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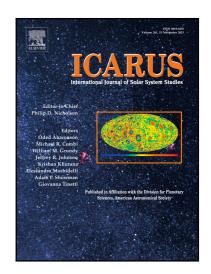
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Highlights

- A 1-D analytical model is presented to examine dusk/dawn asymmetries in tidally-locked satellites.
- The atmospheric shift towards dusk for a thermally-dependent source depends on the ratio between the rotation rate and the atmospheric loss rate
- \bullet A simple thermally-dependent ${\rm O}_2$ source is identified at Europa and possibly Ganymede.
- \bullet At Europa, the thermally-dependent O_2 source implies a large O_2 reservoir embedded in the porous ice.
- If this large O₂ reservoir exists it could be oxidizing the subsurface ocean as suggested by previous works.

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DUSK/DAWN ATMOSPHERIC ASYMMETRIES ON TIDALLY-LOCKED SATELLITES:

O₂ AT EUROPA

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ABSTRACT

We use a simple analytic model to examine the effect of the atmospheric source properties on the spatial distribution of a volatile in a surface-bounded atmosphere on a satellite that is tidally-locked to its planet. Spatial asymmetries in the O₂ exosphere of Europa observed using the Hubble Space Telescope appear to reveal on average a dusk enhancement in the near-surface ultraviolet auroral emissions. Since the hop distances in these ballistic atmospheres are small, we use a 1-D mass conservation equation to estimate the latitudinally-averaged column densities produced by suggested O₂ sources. Although spatial asymmetries in the plasma flow and in the surface properties certainly affect the spatial distribution of the near-surface aurora, the dusk enhancements at Europa can be understood using a relatively simple thermally-dependent source. Such a source is consistent with the fact that radiolytically produced O₂ permeates their porous regoliths. The size of the shift towards dusk is determined by the ratio of the rotation rate and atmospheric loss rate. A thermally-dependent source emanating from a large reservoir of O₂ permeating Europa's icy regolith is consistent with the suggestion that its subsurface ocean might be oxidized by subduction of such radiolytic products.

1. INTRODUCTION

The near-surface, far-ultraviolet oxygen aurorae, observed
for decades at Europa by the Hubble Space Telescope (Hall
tet al. (1995), Hall et al. (1998); McGrath et al. (2013) Roth
tet al. (2015)), appear to exhibit an enhancement in the emission intensity at the dusk observing longitudes as compared to
the emission intensity seen at the dawn observing longitudes
throughout Europa's orbit. A similar asymmetry was sug-

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gested by simulations at Ganymede (Leblanc et al. (2017)) yet to be clearly identified in the HST images (McGrath et al. (2013); Saur et al. (2015); Musacchio et al. (2017)). Although such asymmetries depend on the electron density and temperature, they also depend on the O₂ column density and could suggest that the O₂ column density in these thin atmospheres peak near dusk. If that is the case, this is opposite to what is seen for radiolytic argon on the Moon where peak column density occurs near dawn, due to its rapid release from the

uppermost layer of the regolith after being condensed at night 23 (Hodges & Hoffman (1974); Hodges et al. (1974); Grava et al. 24 (2015)). Molecules in these atmospheres undergo frequent 25 interactions with the surface such that the hop distances are 26 typically much less than the planetary radius, and the ballistic 27 hop time is much smaller than the volatile lifetime. Therefore, 28 the volatiles respond primarily to the local surface properties 29 and are typically assumed to be thermally accommodated to the local surface temperature. The stellar insolation of the local surface therefore should play a critical role in the shaping of the exospheres of these tidally-locked satellites, as was also recently evidenced in-situ by Cassini's INMS sampling of CO₂ & O₂ on Rhea and Dione Teolis & Waite (2016). O2 on the icy satellites of the outer planets have a far differ-36 ent origin than on Earth. These satellites are covered in water 37 ice and are embedded in a gas giant's magnetosphere. The 38 magnetic fields accelerate charged particles which bombard 39 the surface and eject O₂, H₂, H₂O, as well as their dissociation 40 products and trace species (e.g., Leblanc et al. (2017)). This 41 is often referred to as magnetospheric ion sputtering Johnson 42 et al. (1982) with the production of O₂ and H₂ a result of 43 chemical processes initiated in the ice by the incident charged 44 particles, a process also known as radiolysis (Johnson 1990; 54 45 Teolis et al. (2017) in press). Laboratory studies indicate that 55 the product yields have a strong thermal dependence when ejected from the ice which is also the case when they are absorbed and re-emitted. In this way the spatial distribution of the radiolytically produced O2 can, in principle, depend strongly on the local surface temperature. Therefore, the near-

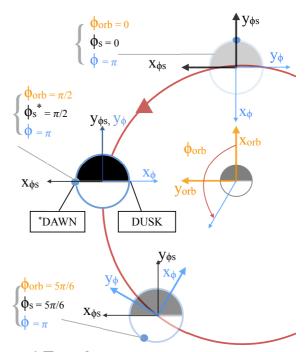


Figure 1. Satellite coordinate system for a rotating satellite at two positions with respect to its parent planet at the center. Black vectors represent the fixed, observers frame, where y_{ϕ_S} in black indicates the incoming solar radiation flux vector. ϕ_S is the anti-stellar insolation vector, defined as such in order to synchronize the two frames, where Dusk: $\phi_S = 3\pi/2$ and Dawn: $\phi_S = \pi/2$. Blue vectors represent the rotating, satellite frame where ϕ is the satellite longitude whose origin is the subplanetary point. We define the origin of the satellite system (blue circle) as $\phi_0 = \phi + \pi$ to effectively compare to observations such that the subobserver longitude is synchronized with the planetary longitude at midnight during satellite eclipse. In a time Δt , the satellite will have rotated a $\Delta \phi_{OBS} = \Omega \Delta t$.

surface O_2 atmosphere is capable of directly responding to the stellar insolation.

As the auroral process is complex, involving the production, loss, transport and excitation rates, we focus here on the role of the O_2 source rate. Therefore, we construct atmospheres using a simple analytic model for a number of possible sources driven by the stellar insolation and the magnetospheric plasma bombardment.

2. ATMOSPHERIC EVOLUTION MODEL

Here we describe the orbital evolution of the local exo-90 61 spheric column density on a tidally-locked satellite. Its mor- 91 62 phology depends on the spatial distributions of the source and 92 63 loss rates as well as on the satellite's rotation rate Ω . For the 64 satellites considered the diffusion time across the surface is 94 65 longer than the volatile's lifetime and their average thermal 95 hop distance is much less than the planetary radius simplify- 96 67 ing the analysis. For a satellite of mass m_s , radius r_s , syn-97 chronously rotating about a planet of mass M_p , a distance a_s 98 away, the rotation rate is $\Omega = (\frac{GM_p}{a_s^2})^{1/2}$ with an orbital period $\tau_{orb} = \frac{\Omega}{2\pi}$. The rotation produces both centrifugal and 100 Coriolis forces, treated in detail in 3-D Monte Carlo simu- 101 72 lations (Oza et al. (2017); Leblanc et al. (2017)), which we 102 73 initially ignore but discuss later. Figure 1 illustrates our equa- 103 74 torial coordinate system, where the stellar flux is fixed and 104 arrives from the bottom of the page, along the radiation vec- 108 76 tor y_{ϕ_s} (black vertical line). The inertial reference frame is 107 77 represented by the black vectors indicating the time in the 108 78 planet-satellite system defined by the anti-stellar insolation 109 79 angle ϕ_s , where midnight corresponds to 0 and the substel-80 lar point is at π . The sub-observer longitude ϕ_{orb} (orange 111 81 axes) is the star-planet-satellite angle and keeps track of the 112 82 satellite's rotation around the planet. Because the satellite is 113 83 phase-locked to its parent planet, it is also the longitude on 114 the body with respect to the plasma ram along the corota- 115 tion axis ϕ' (red). The satellite's rotational reference frame (blue axes) is the longitude measured counter-clockwise with its origin at the subplanetary point where $\phi = \phi' + \frac{\pi}{2}$. As we will integrate over time, we will use an origin synchro- 116

nizing these reference frames. If we begin by evaluating the anti-planetary point, $\phi = \pi$, represented by the blue dot at midnight, $\phi_{orb} = \phi_s = 0$, then the observer's clock is synchronized with the satellite's clock. Therefore, in a time interval t, the satellite rotates about its axis such that $\phi_{orb} = \Omega t$. The translation between the inertial and non-inertial reference frames is $\phi_s = \phi_{orb} + \phi - \pi = \Omega t + \phi - \pi$. Hence, after a time $\frac{\tau_{orb}}{4}$, the satellite arrives at the *sunlit* leading hemisphere corresponding to $\phi_{orb} = \pi/2$ for the observer, and $\phi_s = \pi/2$ corresponding to dawn local time as indicated in Figure 1. By rotating an additional $\tau_{orb}/6$, the fixed blue points on the satellite rotate towards substellar so that in Figure 1: $\phi_s = 5\pi/6$. Lastly, not shown in the figure is the blue point reaching the sunlit trailing hemisphere orbital longitude where $\phi_{orb} = 3\pi/2$ and $\phi_s = 3\pi/2$ corresponding to dusk local time. Since our interest is in the longitudinal variation (e.g., dusk/dawn asymmetries), we consider a latitudinally averaged column of gas, N. For a given source, the atmosphere reaches approximate steady-state after a number of orbits, such that the source flux, Φ , balances the loss flux νN . Here ν is the loss rate which can depend on the longitude ϕ through the local properties of the plasma and on the stellar flux $\Phi_T(\phi_s)$. Writing the rate of change of N as a simple balance between an atmospheric source and loss rate, the orbital evolution is determined from

$$\frac{dN(\phi,t)}{dt} = \Phi(\phi,\phi_s) - \nu(\phi,\phi_s)N(\phi,t) \tag{1}$$

As the parameter space is large, we assume that the variations in loss rate are much smaller than those in the source

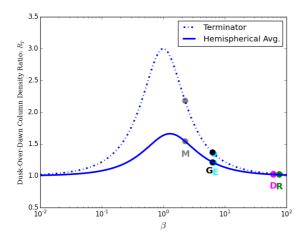


Figure 2. Dusk-over-dawn asymmetry ratio, R, versus $\beta = \frac{\Omega}{V}$ for the thermal source in Eqn. 3. For $\Omega \sim V$ the asymmetry is a maximum. The dash-dotted blue line roughly represents the terminator ratio following Eqn. 6 whereas the solid blue line is the hemispherical average of the ratio following Eqn. 7. The circles represent various satellites from Table 1: cyan & black are the Galilean satellites Europa (E) and Ganymede (G) respectively. The magenta & green points are the Saturnian satellites Dione (D) and Rhea (R) respectively. The gray point is the Moon (M) to demonstrate Lunar argon's 14 natural tendency to peak at dusk-over-dawn should there be 142 negligible condensation and diffusion.

distribution. By assuming that ν is roughly constant in space and time (e.g., Saur et al. (2011), we focus on the role of the source in determining the distribution of gas across the surface. In this case, the solution to Eq.1 is

$$N(\phi,t) = exp(-vt) \left(\int_0^t exp(vt') \Phi(\phi,\phi_s) dt' + N(\phi)_0 \right)$$
(2)

where $N(\phi)_0$ is the initial column density of the gas at time,

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t=0. In this approximation, the latitudinal distribution of the radial column density from Eqn. 2 is determined by primarily the source, Φ .

An asymmetry in N between the dusk and dawn terminators ($\phi_s = 3\pi/2$; $\pi/2$ respectively) can be shown by the ratio of their column densities, R. For R>1, N is larger toward dusk. Since the difference in the scale heights is small, R is roughly proportional to the line-of-sight (LOS) column densities at the terminators. Because remotely observed LOS emission intensities probe the sunlit hemispheres, we also compute a hemispherically-averaged ratio, $\langle R \rangle$, ignoring the small contribution at the terminators from the nightside atmosphere. Averaging N across longitude on the dusk quadrant, $\phi_s = \pi \rightarrow 3\pi/2$, $\langle N_{DUSK} \rangle$, and the dawn quadrant, $\phi_s = \pi \rightarrow \pi/2$ $\langle N_{DAWN} \rangle$ we calculate a ratio that can be more readily compared to the observations.

Below we use three source functions that have been discussed. For each we derive the column density as a function of ϕ and t, first computing R, the dusk/dawn ratio and then the hemispherically averaged ratio, $\langle R \rangle$. These are then discussed based on the observed auroral emission ratios.

2.1. Solar Radiation-Driven Source

We first consider a source solely dependent on the stellar insolation. As an example, we construct a simple sublimation-like source that only depends on ϕ_s and peaks at noon

$$\Phi_T(\phi_s) = \frac{\Phi_0}{2} \left(1 - \cos(\phi_s) \right) \tag{3}$$

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where Φ_0 is the maximum flux. Substituting into and integrating Eqn. 2 we obtain

$$N_T(\phi, t) = \frac{\Phi_0}{2\nu} \left[[1 - exp(-\nu t)] - N_{rot}(\phi_s, t) \right]$$
 (4)

Writing
$$\beta = \Omega/\nu$$
 and $\alpha = \phi - \pi$, then $N_{rot}(\phi_s, t) = (1 + \frac{1}{148} \beta^2)^{-1} ([cos(\phi_s) + \beta sin(\phi_s)] - [cos(\alpha) + \beta sin(\alpha)]exp(-\nu t))$.

As $t \to \infty$, the spatial distribution of the steady-state column density is seen to depend solely on the stellar insolation angle ϕ_s giving

$$N_T(\phi) = \frac{\Phi_0}{2\nu} \left(1 - (1 + \beta^2)^{-1} [\cos(\phi_s) + \beta \sin(\phi_s)] \right)$$
 (5)

The ratio R_T , the dusk-over-dawn column densities evaluated at the terminators, is

$$R_T \sim \frac{1+\beta+\beta^2}{1-\beta+\beta^2}$$
 (6) 180

corresponding to the dashed blue curve shown in Figure 2. 154 Because Φ_0 cancels, it is seen that for such a source with 155 roughly constant loss rate, R_T is greater than unity and de-156 pends only on the ratio β . In this approximation the maxi-157 mum column density occurs past noon simply due to the rotation of the source peak towards dusk (e.g., Hodges & Johnson (1968)). As seen in Oza et al. (2017) for Europa, the peak 185 temperature actually occurs just past noon so that the shift to- 186 wards the dusk terminator would be somewhat larger. 162 The dusk/dawn ratio averaged over the the half-163

hemispheres is

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$$\langle R_T \rangle = \frac{1 + \frac{2}{\pi}(\beta + 1) + \beta^2}{1 - \frac{2}{\pi}(\beta - 1) + \beta^2}$$
 (7) 189

which is given by the solid blue curve in Figure 2. The values of β in Table 1 for the various satellites are indicated by colored dots. It can be seen that for the assumed source and loss process, the atmospheres of Europa and Ganymede would possess column densities which are at least 50 % thicker at dusk than at dawn. For the Saturnian satellites the loss rates are roughly a factor of 10 smaller (Table 3; Teolis & Waite (2016)) consistent with the absence of a dusk-over-dawn asymmetry.

2.2. Magnetospherically-Driven Source

Icy satellite O₂ exospheres embedded in large planetary magnetospheres are thought to be generated by plasma flow along the satellite's orbit. Such a source depends on the satellite corotation direction. Since the plasma has a thermal and an energetic component, and the ions have gyromotion determined by their energy and the local fields, how strongly the O₂ source peaks at plasma ram direction has been discussed (e.g., Cassidy et al. (2013).; Teolis et al. (2005)). Here we only assume the source rate is a function of the angle from the corotation and can be approximated as:

$$\Phi(\phi)_{mag} = \frac{\Phi_0}{2} \left(1 + \sin(\phi) \right) \tag{8}$$

Again assuming a nearly constant loss rate, Eqn. 2 yields, not surprisingly, an expression with no rotational term

$$N_{mag}(\phi) = \frac{\Phi(\phi)_{mag}}{v} \left(\left[1 - exp(-vt) \right] \right)$$
 (9)

Therefore, rather than an enhancement at dusk over the full orbit, the dusk-over-dawn ratio oscillates with an orbit average R of about one.

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S	τ_{orb} [hrs]	$\tau_i \text{ [hrs] } (\tau_{orb})$	β	R_T	$\langle R_T \rangle$
Е	85	90 (1.06)	6.78	1.34	1.20
G	172	170 (1.0)	6.28	1.37	1.22
D	66	604 (9.2)	61	1.03	1.02
R	108	1262 (12)	77	1.02	1.02
М	656	2669 (4.1)	2.3	2.18	1.54

Table 1. Atmospheric evolution parameters and values for 199 the various surface-bounded satellite exospheres described in Figure 2. τ_{orb} is the orbital period in hours, τ_i the atmospheric lifetime of the species i in question, scaled to the orbital period in parentheses. $\beta = \frac{\Omega}{\nu}$ a parameter indicating the magnitude of the asymmetry, depending only on the rotational rate of the satellite Ω and the atmospheric loss rate ν . The loss rates for Europa and Ganymede are calculated from photon and electron ionization and dissociation rates provided in Turc et al. (2014) Table 2, for an isotropic electron density of $n_e = 70cm^{-3}$ (e.g. Marconi (2007)). Dione and Rhea's O₂ loss rates are estimated by Teolis & Waite (2016) Table 3, whereas the Lunar argon loss rate Grava et al. (2015) Table 1. R_T and $\langle R_T \rangle$ are the latitudinally averaged dusk-overdawn ratios for a satellite with a radiation-driven source estimated using Eqn. 6 (evaluated at the terminator) and Eqn. 7 (hemispherically-averaged) respectively.

2.3. Solar Radiation & Magnetosphere-Driven Source

Laboratory data indicate that the plasma-induced source of O_2 depends on the ice temperature. Although it has been ²¹³ argued that on the icy satellites this dependence is averaged ²¹⁴ out due to delayed emission of the O_2 (Teolis et al. (2005)), ²¹⁵ a predominately trailing hemisphere source $\Phi(\phi)_{mag}$ with a ²¹⁶

thermal enhancement has been used often (e.g., Cassidy et al. (2007); Plainaki et al. (2012); Plainaki et al. (2013); Oza et al. (2017)). Here we roughly represent such a source by:

$$\Phi(\phi, \phi_s)_{Tmag} = \frac{\Phi_0}{2} (1 + \sin(\phi)) * [3 - \cos(\phi_s)]/2$$
 (10)

This results in a steady-state column

$$N_{Tmag}(\phi) = \frac{\Phi(\phi)_{mag}}{2\nu} \left(3 - \frac{1}{1 + \beta^2} (\cos(\phi_s) + \beta \sin(\phi_s))) \right)$$
(11)

Based on the above expression, the rotation shifts the maximum and minimum column densties from noon and midnight by $tan^{-1}(\beta)$. For Europa and Ganymede, with β values near 2π , this shift corresponds to $\sim 81^{\circ}$. In satellite local time, the column increases from its minimum value near dawn, $\sim 6h30$, to its maximum column near dusk, $\sim 17h30$. The magnitude of the dusk-over-dawn asymmetry at the terminators is

$$R_{Tmag} = \frac{1 + \beta/3 + \beta^2}{1 - \beta/3 + \beta^2} \tag{12}$$

The ratio of the averaged column densities, $\langle R \rangle$, results in a more complex expression and is given in Table 2 at a few orbital positions. It is seen that although R_{Tmag} in Eqn. 12 is greater than one, the averaged ratio is not always greater than one and varies significantly from the sunlit leading hemisphere to the sunlit trailing hemisphere.

3. ICY SATELLITE O₂ ATMOSPHERES

Although variations on these overly simplified source terms have been used, here we compare the predicted ratios. We start by considering Europa for which its near-surface

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oxygen aurorae observations are the most complete. Since 217 the O₂ exosphere is produced radiolytically by the incident 218 plasma, uncertainties regarding the production, loss, and 219 excitation rates, as well as the O₂ residence in the regolith, 220 have been an obstacle in describing the observed emission 221 asymmetries. Focusing here on the source process, values of 222 R and < R > are given in Table 2 for the above three cases at 223 four orbital positions: sunlit trailing eclipse, sunlit leading, and substellar. Also shown are the ratios obtained from the HST observations.

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These results are also shown in Figure 3 where they are 228 compared to the HST observations (black points) in order 229 to illustrate the effect of the source process on the near sur-230 face aurora. Also shown are the Monte Carlo simulations of 231 Europa's exosphere from Oza et al. (2017) for a thermally-232 dependent sputter source (EGM sim.; blue points). Such sim-233 ulations account for diffusive hopping as well as the centrifu-234 gal and Coriolis forces. It is seen that the observed dusk/ 236 dawn emission ratio does not vary significantly over an orbit. 237 In contrast to that, the simulations in Oza et al. (2017) and 238 the magnetosphere models $(\langle R \rangle_{mag}; \langle R \rangle_{Tmag})$ differ signifi-239 cantly from the observed emission ratio. It is also clear that a 240 temperature-independent magnetospheric sputter source, suggested by the delay times in the O₂ emission (Teolis et al. 242 (2017)) and roughly approximated here by Eqn. 8, would 247 not by itself give the observed dusk/dawn ratio. The lack 248 244 of agreement of the observations with detailed simulations in 249 Oza et al. (2017) and with the approximate magnetospheric 250

ϕ_{orb}	$\langle R(T)_{EGM} \rangle$	$\langle R \rangle_{mag}$	$\langle R(T) \rangle_{Tmag}$	$\langle R_{HST} \rangle$
270	1.58	1.0	1.06	1.64 ±0.14
0	1.35	4.5	4.8	1.73 ±0.7;
				1.53 ±0.2
90	6.87	1.0	1.09	1.29 ±0.1
180	6.06	0.22	0.24	1.57 ±0.2;
				1.92 ±0.2

Table 2. Dusk/Dawn ratios of atmospheric O₂ bulges on Europa, calculated as hemispherically averaged column-density ratios $\langle R \rangle \sim N_{Dusk}/N_{Dawn}$ over the four major orbital phases: sunlit trailing (270), eclipse (0), sunlit leading (90), and substellar (180). The fixed sublimation source maintains a constant dusk-over-dawn asymmetry of: $\langle R_T \rangle = 1.2$ for all orbital longitudes. The four R columns represent the latter two magnetospheric plasma cases in section §2.2, and 2.3. The second column is the EGM output of Oza et al. (2017) for a non-adsorbing, thermally-dependent O2 case. The final column is the average auroral intensity ratio, I_{Dusk}/I_{Dawn} , observed by HST at 1356 Å potentially indicative of the exospheric asymmetry. The HST eclipse values are averaged around a $\sim 23^{\circ}$ interval in orbital longitude, whereas the substellar values are averaged around a larger ~ 100° interval. For these latter two cases we provide pre and post transit values.

sources suggests that the direct use of the laboratory measurements could be problematic. The comparison with the simulations improves somewhat by accounting for the O_2 residence time in Europa's regolith (Oza et al. (2017)). And, of course, significant spatial variations in the plasma-induced loss and in

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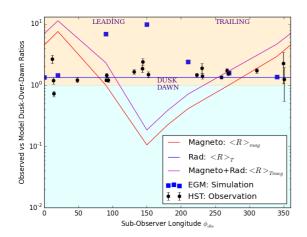


Figure 3. Average dusk-over-dawn asymmetry ratio, $\langle R \rangle$, versus sub-observer longitude ϕ_{orb} indicating dusk enhancements shaded in orange, and dawn enhancements in blue. The 3-D Monte Carlo exosphere simulations from Oza et al. (2017) representing a non-adsorbing O_2 population (blue squares) as well as the oxygen emission data from Roth et al. (2015) by HST (black points) dominate the dusk-enhanced region. Our three source cases: radiation-driven $(\langle R \rangle_T)$, magnetospherically-driven $(\langle R \rangle_{mag})$, and a temperature-enhanced version of the radiation & magnetospherically-driven $(\langle R \rangle_{mag})$ O₂ sources are represented by the solid blue, red, and magenta lines respectively. It can be seen that only the stellar radiation-driven O₂ source provides a reasonable fit to the HST observations.

the excitation rates could further improve these comparisons.

What is striking about the comparison in Fig. 3 is that the
ratio obtained for the hemispherically averaged O₂ sublimation source (blue line), as in Eqn. 3, gives a result that is surprisingly close to that observed. Since the O₂ is beyond doubt
produced by the impacting plasma this comparison could suggest that there is indeed a temperature-dependent source rate.

However, the rate is not like that modeled by a direct use of the laboratory data, roughly approximated here in Eqn. 10 or treated in much more detail in the simulations. Since the residence time in the regolith of the returning O₂ clearly has an effect Oza et al. (2017), we suggest the agreement is likely due to the fact that the O₂ produced over geological time periods permeates the regolith on Europa (e.g., Johnson et al. (2003); Hand et al. (2007); Greenberg (2010); Teolis & Waite (2016)). If that is the case, then the direct production rate, whether enhanced on the trailing hemisphere or thermallydependent, contributes marginally per orbit to a large reservoir of O2 bound in the regolith. This O2 likely exhibits a rough vapor pressure equilibrium determined by its binding in the porous icy regolith. Magnetospheric-ion production acting over long time periods should result in a significant amount of O₂ trapped in the ice which is able to thermally diffuse and populate the exosphere. In this model the maximum in the O₂ atmospheric source is shifted by the satellite rotation generating an asymmetric atmosphere that peaks toward dusk.

The highly simplified model presented here does not account for O₂ transport, although it has been included in the 3-D exosphere simulations (e.g., Leblanc et al. (2017); Oza et al. (2017)). Latitudinal transport is strongly suggested by the observation of equatorial O₂ bulges throughout Ganymede's orbit (Leblanc et al. (2017)). That is, O₂ is produced primarily in the polar regions (McGrath et al. (2004)) but is found to have migrated towards the equator on orbital timescales (Leblanc et al. (2017)). This latitudinal motion

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is driven by the centrifugal force treated in the simulations $_{317}$ but ignored here. The primary effect of this transport is very $_{318}$ roughly accounted for here by latitudinally-averaging the O_2 $_{319}$ column densities.

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4. CONCLUSION & SUMMARY

Because numerous processes affect the auroral observations 324 294 of the exospheres at Europa and Ganymede, we used a simple analytic model to focus on the asymmetry produced only by the atmospheric source. In doing this we assumed that the 327 effect of the spatial asymmetries in both the plasma-induced loss and the auroral excitation rates are much smaller than 399 those produced by the source. Although this is a significant 330 300 assumption, at present it does not contradict models of ion-301 ization and excitation rates which depend on both the local 302 electron density and temperature. In addition, we ignored 333 303 were small 334 thermal transport, as the molecular hop distances 304 compared to the satellite radius so that the latitudinally aver-305 aged column is very roughly synchronized with the surface 336 306 rotation. It is seen that for Europa a plasma-induced source 337 307 enhanced on the trailing hemisphere, whether thermally de-308 pendent or not, does not produce spatial asymmetries consistent with the emission observations. That comparison could, 340 of course, be considerably improved by the presence of a significant spatial variation in the plasma-induced loss and emission rates. However, the orbital dependence of the half hemi-313 sphere average of the dusk/dawn emission ratio at Europa is 344 seen to compare favorably with a simple thermal dependent, 345 315 solar heating source on these tidally locked icy satellites. This

might also be the case at Ganymede, although the observations are less extensive and its magnetic field can complicate the comparison. The magnitude of the shift towards dusk produced by such a source was seen to depend on the ratio of the rotation rate to the loss rate, in addition to any shift in the thermal peak towards dusk as seen in Fig. 5 in Oza et al. (2017). Although the comparison could be fortuitous we note that such a source is not unreasonable. That is, the O₂ produced over long time periods likely permeates the porous icy regolith, even to the point where it has been suggested that it is a viable source of O2 for Europa's ocean (e.g. Johnson et al. (2003); Hand et al. (2007); Vance et al. (2016)), so that the daily varying production is a small fraction of the available O2. Therefore, we suggest that the thermally desorbed O₂ populates the atmosphere at a rate that primarily depends on the local temperature and its binding properties in these porous regoliths. Since the well known hemispherical differences in composition and albedo affect both the surface temperature and the binding we suggest these are secondary effects, a point that needs further testing. Of course, spatial differences in the local source rate will exist at smaller spatial scales than that considered due to the variation in the local surface properties and indeed, in the absence of averaging over the half hemispheres, the observations indicate such spatial variations. While the mechanism that generates the O₂ atmosphere on the icy satellites has been identified, namely plasma-driven production, a good simulation of the orbital observations is still not available. Here we suggest that lack of agreement might be due to the assumptions about the na-

ture of the O2 source. If the observed O2 is indeed due to the 365 Hand, K. P., Carlson, R. W., & Chyba, C. F. 2007, Astrobiology, 7, 1006 346 Hodges, R. R., Hoffman, J. H., & Johnson, F. S. 1974, Icarus, 21, 415 Hodges, Jr., R. R., & Hoffman, J. H. 1974, in Lunar and Planetary Science presence of a large reservoir, the mechanism suggested for 367 347 Conference Proceedings, Vol. 5, Lunar and Planetary Science Conference Proceedings, 2955 oxidizing Europa's subsurface ocean, Johnson et al. (2003), 348 Hodges, Jr., R. R., & Johnson, F. S. 1968, J. Geophys. Res., 73, 7307 Johnson, R. E. 1990, Energetic Charged-Particle Interactions with Hand et al. (2007) may be more credible. 349 372 Atmospheres and Surfaces, 84 Johnson, R. E., Lanzerotti, L. J., & Brown, W. L. 1982, Nuclear Instruments 373 and Methods, 198, 147 374 Johnson, R. E., Quickenden, T. I., Cooper, P. D., McKinley, A. J., & Freeman, C. G. 2003, Astrobiology, 3, 823 Acknowledgments: The authors acknowledge helpful 376 350 Leblanc, F., Oza, A. V., Leclercq, L., et al. 2017, carus Marconi, M. L. 2007, Icarus, 190, 155 comments from Dr. T.A. Cassidy. AVO and FL would like 351 McGrath, M. A., Jia, X., Retherford, K., et al. 2013, Journal of Geophysical Research (Space Physics), 118, 2043 to acknowledge LabEx/ESEP for their support, as well as the 381 McGrath, M. A., Lellouch, E., Strobel, D. F., Feldman, P. D., & Johnson, 352 R. E. 2004, Satellite atmospheres, ed. F. Bagenal, T. E. Dowling, & W. B. CNES Système Solaire program. REJ acknowledges support 383 McKinnon 457 353 Musacchio, F., Saur, J., Roth, L., et al. 2017, Journal of Geophysical Research (Space Physics 122, 28 385 from NASA's Planetary Data Systems Program. Oza, A., Leblanc, F., Johnson, R., et al. 2017, submitted Plainaki, C., Milillo, A., Mura, A., et al. 2012, Icarus, 218, 956 387 . 2013, Planet Roth, L., Saur, J., Retherford, K. D., et al. 2015, J. Geophys, 261, 1 389 REFERENCES Saur, J., Feldman, P. D., Roth, L., et al. 2011, ApJ, 738, 153 391 Saur, J., Duling, S., Roth, L., et al. 2015, Journal of Geophysical Research sics), 120, 1715 ace Phy Teolis, B., Plainaki, C., Cassidy, T., & Raut, U. 2017, in Press 355 Cassidy, T. A., Johnson, R. E., McGrath, M. A., Wong, M. C., & Cooper, Teolis, B. D., Vidal, R. A., Shi, J., & Baragiola, R. A. 2005, Phys. Rev. B, 394 J. F. 2007, Icarus, 191, 755 356 245422 Cassidy, T. A., Paranicas, C. P., Shirley, J. H., et al. 2013, Planet. Space Sci. 357 358 Teolis, B. D., & Waite, J. H. 2016, Icarus, 272, 277 Grava, C., Chaufray, J.-Y., Retherford, K. D., et al. 2015, Icarus, 255, 135 359 Turc, L., Leclercq, L., Leblanc, F., Modolo, R., & Chaufray, J.-Y. 2014, 360 Greenberg, R. 2010, Astrobiology, 10, 275 Icarus, 229, 157 Hall, D. T., Feldman, P. D., McGrath, M. A., & Strobel, D. F. 1998, Application of the control o 361 Vance, S. D., Hand, K. P., & Pappalardo, R. T. 2016, Geophys. Res. Lett., 362 Hall, D. T., Strobel, D. F., Feldman, P. D., McGrath, M. A., & Weaver, H. A. 43, 4871 363 364 1995, Nature, 373, 677