the Amundsen Gulf). Within each transect, station numbers were listed in descending order from south to north. Because our goal was to sample in open waters, the order in which the transects were visited depended on the ice cover. On 20 July 2009, just before the mission, a relatively large portion of the shelf was still covered by sea ice (Fig. 2b). Soon after the beginning of the cruise, most of the shelf area

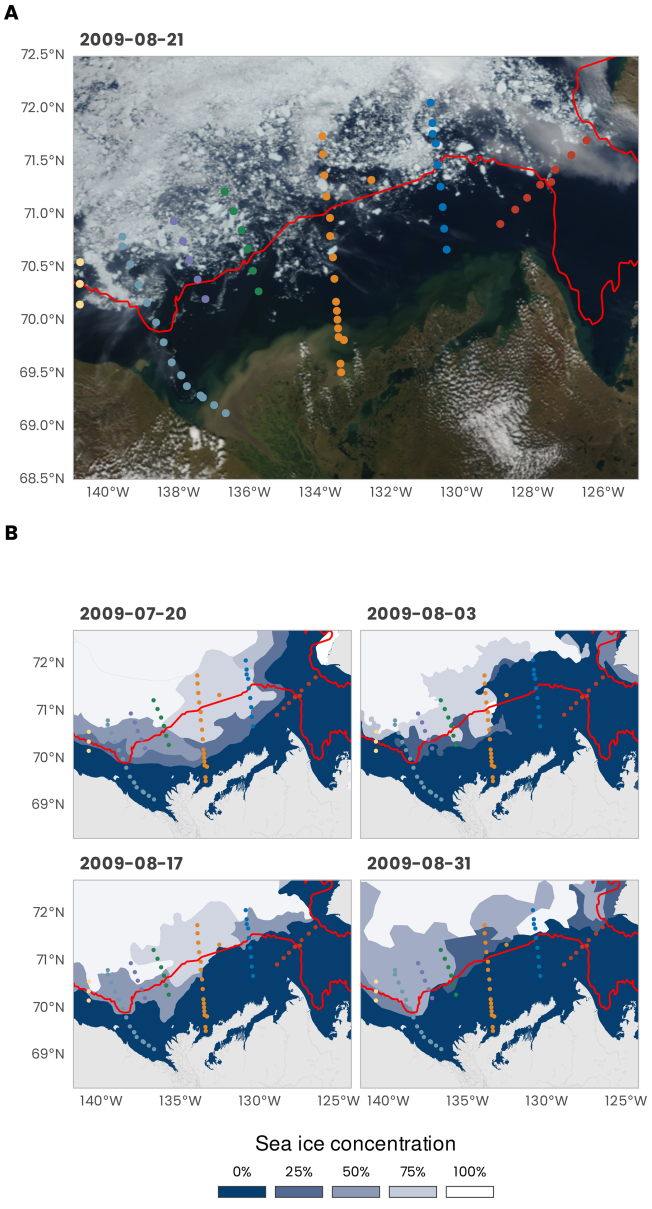

Figure 2. **(a)** True-color image from MODIS Terra (data from https://wvs.earthdata.nasa.gov, last access: 1 February 2021). **(b)** Weekly sea ice concentration from the US National Ice Center (U.S. National Ice Center, 2020). The red line shows the 200 m isobath (data from http://www.naturalearthdata.com, last access: 1 February 2021). The dots represent the stations (see Fig. 1 for the legend).

was ice-free. The shelf region was not ice-free before mid-August. The bathymetry at the sampling stations varied between 2 and 1847 m (394 ± 512 m, mean ± standard deviation). The stations located in the Beaufort Sea were sampled onboard the Canadian research icebreaker CCGS *Amundsen*. Biological, chemical and optical water column sampling was almost always restricted to the first 400 m of the water column during daytime. Deeper profiles for sampling the whole water column and bottom sediment were usually repeated during nighttime at the same stations. Sediment sampling for fauna and biogeochemistry was conducted at eight stations (110, 140, 235, 260, 345, 390, 680, 690). Two transects (600 and 300) were extended to very shallow waters on the shelf and sampled from either a zodiac or a barge (the bathymetry profiles are shown in Fig. 1b). In the context of this data paper, these two transects were chosen to present an overview of the principal variables measured during the MALINA campaign. A summary of the various sampling strategies is presented below.

2.3 CTD and rosette deployment

Onboard the CCGS *Amundsen*, a general oceanic rosette equipped with a CTD (instrument for measuring conductivity, temperature, and depth; Sea-Bird SBE-911+) was deployed at each sampling station (Fig. 1). The rosette was equipped with twenty-four 12 L Niskin bottles. The rosette was also equipped with a transmissometer sensor (Wetlabs), a photosynthetic active radiation (PAR) sensor (Biospherical), an oxygen sensor (SBE-43), a pH sensor (SBE-18), a nitrate sensor (Satlantic ISUS), a fluorometer (Seapoint) and an altimeter (Benthos). A surface PAR (Biospherical) was also installed on the roof of the rosette control laboratory. A UVP5 (underwater vision profiler, Hydroptics) was also mounted on the rosette frame, providing size and abundance of particles above 200 µm and plankton above 700 µm. The rosette data processing and quality control are described in detail in Guillot and Gratton (2010). Data processing included the following steps: validation of the calibration coefficients, conversion of data to physical units, alignment correction and extraction of useless data. Oxygen sensor calibration was done using Winkler titrations, and salinity data were compared with water samples analyzed with a Guildline 8400B Autosal. The quality control tests were based on the International Oceanographic Commission's suggested procedures and UNESCO's algorithm standards (Commission of the European Community, 1993). The recorded data were averaged every decibar. On 5 August, the pH sensor was replaced by a chromophoric dissolved organic matter (CDOM) fluorometer (excitation: 350–460 nm; emission: 550 nm; half-width (HW) 40 nm; Dr. Haardt Optik Mikroelektronik). The rosette depth range was restricted to the first 1000 m when carrying the pH, PAR and nitrate sensors because of their rating.

2.4 Sediment sampling

Surface sediments were sampled using an USNEL box corer (50 × 50 × 40 cm). Box cores with undisturbed surfaces were subsampled for (a) lipids (Rontani et al., 2012b) and isotopic signature of lipid biomarkers (Tolosa et al., 2013), stable isotopes (C, N), and manganese and iron oxides (Link et al., 2013a) in the 1 cm surface layer; (b) sediment pigment profiles down to 8 cm; and (c) fluxes at the sediment–water

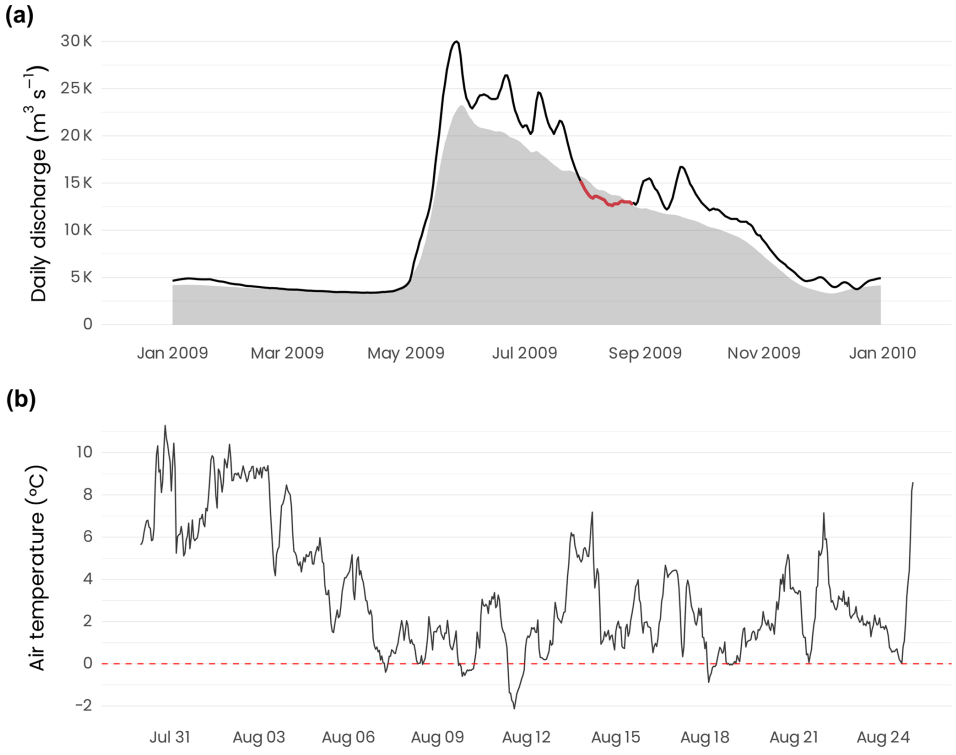

Figure 3. **(a)** Daily discharge of the Mackenzie River at the Arctic Red River junction (station 10LC014). The black line corresponds to the 2009 discharge, whereas the colored segment identifies the period of the MALINA campaign. The shaded area is the mean discharge calculated between 1972 and 2016. Discharge data from the government of Canada (https://wateroffice.ec.gc.ca/search/historical_e.html, last access: 6 September 2019). **(b)** Hourly air temperature recorded from the foredeck meteorological tower of the *Amundsen* during the campaign.

interface using onboard microcosm incubations on subcores (10 cm diameter, 20 cm deep) (Link et al., 2013a, 2019). At three stations (140, 345, 390), macrofauna abundance and diversity were determined from sieved and conserved samples (Link et al., 2019). At station 680, a core was subsampled at 1 cm to determine dinocyst abundance and dated using ^{210}Pb and ^{137}Cs (Durantou et al., 2012). Samples for (a)–(b) were stored frozen until analysis in the respective home labs.

3 Data quality control and data processing

Different quality control procedures were adopted to ensure the integrity of the data. First, the raw data were visually screened to eliminate errors originating from the measurement devices, including sensors (systematic or random) and errors inherent from measurement procedures and methods. Statistical summaries such as average, standard deviation and range were computed to detect and remove anomalous values in the data. Then, data were checked for duplicates and remaining outliers. The complete list of variables is presented in Table 1.

4 Data description: an overview

The following sections present an overview of a subset of selected variables from the water column. For these selected variables, a brief description of the data collection methods is presented along with general results.

4.1 Water mass distribution

According to previous studies (Carmack et al., 1989; Macdonald et al., 1989), five main source water types can be distinguished in the southeastern Beaufort Sea: (1) meteoric water (MW; Mackenzie River plus precipitation), (2) sea ice meltwater (SIM), (3) winter polar mixed layer (wPML), (4) upper-halocline water (UHW; modified Pacific water with core salinity of 33.1 PSU) and (5) lower-halocline water (LHW; water of Atlantic origin). In this study, we used the optimum multiparameter (OMP) algorithm to quantify the relative contributions of the different source water types to the observed data (https://omp.geomar.de/, last access: 1 February 2021). We used salinity, TA and δ^{18}O as conservative tracers as well as temperature and O_2 concentration as non-conservative tracers, to constrain the water mass analysis following Lansard et al. (2012). Briefly, the method finds

Table 1. Parameters measured during the MALINA oceanographic expedition. Parameters are ordered alphabetically.

Parameters	Sampling	Principal	Included in the investigators	Reference data repository	
^{137}Cs dating of core samples	Gamma spectrometer	Calypso Square (CASQ) corer	Rochon A./Schmidt	No	1
^{137}Cs dating of core samples	Gamma spectrometry	Box corer	Rochon A., Schmidt	No	1
^{14}C dating of core samples	Accelerator mass spectrometry	Box corer	Rochon A.	No	1
^{14}C dating of core samples	Accelerator mass spectrometry	CASQ corer	Rochon A.	No	1
^{15}N ammonium assimilation	^{15}N spiking – incubation – mass spectrometry	Rosette – deck incubations	Tremblay J. E., Raimbault P.	Yes	2, 3, 4
^{15}N ammonium assimilation	^{15}N spiking – incubation – mass spectrometry	Rosette in situ production line	Tremblay J. E., Raimbault P.	Yes	2, 3, 4
^{15}N ammonium oxidation (nitrification)	^{15}N spiking – incubation – mass spectrometry	Rosette – deck incubations	Tremblay J. E., Raimbault P.	Yes	2, 3, 4
^{15}N ammonium oxidation (nitrification)	^{15}N spiking – incubation – mass spectrometry	Rosette in situ production line	Tremblay J. E., Raimbault P.	Yes	2, 3, 4
^{15}N ammonium primary production (^{13}C)	^{15}N spiking – incubation – mass spectrometry	Rosette – deck incubations	Tremblay J. E., Raimbault P.	Yes	2, 3, 4
^{15}N ammonium regeneration	^{15}N spiking – incubation – mass spectrometry	Rosette – deck incubations	Tremblay J. E., Raimbault P.	Yes	2, 3, 4
^{15}N ammonium regeneration	^{15}N spiking – incubation – mass spectrometry	Rosette in situ production line	Tremblay J. E., Raimbault P.	Yes	2, 3, 4
^{15}N N_2 fixation	^{15}N spiking – incubation – mass spectrometry	Rosette water sample	Tremblay J. E., Raimbault P.	No	2, 3, 4
^{15}N nitrate assimilation	^{15}N spiking – incubation – mass spectrometry	Rosette – deck incubations	Tremblay J. E., Raimbault P.	Yes	2, 3, 4
^{15}N nitrate assimilation	^{15}N spiking – incubation – mass spectrometry	Rosette in situ production line	Tremblay J. E., Raimbault P.	Yes	2, 3, 4
^{210}Pb geochronology of core samples	^{209}Po alpha spectrometry	Box corer	Rochon A.	No	1
^{210}Pb geochronology of core samples	^{209}Po alpha spectrometry	CASQ corer	Rochon A.	No	1
^{234}Th (1 µm < particles > 70 µm)	Beta-counting	Foredeck in situ pump	Gasser B.	Yes	
^{234}Th (particles > 70 µm)	Beta-counting	Foredeck in situ pump	Gasser B.	Yes	
^{234}Th (Particulate)	Beta-counting	Drifting sediment trap	Gasser B.	Yes	5
AAPB (abundance)	IR microscopy, fluorimetry; FISH	Rosette water sample	Jeanthon C., Boeuf D.	Yes	6
AAPB (abundance)	IR microscopy, fluorimetry; FISH	Zodiac water sample	Jeanthon C., Boeuf D.	Yes	6
Absorption (particulate)	Point-source integrating-cavity absorption meter (PSICAM)	Barge water sample	Leymarie E.	Yes	
Absorption (particulate)	PSICAM	Rosette water sample	Leymarie E.	Yes	
Absorption (particulate)	Spectrophotometer (filters)	Barge water sample	Belanger S.	Yes	7, 8
Absorption (particulate)	Spectrophotometer (filters)	Continuous on way	Belanger S.	Yes	7, 8
Absorption (particulate)	Spectrophotometer (filters)	Rosette water sample	Belanger S.	Yes	7, 8
Absorption (particulate)	Spectrophotometer (filters)	Zodiac profiler	Belanger S.	Yes	7, 8
Absorption (total)	PSICAM	Barge water sample	Leymarie E.	Yes	
Absorption (total)	PSICAM	Rosette water sample	Leymarie E.	Yes	
Absorption coefficient (total) (nine wavelengths)	Wetlabs AC9 serial no. 156	Rosette profiler	Ehn J.	Yes	
Absorption coefficient (total) (nine wavelengths in IR)	Wetlabs AC9 serial no. 303	Barge profiler	Doxaran D.	Yes	9
Absorption coefficient (total) (nine wavelengths)	Wetlabs AC9 serial no. 279	Barge profiler	Doxaran D.	Yes	10, 11
Air relative humidity	Humidity sensor	Foredeck meteorological tower	Papakyriakou T.	Yes	
Alkalinity total (TA)	Potentiometry	Barge water sample	Mucci A., Lansard B.	Yes	12, 13, 14
Alkalinity total (TA)	Potentiometry	Rosette	Mucci A., Lansard B.	Yes	12, 13, 14
Alkalinity total (TA)	Potentiometry	Zodiac water sample	Mucci A., Lansard B.	Yes	12, 13, 14
Alkanes	Gas chromatography–mass spectrometry (GC–MS)	Box corer	Bouloubassi I.	Yes	
Alkanes	GC–MS	CASQ corer	Bouloubassi I.	Yes	
Ammonium (NH_4^+) photoproduction apparent quantum yield (AQY)	Sun simulator – fluorimetry	Rosette water sample	Xie H., Tremblay J. E.	Yes	15
Ammonium (NH_4^+) photoproduction apparent quantum yield (AQY)	Sun simulator – fluorimetry	Zodiac water sample	Xie H., Tremblay J. E.	Yes	15
Aragonite: saturation state	Derived parameter	Barge water sample	Mucci A., Lansard B.	Yes	16
Aragonite: saturation state	Derived parameter	Rosette water sample	Mucci A., Lansard B.	Yes	16
Aragonite: saturation state	Derived parameter	Zodiac water sample	Mucci A., Lansard B.	Yes	16
Attenuation coefficient (total) (nine wavelengths in IR)	Wetlabs AC9 serial no. 0303	Barge profiler	Doxaran D.	Yes	9
Attenuation coefficient (total) (nine wavelengths)	Wetlabs AC9 serial no. 156	Rosette profiler	Ehn J.	Yes	
Attenuation coefficient (total) (nine wavelengths)	Wetlabs AC9 serial no. 279	Barge profiler	Doxaran D.	Yes	10, 11
Attenuation coefficient at 660 nm	Wetlabs (CRover) transmissometer	Drifting profiling float	Doxaran D.	Yes	17, 18
Backscattering 532 nm	Wetlabs (ECO^3) backscatterometer	Drifting profiling float	Doxaran D.	Yes	17, 18
Backscattering coefficient (three wavelengths in IR)	Wetlabs ECO-BB3 serial no. 538	Barge profiler	Doxaran D.	Yes	19, 20
Backscattering coefficient (three wavelengths)	Wetlabs ECO-BB3 serial no. 028	Barge profiler	Doxaran D.	Yes	19, 20
Backscattering coefficient (eight wavelengths, spectral)	HydroScat-6 (serial no. 97074) and two a-Beta (HOBI Labs)	Barge profiler	Reynolds R.	Yes	21, 22
Backscattering coefficient (eight wavelengths, spectral)	HydroScat-6 (serial no. 97074) and two a-Beta (HOBI Labs)	Foredeck	Reynolds R.	Yes	21, 22
Backscattering coefficient (eight wavelengths)	Wetlabs ECO-BB9 serial no. 274	Rosette profiler	Ehn J.	Yes	
Bacteria (abundance)	Flow cytometry	Rosette water sample	Vaulot D.	Yes	23
Bacterial bio-volume	Epifluorescence microscopy	Rosette water sample	Joux F., Ortega-Retuerta E.	No	24, 25

Table 1. Continued.

Parameters	Method	Sampling	Principal investigators	Included in the data repository	Reference
Bacterial diversity	CE-SSCP and 454 tag-pyrosequencing	Rosette and zodiac water sample	Joux F., J.F. Ghiglione	No	26, 27
Bacterial growth (limitation by nutrients)	Leucine-^{3}H incubations – cells counts	Rosette water sample	Joux F., Jeffrey W., Ortega-Retuerta E.	No	25, 28
Bacterial production	Leucine-^{3}H incorporation	Rosette water sample	Joux F., Jeffrey W.	Yes	25, 29, 30
Bacterial production	Leucine-^{3}H incorporation	Zodiac water sample	Joux F., Jeffrey W.	Yes	25, 29, 30
Bacterial respiration (whole community)	O_2 consumption – Winkler – Incubations	Rosette water sample	Joux F., Ortega-Retuerta E.	Yes	25, 31
Benthic ammonium flux	Incubations – colorimetry	Box corer	Link H., Archambault P., Chaillou G.	Yes	32, 33
Benthic macrofauna abundance	Microscopy	Box corer	Link H., Archambault P., Chaillou G.	No	33, 34, 35
Benthic macrofauna diversity	Microscopy	Box corer	Link H., Archambault P., Chaillou G.	No	33, 34, 35
Benthic nitrate flux	Incubations – colorimetry, autoanalyzer	Box corer	Link H., Archambault P., Chaillou G.	Yes	32, 33, 34
Benthic nitrite flux	Incubations – colorimetry, autoanalyzer	Box corer	Link H., Archambault P., Chaillou G.	Yes	32, 33, 34
Benthic phosphate flux	Incubations – colorimetry, autoanalyzer	Box corer	Link H., Archambault P., Chaillou G.	Yes	32, 33, 34
Benthic respiration	Incubations – optic – oxygen probe	Box corer	Link H., Archambault P., Chaillou G.	Yes	32, 33, 34
Benthic silicic acid flux	Incubations – colorimetry, autoanalyzer	Box corer	Link H., Archambault P., Chaillou G.	Yes	32, 33, 34
Calcite: saturation state	Derived parameter	Rosette water sample	Mucci A., Lansard B.	Yes	16
Calcite: saturation state	Derived parameter	Barge water sample	Mucci A., Lansard B.	Yes	16
Calcite: saturation state	Derived parameter	Zodiac water sample	Mucci A., Lansard B.	Yes	16
Campesterol, cholesterol, sitosterol and products of degradation	GC–MS	Rosette water sample	Sempere R.	Yes	36, 37, 38
Chromophoric dissolved organic matter (CDOM) absorption	PSICAM	Barge water sample	Leymarie E.	Yes	
CDOM absorption	PSICAM	Rosette water sample	Leymarie E.	Yes	
CDOM absorption	PSICAM	Zodiac water sample	Leymarie E.	Yes	
CDOM absorption	Spectrophotometer	Barge water sample	Matsuoka A., Bricaud A.	No	41, 42
CDOM absorption	Spectrophotometer	Rosette water sample	Matsuoka A., Bricaud A.	No	41, 42
CDOM absorption	Spectrophotometer	Zodiac water sample	Matsuoka A., Bricaud A.	No	41, 42
CDOM absorption	Ultrapath	Barge water sample	Bricaud A.	Yes	39, 40
CDOM absorption	Ultrapath	Rosette water sample	Bricaud A.	Yes	39, 40
CDOM absorption	Ultrapath	Zodiac water sample	Bricaud A.	Yes	39, 40
CDOM fluorescence	Haardt fluorometer	Rosette profiler	Benner R., Belanger S., Amon R., Sempere R.	Yes	43, 44
CDOM fluorescence	Wetlabs (ECO^3) fluorometer	Drifting profiling float	Doxaran D.	Yes	18
CDOM fluorescence	Wetlabs WetStar WSCD	Barge profiler	Doxaran D.	Yes	18
CDOM fluorescence EEM (excitation emission matrix)	Spectrofluorometry	Rosette water sample	Sempere R.	No	45
CDOM fluorescence EEM (excitation emission matrix)	Spectrofluorometry	Zodiac water sample	Sempere R.	No	45
Chlorophyll *a* and pheopigment (concentration)	Fluorimetry size fractionated	Rosette water sample	Gosselin M., Belanger S.	Yes	46
Chlorophyll *a* and pheopigment (benthic)	Fluorometric analysis	Box corer	Link H., Archambault P., Chaillou G.	Yes	32, 33, 34
Chlorophyll *a* fluorescence [Fchla (z)]	Chelsea Mini-Track a II fluorometer	Barge profiler	Doxaran D.	Yes	18
Chlorophyll *a* fluorescence [Fchla (z)]	Seapoint fluorometer	Rosette profiler	Gratton Y., Prieur L., Tremblay J. E.	Yes	
Chlorophyll *a* fluorescence [Fchla (z)]	Wetlabs (ECO^3) fluorometer	Drifting profiling float	Doxaran D.	Yes	18
CO photoproduction apparent quantum yield for CDOM	Sun simulator – reduction gas analyzer	Rosette water sample	Xie H.	Yes	47
CO photoproduction apparent quantum yield for CDOM	Sun simulator – reduction gas analyzer	Zodiac water sample	Xie H.	Yes	47
CO photoproduction apparent quantum yield for particulate matter	Sun simulator – reduction gas analyzer	Rosette water sample	Xie H.	Yes	47
CO photoproduction apparent quantum yield for particulate matter	Sun simulator – reduction gas analyzer	Zodiac water sample	Xie H.	Yes	47
CO_2 (atm) concentration	Infrared	Foredeck meteorological tower	Papakyriakou T.	Yes	
CO_3^{2-} concentration	Derived parameter	Barge water sample	Mucci A., Lansard B.	Yes	16
CO_3^{2-} concentration	Derived parameter	Rosette water sample	Mucci A., Lansard B.	Yes	16
CO_3^{2-} concentration	Derived parameter	Zodiac water sample	Mucci A., Lansard B.	Yes	16
Coccolithophorids	Microscopy	Rosette water sample	Coupel P.	Yes	
Conductivity (z)	Sensor on SBE FastCAT CTD serial no. 175–217	Barge profiler	Doxaran D.	Yes	48
Conductivity (z)	Sensor Sea-Bird 4c on CTD SBE-911	Rosette profiler	Gratton Y., Prieur L.	Yes	
CTD	Sea-Bird	Drifting profiling float	Doxaran D.	Yes	48
Cultures of sorted populations	Sorted by flow cytometry, serial dilution and single-cell pipetting	Rosette water sample	Vaulot D.	No	49
δ^{13}C – DIC	Mass spectrometry (IRMS)	Barge water sample	Mucci A., Lansard B.	Yes	
δ^{13}C – DIC	Mass spectrometry (IRMS)	Rosette water sample	Mucci A., Lansard B.	Yes	
δ^{13}C – DIC	Mass spectrometry (IRMS)	Zodiac water sample	Mucci A., Lansard B.	Yes	
δ^{13}C on suspended particulate matter	Mass spectrometry	Rosette water sample	Tremblay J. E., Raimbault P.	No	50
δ^{18}O – water	Mass spectrometry (IRMS)	Barge water sample	Mucci A., Lansard B.	Yes	13

Table 1. Continued.

Parameters	Method	Sampling	Principal investigators	Included in the data repository	Reference
$\delta^{18}O$ – water	Mass spectrometry (isotope ratio mass spectrometry, IRMS)	Rosette water sample	Mucci A., Lansard B.	Yes	13
$\delta^{18}O$ – water	Mass spectrometry (IRMS)	Zodiac water sample	Mucci A., Lansard B.	Yes	13
Diacids composition	GC–MS	Rosette water sample	Sempere R.	No	51
Diacids composition	GC–MS	Zodiac water sample	Sempere R.	No	51
Diacids photoproduction apparent quantum yield (AQY)	Sun simulator – GC–MS	Zodiac water sample	Sempere R.	No	52
Dinocyst, pollen and spore abundance	Microscopy	Box corer	Rochon A.	No	1
Dinocyst, pollen and spore abundance	Microscopy	CASQ corer	Rochon A.	No	1
Dinocyst, pollen and spore identification	Microscopy	Box corer	Rochon A.	No	1
Dinocyst, pollen and spore identification	Microscopy	CASQ corer	Rochon A.	No	1
Dinoflagellate cyst abundance	Microscopy	Box corer	Rochon A.	No	1
Dinoflagellate cyst abundance	Microscopy	CASQ corer	Rochon A.	No	1
Dinoflagellate cyst identification	Microscopy	Box corer	Rochon A.	No	1
Dinoflagellate cyst identification	Microscopy	CASQ corer	Rochon A.	No	1
Dissolved-inorganic-carbon photoproduction apparent quantum yield	Sun simulator – infrared CO_2 analyzer	Rosette water sample	Xie H., Belanger S.	Yes	53
Dissolved-inorganic-carbon photoproduction apparent quantum yield	Sun simulator – infrared CO_2 analyzer	Zodiac water sample	Xie H., Belanger S.	Yes	53
Dissolved nitrogen (total) (TDN)	High-temperature catalytic oxidation	Rosette water sample	Benner R.	Yes	54
Dissolved nitrogen (total) (TDN)	High-temperature catalytic oxidation	Zodiac water sample	Benner R.	Yes	54
Dissolved organic carbon (DOC)	High-temperature catalytic oxidation	Rosette water sample	Benner R.	Yes	54
Dissolved organic carbon (DOC)	High-temperature catalytic oxidation	Rosette water sample	Sempere R.	Yes	45, 55
Dissolved organic carbon (DOC)	High-temperature catalytic oxidation	Zodiac water sample	Benner R.	Yes	54
Dissolved organic carbon (DOC)	High-temperature Catalytic Oxidation	Zodiac water sample	Sempere R.	Yes	45, 55
Dissolved organic carbon (DOC)	Wet oxidation	Rosette water sample	Tremblay J. E., Raimbault P.	Yes	56
Dissolved organic nitrogen (DON)	Wet oxidation	Rosette water sample	Tremblay J. E., Raimbault P.	Yes	56
Dissolved organic phosphorus (DOP)	Wet oxidation	Rosette water sample	Tremblay J. E., Raimbault P.	Yes	56
Ed, Lu, Eu, Es	Compact optical profiling system (C-OPS) package (320, 340, 380, 395 nm)	Barge profiler	Hooker	Yes	57
Electric resistivity (sediment core physical properties)	Geotek multi-sensor core logger	Box corer	Rochon A.	No	1
Electric resistivity (sediment core physical properties)	Geotek multi-sensor core logger	CASQ corer	Rochon A.	No	1
Eukaryotes (abundance)	DAPI epifluorescence microscopy	Rosette water sample	Lovejoy C.	No	58
Eukaryotes (biomass)	DAPI epifluorescence microscopy	Rosette water sample	Lovejoy C.	No	58
fCO_2	Derived parameter	Barge water sample	Mucci A., Lansard B.	Yes	12
fCO_2	Derived parameter	Rosette water sample	Mucci A., Lansard B.	Yes	12
fCO_2	Derived parameter	Zodiac water sample	Mucci A., Lansard B.	Yes	12
Foraminifera abundance	Microscopy	Box corer	Rochon A.	No	1
Foraminifera abundance	Microscopy	CASQ corer	Rochon A.	No	1
Foraminifera identification	Microscopy	Box corer	Rochon A.	No	1
Foraminifera identification	Microscopy	CASQ corer	Rochon A.	No	1
Gamma density (sediment core physical properties)	Geotek multi-sensor core logger	Box corer	Rochon A.	No	1
Gamma density (sediment core physical properties)	Geotek multi-sensor core logger	CASQ corer	Rochon A.	No	1
H_2O (atm) concentration	Infrared gas analyzer	Foredeck meteorological tower	Papakyriakou T.	Yes	
HCO_3^- concentration	Derived parameter	Barge water sample	Mucci A., Lansard B.	Yes	12, 16
HCO_3^- concentration	Derived parameter	Rosette water sample	Mucci A., Lansard B.	Yes	12, 16
HCO_3^- concentration	Derived parameter	Zodiac water sample	Mucci A., Lansard B.	Yes	12, 16
Hydro SCAMP (temperature, salinity, chlorophyll, turbidity)	Self-Contained Autonomous Micro-Profiler (SCAMP)	In-water profiler	Gratton Y.	Yes	
Hydrolyzable amino acids (total) (THAAs)	High-performance liquid chromatography (HPLC)	Rosette water sample	Benner R.	Yes	54
Hydrolyzable amino acids (total) (THAAs)	HPLC	Zodiac water sample	Benner R.	Yes	54
Hydroxyl radicals (OH)	HPLC	Rosette water sample	Sempere R.	Yes	
Hydroxyl radicals (OH)	HPLC	Zodiac water sample	Sempere R.	Yes	
IP25 (C25 monounsaturated hydrocarbon)	GC	Box corer	Masse G.	Yes	
IP25 (C25 monounsaturated hydrocarbon)	GC	CASQ corer	Masse G.	Yes	
Irradiance	Satlantic (PUV) (305, 325, 340, 380,. . .)	Foredeck	Sempere R.	Yes	45
Irradiance (412, 490, 555 nm)	Satlantic (OCR) radiometer	Drifting profiling float	Doxaran D.	Yes	18
Lignin phenols (dissolved)	GC–MS	Rosette water sample	Benner R.	Yes	59
Lignin phenols (dissolved)	GC–MS	Zodiac water sample	Benner R.	No	59
Lipid biomarkers	GC flame ionization detection, GC–MS	Box corer	Tolosa I.	Yes	60, 61
Lipid biomarkers	GC flame ionization detection, GC–MS	Foredeck in situ pump	Tolosa I.	Yes	60, 61
Lipid biomarkers $d^{13}C$	GC–combustion–isotope ratio MS	Box corer	Tolosa I.	Yes	60, 61
Lipid biomarkers $d^{13}C$	GC–combustion–isotope ratio MS	Foredeck in situ pump	Tolosa I.	Yes	60, 61
Longwave radiation (Lwin)	Pyrgeometer	Wheelhouse radiation platform	Papakyriakou T.	Yes	
Magnetic susceptibility (sediment core physical properties)	Geotek multi-sensor core logger	Box corer	Rochon A.	No	1
Magnetic susceptibility (sediment core physical properties)	Geotek multi-sensor core logger	CASQ corer	Rochon A.	No	1

Table 1. Continued.

Parameters	Method	Sampling	Principal investigators	Included in the data repository	Reference
Major and minor elements	X-ray fluorescence spectroscopy (XRF) core scanner	CASQ corer	Martinez P.	Yes	
Nanoeukaryotes (abundance)	Flow cytometry	Rosette water sample	Vaulot D.	Yes	62
NH_4^+	Fluorescence	Rosette water sample	Tremblay J. E., Raimbault P.	Yes	63
Nitrate (concentration)	Satlantic ISUS	Rosette profiler	Gratton Y., Prieur L., Tremblay J. E.	Yes	
NO_2^-	Colorimetry, autoanalyzer	Rosette water sample	Tremblay J. E., Raimbault P.	Yes	64
NO_3^-	Colorimetry, autoanalyzer	Rosette water sample	Tremblay J. E., Raimbault P.	Yes	64
Oxygen (dissolved)	Discrete samples, Winkler method	Rosette water sample	Prieur L.	Yes	
Oxygen (dissolved)	Idronaut Ocean Seven O_2 sensor	Continuous horizontal	Papakyriakou T.	Yes	
Oxygen (dissolved)	Sea-Bird SBE-43 sensor	Rosette profiler	Gratton Y., Prieur L.	Yes	
P-wave speed (sediment core physical properties)	Geotek multi-sensor core logger	Box corer	Rochon A.	No	1
P-wave speed (sediment core physical properties)	Geotek multi-sensor core logger	CASQ corer	Rochon A.	No	1
Paleomagnetism	Cryogenic magnetometer	Box corer	Rochon A.	No	1
Paleomagnetism	Cryogenic magnetometer	CASQ corer	Rochon A.	No	1
PAR	Biospherical sensor	Barge profiler	Wright V., Hooker S.	No	57
PAR	Biospherical sensor	Rosette profiler	Gratton Y., Prieur L., Tremblay J. E.	Yes	
PAR	PARLite sensor	Wheelhouse radiation platform	Papakyriakou T.	Yes	
Particle size distribution	Coulter counter	Barge water sample	Reynolds R.	Yes	21, 67
Particle size distribution	LISST-100X	Barge profiler	Reynolds R.	Yes	21, 66
Particle size distribution	LISST-100X	Rosette profiler	Reynolds R.	Yes	21, 66
Particle size distribution	UVP-5 (underwater vision profiler)	In-water profiler	Picheral M.	Yes	65
Particle size distribution	Coulter counter	Rosette water sample	Reynolds R.	Yes	21, 67
Particulate organic carbon (POC)	CHN analyzer on SPM filters	Barge water sample	Doxaran D., Ehn J., Babin M.	Yes	68
Particulate organic carbon (POC)	CHN analyzer on SPM filters	Rosette water sample	Doxaran D., Ehn J., Babin M.	Yes	68
Particulate organic carbon (POC)	CHN analyzer on SPM filters	Zodiac water sample	Doxaran D., Ehn J., Babin M.	Yes	68
Particulate organic carbon (POC)	Wet oxidation	Rosette water sample	Tremblay J. E., Raimbault P.	Yes	69
Particulate organic nitrogen (PON)	Wet oxidation	Rosette water sample	Tremblay J. E., Raimbault P.	Yes	69
Particulate organic phosphorus (POP)	Wet oxidation	Rosette water sample	Tremblay J. E., Raimbault P.	Yes	69
pH (National Bureau of Standards (NBS) scale)	Sea-Bird SBE-18 sensor	Rosette profiler	Gratton Y., Prieur L., Tremblay J. E.	Yes	
pH (total proton scale)	Derived parameter	Barge water sample	Mucci A., Lansard B.	Yes	12, 16
pH (total proton scale)	Derived parameter	Rosette water sample	Mucci A., Lansard B.	Yes	12, 16
pH (total proton scale)	Derived parameter	Zodiac water sample	Mucci A., Lansard B.	Yes	12, 16
pH (total proton scale)	Spectrophotometry	Barge water sample	Mucci A., Lansard B.	Yes	12, 16
pH (total proton scale)	Spectrophotometry	Rosette water sample	Mucci A., Lansard B.	Yes	12, 16
pH (total proton scale)	Spectrophotometry	Zodiac water sample	Mucci A., Lansard B.	Yes	12, 16
Photoheterotrophs (DNA diversity)	DNA clone library	Rosette water sample	Jeanthon C., Boeuf D.	Yes	6, 70
Photosynthetic eukaryotes (diversity)	DNA clone library and TRFLP of sorted populations	Rosette water sample	Vaulot D.	No	71
Photosynthetic eukaryotes (diversity)	DNA from filters	Rosette water sample	Vaulot D.	No	71
Photosynthetic eukaryotes (morphology)	Scanning electron microscopy	Rosette water sample	Vaulot D.	No	72
Photosynthetic parameters	^{14}C incubations	Rosette water sample	Huot Y.	Yes	73
Phytoplankton (abundance)	Inverted microscope	Rosette water sample	Gosselin M., Belanger S.	Yes	46, 74
Phytoplankton (taxonomy)	Inverted microscope	Rosette water sample	Gosselin M., Belanger S.	Yes	46, 74
Phytoplankton pigments	HPLC	Barge water sample	Wright V., Hooker S.	Yes	
Phytoplankton pigments	HPLC	Rosette water sample	Ras J., Claustre H.	Yes	
Picoeukaryotes (abundance)	Flow cytometry	Rosette water sample	Vaulot D.	Yes	62
Picoplankton (diversity)	DNA amplicon library	Rosette water sample	Lovejoy C.	No	75, 82, 83
Picoplankton (diversity)	RNA amplicon library	Rosette water sample	Lovejoy C.	No	75, 82, 83
Plankton taxonomy	UVP-5	In-water profiler	Picheral M., Marec C.	Yes	
PR-containing bacteria (abundance)	Q-PCR	Rosette water sample	Jeanthon C., Boeuf D.	Yes	70
Pressure (barometric)	Pressure sensor	Foredeck meteorological tower	Papakyriakou T.	Yes	
Radiance	Camera luminance	Profile mode	Antoine D., Leymarie E.	Yes	76
Radiance	Camera luminance	Surface mode	Antoine D., Leymarie E.	Yes	76
Radiance: subproduct: average cosines	Camera luminance	Profile mode	Antoine D., Leymarie E.	Yes	76
Radiance: subproduct: average cosines	Camera luminance	Surface mode	Antoine D., Leymarie E.	Yes	76
Radiance: subproduct: irradiance (E)	Camera luminance	Profile mode	Antoine D., Leymarie E.	Yes	76
Radiance: subproduct: irradiance (E)	Camera luminance	Surface mode	Antoine D., Leymarie E.	Yes	76
Radiance: subproduct: Lnadir	Camera luminance	Profile mode	Antoine D., Leymarie E.	Yes	76
Radiance: subproduct: Lnadir	Camera luminance	Surface mode	Antoine D., Leymarie E.	Yes	76
Radiance: subproduct: Qnadir	Camera luminance	Profile mode	Antoine D., Leymarie E.	Yes	76
Radiance: subproduct: Qnadir	Camera luminance	Surface mode	Antoine D., Leymarie E.	Yes	76
Radiance: subproduct: scalar irradiance (Escal)	Camera luminance	Profile mode	Antoine D., Leymarie E.	Yes	76

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Table 1. Continued.

Parameters	Method	Sampling	Principal investigators	Included in the data repository	Reference
Radiance: subproduct: scalar irradiance (Escal)	Camera luminance	Surface mode	Antoine D., Leymarie E.	Yes	76
Radiance (surface leaving radiance)	Biospherical Surface Ocean Reflectance System (BIOSORS)	Foredeck	Hooker	Yes	
Radiance (surface leaving radiance)	Satlantic HyperSAS (hyperspectral surface acquisition system)	Foredeck	Belanger S.	No	77
Radiance (surface leaving radiance)	TriOS above-water sensor	Foredeck	Doxaran D.	No	78
Salinity	Salinometer	Barge water sample	Gratton Y., Prieur L.	Yes	
Salinity	Salinometer	Rosette water sample	Gratton Y., Prieur L.	Yes	
Salinity [S (z)]	Derived parameter	Rosette profiler	Gratton Y., Prieur L., Tremblay J. E.	Yes	
Salinity [S (z)]	Derived parameter from SBE FastCAT LOC IOP package	Barge profiler	Doxaran D.	Yes	48
Shortwave radiation (Swin)	Pyranometer	Wheelhouse radiation platform	Papakyriakou T.	Yes	
Si $(OH)_4$	Colorimetry, autoanalyzer	Rosette water sample	Tremblay J. E., Raimbault P.	Yes	64
Soluble reactive phosphate	Colorimetry, autoanalyzer	Rosette water sample	Tremblay J. E., Raimbault P.	No	64
SPM (suspended particulate material)	Dry weight (gravimetry)	Barge water sample	Doxaran D., Ehn J., Babin M.	Yes	68
SPM (suspended particulate material)	Dry weight (gravimetry)	Rosette water sample	Doxaran D., Ehn J., Babin M.	Yes	68
SPM (suspended particulate material)	Dry weight (gravimetry)	Zodiac water sample	Doxaran D., Ehn J., Babin M.	Yes	68
Sugars	HPLC	Rosette water sample	Sempere R.	No	79
Sugars	HPLC	Zodiac water sample	Sempere R.	No	79
Synechococcus (abundance)	Flow cytometry	Rosette water sample	Vaulot D.	Yes	62
Temperature (air)	Temperature sensor	Foredeck meteorological tower	Papakyriakou T.	Yes	
Temperature (surface skin)	IR transducer	Foredeck meteorological tower	Papakyriakou T.	Yes	
Temperature [T (z)]	Sensor Sea-Bird 3plus on CTD SBE-911	Rosette profiler	Gratton Y., Prieur L., Tremblay J. E.	Yes	
Temperature [T (z)]	Temperature sensor on SBE FastCAT CTD serial no. 175–217	Barge profiler	Doxaran D.	Yes	48
Total inorganic carbon (TIC)	Derived parameter	Barge water sample	Mucci A., Lansard B.	Yes	
Total inorganic carbon (TIC)	Derived parameter	Rosette water sample	Mucci A., Lansard B.	Yes	
Total inorganic carbon (TIC)	Derived parameter	Zodiac water sample	Mucci A., Lansard B.	Yes	
Total organic carbon (TOC)	Wet oxidation	Rosette water sample	Tremblay J. E., Raimbault P.	Yes	56
Total organic nitrogen (TON)	Wet oxidation	Rosette water sample	Tremblay J. E., Raimbault P.	Yes	56
Total organic phosphorus (TOP)	Wet oxidation	Rosette water sample	Tremblay J. E., Raimbault P.	Yes	56
Trace metals	X-ray fluorescence spectroscopy	CASQ corer	Martinez P.	Yes	
Urea (concentration)	Spectrophotometry	Rosette water sample	Tremblay J. E., Raimbault P.	No	64
Volume scattering function (VSF)	Benchtop use of POLVSM (Polarized Volume Scattering Meter)	Barge water sample	Chami M.	No	80, 81
Volume scattering function (VSF)	Benchtop use of POLVSM	Rosette water sample	Chami M.	No	80, 81
Volume scattering function (VSF)	Benchtop use of POLVSM	Zodiac water sample	Chami M.	No	80, 81
Wind direction	Vane	Foredeck meteorological tower	Papakyriakou T.	Yes	
Wind speed	Anemometer	Foredeck meteorological tower	Papakyriakou T.	Yes	

(1) Durantou et al. (2012); (2) Raimbault et al. (1999c); (3) Raimbault and Garcia (2008); (4) Fernández and Raimbault (2007); (5) Miquel et al. (2015); (6) Boeuf et al. (2013); (7) Bélanger et al. (2013a(@); (8) Röttgers and Gehnke (2012); (9) Doxaran et al. (2007); (10) Pegau et al. (1997); (11) Sullivan et al. (2006); (12) Mucci et al. (2010); (13) Lansard et al. (2012); (14) Dickson et al. (2007); (15) Xie et al. (2012); (16) Beaupré-Laperrière et al. (2020); (17) Poteau et al. (2017); (18) Xing et al. (2012); (19) Doxaran et al. (2016); (20) Boss and Pegau (2001); (21) Reynolds et al. (2016); (22) Neukermans et al. (2016); (23) Marie et al. (1997); (24) Cottrell and David (2003); (25) Ortega-Retuerta et al. (2012a); (26) Ortega-Retuerta et al. (2013); (27) Ghiglione et al. (2008); (28) Ortega-Retuerta et al. (2012b); (29) Kirchman (2018); (30) Smith and Azam (1992); (31) Carignan et al. (1998); (32) Link et al. (2013b); (33) Link et al. (2019); (34) Link et al. (2013c); (35) Wei et al. (2020); (36) Rontani et al. (2014); (37) Rontani et al. (2012a); (38) Rontani et al. (2012b); (39) Bricaud et al. (2010); (40) Matsuoka et al. (2012b); (41) Matsuoka et al. (2011); (42) Matsuoka et al. (2014); (43) Amon (2003); (44) Cooper et al. (2005); (45) Para et al. (2013); (46) Blais et al. (2017); (47) Song et al. (2013); (48) Sea-Bird Scientific. (2017); (49) Balzano et al. (2012a); (50) Andrisoa et al. (2019); (51) Sempéré et al. (2019); (52) Tedetti et al. (2007); (53) Bélanger et al. (2006b); (54) Shen et al. (2012); (55) Sohrin and Sempéré (2005); (56) Raimbault et al. (1999b); (57) Hooker et al. (2020); (58) Vaqué et al. (2008); (59) Fichot et al. (2016); (60) Tolosa et al. (2013); (61) Tolosa and de Mora (2004); (62) Marie et al. (2014); (63) Holmes et al. (1999); (64) Aminot and Kérouel (2007); (65) Picheral et al. (2010); (66) Reynolds et al. (2010); (67) Runyan et al. (2020); (68) Doxaran et al. (2012b); (69) Raimbault et al. (1999a); (70) Boeuf et al. (2016); (71) Balzano et al. (2012b); (72) Balzano et al. (2017); (73) Marcel et al. (1994); (74) Ardyna et al. (2017b); (75) Comeau et al. (2012); (76) Antoine et al. (2013b); (77) Mobley (1999); (78) Ruddick et al. (2006); (79) Panagiotopoulos et al. (2014); (80) Chami et al. (2014); (81) Harmel et al. (2016); (82) Monier et al. (2014); (83) Monier et al. (2015)

the best-fitting fraction (x) of ($n+1$) source water types that contribute to the n observed values of the selected tracers in a parcel of water via a solution of an overdetermined system of linear equations that minimizes the residual error. Boundary conditions were applied to the method to guarantee that all fractions calculated were positive and that the sum of all fractions was 100 % (mass conservation).

During MALINA, the Mackenzie Shelf was for the most part entirely ice-free, and the ice pack was located beyond the shelf break (Fig. 2). The transition zone was characterized by different expanses of drifting sea ice. Significant contributions of meteoric water (> 25 %) to the surface mixed layer (SML) were only observed close to the Mackenzie River mouth and on the inner shelf (Fig. 4). A relatively small fraction of sea ice meltwater was detected beyond the shelf break, mostly along the transect 600. Below the SML, the wPML was the predominant water mass down to 100 m depth. The UHW extends from the interior ocean onto the outer shelf from 120 to 180 m depth. Relatively high fractions of UHW were also found at 50 m depth along the Mackenzie and Kugmallit canyons, which are recognized sites of enhanced shelf-break upwelling caused by wind- and ice-driven ocean surface stresses. Below 200 m depth, the LHW with an Atlantic origin was always the prevailing water mass.

4.2 Temperature and salinity from the CTD

Temperature and salinity for the first 100 m of transects 600 and 300, the two transects originating from the Mackenzie River delta, are presented in Fig. 5. They confirm what was found by the water mass analysis (Sect. 4.1): most of the freshwater is coming from the western part of the Mackenzie River delta. This is also in accordance with many studies that documented that during the summer, a combination of ice melting and river runoff was generating a highly stratified surface layer (Carmack and Macdonald, 2002; Forest et al., 2013). The signature of an eddy may be observed at 75 m in the salinity data at 70° N, approximately 70 km from shore (Fig. 5b).

4.3 Underwater bio-optical data

4.3.1 Inherent optical property (IOP) profiling from the ship, the barge and the zodiac

The total, non-water, spectral absorption (a), attenuation (c) and backscattering (b_b) coefficients were measured using an AC9 attenuation and absorption meter and a BB9 scattering meter (Wetlabs), a HydroScat-6, and a-Beta sensors (HOBI Labs) either attached to the CTD rosette frame onboard the CCGS *Amundsen* or deployed separately from the barge or the zodiac tender. These devices were using either 10 or 25 cm optical path lengths, depending on the turbidity of the water sampled. Detailed information about the deployment and the data processing of the IOP data can be found in Doxaran et al. (2012a).

Figure 6 shows cross-sections of the total absorption and backscattering coefficients at 440 nm ($a(440)$ and $b_b(440)$), derived as $b_b = b_{bp} + b_{bw}$, where b_{bw} is the backscattering coefficient of pure seawater (Morel, 1974). Both $a(440)$ and $b_b(440)$ showed the same patterns along the transects 600 and 300. Close to the estuary, higher absorption (Fig. 6a) and total scattering (Fig. 6b) can be observed at the surface, likely reflecting the important quantities of dissolved and particulate organic matter delivered by the Mackenzie River. Higher values are also observed in transect 600 compared to transect 300, which is farther away from the mouth of the Mackenzie River. Both $a(440)$ and $b_b(440)$ decreased rapidly toward higher latitudes, where the water of the Mackenzie River mixes with seawater from the Beaufort Sea.

4.3.2 Particulate and CDOM absorption

Chromophoric dissolved organic matter absorption (a_{CDOM}) was measured from water samples filtered with 0.2 µm GHP (glass fiber hydrophilic polypropylene) filters (Acrodisc Inc.) using an UltraPath (World Precision Instruments Inc.) between 200 and 735 nm. In most cases, a 2 m optical path length was used for the measurement, except for coastal waters near the Mackenzie River mouth (Fig. 1), where a 0.1 m optical path length was used. Particulate absorption (ap) was measured using a filter-pad technique modified from Röttgers and Gehnke (2012). Briefly, seawater was filtered through a 25 mm Whatman GF/F (glass fiber filters) less than 3 h after sampling. Filters were placed in the center of a 150 mm integrating sphere equipped with a handmade Spectralon filter holder. The spectral optical density (OD(λ)) of the particles retained on the filter was then measured using a PerkinElmer Lambda-19 spectrophotometer from 300 to 800 nm at 1 nm resolution. More details about particulate- and dissolved-absorption measurements can be found in Röttgers and Gehnke (2012), Bélanger et al. (2013b), and Matsuoka et al. (2012a).

Examples of a_{CDOM} spectra measured at the surface for the northernmost and the southernmost stations of transects 600 and 300 are presented in Fig. 7a. The marked influence of the organic matter of terrestrial origin can be observed for the stations located at the mouth of the Mackenzie River (697 and 398). Because the organic matter delivered by the river is highly humic and colored, the absorption at 254 nm was approximately 15 times higher at the southern shelf stations for both transects compared to the northern stations (620 and 320). Likewise, the specific UV absorbance of dissolved organic carbon at 254 nm ($SUVA_{254}$), a metric commonly used as a proxy for assessing both chemical (Weishaar et al., 2003; Westerhoff et al., 2004) and biological reactivity (Berggren et al., 2009; Asmala et al., 2013) of the dissolved organic matter (DOM) pool in natural aquatic ecosystems, decreased rapidly along the south–north gradient in both transects 600 and 300 (Fig. 7c). This observation is in accordance with a previous study that showed that $SUVA_{254}$ was higher in in-

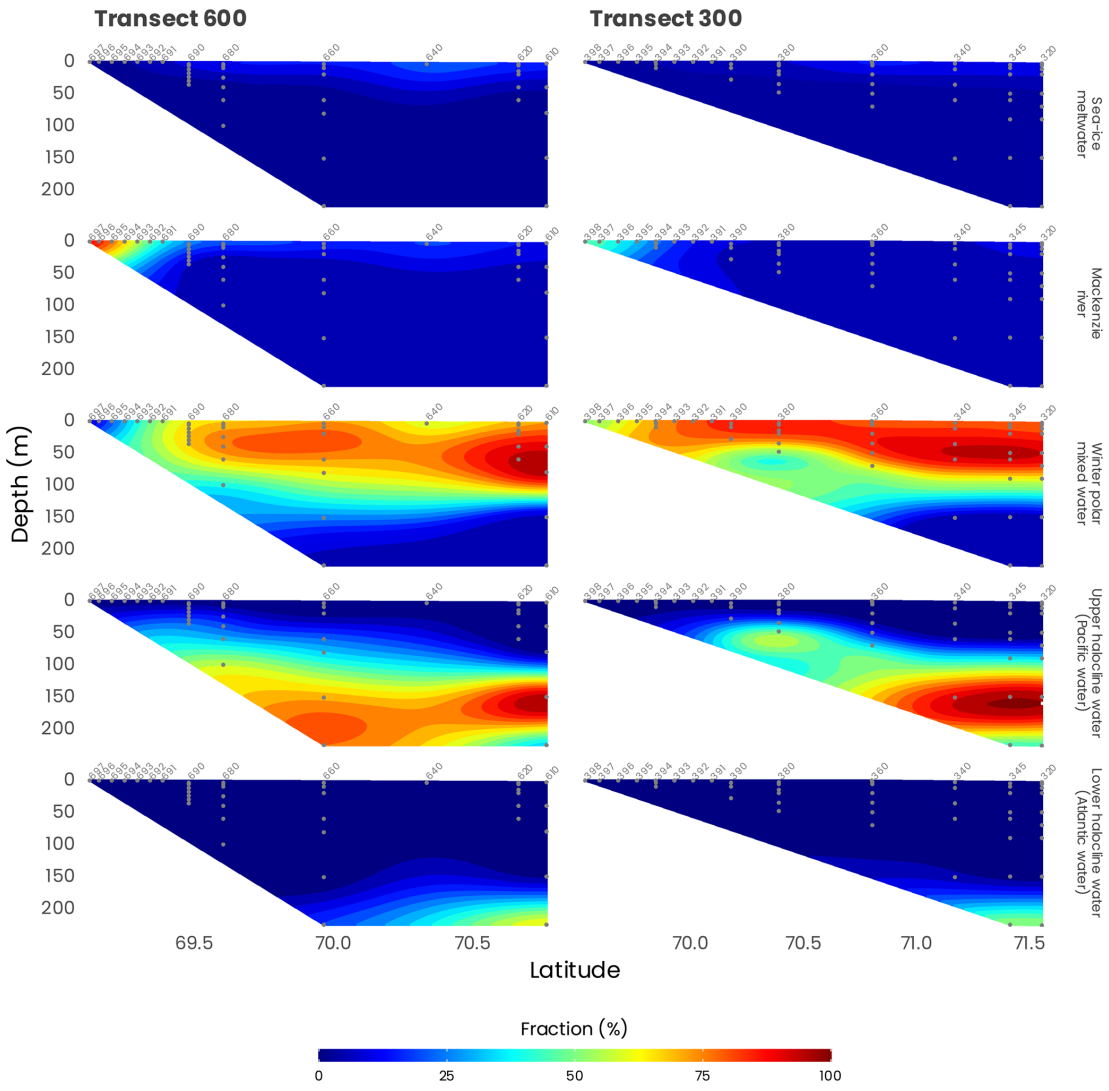

Figure 4. Distribution of source water types along transects 600 and 300 (see Fig. 1). Station numbers are identified in light gray on top of each panel.

land ecosystems due to elevated lateral connectivity with the surrounding terrestrial landscape and organic-matter inputs from the tributaries (Massicotte et al., 2017). The decrease in $SUVA_{254}$ toward northern stations (Fig. 7c) suggests that terrestrially derived DOM transiting toward the ocean is gradually degraded into smaller and more refractory molecules.

Particulate-absorption spectra (a_p) for the northernmost and the southernmost stations of transects 600 and 300 are presented in Fig. 7b. Particulate absorption at the stations located in the estuary (697 and 398) was much higher than that measured at the open-water stations (620 and 320). For instance $a_p(443)$ measured at stations 620 ($0.03\,m^{-1}$) and 697 ($8.62\,m^{-1}$), the northernmost and the southernmost stations of transects at the mouth of the Mackenzie River, shows that a_p decreases rapidly along the latitudinal axes. This can be possibly explained because the drained organic and inorganic material from the surrounding landscape of the Mackenzie's watershed is degraded or sediments rapidly as it is transferred to the ocean.

4.3.3 Other optical measurements and radiometric quantities

Other optical instruments were attached to the rosette sampler. These include a transmissometer (Wetlabs C-Star, path 25 cm) for beam attenuation measurement, a chlorophyll fluorometer (Seapoint) and a CDOM fluorometer (Optic & Mikro Elektronik, Germany; see Amon, 2003). Additionally, a LISST-100X (laser in situ scattering and transmissometry, Sequoia Scientific) was attached to the rosette and provided beam attenuation (532 nm) and forward light scattering measurements at 32 angles, from which particle size distribution was estimated. Various optical measurements were also made

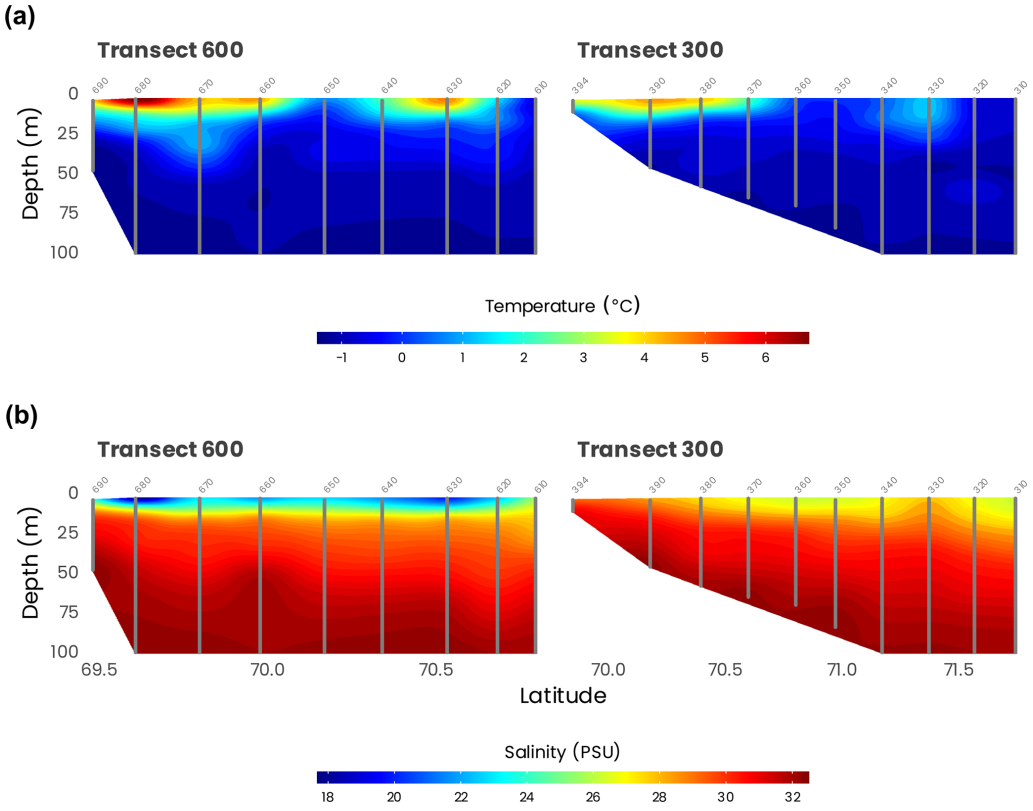

Figure 5. Cross-sections of temperature **(a)** and salinity **(b)** measured by the CTD (gray dots) along transects 600 and 300. Station numbers are identified in light gray on top of each panel.

in the laboratory to determine other IOPs. These include the absorption of colored dissolved (a_{CDOM}) and particulate (a_p) organic matter and the absorption coefficients of non-algal particles (a_{NAP}) and phytoplankton (a_{phi}). Apparent-optical-property (AOPs) measurements included light transmittance (T), photosynthetically available radiation (PAR), downward irradiance (E_d), upwelled radiance (L_u) and global solar irradiance (E_s). The latter three radiometric quantities were measured simultaneously using a compact optical-profiling system (C-OPS) manufactured by Biospherical Instruments Inc. (San Diego, California) that was deployed during MALINA Leg 2b. The principal data products obtained from the C-OPS data were the diffuse attenuation coefficient (K_d) plus the water-leaving radiance (L_W), including all normalized forms. Detailed methodology and results derived from C-OPS measurements can be found in Doxaran et al. (2012a), Antoine et al. (2013a), Bélanger et al. (2013b) and Hooker et al. (2013).

4.4 Nutrients

Samples for nitrate, nitrite, soluble reactive phosphorus and silicate determination were collected into 20 mL polyethylene flasks, immediately poisoned with mercuric chloride (Kirkwood, 1992) and stored for subsequent laboratory analysis according to Raimbault et al. (1990) and Aminot and Kérouel (2007). Ammonium concentrations (40 mL collected into 60 mL polycarbonate tubes) were measured onboard using the sensitive method of Holmes et al. (1999), with a detection limit of 5 nmol L^{-1}. Samples for organic-matter determination were collected into 50 mL glass Schott bottles, immediately acidified with 100 µL of 0.5 N H_2SO_4 and stored in the dark at 5 °C. Dissolved organic carbon (DOC), dissolved organic nitrogen (DON) and dissolved organic phosphorus (DOP) were determined at the laboratory using the wet-oxidation procedure according to Raimbault et al. (1999b).

Nitrate levels were always very low at the surface, with concentration generally lower than 0.01 µmol L^{-1}, except in the Mackenzie plume (Fig. 8). It is interesting to note that nitrate was never entirely depleted, and some traces (0.005 to 0.01 µmol L^{-1}) were always detectable in surface waters (Fig. 8a). Ammonium distribution showed the same pattern. Even if concentrations were very low (generally < 0.03 µmol L^{-1}), this nutrient, like nitrate, was always detected, suggesting that in situ sources of nitrate and ammonium exist offshore, certainly due to biological processes. Phosphate concentrations showed the opposite distribution (Fig. 8b). Despite nitrogen depletion, surface waters were always phosphate-replete. The highest concentrations, around

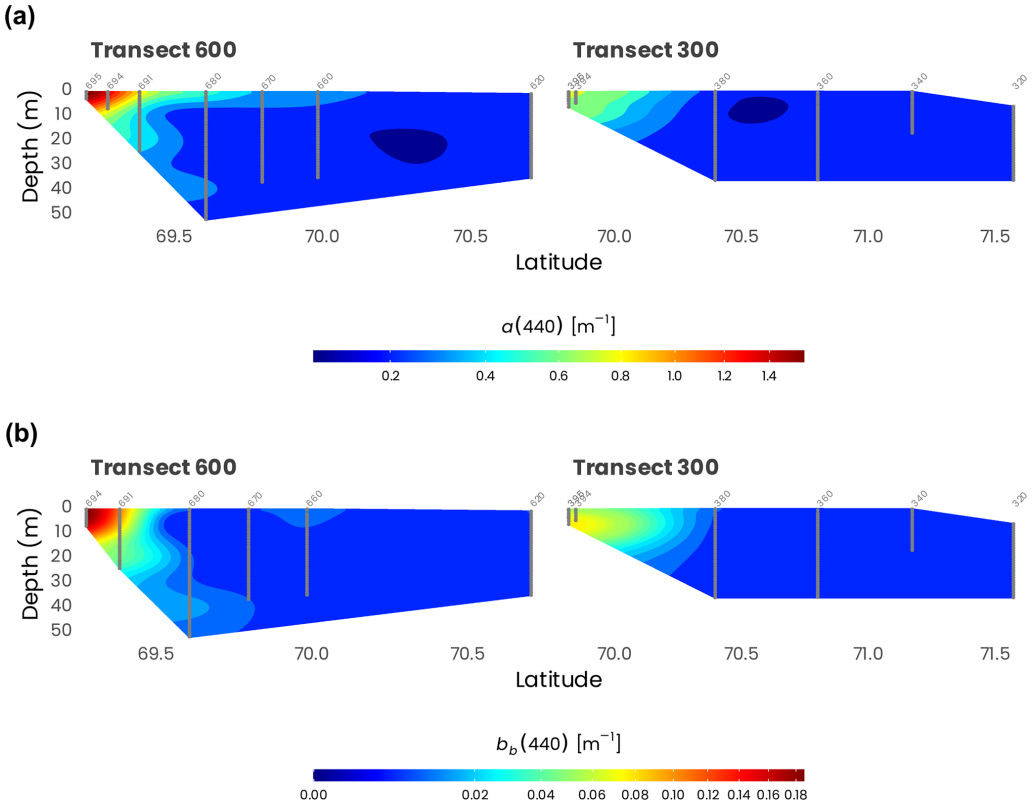

Figure 6. Cross-sections of **(a)** absorption ($a(440)$) and **(b)** total scattering ($b_b(440)$) measured from the barge at 440 nm with an AC9 and BB9, respectively, along transects 600 and 300. Station numbers are identified in light gray on top of each panel. Note that the data have been square-root-transformed for the visualization.

$0.5\,\mu mol\,L^{-1}$, were observed far from Mackenzie's mouth, revealing a clear west–east gradient. The silicate distribution was similar to that of nitrate. But surface waters were always silicate-replete, with concentrations largely above the detection limit ($> 4\,\mu mol\,L^{-1}$). The impact of the Mackenzie River was clear close to the coast for inorganic nutrients and farther offshore for dissolved organic nutrients. A quarter of the estimated annual nutrient supply by the Mackenzie River occurred during July–August. The supply of DON was 8 times larger than that of nitrate-N. By contrast, the amount of DOP supplied was only 2.5 times higher than the amount of phosphate (Tremblay et al., 2014). The Mackenzie River enriched the western Canadian Beaufort Shelf with inorganic and organic N, potentially supporting most of the primary production but not with phosphate or ammonium. Large deliveries of N relative to P by rivers relax coastal communities from N limitation, allowing them to tap into the excess P originating from the Pacific Ocean. Then, river inputs locally rectified the strong regional deficit of inorganic N, i.e., negative N^* (Tremblay et al., 2014).

4.5 Dissolved organic carbon, total dissolved nitrogen, total hydrolyzable amino acids and total dissolved lignin phenols

Water samples were collected at selected stations and water masses for analyses of dissolved organic carbon (DOC), total dissolved nitrogen (TDN), total hydrolyzable amino acid (THAA) and total dissolved lignin phenol ($TDLP_9$) concentrations. Samples for DOC, TDN and THAAs were gravity-filtered from Niskin bottles using pre-combusted (GF/F) filters (0.7 µm pore size) and stored frozen (−20 °C) immediately after collection in pre-combusted borosilicate glass vials (Shen et al., 2012). Samples for $TDLP_9$ analysis (between 1 and 10 L) were gravity-filtered from Niskin bottles using Whatman Polycap AS cartridges (0.2 µm pore size), acidified to pH between 2.5 and 3 with sulfuric acid, and extracted within a few hours using C-18 cartridges (Louchouarn et al., 2000; Fichot et al., 2013). The C-18 cartridges were stored at 4 °C until elution with 30 mL of methanol (high-precision-liquid-chromatography-grade, HPLC-grade), and the eluent was stored in sealed, pre-combusted glass vials at −20 °C until analysis. DOC and TDN concentrations were measured by high-temperature combustion using a Shimadzu total-organic-carbon analyzer

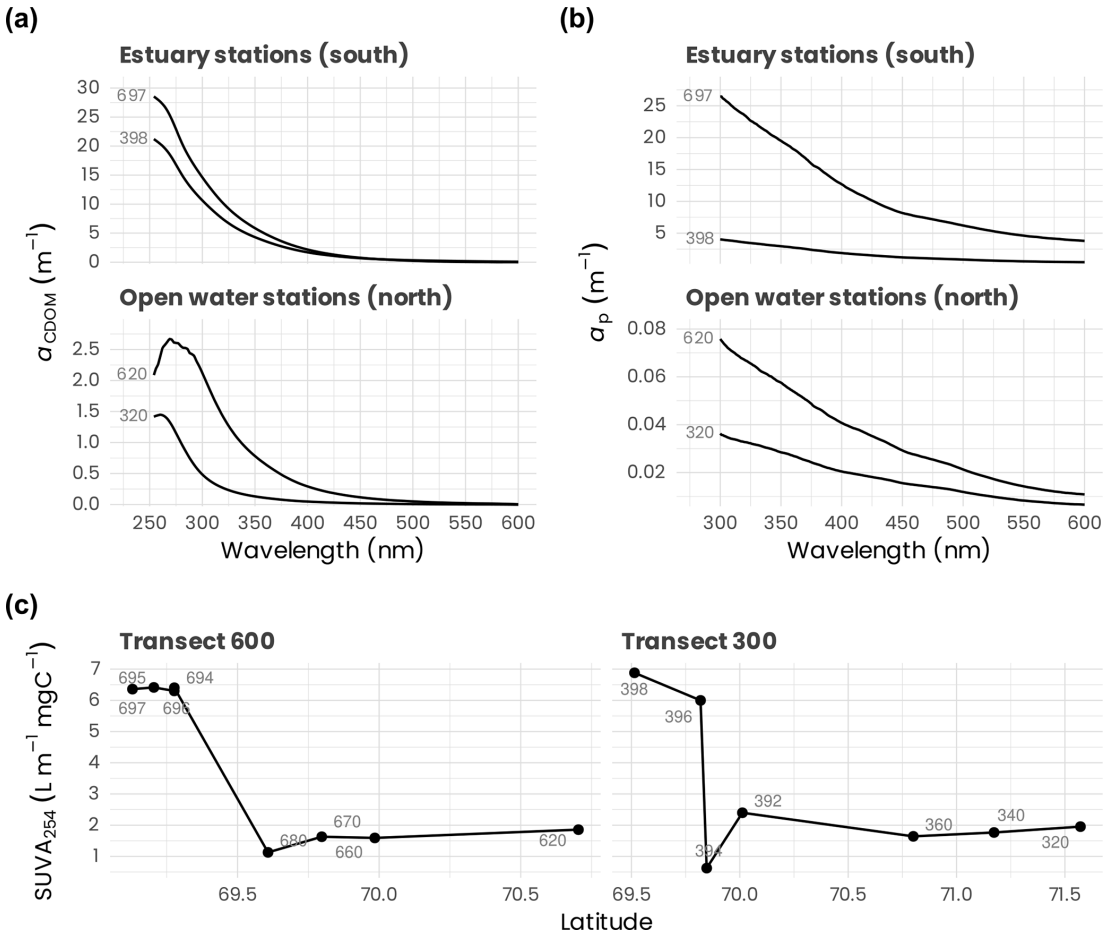

Figure 7. **(a)** Absorption spectra between 254 and 600 nm of chromophoric dissolved organic matter (a_{CDOM}) measured at the surface for the northern (620, 320) and southern (697, 398) stations of the transects 600 and 300. **(b)** Particulate-absorption spectra (a_p) measured between 300 and 600 nm measured at the surface for the northernmost and the southernmost stations of the transects 600 and 300. **(c)** Specific UV absorbance at 254 nm ($SUVA_{254}$, i.e., absorption of light at 254 nm per unit of carbon) at surface for stations along transects 600 and 300. Stations are identified in light gray (see Fig. 1 for an overview of the station locations). Note the difference in the y axes used in panels **(a)** and **(b)**, which highlights the important differences in dissolved and particulate absorption between stations in the estuary and those offshore.

(TOC-V) equipped with an inline chemiluminescence nitrogen detector and an autosampler (Benner and Strom, 1993). Blanks were negligible, and the coefficient of variation between injections of a given sample was typically less than 1 %. Analysis of a deep-seawater reference standard (University of Miami) every sixth sample was used to check the accuracy and consistency of measured DOC and TDN concentration. THAAs were determined as the sum of 18 dissolved amino acids using an Agilent high-performance liquid chromatography system equipped with a fluorescence detector (excitation: 330 nm; emission: 450 nm). Samples (100 µL) of filtered seawater were hydrolyzed with 6 mol L^{-1} hydrochloric acid using a microwave-assisted vapor-phase method (Kaiser and Benner, 2005). Free amino acids liberated during the hydrolysis were separated as o-phthaldialdehyde derivatives using a LiCrospher RP-18 or Zorbax SB-C18 column (Shen et al., 2012). Detailed methodological information can be found in Fichot et al. (2013) and Shen et al. (2012).

Surface DOC concentrations along the transects 300 and 600 behaved approximately conservatively with salinity, decreasing from 458 µmol L^{-1} in the Mackenzie River end-member (salinity = 0.2 PSU) to 123 µmol L^{-1} at a salinity of 26.69 PSU (Fig. 9a). DOC concentrations in surface waters further decreased to minimum values of ≈ 66 µmol L^{-1} offshore (Fichot and Benner, 2011). Concentrations generally increased by a few µmol L^{-1} in the upper halocline relative to surface values but then generally decreased with depth, reaching 53–57 µmol L^{-1} in the lower halocline and ≈ 43–50 µmol L^{-1} in deep water masses (depth > 1000 m). Similarly to DOC, surface $TDLP_9$ concentrations along transects 600 and 300 behaved approximately conservatively with salinity, decreasing from ≈ 93–96 nmol L^{-1} in the Mackenzie River end-member (salinity = 0.2 PSU) to ≈ 12 nmol L^{-1}

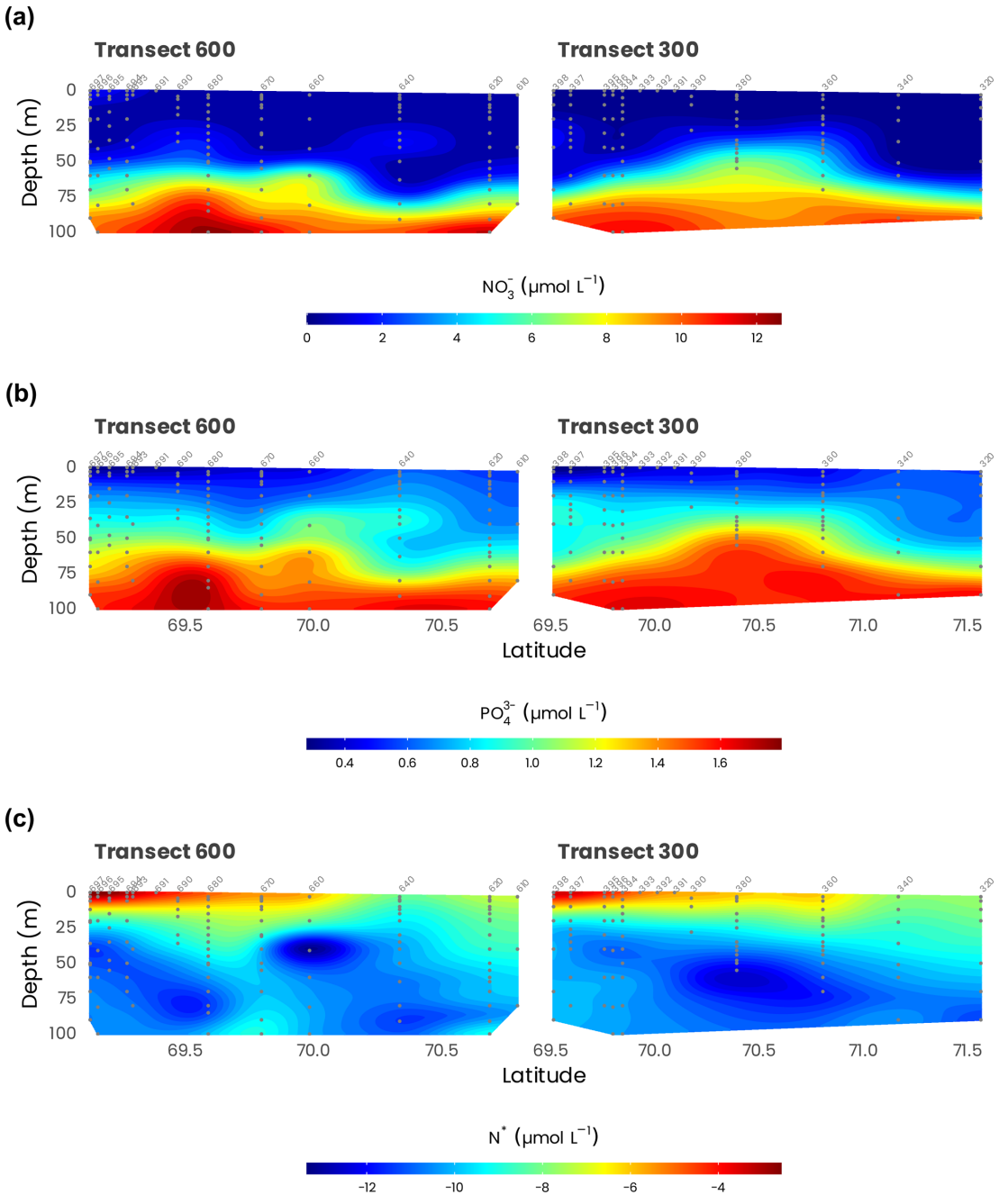

Figure 8. Cross-sections of **(a)** NO_3^- and **(b)** PO_4^{3-} measured from Niskin bottles (gray dots) along transects 600 and 300. **(c)** N^* defined as $N - rP$, with $r = N/P = 13.1$ (see the text for the details). Station numbers are identified in light gray on top of each panel.

at a salinity of 26.69 PSU (Fig. 9b). Surface concentrations reached minimum values of $\approx 2.5\,\text{nmol}\,\text{L}^{-1}$ offshore (Fichot et al., 2016). $TDLP_9$ concentrations generally decreased with depth, reaching minimum values of $< 1.5\,\text{nmol}\,\text{L}^{-1}$ below the halocline. Surface concentrations of THAAs along the transects 600 and 300 decreased from $576\,\text{nmol}\,\text{L}^{-1}$ in the Mackenzie River end-member (salinity = 0.2 PSU) to $317\,\text{nmol}\,\text{L}^{-1}$ at a salinity of 26.69 PSU (Fig. 9c). Unlike DOC and $TDLP_9$'s, concentrations of THAAs did not follow a conservative mixing line along the salinity gradient. Elevated concentrations of THAAs were observed in mid-salinity waters in both transects, suggesting plankton production in these regions. In comparison, THAA concentrations in the slope and basin waters were lower and decreased with depth, reaching minimal values of $\approx 70\,\text{nmol}\,\text{L}^{-1}$ below the halocline (Shen et al., 2012).

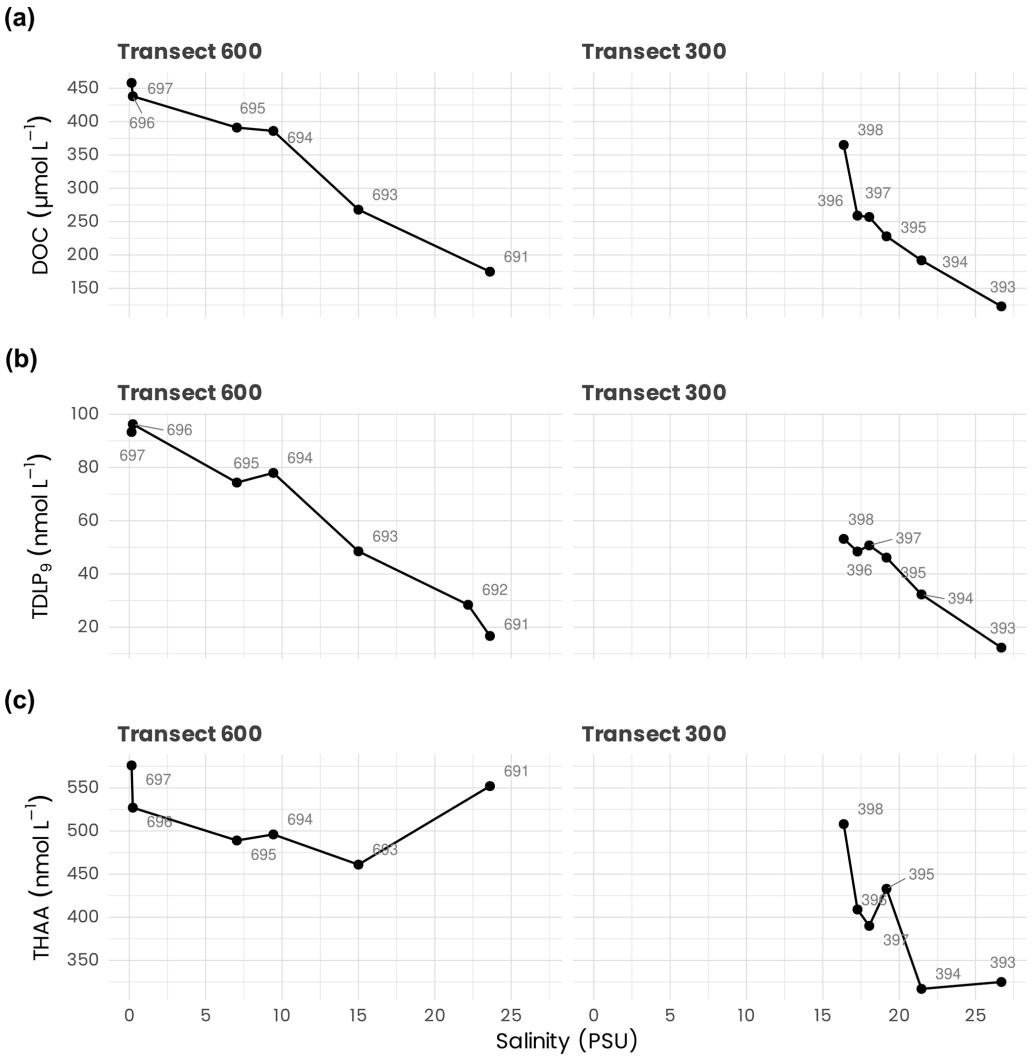

Figure 9. Surface concentrations of **(a)** dissolved organic carbon (DOC), **(b)** total dissolved lignin phenols ($TDLP_9$'s) and **(c)** total hydrolyzable amino acids (THAAs) measured along transects 600 and 300 and plotted against salinity.

4.6 Pigments

Water samples (volumes between 0.25 and 2.27 L) were filtered through glass fiber GF/F filters (25 mm diameter, particle retention size 0.7 µm). They were immediately frozen at −80 °C, transported in liquid nitrogen and then stored at −80 °C until analysis on land. Samples were extracted in 3 mL HPLC-grade methanol for 2 h minimum. After sonication, the clarified extracts were injected (within 24 h) onto a reversed-phase C8 Zorbax Eclipse column (dimensions: 3 × 150 mm; 3.5 µm pore size). The instrumentation comprised an Agilent Technologies 1100 series HPLC system with diode array detection at 450, 667 and 770 nm of phytoplankton pigments (carotenoids; chlorophylls *a*, *b* and *c*; and bacteriochlorophyll *a*). A total of 22 pigments were analyzed and quantified. Details of the HPLC analytical procedure can be found in Ras et al. (2008).

As illustrated in Fig. 10, the phytoplankton biomass, indicated by total chlorophyll *a* concentrations, was the highest at the coast (up to 3.5 mg m^{-3}), decreasing offshore (to about 0.010 mg m^{-3}) with the formation of a subsurface chlorophyll maximum (SCM) around 60 m. In terms of biomass integrated over the sampled depth, values range from 6.2 and 8.9 mg m^{-2} at the coast to 14.3 and 13.2 mg m^{-2} offshore for transects 300 and 600, respectively. In general, the most predominant accessory pigment was fucoxanthin, indicating that diatoms constitute a large proportion of the phytoplankton assemblage. However, in offshore waters and around the SCM, 19'-hexanoyloxyfucoxanthin concentrations were equivalent to or sometimes higher than fucoxanthin, suggesting that, in these waters, haptophytes can predominate over diatoms. Other pigments such as chlorophyll *b* and prasinoxanthin suggest the presence of green algae and probably *Micromonas*-type cells, especially in coastal waters and at

the surface. For more detailed information, see Coupel et al. (2015), who used this data set applied to the CHEMTAX (CHEMical TAXonomy) chemotaxonomic tool to assess the distribution of phytoplankton communities.

4.7 Phytoplankton abundance and diversity

The abundance of the eukaryotic pico- and nano-phytoplankton was measured by flow cytometry onboard the *Amundsen* with a FACSAria instrument (Becton Dickinson, San Jose, CA, USA) following the protocol of Marie et al. (1999).

In transect 300 and 600 (Fig. 11), the abundance of pico- and nano-phytoplankton reached maximal values around 5000 and 3000 cells mL^{-1}, respectively. On transect 600, pico-eukaryotes' higher abundances were restricted to the surface layer, with a 5- to 10-fold drop at 30 m. In contrast, nano-eukaryotes formed clear deep maxima, especially at stations 610 and 680. On transect 300, pico-eukaryotes were also abundant in the surface at the more off-shore stations. Still, they decreased sharply near shore, while nano-eukaryotes' highest concentrations were near the river mouth, linked to high diatom concentrations (Balzano et al., 2012a). The composition of eukaryotic phytoplankton was determined with two different approaches. We isolated 164 cultures using a range of techniques (single-cell isolation, serial dilution, flow cytometry sorting) that have been characterized morphologically and genetically (Balzano et al., 2012a, 2017) and deposited to the Roscoff Culture Collection (http://www.roscoff-culture-collection.org, last access: 16 March 2021). Among these cultures, several new species have been discovered such as the new species of green algae *Mantoniella beaufortii* (Yau et al., 2020) or the diatom *Pseudo-nitzschia arctica* (Percopo et al., 2016), but more await description, in particular among *Pelagophyceae*. One of the strains isolated (RCC2488, *Chlamydomonas malina* nomen nudum) has been recently found to be suitable for biotechnology applications (Morales-Sánchez et al., 2020). We also used molecular approaches by sorting pico- and nano-eukaryotic communities and characterizing their taxonomic composition by TRFLP (terminal-restriction fragment length polymorphism) analysis and cloning and sequencing of the 18S ribosomal RNA gene (Balzano et al., 2012a). While the pico-phytoplankton was dominated by the species *Micromonas polaris*, the nano-phytoplankton was more diverse and dominated by diatoms mostly represented by *Chaetoceros neogracilis* and *C. gelidus*, with the former mostly present at surface waters and the latter prevailing in the SCM (Balzano et al., 2012a). Furthermore, *C. neogracilis* sampled from the Beaufort Sea consists of at least four reproductively isolated genotypes (Balzano et al., 2017). The comparison between the taxonomy of natural communities and isolated cultures (Fig. 12) reveals that, although we succeeded at isolating some dominant species in the field such as *M. polaris*, *C. neogracilis* and *C. gelidus*, some other important taxa such as the diatom *Fragilariopsis* or the haptophyte *Chrysochromulina* were not recovered.

4.8 Carbon fluxes

In the context of climate change, the main objective of the MALINA oceanographic expedition was to determine how (1) primary production, (2) bacterial activity and (3) photodegradation influence carbon fluxes and cycling of organic matter in the Arctic. In the following sections, we present an overview of these processes in the water column that are detailed in Ortega-Retuerta et al. (2012a), Xie et al. (2012) and Tremblay et al. (2014) and refer to Link et al. (2013a), Tolosa et al. (2013) and Rontani et al. (2012b) for the related processes at the sediment–water interface.

4.8.1 Phytoplankton primary production

At each station, when productivity was quantified, rates of carbon fixation (primary production) were determined using a ^{13}C isotopic technique (Raimbault and Garcia, 2008). For this purpose, three 580 mL samples were collected at minimum sun elevation or before sunrise at six to seven depths between the surface and the depth where irradiance was 0.3 % of the surface value and poured into acid-cleaned polycarbonate flasks. Incubations were carried out immediately following the tracer addition in an on-deck incubator. This consisted of six to seven opaque boxes, each with associated neutral and blue screens, allowing around 50 %, 25 %, 15 %, 8 %, 4 %, 1 % and 0.3 % light penetration. At five stations, incubations were also performed in situ on a drifting rig with incubation bottles positioned at the same depth where samples for on-deck incubations were collected. After 24 h, samples were filtered through pre-combusted (450 °C) Whatman GF/F filters (25 mm diameter). After filtration, filters were placed into 2 mL glass tubes, dried for 24 h in a 60 °C oven and stored dry until laboratory analysis. These filters were used to determine the final ^{13}C enrichment ratio in the particulate organic matter on an Integra CN mass spectrometer. Filtrates were poisoned with $HgCl_2$ and stored to estimate ammonium regeneration and nitrification rates. The isotopic enrichment of particulate organic matter and dissolved NH_4^+ and NO_3^- at the end of incubations were used to calculate net C and N uptake and the recycling of NH_4^+ and NO_3^- (Raimbault et al., 1999c).

Daily rates of primary production at the surface were generally very low across the survey area, ranging from 0.1 mg C m^{-3} d^{-1} offshore to a maximum of 545 mg C m^{-3} d^{-1} in Kugmallit Bay (Fig. 13), associated with the Mackenzie River discharge (Tremblay et al., 2014). Ammonification and nitrification followed the same coastal-offshore pattern, with rates driving most, if not all, of the NH_4^+ and NO_3^- consumption in the surface layer. Primary production was generally maximum at the surface, but high rates were often observed at depth in the nitracline layer, as-

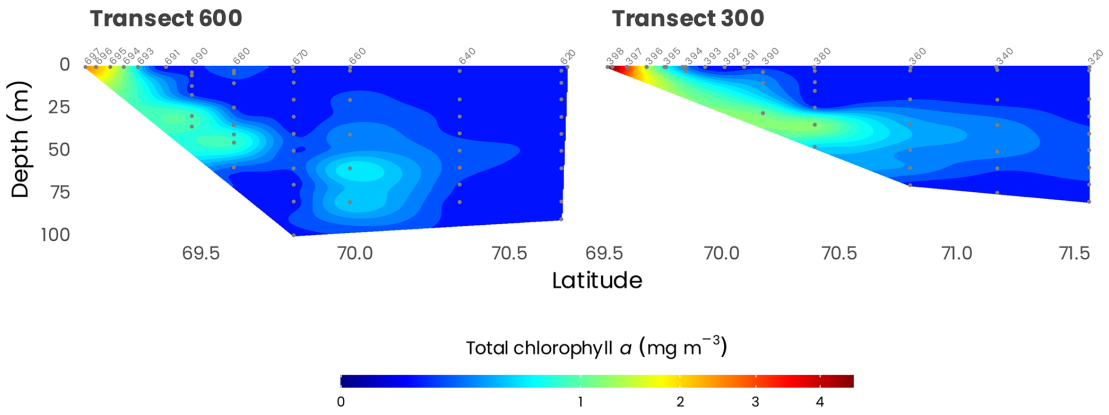

Figure 10. Cross-sections of total chlorophyll *a* measured from HPLC (gray dots) along transects 600 and 300. Station numbers are identified in light gray on top of each panel. Note that the data have been square-root-transformed for the visualization.

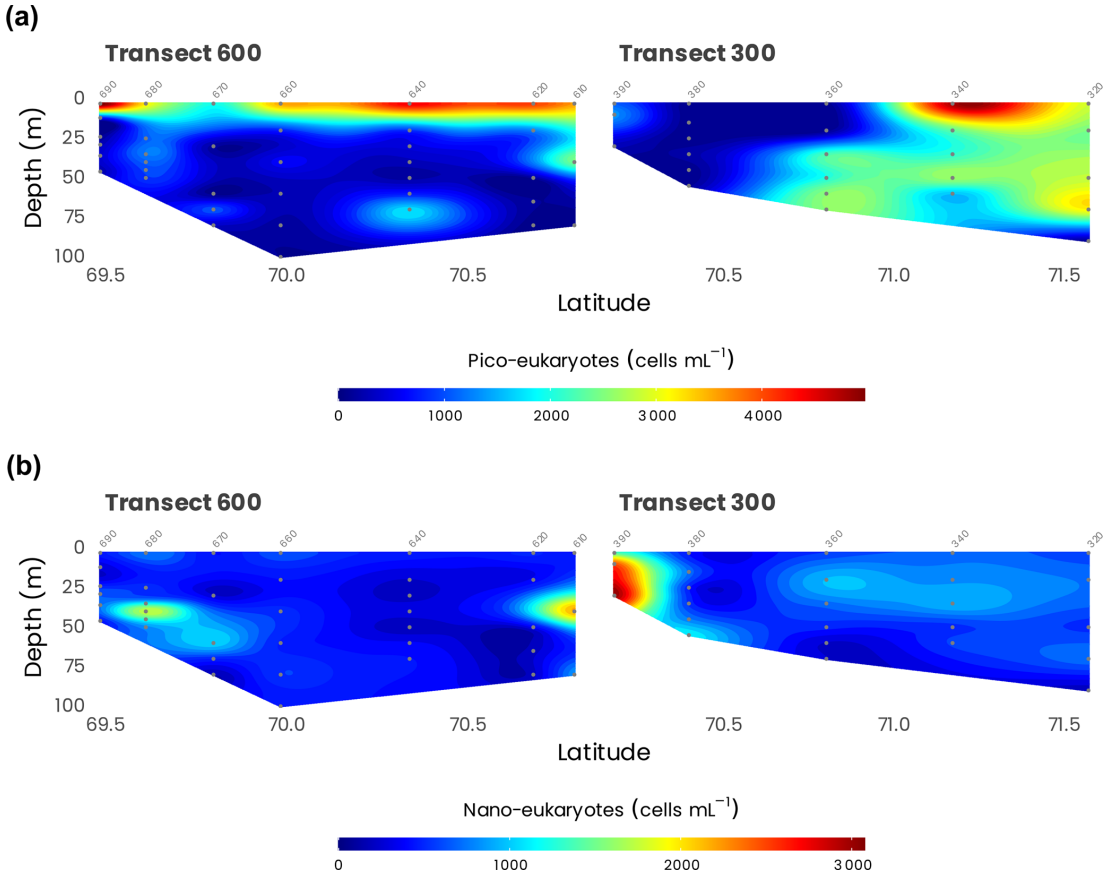

Figure 11. Concentrations of photosynthetic **(a)** pico- and **(b)** nano-eukaryotes measured by flow cytometry during the MALINA cruise on transects 600 and 300.

sociated with a chlorophyll maximum. The range of uptake rates of ammonium at the surface generally overlapped with the range of nitrate uptake rates. Nitrate uptake below the surface amounted to 40 %–60 % of total nitrogen uptake, a proportion that is approximately 2 times greater than at the surface (Ardyna et al., 2017a).

Nitrification and ammonium regeneration was detectable over the whole water column, ranging from 2 to 20 nmol L^{-1} d^{-1}. The highest rates were generally located at the base of the euphotic zone, leading to the formation of subsurface ammonium- and nitrite-maximum layers. Surface communities and especially the accumulation of large

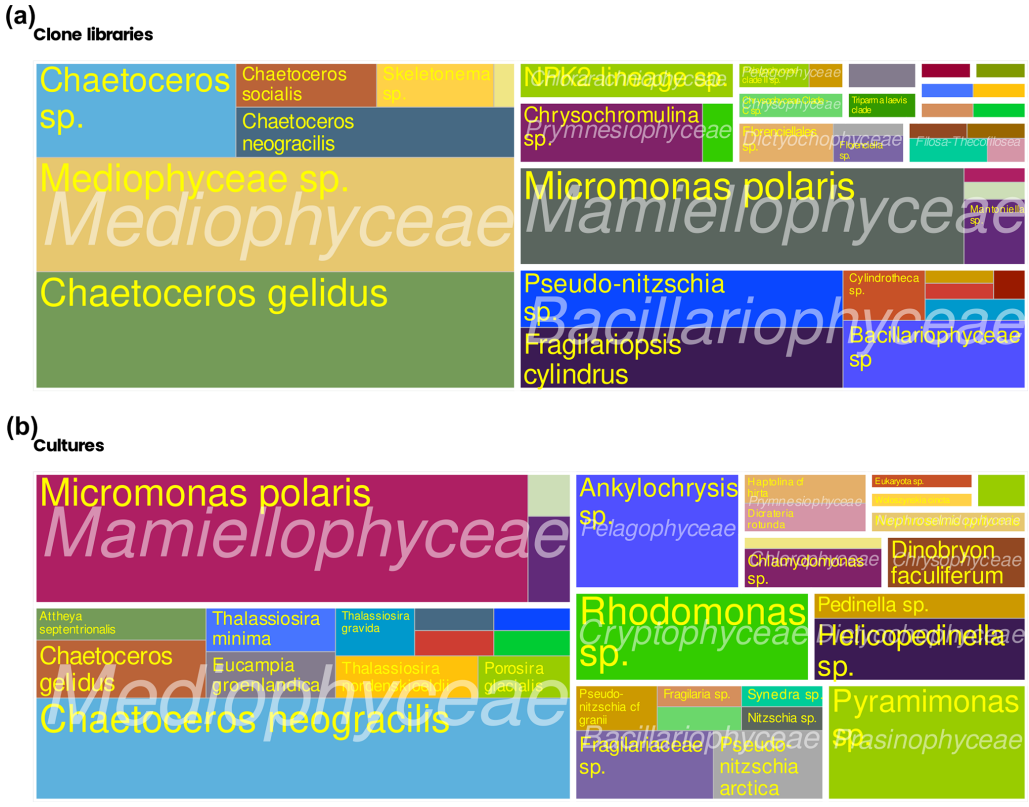

Figure 12. **(a)** Taxonomic composition of populations of photosynthetic pico- and nano-eukaryotes sorted by flow cytometry from clone library sequences (Balzano et al., 2012a). **(b)** Taxonomic composition of cultures of phytoplankton isolated during the MALINA cruise (Balzano et al., 2012a).

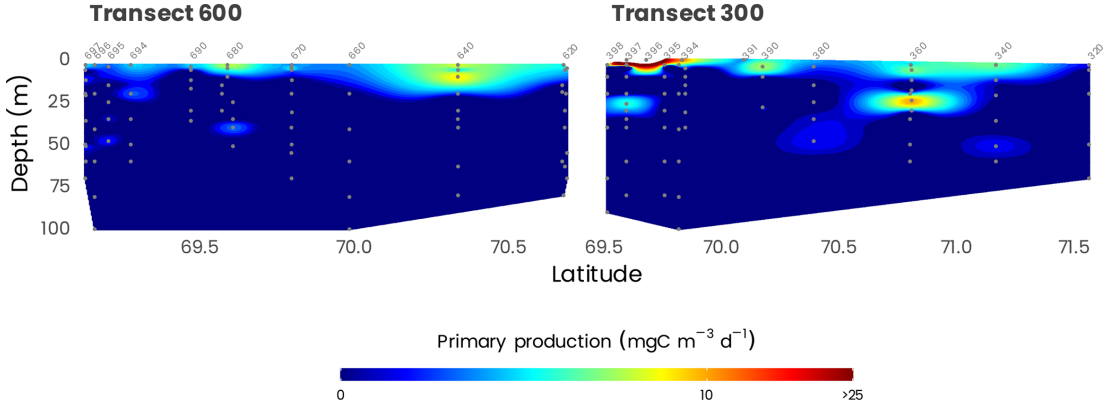

Figure 13. Cross-sections of primary production (gray dots) along transects 600 and 300. Station numbers are identified in light gray on top of each panel. Note that the color scale is presented on a log10 scale.

cells thrived mostly on regenerative NH_4^+, and their reliance on NO_3^- increased with depth to reach a maximum in the subsurface chlorophyll maximum, where substantial levels of primary production occurred (Ardyna et al., 2017a). This is consistent with Ortega-Retuerta et al. (2012a), who reported elevated bacterial abundance and bacterial production rates in association with photoammonification of riverine organic matter (Le Fouest et al., 2013). Nitrification accounted for a variable and sometimes a large share of the NO_3^- demand, consistent with the persistence of trace amounts of NO_3^- at the surface. Collectively, the data indicate that the coastal Beaufort Sea is an active regenerative system during summer, probably fueled by large pools of organic matter brought by rivers. Consequently, new production was very low and of-

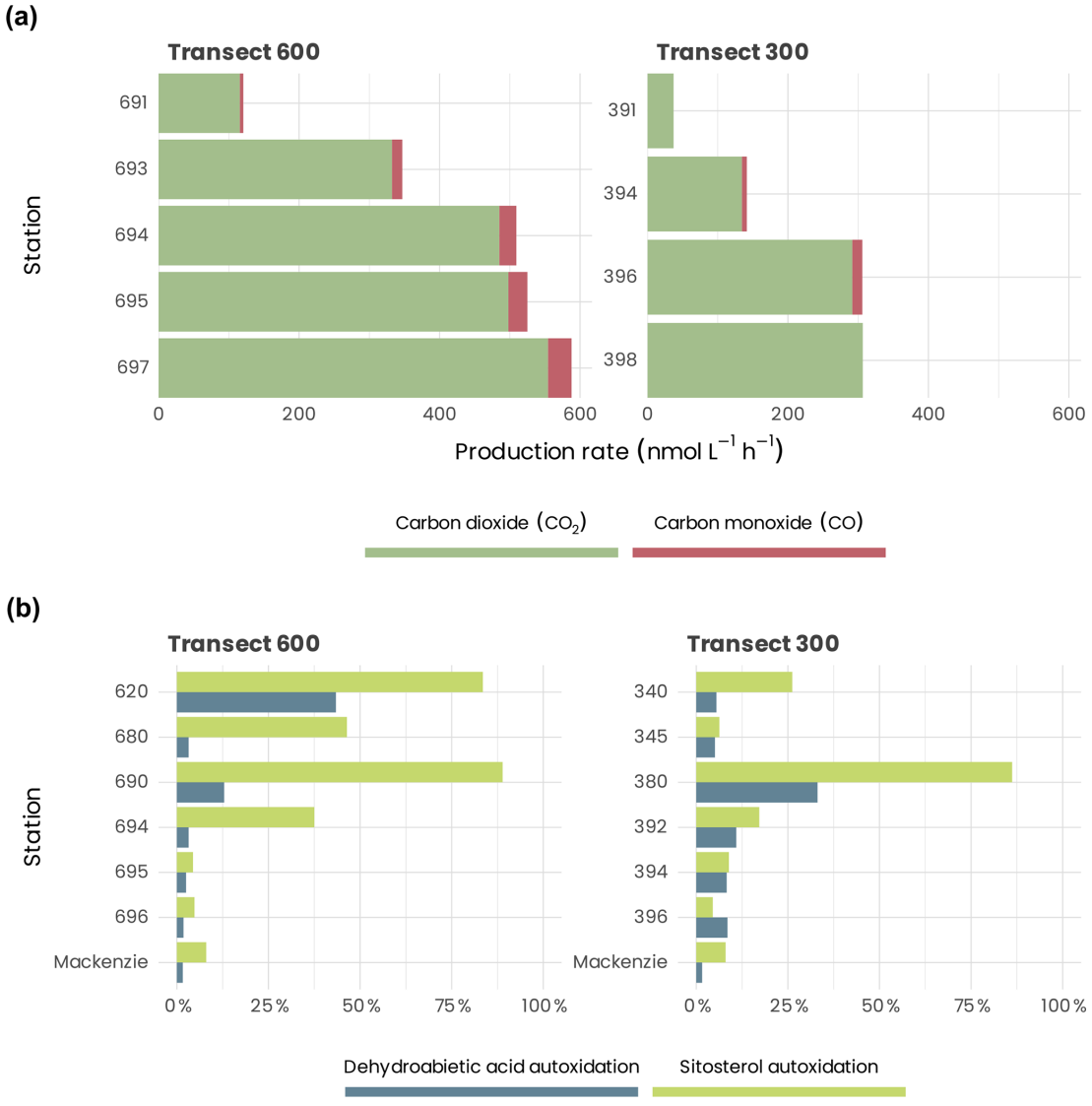

Figure 14. **(a)** CO and CO_2 production measured at 295 nm at the surface for stations of transects 600 and 300. **(b)** Autoxidation of suspended particulate material for stations of transects 600 and 300.

ten close to 0 in the 0–40 m layer. But high nitrate uptake rates can be observed at depth (Station 135), often associated with high primary production located in the chlorophyll-maximum layer, which is the place of significant new production. The impact of the Mackenzie River on shelf productivity during summer is moderate and associated mostly with localized nutrient recycling in the near-shore estuarine transition zone (Tremblay et al., 2014).

4.8.2 Photo-degradation

CO and CO_2 production from dissolved organic matter

Surface water samples were gravity-filtered upon collection through a pre-cleaned Pall AcroPak 1000 filtration capsule sequentially containing 0.8 and 0.2 µm polyethersulfone membranes. Filtered water was stored in clear glass bottles at 4 °C in darkness. CO photoproduction rates (P_{CO}, nmol L^{-1} h^{-1}) were determined aboard the CCGS *Amundsen* immediately after sample collection, whereas CO_2 photoproduction rates (P_{CO_2}, nmol L^{-1} h^{-1}) were measured in a land-based laboratory in Rimouski, Québec, within 3 months of sample collection. The sample-pretreatment and irradiation procedures followed those reported previously (Bélanger et al., 2006b; Song et al., 2013). Briefly, after minimizing the background CO and CO_2 concentrations, samples were transferred into combusted, quartz-windowed cylindrical cells (CO: interior diameter of 3.4 cm, length of 11.4 cm; CO_2: interior diameter of 2.0 cm, length of 14 cm) and irradiated at 4 °C using a SUNTEST XLS+ solar simulator equipped with a 1.5 kW xenon lamp. The radiation emitted

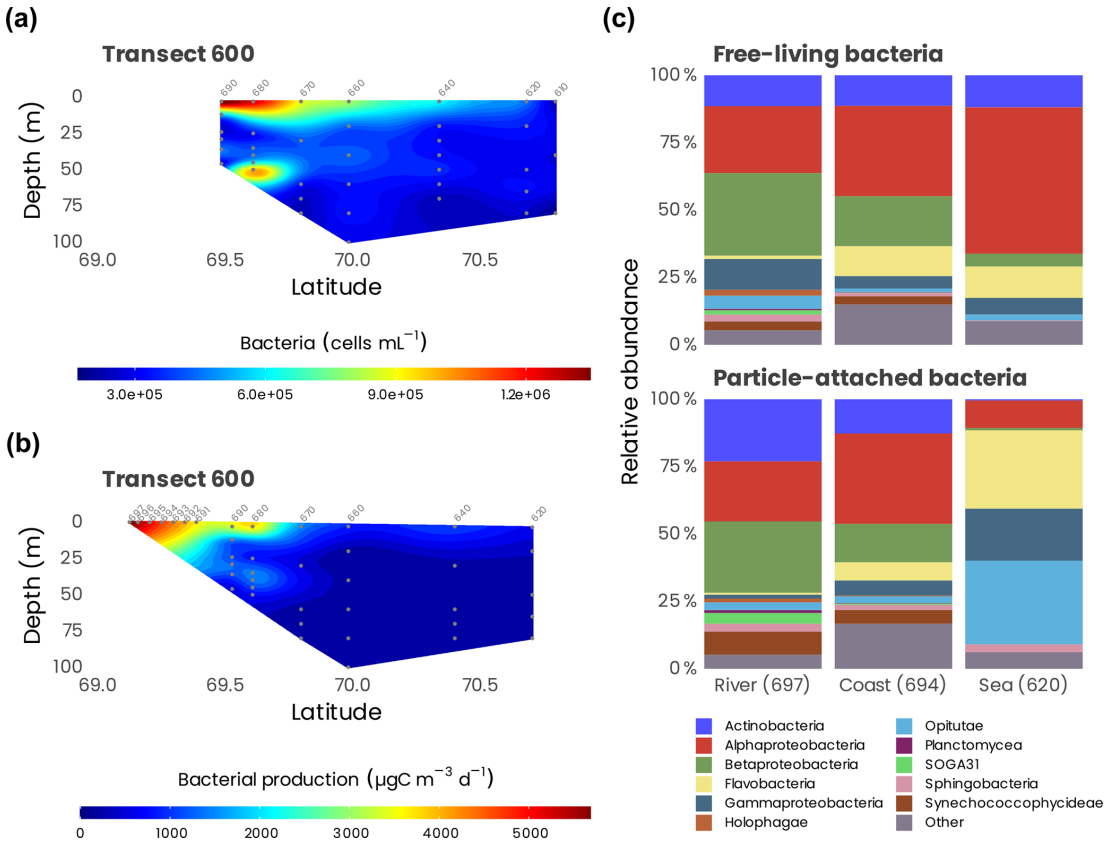

Figure 15. **(a)** Cross-sections of bacterial abundance measured from flow cytometry and **(b)** bacterial production measured along transect 600. Station numbers are identified in light gray on top of each panel. **(c)** Cumulative bar charts comparing the relative class abundances in particle-attached (PA) and free-living (FL) bacteria for a selected number of samples in transect 600.

from the solar simulator was screened with a Schott long-pass glass filter to remove UV radiation < 295 nm. The irradiations lasted for 10 min to 2 h for CO and 24 to 48 h for CO_2. The photon flux reaching the quartz windows of the cells was measured to be 835 µmol m^{-2} s^{-1} for CO and 855 µmol m^{-2} s^{-1} for CO_2 over the wavelength range from 280 to 500 nm.

Both P_{CO_2} and P_{CO} increased landward, with the difference between the most and least saline samples reaching a factor of ≈ 5 along transect 300 and ≈ 8 along transect 600 for P_{CO_2} and ≈ 7 along transect 600 for P_{CO} (Fig. 14a). This landward increase in P_{CO_2} and P_{CO} was due principally to the parallel augmentation in CDOM absorption, as demonstrated by the linear relationships between these two rates with CDOM absorption: $P_{CO_2} = 279.1 \times a_{CDOM}(412) - 17.0$ ($R^2 = 0.964, n = 9$) and $P_{CO} = 17.5 \times a_{CDOM}(412) - 4.8$ ($R^2 = 0.966, n = 7$), where $a_{CDOM}(412)$ (m^{-1}) is the CDOM absorption coefficient at 412 nm published previously (Song et al., 2013; Taalba et al., 2013). The irradiance-normalized P_{CO_2}/P_{CO} ratio gradually decreased landward along transect 600, from 23.5 at station 691 to 16.2 at station 697, suggesting that the near-shore samples were more efficient at CO photoproduction relative to CO_2 photoproduction than the shelf samples. The P_{CO_2}/P_{CO} ratios at the two stations on transect 300 were, however, similar (18.9 for station 394 and 20.1 for station 396). Combining the P_{CO_2}/P_{CO} ratios from both transects results in an average ratio of 19.8 (± 2.5 SD), with a rather small relative standard deviation of 12.5 %.

It should be pointed out that extrapolating the lab-determined CO_2 and CO photoproduction rates to the sampling area is practically infeasible due to the very different laboratory and real-environment conditions. For instance, the water column in the Mackenzie estuary and shelf areas contains large numbers of particles (Doxaran et al., 2012a), which are also optically active, whereas the irradiated samples were particle-free. Furthermore, the photoproduction rates in the water column would decrease rapidly with depth because of the strong light attenuation by CDOM and particles, while the laboratory radiation at best simulated the radiation of the top 1–2 cm layer of the water column even without considering the constant vs. varying irradiance from the solar simulator and natural sunlight, respectively. Estimating the areal photoproduction rates in the water column from lab-derived data often requires coupled optical–photochemical modeling that incorporates spectral apparent quantum yields

of the photoproduct of interest (Bélanger et al., 2006b; Xie et al., 2009; Fichot and Miller, 2010). Using this approach and CO data from the MALINA cruise, Song et al. (2013) estimated a yearly averaged areal CO photoproduction rate of 9.6 µmol m^{-2} d^{-1} in the Mackenzie estuary and shelf areas, which implies a yearly averaged areal CO_2 photoproduction rate of 191.1 µmol m^{-2} d^{-1} based on the average P_{CO_2}/P_{CO} ratio of 19.8 obtained above. Aggregating the CO_2 and CO rates gives a total photomineralization rate of 199.7 µmol C m^{-2} d^{-1}.

Autoxidation of suspended particulate material

Water samples were filtered immediately after collection through a pre-combusted glass fiber filter (Whatman GF/F, 0.7 µm) under a low vacuum. The filters were frozen immediately at −20 °C until analysis and transported to the laboratory. Treatment of the filters involved $NaBH_4$ reduction and classical alkaline hydrolysis (Rontani et al., 2012a). Reduction of labile hydroperoxides to alcohols is essential for estimating the importance of autoxidative degradation in natural samples by gas chromatography–electron impact mass spectrometry (GC–EIMS) (Marchand and Rontani, 2001). Autoxidative degradation of terrigenous particulate organic matter (POM) discharged by the Mackenzie River was monitored thanks to specific oxidation products of sitosterol (main sterol of higher plants) and dehydroabietic acid (a component of conifers).

The autoxidation state of these tracers increases strongly at the offshore stations (Fig. 14b) (reaching 89 % and 86 % at station 690 and station 380, respectively, in the case of sitosterol; see Rontani et al., 2014). These results allowed us to demonstrate that in surface waters of the Beaufort Sea, autoxidation strongly affects vascular plant lipids and probably also the other components of terrestrial organic matter (OM) delivered by the Mackenzie River. Initiation of these abiotic oxidation processes was attributed to the involvement of some enzymes producing radicals (lipoxygenases), which are present in higher plant debris and whose activity is enhanced at high salinities (Galeron et al., 2018).

Bacterial production and respiration

Bacterial production (BP; assessed by ^{3}H-leucine incubations, $n = 171$) and respiration (BR; assessed by changes in O_2 by Winkler titration, $n = 13$) were measured from the surface to 200 m at 44 sampling locations. Bacterial production ranged from 8.8 to 7078 µg C m^{-3} d^{-1} and showed a marked decreasing pattern from the mouth of the Mackenzie to the open Beaufort Sea and from the surface to deep waters (Fig. 15). Temperature and labile dissolved organic matter (indicated as dissolved amino acids) controlled BP variability (Ortega-Retuerta et al., 2012a), and the nitrogen limitation of surface BP during the summer period was demonstrated experimentally (Ortega-Retuerta et al., 2012b). BR ranged from 5500 to 45 500 µg C m^{-3} d^{-1}, leading to a bacterial growth efficiency of 8 % on average. BP and BR were low with respect to lower latitudes but within the range of those in polar ecosystems, suggesting the role of low temperatures driving carbon fluxes through bacteria (Kirchman et al., 2009). Bacterial carbon demand (BP + BR), which averaged 21 500 ± 14 900 µg C m^{-3} d^{-1}, was higher than primary production in the whole study area, indicating that the Mackenzie River platform and the Beaufort Sea are net heterotrophic during summer. This may suggest a temporal decoupling between carbon fixation and remineralization in the area.

4.8.3 Bacterial diversity

Spatial variations in bacterial-community structure were explored in surface waters from the Mackenzie River to the open Beaufort Sea ($n = 20$). By using 16S rRNA-based analysis, we investigated both particle-attached (PA; > 3 µm size fraction) and free-living (FL; size fraction between 3 and 0.2 µm) bacteria along a river-to-open-sea transect. Multivariate statistical analysis revealed significant differences in community structure between the river, coastal and open-sea waters, mainly driven by salinity, particle loads, chlorophyll *a* and amino acid concentration (Ortega-Retuerta et al., 2013). Bacterial communities differed between PA and FL fractions only at open-sea stations, likely due to the higher organic-carbon content in particles with respect to particles from the river and coast, which were enriched in minerals. Alphaproteobacteria dominated in FL open-sea samples, while the PA fraction was mainly composed of Gammaproteobacteria, Opitutae (Verrucomicrobia) and Flavobacteria. The coastal and river samples were dominated by Betaproteobacteria, Alphaproteobacteria and Actinobacteria in both the PA and FL fractions (Fig. 15c). These results highlight the importance of particle quality, a variable that is predicted to change along with global warming, in influencing bacterial-community structure and thus likely altering the biogeochemical cycles that they mediate.

5 Code and data availability

Metadata and detailed information about measurements can be found in associated MALINA papers presented in Table 1. Data are provided as a collection of comma-separated values (CSV) files that are regrouping measurements associated with a particular type of measure. To aid the user in merging these files, there is a lookup table file called stations.csv that can serve as a table to join all the data together based on date, time, station, cast, depth, longitude and latitude. Additionally, original data provided by all the researchers as well as additional metadata are available on the LEFE-CYBER repository (http://www.obs-vlfr.fr/proof/php/malina/x_datalist_1.php?xxop=malina&xxcamp=malina, last access: 8 Febru-

ary 2021). The processed and tidied version of the data is hosted at SEANOE (SEA scieNtific Open data Edition) under the CC-BY license (https://doi.org/10.17882/75345; Massicotte et al., 2020). The raw UVP5 large-particulate data and images are all available from the EcoPart and Ecotaxa website (https://ecotaxa.obs-vlfr.fr/, last access: 25 February 2021, Picheral et al., 2017). Note that Table 1 also indicates whether the measured variables are directly available in the data files or by contacting the responsible principal investigators. For specific questions, please contact the principal investigator associated with the data (see Table 1). If more data become available, they will be added to the SEANOE repository. The code used to produce the figures and the analysis presented in this paper is available under the GNU GPLv3 license (https://doi.org/10.5281/zenodo.4518943, Massicotte, 2021).

6 Conclusions

The comprehensive data set assembled during the MALINA oceanographic cruise has given unique insights into the stocks and the processes controlling carbon fluxes in the Mackenzie River and the Beaufort Sea. In this paper, only a handful of variables have been presented. The reader can find the complete list of measured variables in Table 1, all of which are also fully available in the data repository. The uniqueness and comprehensiveness of this data set offer more opportunities to reuse it for other applications.

Author contributions. MB and SB designed the MALINA project, including the scientific objectives and the sampling strategy. PM prepared the initial draft of the manuscript. SB was responsible for spectrophotometric measurement of particle absorption and fluorometric chlorophyll *a* determination onboard and with above-water radiometry (e.g., continuous incident irradiance measurements). JE was the logistic coordinator for the cruise and was responsible for IOP measurements from the CTD rosette and contributed to SPM, POC and particulate absorption measurements and processing. PF performed the GC–MS experiments on organic molecular compositions. AM and BL were in charge of determining carbonate system parameters (pH and total alkalinity) as well as the stable oxygen isotopic composition of the water and stable carbon isotopic composition of the dissolved inorganic carbon. BL participated in the cruise, collected the samples and compiled and processed most of the data. Carbonate system parameters were used to determine the saturation of waters concerning aragonite as well as calculate the surface water pCO_2. These parameters and others were used by the proponents to identify and estimate the contribution of parental waters to the structure of the water column in the study area. CM, LP and YG operated the rosette during the MALINA cruise. CM was in charge of all the physical instruments during the cruise. She also operated the UVP5 sensor. YG was in charge of the physical data processing during and after the cruise. CM and LP processed the LADCP data during the cruise. MP developed and operated the UVP5 sensor and processed the data. GG contributed to the UVP development. CJ processed samples during the MALINA cruise, designed the experiments on photoheterotrophic bacteria and wrote the related manuscripts. DB designed and performed the experiments on photoheterotrophic bacteria and wrote the related manuscripts. DV and DM participated in the MALINA cruise and sampled for flow cytometry, DNA and cultures. DM performed flow cytometry measurements and cell sorting. DV also coordinated the work on phytoplankton culture and molecular analyses. SB performed molecular analyses of flow-cytometry-sorted photosynthetic pico- and nano-eukaryotes as well as phytoplankton cultures. PG isolated and characterized phytoplankton cultures. WJ, EOR and FJ participated in the MALINA cruise, collected samples for bacterial production and respiration and prokaryotic diversity, and performed nutrient limitation experiments. RB participated in the MALINA cruise and collected the samples for DOC, TDN, THAA and TDLP9. CF analyzed the samples for DOC, TDN and TDLP6. YS analyzed the samples for THAA. RA contributed advice for the leads and the measurements of CDOM fluorescence. PG was in charge of the processing and quality control of the CTD data. CL contributed material, protocols and advice for DNA sampling for planktonic material. HJH contributed to the data interpretation regarding bacterial diversity and activity. HX and GS participated in the MALINA cruise, designed and conducted the onboard CO photoproduction experiment, and collected samples for the CO_2 photoproduction experiment. TP oversaw the measurement of meteorological elements from an instrumented tower on the ship's foredeck and the measurement of incoming solar shortwave, longwave, PAR and UV radiation from sensors mounted on top of the ship's wheelhouse. FB and YH participated in the MALINA cruise, performed the P vs. E curve and the FIRe fluorometer measurements (phytoplankton photosynthesis efficiency), and sampled for the photosynthetic protein analysis. FB also helped gather and merge the initial form of the MALINA database with the help of CAB, who coded for SQL formatting, and MHF, who coordinated the effort. DD participated in the MALINA field campaign; measured from the barge the water temperature, salinity and inherent optical properties as a function of depth; and processed the data. He was also in charge of measuring the suspended particulate matter and particulate organic carbon concentrations in collected water samples. RR and GZ participated in the MALINA field program and collected data on seawater optical properties and particle characteristics. Together with DS, they conducted processing and curation of these data, formal analysis, and publication of results. All three authors have reviewed and edited the current paper for accuracy. DA participated in MALINA Leg 1. He carried out radiometry measurements, including from the underwater radiance cameras, and then processing, analysis and publication of the results. He has contributed to editing the current paper. PA contributed material, protocols and advice of benthic sampling and participated in the interpretation of benthic biodiversity and biodiversity–ecosystem functioning through his ArcticNet funding. HL participated in the MALINA cruise; was in charge of sediment sampling with the box corer; designed and performed benthic incubation experiments onboard; analyzed sediment samples for pigments and macrofauna; processed, analyzed, interpreted and published results obtained from benthic incubations (benthic nutrient fluxes and respiration, macrofauna, TOC, pigments); and contributed to writing this paper. EL participated in the MALINA cruise and was involved in the radiance camera measurements and data processing as well as the POLVSM instrument. CS gathered

data and metadata for the MALINA cruise through INSU-CNRS (Institut National des Sciences de l'Univers – Centre National de la Recherche Scientifique) funding. MC and AT participated in the MALINA cruise; they operated the POLVSM instrument to measure the particulate scattering phase function of the water sample. BG and JaM participated in the MALINA cruise, performed sampling of suspended particles with foredeck in situ pumps and assisted CM in the deployment of the drifting sediment trap line. Both performed sample treatment and analysis of ^{234}Th, and JCM participated in the interpretation of the data and coordination of these operations prior to the cruise. IT performed data analysis and interpretation of lipid biomarkers and their carbon isotopic ratios in suspended particulate matter and sediment and sampled respectively with the foredeck in situ pumps and box corer. VM participated in the MALINA cruise, mainly on the barge to collect inherent optical properties, deploying the optical package to obtain vertical profiles. NG, PR and PC participated in the MALINA cruise by being in charge of experimental incubations for the determination of primary production and nitrogen assimilation rates as well as the preservation of samples for inorganic and organic nutrients in the euphotic zone. NG and PR performed the laboratory nutrient analysis and isotopic measurements of preserved samples, respectively. PR has compiled and processed the data and has contributed to editing the current paper. JÉT and JoM were responsible for the collection of inorganic nutrient samples in the entire water column and their analysis immediately on board (fresh). JFR carried out lipid analyses (lipid extraction, derivatization and GC–MS analyses) on SPM samples collected during the MALINA cruise, interpreted the results obtained and participated in the writing of the present paper. BC participated in the MALINA cruise, being in charge of sampling and analysis for DOC, CDOM, monosaccharides, diacids, and atmospheric underwater PAR and UV radiation. AM and AB participated in the MALINA cruise, measured CDOM absorption spectra, processed and analyzed the data, and published the results. RS defined the sampling strategy of DOC, CDOM, monosaccharides, diacids, and atmospheric underwater PAR and UV radiation; interpreted the results obtained; and wrote the related manuscripts. CP performed monosaccharide analysis and wrote the related manuscript. MV performed diacid analysis and wrote the related manuscript. SH interpreted the mass flux data and wrote the related paper. MG contributed material, equipment, protocols and advice for the determination of chlorophyll *a* and phaeopigments in the water column (fluorometric method) and participated in the interpretation of the phytoplankton biomass and community composition data. ND performed sediment trap sample preparation. AR and GM designed the MSc. project to study the evolution of sea surface conditions in the Beaufort Sea. SS performed the analysis and interpretation of sedimentary ^{210}Pbxs and ^{137}Cs profiles. SBH co-developed the C-OPS instrumentation, designed the experiments, co-developed the software to acquire the data by launching a small vessel from the much larger icebreaker and co-developed the software to process the data. HC and JR were in charge of HPLC measurements and participated in the interpretation of data. AM collected the samples for the HPLC analysis during the MALINA cruise. WLM advised on in situ optical sampling and photochemical experiments. JFG performed the data analysis on prokaryotic diversity (see the text above; together with the team WJ, EOR and FJ). FV was in charge of lipid extractions. JKV participated in the interpretation of lipid data. GC contributed materials and protocols to benthic incubations. She performed the analyses of diagenetic species and participated in the interpretation of redox conditions and benthic oxygen demand. She contributed to writing the related manuscripts. KK, KO and ET contributed to the analysis of the aerosol filter samples for organic species. JG analyzed all the nutrient samples and managed the data. JC performed the analysis of organic aerosols by GC–MS.

Competing interests. The authors declare that they have no conflict of interest.

Acknowledgements. This work is dedicated to the memory of Marc Thibault (commanding officer of the CCGS *Amundsen*, Canadian Coast Guard), Daniel Dubé (CCGS *Amundsen* helicopter pilot, Transport Canada) and Klaus Hochheim (research scientist at the Centre for Earth Observation Science, University of Manitoba), who died in the CCGS *Amundsen* helicopter crash on the evening of 9 September 2013 in the icy waters of McClure Strait in the Canadian Arctic. We are very grateful to the captain (Marc Thibault) and crews of the Canadian research icebreaker CCGS *Amundsen* during the MALINA cruise in the Beaufort Sea. The International Atomic Energy Agency is grateful to the government of the Principality of Monaco for the support provided to its Environment Laboratories. Abderrahmane Nassim Taalba performed the CO_2 photoproduction experiment. Lise Durantou carried out the analyses and interpreted the data. Claude-Anne Bouin helped gather and merge the initial form of the MALINA database. ArcticNet and Québec-Océan funded several students and analyses for their project.

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