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The effect of tile-drainage on nitrous oxide emissions from soils and drainage streams in a cropped landscape in Central France

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ABSTRACT

Tile drainage may have contrasting effects on soil nitrous oxide (N₂O) emission. Because drainage decreases anoxic periods in soils, it could reduce N₂O production via denitrification and also limit the reduction of N₂O into nitrogen gas (N₂). Moreover, drainage accelerates the discharge of water enriched in dissolved N₂O and mineral nitrogen. Thus, nitrogen losses and N₂O releases from discharged surface water need to be quantified to assess the total effect of drainage on N₂O emissions. Thus, the objectives of this study were two-fold: (1) to assess the effect of tile-drainage on soil N₂O emissions in an agricultural area in Central France (direct emissions) and (2) to compare emissions from soils and from the stream draining the area (indirect emissions). The emissions of N₂O by soils were measured using static chambers in two drained and two undrained cereal plots over two growing seasons. A rule-based model was fitted to identify the influence of drainage and ancillary variables. Stream N₂O emissions were measured with a floating chamber during one growing season. The mean direct N₂O emissions were 0.071 mg N m⁻² h⁻¹ and were larger in the undrained plots than in the drained plots in both growing seasons ($p < 0.001$). The rule-based model showed that the drainage effect on N₂O emissions was dominant over the permanent soil variables. The mean stream N₂O emissions were 0.190 mg N m⁻² h⁻¹. The surface water emissions represented 31 kg N during the flow period (7 months) while direct emissions were 1846 kg N during the same period. Thus, indirect emissions accounted for <2% of the total N₂O emissions in the study site. While tile-drainage did not result in significant indirect emissions at this local site scale, it was identified as the dominant factor controlling the direct soil N₂O emissions. Thus, drainage should be taken into account in greenhouse gas emission inventories for larger areas.

Keywords:

Greenhouse gas
Artificial drainage
Indirect emission
Surface water

1. Introduction

Artificial drainage of hydromorphic soils is common in cropped areas to remove excess water from soils. Drainage has been installed in approximately 11% of France's agricultural areas, with 23% of the drained areas in the Centre region (Agreste 2012AR26, recensement agricole 2010). Soil nitrous oxide (N₂O) emissions are highly dependent on soil aeration status. N₂O is produced in soil by different microbial processes, mainly nitrification, i.e., oxidation of ammonium (NH₄⁺) to nitrite (NO₂⁻) and then to nitrate (NO₃⁻), and denitrification, i.e., reduction of nitrate to N₂O and ultimately to nitrogen gas (N₂) (Smith et al., 2003). Anoxia promotes

denitrification, which may dominate N₂O production in the soils (Dobbie and Smith, 2001). Thus, installation of artificial drainage, which decreases anoxic periods, is expected to decrease the total N₂O emission (Bouwman, 1996; Dobbie and Smith, 2006).

However, maximum N₂O emission rates are generally observed for water-filled pore space (WFPS) <90% (Davidson and Verchot, 2000; Castellano et al., 2010), i.e., for unsaturated soils, because of the slower diffusion of gas and the possible further reduction of N₂O at higher WFPS (Smith et al., 2003). This may explain why larger N₂O emissions have been measured in drained soils than in undrained soils under grassland (van Beek et al., 2010) or under forest (von Arnold et al., 2005; Jungkunst and Fielder, 2007). A review of the N₂O emissions studies from German soils has also suggested that under cultivation, regularly water-logged soils showed lower emissions than well-aerated soils (Jungkunst et al., 2006). Direct comparisons of N₂O emissions from drained and

undrained cropped plots are scarce, but those that are available suggest contrasting effects of drainage, with larger emissions in undrained soils in dry periods but smaller emissions in wet periods (Venterea et al., 2008; Colbourn and Harper, 1987). Thus, drainage may have contrasting effects on local N₂O emissions from soils depending on climatic conditions and the hydric history of the soils, and it is important to assess these overall effects.

Determining controls on soil N₂O emissions is often difficult in field studies because of the many factors involved and the complex interactions between these factors. N₂O production, at the microsite scale, is controlled by nitrogen, carbon substrates and oxygen availability, soil temperature and pH (Stehfest and Bouwman, 2006; Saggar et al., 2013). At larger scales, N₂O emissions depend on distal factors (Williams et al., 1992) that influence the local factors, such as soil type (Van Groenigen et al., 2004), top soil texture (Skiba and Ball, 2002; Stehfest and Bouwman, 2006), topographic attributes (Vilain et al., 2010) and soil drainage class (Bouwman et al., 2002). Thus, ideally, it is necessary to assess the influence of the climate, crop management and soil factors to distinguish the effect of a single factor. Drainage may interact with other distal factors, especially because it deeply modifies soil hydrodynamics and a comprehensive understanding of the control of N₂O emissions in both drained and undrained situations may be useful for proposing strategies to mitigate the emissions of greenhouse gases.

Moreover, drainage can also have an effect at larger scales because of possible nitrogen (N) transfer by water. Drainage accelerates water discharge, and this water may entrain nitrate and dissolved N₂O, i.e., there could be pollution swapping with an increase of indirect N₂O emissions from surface water. Large N₂O emissions have been measured in drainage water (Reay et al., 2009; Beaulieu et al., 2009) and several processes have been identified to explain these indirect emissions: leaching of dissolved N₂O produced in subsoil by denitrification (Reay et al., 2009), or leaching of nitrate followed by direct N₂O production via denitrification in surface water, such as drainage ditches, streams and rivers in agriculture landscapes (Beaulieu et al., 2009; Boehlke et al., 2009; Garnier et al., 2010). Compared to direct N₂O emissions, the relative importance of indirect N₂O emissions from streams is still very uncertain; while some studies report that they represent a small fraction of the total N₂O emissions (Reay et al., 2009; Vilain et al., 2012), they may also contribute significantly in other regions (Outram and Hiscock, 2012). In a recent study, Turner et al. (2015) showed that indirect N₂O emissions represented 32% of the total N₂O emission in the US Corn Belt and that headwater streams dominated this indirect contribution (60%). This study also highlighted the severe lack of information on N₂O emissions from non-permanent water flows, such as small drainage channels; taking these sources into account may double the total N₂O emissions when compared to the estimations obtained from applying the standard Tier I IPCC methodology, which relies on a constant proportion of nitrogen supply emitted as N₂O (Turner et al., 2015). This points to the need for new measurements in large agricultural areas around the world. Thus, assessing the overall impact of artificial drainage on N₂O emissions requires determining whether drainage enhances indirect emissions. At present, answering this question is extremely complex: dissolved N₂O measured in underground monitoring wells may further react before discharge depending on the residence time of underground water and the speed of denitrification (Well and Butterbach-Bahl, 2010), and estimating emissions from dissolved N₂O in drainage water and gas exchange models can lead to an underestimation of emissions (Turner et al., 2015). Last, as water from field drains is generally conducted to collectors before reaching the atmosphere, this integrates N₂O emissions from large surfaces and requires assessments at the landscape scale.

The objective of this work was to study the effect of artificial drainage on direct and indirect N₂O emission from soils within a small agricultural region (~20 km²) with naturally hydromorphic soils (Gu et al., 2011, 2013). Two questions were assessed. First, the influence of drainage as well as ancillary variables (climate conditions, soils properties) on direct N₂O emissions was assessed by comparing the N₂O emissions in drained and undrained plots. Second, as drainage is extensively used in the studied region, the indirect N₂O emissions from the non-permanent stream that drains this region were investigated to evaluate the potential for pollution swapping induced by drainage and to assess the indirect N₂O emissions as a proportion of the total N₂O emissions.

2. Materials and methods

2.1. Study site

The study was carried out within an experimental site (48°23'N, 1°11'E; elevation 202 m a.s.l) in the Loir River valley, approximately 120 km southwest of Paris, France. The climate records (1971–2000) in the closest meteorological station in Chartres (48°27'N, 1°30'E, elevation 155 m a.s.l.) showed a mean annual temperature of 10.6° C, precipitation of 598 mm and potential evaporation of 740 mm. The study area (20 km²) is cultivated by commercial farms (not for scientific goals) and dominated by winter wheat (*Triticum aestivum*), barley (*Hordeum vulgare*) and rapeseed (*Brassica napus*) crops. The soils are LUVISOLS (WRB, 2015) with a loamy texture and poor natural drainage. However, tile-pipes have been installed in most fields to attenuate soil hydromorphy. Streams and main drainage ditches represent only a small surface of the studied site (approximately 0.1%). Measurements of the direct N₂O emissions were made during two growing seasons, from seeding to harvest (typically Nov–July), in 2010–2011 and in 2012–2013. The direct N₂O emissions were measured on two undrained plots (referred to as ND) and two drained plots (referred to as D) during these two periods. ND plots were situated at the shoulder and the footslope of a very gently sloping field (slope 1.6%) at a 208 m distance from each other. As climatic conditions are considered an influencing factor for N₂O emissions, these two replicates will be hereafter referred to as ND1 (shoulder) and ND2 (footslope) in 2010–2011 and ND3 (shoulder) and ND4 (footslope) in 2012–2013 (Fig. 1). To maintain the same management practices as for the ND plots, the locations of the D plots were changed from the first year (referred as D1 and D2) to the second year (referred as D3 and D4). This is because the studied plots were managed by commercial farms, and crop management for D1 and D2 was not the same as that of ND1 and ND2 during the second growing period. The crops were winter barley in 2010–2011 (total fertilization 140 kg N ha⁻¹ during spring in D1, D2, ND1 and ND2) and winter wheat in 2012–2013 (total fertilization 166 kg N ha⁻¹ in D3, D4, ND3 and ND4). All plots were tilled down to 30 cm depth with burial of previous crop residues. Details on the management practices are included in the supplementary material.

The experimental site was very close to the Loir headwater, where the stream is still non-permanent. The indirect N₂O emissions from the surface water were thus measured during the flow periods (Nov–May) during the second growing season 2012–2013. The water drained from the D3 and D4 plots was routed to a main under-ground pipe reaching the surface at a distance of 1 km from the field where the collector discharged into a drainage ditch connecting to the Loir stream. To determine appropriate sites for the indirect N₂O measurements, prospective measurements were made at a few dates in the drainage ditches and in the Loir stream channel (acting itself as a drainage ditch). The drainage ditches and the stream channel showed very similar N₂O emissions. Thus, measurements were finally made at two

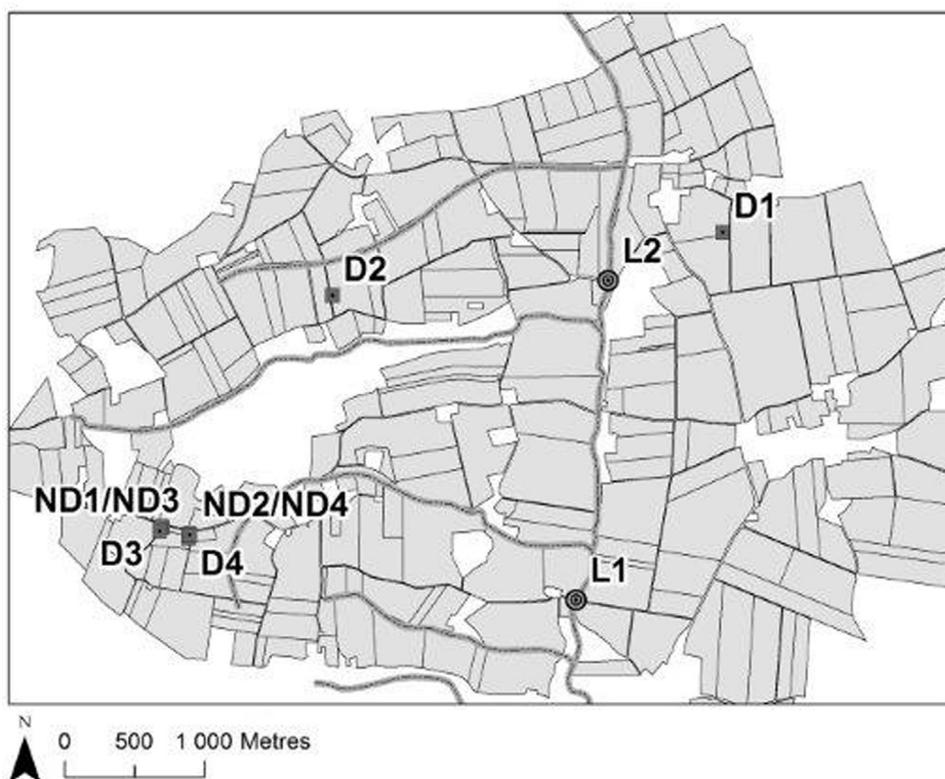


Fig. 1. Position of experimental plots in the studied region. Grey areas indicate cropped fields; white areas indicate forest or villages. D are drained plots; ND undrained plots and L stream positions. ND1/ND3 and ND2/ND4 are situated on the same field but correspond to 2 different years (see Section 2.1 for details).

positions on the Loir stream channel, upstream and downstream of the site (Fig. 1), referred as L1 and L2, respectively.

2.2. Measurements of N_2O emissions

Direct N_2O emissions were measured with five static chambers per plot. Uncertainty was given by the standard error on these 5 replicates. During the first growing season (2010–2011), a non-flow through non-steady state method was used (Rochette and Bertrand, 2008). Sampling frequency was generally weekly, except in Oct–Dec 2010 and June 2011, when it was monthly. Square stainless-steel frames ($50 \times 50 \times 25$ cm height) were inserted to 10 cm depth into the soil at least two days before the first sampling date and remained until the end of the growing season. The distance between two frames was always >1 m, and the maximum distance between two frames on one plot was 15 m. During measurements, the chambers were tightly closed with a vented PVC cover insulated by rubber foam. Four samples of air headspace were taken: just after closure and after 10, 20 and 30 min of accumulation. In the late spring, a supplementary 50 cm high frame was used to allow for plant height. Air samples were then taken 0, 50 and 100 min after closing the chambers because of the higher volume. At each measurement time, three replicate gas samples were drawn using a syringe (20 mL) and injected into pre-evacuated vials (12 mL). In each vial, 3 mg of magnesium perchlorate was placed to absorb the water vapour. The concentrations of N_2O were analysed in the laboratory within two weeks of sampling using a gas chromatograph (Model 3800, Varian Inc., Walnut Creek, CA, USA) equipped with an electron capture detector (GC-ECD) and a headspace auto-injector (Combi Pal, CTC Analytics, Zurich, CH). N_2O fluxes are calculated following Gu et al. (2011) and Rochette and Bertrand (2008).

During the second year, a new instrument (QCL spectrometer) was available for gas analysis, enabling an on-line analysis with a

higher sensitivity. This analyser is a laboratory-built instrument called SPIRIT (Robert, 2007; Guimbaud et al., 2011; Gogo et al., 2011). It was equipped for this study with a laser emitting between 2238.85 and 2239.20 cm^{-1} giving a standard deviation on a signal of 0.15 ppb for N_2O at 0.7 Hz. A cover was adapted on top of the static chambers to recirculate air from the chamber headspace to the SPIRIT analyser and back to the chamber headspace with PTFE tubing (1/4-in. diameter). The air flow inside the headspace was not high enough to avoid air stratification, so the cover was equipped with a fan (SUNON KD 1209 PTS 3, size 90×90 mm, 2500 rpm under 12 V). The fan was powered under low voltage and provided a slow mixing of air in the chamber to avoid pressure disturbance on the soil surface. The high sensitivity of the SPIRIT analyser enabled measurements 3 min after the chamber closing. Once again, supplementary frames of 15 then 50 cm height were used when the plants grew higher with a longer accumulation time (4 then 5 min). A preliminary test was conducted on 10 chambers to ensure that results obtained with the QCL spectrometer were consistent with the previously used method, i.e., the analysis of gas samples by GC-ECD in the lab (Grossel et al., 2014). The higher temporal resolution of the SIRIT analyser permitted a larger number of concentration measurements, so the concentration increase could be fitted by the HMR model to account for possible non-linear effects (Pedersen et al., 2010). In this study, the HMR non-linear model was applied to 11% of the cases, that is, only when it provided a better fit of the observed increase of concentration, most likely because the linear model is appropriate enough for very short time of gas accumulation.

The SPIRIT analyser also enabled indirect N_2O flux measurements on the Loir stream. Indirect N_2O emissions are generally derived from dissolved N_2O measurements in the water flow coupled to simple gas exchange models, but the actual gas exchange depends on flow velocity; this method leads to underestimation (Turner et al., 2015). Direct measurements by

drifting box were not possible in this study due to the small stream width and the frequent presence of vegetation. We thus coupled the SPIRIT analyser to a non-drifting box to provide indirect N₂O emission measurements. The floating chamber was a 35 cm diameter Plexiglas cylinder of 12.5 cm height. It was equipped with a vent; the chamber was carefully placed over water before vent closure to avoid any pressure disturbance. Measurements were made with the same methodology used over soils: a slow mixing of air was made by a fan and the N₂O accumulation in the headspace was measured during 4 min. Five flux measurements were carried out along a stream cross-section at positions L1 and L2, and the mean flux was estimated as the average of the single fluxes. Nevertheless, as the stream is temporary, flow properties are extremely variable in time. At L1, the stream is channelled in a 4.35 m corridor under a bridge. The flow was measured at a few dates and was typically 0.2–0.6 m s⁻¹. At L2, the streambed is natural (1–5 m width) with a very low flow at the edges (typically 0.1 m s⁻¹) and riffles in the middle. The dissolved nitrate concentration was measured at all sampling dates and varied from 4.6 to 12.2 mg N L⁻¹ (mean 8.6 mg N L⁻¹) at L1 and from 4.4 to 9.0 mg N L⁻¹ (mean 6.8 mg N L⁻¹) at L2. The stream width was measured at a few dates over 6 sections along the stream between L1 and L2 and varied from 1.0 to 4.5 m; the median (2.95 m) was taken as representative of the stream width. The stream depth varied from 6 to 30 cm but was observed at more than 1.3 m during a flooding event. The mean water pH was 7.8 and the mean oxydoreduction potential was 452 mV.

2.3. Ancillary variables

The climate and soil properties were also measured to assess possible controls on the N₂O emissions. The precipitation and air temperature were monitored using an automatic weather station (Campbell Scientific Ltd., Shepshed, UK). The volumetric water content and temperature were measured at the middle of the tilled layer (15 cm) close to each static chamber at the time of flux measurement. For each plot, twelve undisturbed soil cores were sampled using 9 cm diameter cylinders (0.5 L) over the 0–10, 10–20 and 20–30 cm soil layers to measure the soil bulk density. All plots except D1 were also equipped with moisture probes (TDR CS616, Campbell Scientific) and thermocouple sensors (TC Direct, UK)

placed horizontally at the middle of each layer, i.e., 5, 15 and 25 cm deep (three replicates per depth) to record the volumetric water content and temperature at a 2 h time step. The gravimetric soil water contents were determined every 3 weeks in 2011–2012 and at each sampling date in 2012–2013 on the 0–10, 10–20 and 20–30 cm soil layers to calibrate the TDR probes. A 60 cm deep piezometer was installed in the ND1, ND2, D3, ND3, D4 and ND4 plots to measure the water table level at a 2 h time step.

Moreover, three composite samples were made from six soil samples collected from each plot in two soil layers (0–10 and 10–30 cm) to determine the mineral N contents. Fresh soils were extracted with a KCl solution (0.5 M) and the NH₄⁺ and NO₃⁻ contents were determined using an automated discrete photometric analyser (Aquakem 600, Thermo Fisher Scientific Inc., USA).

The soil texture (clay, silt and sand contents), soil organic carbon, total N contents (after dry combustion at 1000°C) and pH in water were measured from the soil samples taken in the 0–20 cm layer in each chamber frame at the end of the measurement campaign. The samples were dried at room temperature, crushed and sieved with a 2 mm mesh, and analysed at LAS (Soil Analysis Laboratory, France). The soil type was identified by a profile examination.

2.4. Data analysis

Statistical data analysis was conducted with R (R Core Team, 2014). The N₂O measurements of every chamber are referred to as N₂O fluxes, and the average of the fluxes per date was used to estimate the mean plot emission. The difference between the N₂O emissions by the drained and undrained soils was tested by applying a non-parametric test (Mann-Whitney at level $p < 0.05$) to the N₂O fluxes from all chambers and all dates because the frequency distributions were highly asymmetric, resulting in a non-normal distribution (Shapiro-Wilk test, $p < 0.001$). The skewness coefficient was 13.1 for all flux data and ranged from 1.8 (D4) to 5.8 (ND3).

To examine possible controls of the environmental covariates on the N₂O emissions by the soils, a rule-based regression model was fitted to the N₂O fluxes using the R Cubist package (Kuhn et al., 2014). Cubist is based on the tree-model algorithm M5 developed by Quinlan (1992) and deals simultaneously with quantitative and

Table 1
Covariates used in the Cubist model to explain direct N₂O emissions in soils. Top: permanent soil variables in the 0–30 cm layer. SOC is the soil organic carbon content, Bd the bulk density. Each value is the mean of 4–5 replicates with standard deviation in parenthesis. Silt and sand contents were discarded from the model fit (see Material and Method). Bottom: non-permanent variables; mean of 78–102 values including spatial and temporal replicates. See text for details.

Plot	Clay (%)	Silt (%)	Sand (%)	SOC (g C kg ⁻¹)	Bd (g cm ⁻³)	pH-H ₂ O	Soil type	Topography
D1	17.3 (0.3)	78.1 (0.2)	4.6 (0.2)	10.4 (0.6)	1.36 (0.02)	6.4 (0.3)	Luvisol	Shoulder
D2	13.7 (0.7)	82.0 (0.7)	4.5 (0.4)	9.9 (0.4)	1.32 (0.04)	6.6 (0.2)	Albeluvisol	Shoulder
D3	17.0 (0.6)	78.2 (0.5)	4.8 (0.5)	8.1 (0.5)	1.41 (0.06)	8.1 (0.1)	Albeluvisol	Shoulder
D4	13.0 (0.2)	82.3 (1.0)	3.8 (0.5)	9.1 (0.3)	1.38 (0.08)	8.1 (0.1)	Colluvic Cambisol	Footslope
ND1	16.7 (1.0)	77.3 (1.6)	6.0 (0.9)	9.4 (0.5)	1.33 (0.05)	5.9 (0.2)	Albeluvisol	Shoulder
ND2	13.5 (0.4)	82.4 (0.6)	4.3 (0.3)	10.3 (0.9)	1.32 (0.04)	6.3 (0.1)	Colluvic Cambisol	Footslope
ND3	16.3 (0.8)	77.6 (1.5)	6.1 (0.7)	8.1 (0.6)	1.50 (0.04)	5.9 (0.1)	Albeluvisol	Shoulder
ND4	13.2 (0.5)	82.6 (0.6)	4.2 (0.4)	10.0 (0.4)	1.37 (0.09)	6.4 (0.1)	Colluvic Cambisol	Footslope
Plot	Crop	WFPS (%)	T (°C)	(NO ₃ ⁻) (mg N kg ⁻¹)	(NH ₄ ⁺) (mg N kg ⁻¹)	Precipitations (mm)		
D1	Barley	49.2 (23.2)	8.4 (3.5)	3.7 (3.0)	2.3 (2.1)	3.8 (6.6)		
D2	Barley	47.2 (20.9)	8.6 (3.4)	3.6 (2.7)	2.6 (2.0)	3.8 (6.6)		
D3	Wheat	75.9 (7.1)	6.2 (4.1)	7.4 (5.0)	1.9 (1.9)	14.6 (15.9)		
D4	Wheat	73.14 (6.3)	6.2 (4.0)	10.5 (8.8)	4.2 (7.3)	14.6 (15.9)		
ND1	Barley	55.3 (24.4)	9.7 (4.2)	4.3 (4.1)	5.9 (6.0)	3.8 (6.6)		
ND2	Barley	61.7 (29.1)	9.6 (4.0)	6.5 (7.1)	8.9 (10.1)	3.8 (6.6)		
ND3	Wheat	84.0 (9.0)	6.7 (4.7)	4.4 (2.9)	3.0 (3.0)	14.6 (15.9)		
ND4	Wheat	83.0 (9.6)	6.6 (4.2)	5.0 (3.6)	3.4 (3.4)	14.6 (15.9)		

qualitative covariates. This algorithm splits the dataset into subsets, following decision rules in the IF-THEN form (namely condition rules) based on the environmental covariates (Kuhn and Johnson, 2013). Within each subset, an ordinary least-squares regression is applied to define a linear regression between the N₂O fluxes and selected covariates (namely regression rules). The N₂O emissions present the non-linear and non-additive influence of the control variables (Schmidt et al., 2000). Thus, the data split between the subdomains enables the approximation of linear rules, and the objective here is not to use the fitted model as a predictive tool but to discriminate the relative importance of the control factors. For this purpose, Cubist provides an explicit model enabling an easy interpretation of predictor importance (Lacoste et al., 2014).

The covariates included qualitative attributes (crop, soil type, topographic position and drainage) and quantitative attributes (nitrate and ammonium content, soil temperature, WFPS, soil bulk density, carbon (C) content, clay, pH and cumulated precipitation over the 5 days preceding the day of measurement) that are described in Table 1. All the soil covariates were averaged over the 0–30 cm soil layer. The data gaps in the soil covariates (WFPS, temperature) were replaced with the mean of the other replicates at each plot at the same date.

To identify possible specific control factors in different situations, Cubist models were fitted to the N₂O flux (i) from the full dataset and (ii) from drained and undrained plots separately. Regression tree models can suffer from instability (Kuhn and Johnson, 2013), i.e., a small change in the calibration dataset may generate a very large change in the predicted rules. The models were first fit to all available N₂O flux data. Then, each N₂O flux dataset was split randomly between a training fraction (75%) used for the model fit and an evaluation fraction (25%). This procedure was repeated 100 times to check the stability of the predicted factor weight and to provide error estimation on the relative importance of the covariates.

The performance of the fitted Cubist models was assessed by considering several statistical indexes. The agreement between the measured and predicted data is characterized by the coefficient of determination (R^2):

$$R^2 = \frac{\sum_{i=1}^n y_i - \bar{x}}{\sum_{i=1}^n x_i - \bar{x}} \quad (1)$$

The accuracy of prediction is given by the root mean square error (RMSE):

$$RMSE = \sqrt{\frac{\sum_{i=1}^n (y_i - x_i)^2}{n}} \quad (2)$$

and the bias is given by the mean absolute error (ME):

$$ME = \frac{1}{n} \sum_{i=1}^n |y_i - x_i| \quad (3)$$

where n is the N₂O flux number, x_i is the one flux measurement, y_i is the corresponding flux prediction, and \bar{x} is the mean measured flux.

The indirect N₂O emissions could not be studied using a rule-based regression model because the study scale was different. However, the direct and indirect N₂O emissions were compared by assessing the total emissions over the high-water period of 2012–2013. The total soil emissions were estimated from a triangular integration on data points, assuming that the mean of the 5 fluxes in each plot is representative of the daily emission level. The uncertainty of cumulative emissions was estimated from triangular integration from the mean value plus/minus the uncertainty at each date. The investigation in main drainage

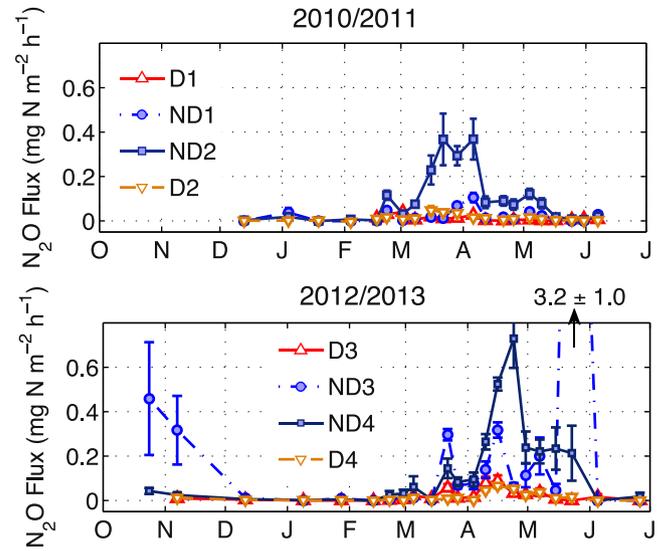


Fig. 2. N₂O emissions from drained and undrained plots. Error bars show the standard error on 5 replicates. For the emission peak observed in ND3, the peak value is written with its associated error.

ditches at a few dates in the same area has shown the same magnitude of N₂O emissions as the Loir stream; thus, the measured emission was supposed to be representative of surface water emissions at the study site scale. The total water emissions were estimated by multiplying the length of the stream and main drainage channels in the study site by a median stream width of 2.95 m and a mean channel width of 0.5 m. The area covered by surface water was estimated roughly to represent 0.1% of the whole study site.

3. Results

3.1. Specific climatic conditions during experiments

The two growing seasons benefitted from very different climatic conditions, the first one corresponding to a rather dry period and the second to a wet one. During the period of Oct 2010–Jan 2011, the cumulative precipitation was 252 mm, whereas it reached 328 mm during Oct 2012–Jan 2013, i.e., 30% more. During the period of Feb–May 2011, the cumulative precipitation was only 83 mm (50% drier than the mean climatological record over 30 years in the region), whereas over the same period in 2013, it was 3 times higher (252 mm).

3.2. Direct N₂O emissions

3.2.1. N₂O emissions

For every plot and the two studied periods, the N₂O emissions were lower during the winter periods with emission peaks following fertilization events during spring in all plots (Fig. 2). These peak N₂O emissions were much larger and lasted longer in the undrained plots than in the drained plots. Large N₂O fluxes were also measured in autumn 2012 after tillage in the undrained soils (especially for ND3). A very large peak was also observed on the 23rd of May 2013 for the ND3 plots: the mean N₂O emission was one order of magnitude larger than the mean peak flux level of the undrained plots.

N₂O emissions were ten-fold smaller on the drained plots than on the undrained ones in both growing seasons ($p \leq 0.001$, Fig. 3). There was no significant difference ($p=0.24$) between N₂O emissions from the drained plots in 2011 (D1 and D2) and in 2013 (D3 and D4); neither was there a difference ($p=0.14$) in the

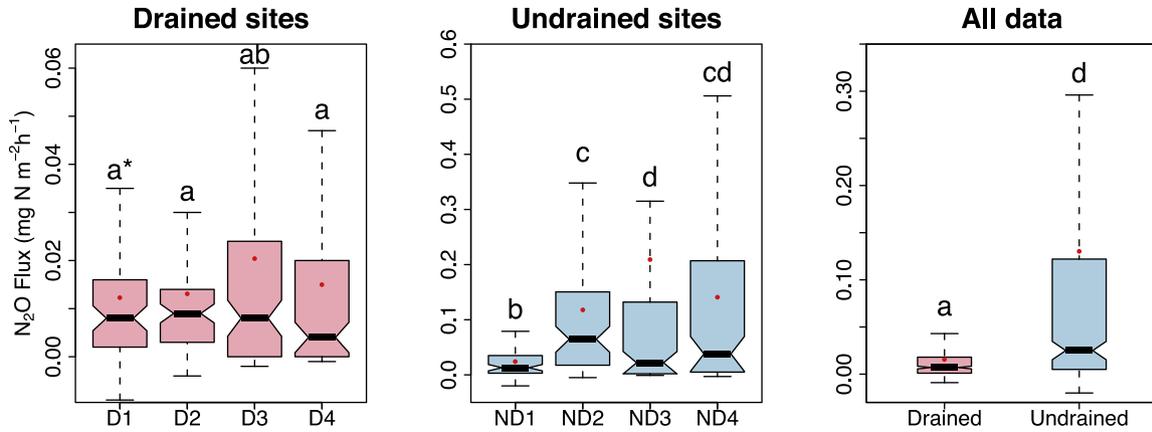


Fig. 3. Box plot of N_2O fluxes for drained sites (left), undrained sites (middle) and merged data from all sites (right). The point indicates the mean value. The letter above boxes indicates whether flux distribution was significantly different at the 0.05 level with a Mann-Whitney test.

N_2O emissions from the undrained plots in both years (ND1 and ND2 compared to ND3 and ND4). The variability between the undrained sites is larger than between the drained sites, and there were significant differences between the ND1, ND2 and both sites in 2013 (ND3 and ND4, Fig. 3, $p < 0.05$).

3.2.2. Soil water status

The soil WFPS was, on average, 9% higher in the undrained plots than in the drained ones (Fig. 4). The soils often reached saturation during winter in the undrained plots (especially ND2, ND3 and ND4) but never in the drained plots. The difference in WFPS between the drained and undrained situations was still visible at the end of the wet season 2012–2013. Similarly, a water table was measured at the surface at the ND2 plot in Feb. 2011 and at the ND3 and ND4 plots during long periods of the wet spring 2013 (Fig. 4). At the end of May 2013, a water table was again measured at the

ND3 plot after a heavy precipitation event (56 mm rain from 17th to 21th of May 2013), whereas no water table was observed above 60 cm at the other plots. Unexpectedly, a high water table level was observed over a long period in 2012–2013 at one of the drained plots (D3), suggesting that artificial drainage did not work efficiently while it was always able to decrease the surface WFPS.

3.2.3. Influence of management, climate and soil variables on N_2O emissions

The fitted models resulted in a good R^2 value but also a rather large inaccuracy with the RMSE equivalent to the mean flux value (Table 2). When a 25% bag fraction was considered, the evaluation of the models on it still gives a satisfactory result. The performances of the fitted model for the data for the undrained condition were slightly better than the data for the drained condition (Table 2). The best result was obtained for the whole

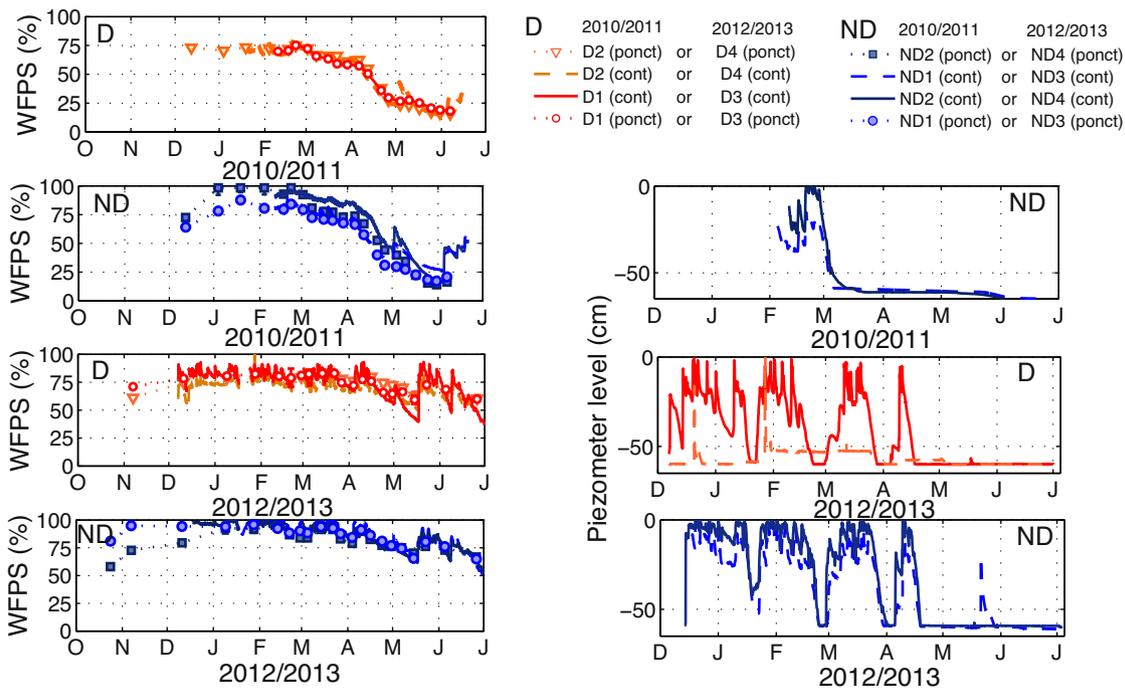


Fig. 4. Left: WFPS measured in drained and undrained plots for both years. Lines indicate measurements at 5 cm with a 2 h time step (referred to as cont. in legend) and symbols with dashed line indicate measurements at 15 cm at each sampling date (referred to as ponct. in legend). Right: piezometer level in drained and undrained plots for both years.

Table 2

Performance of the fitted Cubist models for all data, drained data (D) and undrained data (ND). Mean error (ME), standard error (SDE) and root mean square error (RMSE) of predictions are in $\text{mg N m}^{-2} \text{h}^{-1}$. Models were fitted first on all data, then on a training subset and the fitted model was evaluated on the remaining data (evaluation). See text for details.

Situation		n	Rules number	R ²	ME	RMSE
D and ND	All	714	11	0.78	-0.008	0.138
	Training	536	-	0.77	-0.008	0.135
	Evaluation	178	-	0.63	-0.013	0.197
D	All	353	6	0.71	-0.002	0.014
	Training	264	-	0.64	-0.002	0.015
	Evaluation	89	-	0.40	-0.003	0.020
ND	All	361	8	0.76	-0.015	0.197
	Training	272	-	0.74	-0.002	0.200
	Evaluation	89	-	0.59	-0.010	0.287

dataset (D and ND data) because the R² was larger on the validation dataset. This suggests that the more reliable factors have been identified in this case.

The rule-based model identified drainage as the main factor used in the condition rules (for 97% of the whole dataset; Fig. 5). This finding was stable: when only 75% of the dataset was randomly chosen for calibration, drainage was still identified in the condition rules in 83% of the cases. In the total model (11 rules), one rule included nearly all the data for the drained condition (345 over 353), and 9 rules were dedicated to the data for the undrained condition. The last rule is due to large precipitation events (>33 mm, 16 N₂O flux data) in either the drained or undrained conditions (for the complete model description, see supplementary material) and uses ammonium content as the only variable in the regression. The crop, soil type and topographic attribute were not identified as possible control factors, even when the data from drained, and undrained plots were pooled separately. This may be due to correlation with other factors: for example, topographic attributes influence WFPS and soil nitrate content. Regression rules were mainly based on the nitrate content, temperature, WFPS and precipitation. Permanent soil variables such as pH, clay and carbon (C) content have a smaller influence.

3.3. Indirect N₂O emissions from surface water

The temporal variations of the N₂O emissions emitted by the stream were different from those emitted by the soils (Figs. 2 and 5). High emissions were measured just after the stream began to flow in autumn and during high flow periods in winter and spring. Emissions were high during the whole experiment period and decreased at the end of flow period. The mean stream emission was $0.190 \text{ mg N m}^{-2} \text{h}^{-1}$ (ranging from 0.002 to $1.607 \text{ mg N m}^{-2} \text{h}^{-1}$) with non-significant differences between the L1 and L2 positions ($p > 0.05$). During the high-water period, the total Loir N₂O emission per unit area was estimated to be $15.3 \pm 0.8 \text{ kg N ha}^{-1}$. For comparison, the total direct emission was $0.8 \pm 0.2 \text{ kg N ha}^{-1}$ for the drained soils and $4.9 \pm 0.3 \text{ kg N ha}^{-1}$ for the undrained soils during the same period.

3.4. Comparison between direct and indirect N₂O emission

The N₂O emissions from the stream were much larger than the direct N₂O emissions on a per-unit-area basis. For comparison at the site scale, the mean emissions of the plots D3 and D4 were considered representative of drained (85.7% of site surface) and non-hydromorphic soils (11.1% of site surface); the mean emissions of ND3 and ND4 was considered representative of undrained hydromorphic soils (3.2% of site surface). In this way, the indirect

N₂O emissions were estimated at 31 kg N between mid-Oct and mid-May (flow period), and the soil emissions were estimated at 1846 kg N , resulting in a fraction of 1.6% of the indirect emissions by the surface water compared to whole site emissions.

4. Discussion

The main objectives of this study were two-fold: (1) to compare the soil N₂O emissions in drained and undrained plots and identify the main controls of emissions and (2) to compare the direct and indirect emissions in this drained landscape.

4.1. Effect of drainage on direct N₂O emissions

4.1.1. Effect of drainage, crop and soils properties

The most important drivers found in the rule-based models (Fig. 5) were clearly drainage and nitrate content, followed by soil temperature and WFPS. These variables are well-known local factors controlling N₂O production by soils (Hénault et al., 2005; Laville et al., 2011). These factors actually drive the temporal variation, both in time and intensity, of emissions following seasonal changes, precipitation and fertilization events. The N₂O emission variability was larger in the undrained soils than in the drained ones, with higher inter-annual and inter-plot variability of the N₂O emissions (Fig. 3). This lower N₂O emission variability for the drained soils might be due to a lower heterogeneity in the soil hydromorphic conditions, linked to a drainage effect. The influence of ammonium content on the N₂O emissions was much smaller, which suggests that nitrification played only a minor role in the N₂O production at these sites, while denitrification may dominate.

Permanent soil variables can also be explanatory factors of the spatial variability of emissions (Stehfest and Bouwman, 2006). These variables were indeed used by the regression model, but their influence was smaller than the variables having a temporal dynamic. The clay content was identified as a stronger control factor in the undrained soils than in the drained ones, whereas a previous study conducted in the same region identified clay content as a control factor in the drained plots (Gu et al., 2013). The effect of pH, soil C content and bulk density are either very small in the undrained soils or extremely unstable in the drained soils as shown by the large error bars in Fig. 5, so no clear conclusions can be drawn from the present study.

Except for drainage, none of the qualitative variables (crop, soil type, topographic attribute) was identified by the model as a control factor. A previous study in this region showed no link between the topographic attributes and emissions in the drained soils (Gu et al., 2011), although other studies have reported larger water, carbon and nitrogen contents in soils, denitrification hotspots and larger N₂O emissions in footslopes (Pennock et al., 1992; Vilain et al., 2010). The absence of topography effects in the drained soils in this region may be due to a decrease in runoff because of drainage and the gentle slope. So an effect may still be possible in the undrained plots. Emissions were indeed significantly larger in the undrained footslope in 2011 ($\text{ND1} < \text{ND2}$, $p \leq 0.001$) and between January and mid-May 2013 ($\text{ND3} < \text{ND4}$, $p = 0.02$). The opposite was observed in the periods of heavy precipitations, with much larger emissions in the shoulder of ND3 (autumn 2012, 23rd of May 2013); a competing effect between topographic attributes and soil type may have taken place. Topographic attributes were likely not identified as controlling variables in the Cubist models both because of their opposite effects and also because it is a categorical variable with only 2 levels (shoulder/footslope). The Cubist method is less sensitive to categorical variables having very few levels (Kuhn and Johnson, 2013). Thus, these results do not show that topography is not an important factor to explain the N₂O emissions but only show that

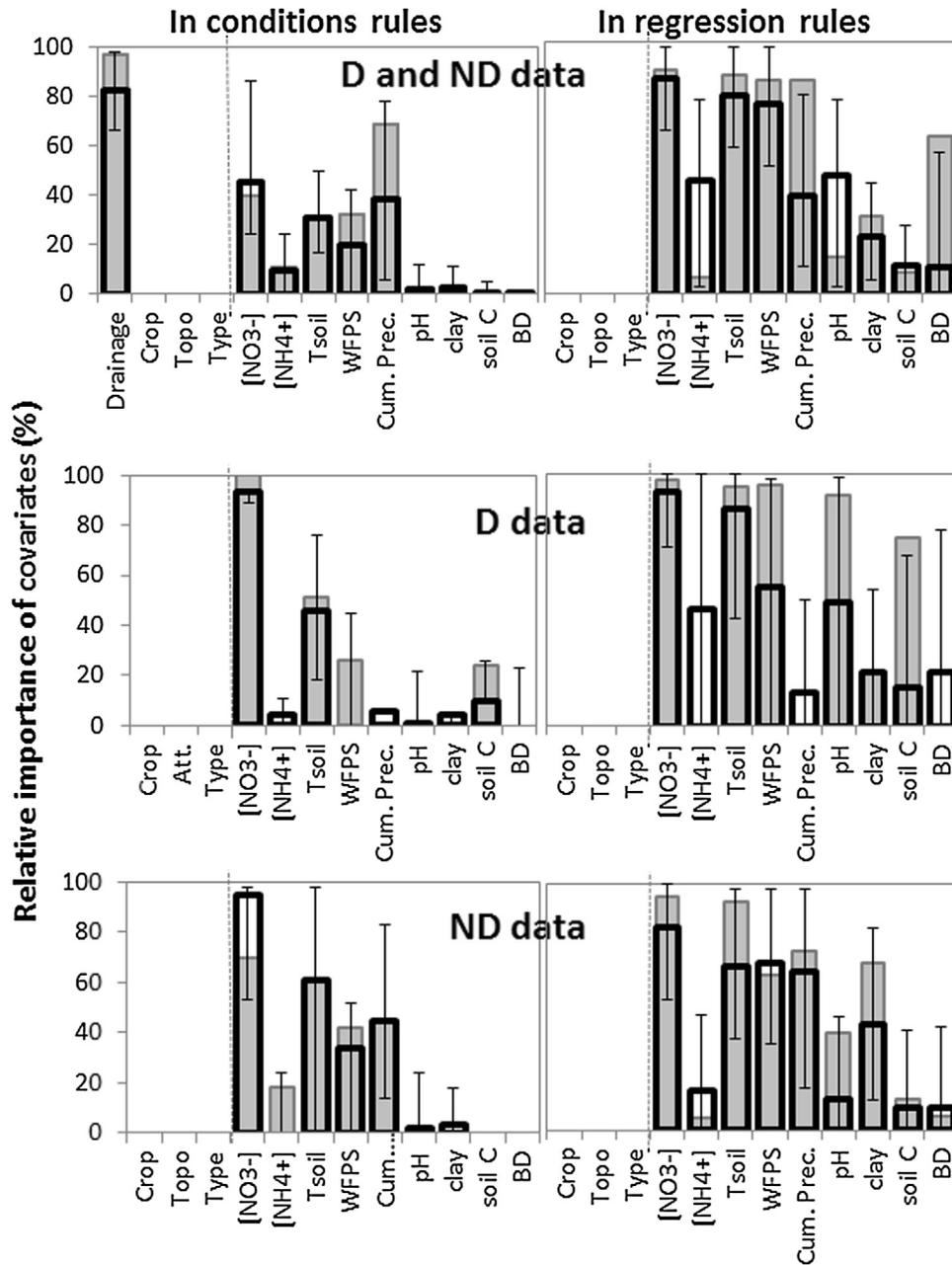


Fig. 5. Relative importance of covariates used by the fitted models in conditions rules (left) and in regression rules (right), for all plots (top), drained plots (middle) and undrained plots (bottom). Variables include qualitative factors (drainage, crop, attribute and soil type) and quantitative factors (nitrate and ammonium content in the 0–30 cm layer, soil temperature, WFPS, cumulated precipitation over 5 days, pH, clay, soil C content and mean bulk density of the 0–30 cm layer). Grey bars indicate the result for fit of all data and open black bars represent the average of 100 fits with a partitioned training dataset (see text for details). Error bars indicate here the first and third quartile of distributions.

the two-level topographic attribute used in the models should not be the most appropriate variable. Moreover, it is likely that the topography influences the soil water conditions, as well as the WFPS, which has been identified as an important factor controlling the N₂O emissions. On the other hand, the drainage variable also had only 2 levels (drained/undrained) and was still identified clearly as a dominant control factor. This is not surprising because much larger N₂O emissions were measured on the undrained plots on nearly all the sampling dates. This suggests that drainage was the dominant soil-controlling factor in this study, playing a role as important as very well-known factors, such as nitrate content. Thus, it would be important to also take it into account for inventories on larger areas and for the prediction of N₂O emission by models.

4.1.2. Effect of drainage and climatic conditions

The a priori hypothesis was that the effect of drainage on the N₂O emissions could be different depending on the climatic conditions because smaller N₂O emissions are expected from saturated soils (Colbourn and Harper, 1987; Venterea et al., 2008). However, the N₂O emissions were measured even in saturated soils in this study. The difference between emissions from the drained and undrained soils was also significant when considering the data from the two growing seasons together. The total N₂O emissions were larger, although not significantly, in the wet spring of 2013 than in the dry spring of 2011. Thus, in this region and in the periods following fertilizations, the N₂O emission levels from the undrained soils were larger than the emissions from the drained soils, both in wet and dry climatic conditions.

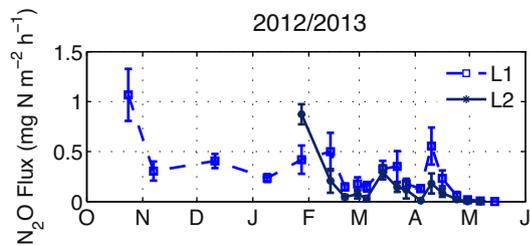


Fig. 6. N₂O emissions in the Loir stream for positions L1 and L2. Sampling frequency varied from one to four times per month.

One interesting result found by using the rule-based models to identify the controls on N₂O emissions was the importance of precipitation in the undrained plots but not in the drained ones (Fig. 6). Tiles can accelerate drainage of heavy precipitations and avoid soil saturation, thus decreasing the N₂O production in anoxic situations. Heavy precipitation events occurred during the second period of the experiment, between 24th of Sept. and 23th of Oct. 2012 (174 mm) and between 17th and 21th of May 2013 (56 mm). The model that was fitted to all data used one rule to partition fluxes following a large amount of precipitation, independently of the presence of drainage. The regression rule in this case actually used only the ammonium content. Actually, large ammonium contents were measured in the undrained situations, likely due to mineralization flushes in opposition to the drained ones, especially on the 23rd of May 2013. At this date, the very large peak measured at ND3 likely corresponds to a rewetting event as reported in numerous studies (Barton et al., 2008; Zona et al., 2011; Kim et al., 2012). The very large peak may be explained by the higher soil bulk density, as compaction enhances emission pulse intensity after rewetting (Beare et al., 2009) and by the rising groundwater table, only observed at this plot (Fig. 2). Emissions were observed two days after the end of the rain, when the soil was draining, and the water table was receding. The release of soil gas can indeed occur as the air entry point is reached when drainage proceeds, whereupon maximum N₂O fluxes occur with declining N₂O emissions also occurring as O₂ enters the soil due to drainage (Saggar et al., 2013; Balaine et al., 2013; Rabot et al., 2014). Thus, drainage may enable the effect of rewetting events to be reduced, which further amplifies the net effect of drainage on N₂O emissions. This is important because rewetting events can represent a large part of the annual N₂O budget (Goldberg et al., 2010).

4.2. N₂O emissions from the surface water

The local decrease of direct N₂O emissions due to drainage could be partly counteracted by an increase in the indirect N₂O emissions. However, the fraction of indirect N₂O emissions by the surface water compared to the whole site emissions is small (1.6%). It is much smaller than the one observed by Turner et al. (2015) in a 7850 km² area within the US Corn Belt (32%), whereas the stream area was 0.16% of that site, which is similar to the present site (0.1%). This might be because the mean indirect N₂O emissions were one order of magnitude smaller in the present study than in the Turner study (17.3 nmol N₂O m⁻² s⁻¹, i.e., 1.74 mg N m⁻² h⁻¹, against 0.190 mg N m⁻² h⁻¹ for our study). Oppositely, the present result is very similar to the 1.8% ratio proposed by Vilain et al. (2012) in France. They estimated indirect N₂O emissions of 0.035 kg N ha⁻¹ year⁻¹ when reported to the whole area of a 45.7 km² watershed; in the present study, reporting indirect N₂O emission to the site area (19.41 km²) gives 0.016 kg N ha⁻¹ year⁻¹, which is smaller but comparable, because indirect emissions occur only during the flow period (mid. Oct. 2012 to mid. May 2013).

Turner et al. (2015) used a floating chamber and measured emissions from surface waters. Oppositely, Vilain et al. (2012) considered only the part of the indirect emission linked to underground waters and used measurements of underground dissolved nitrous oxide. The total indirect emission was estimated from the watershed discharge. The study scale, crops (cereals and beans) and climate in Vilain et al. (2012) were more similar to the present study than those of Turner et al. (2015). Thus, the present study suggests that indirect emissions from temporary streams in France is comparable to the emissions from underground water in a site with permanent streams.

Measuring N₂O fluxes from a temporary stream is difficult and some limitations can be seen in the present method. Non-drifting chambers can enhance the water-atmosphere exchange and lead to flux overestimation. Moreover, the measurements were limited to one year, but the year was wetter than average, and the flow period was longer than usual (7 months instead of 4–5 months, as is usually observed). These facts would actually reinforce the conclusion of the present study about the small influence of indirect emissions at the site scale.

Pollution swapping still cannot be ruled out on the current measurements. Secondary drainage ditches, where water flows only after precipitation events, were neglected in the budget and have to be verified. Large and temporally variable sources of N₂O may exist at drain outlets (Reay et al., 2004). Moreover, high nitrate content is observed in the water in the present study and is directly exported. Nitrate may be denitrified in water sediments (Garnier et al., 2010) and an unknown amount of N₂O may be emitted downstream from the experimental site. This has still to be assessed in future studies. Thus, there remains a need for further studies on indirect N₂O emissions at several scales.

5. Conclusion

The first objective of this study was to assess the effect of artificial drainage on the N₂O emissions from loamy soils. The undrained soils showed significantly larger emissions than drained soils during both dry and wet years. The net effect of artificial drainage may be a large decrease in the direct N₂O emissions at this study site. A rule-based model was fitted to all flux data and clearly split fluxes between the drained and undrained situations. Drainage was the main factor explaining the spatial variability of the N₂O emissions within the studied soils, and its effect was dominant over other permanent soil variables. This strongly suggests that drainage must be taken into account for N₂O emission inventory.

As drainage could also induce pollution swapping by increasing indirect N₂O emissions, the second objective of this study was to investigate the indirect N₂O emissions from the surface water draining the site. The monitoring results suggested that N₂O emissions from streams might represent only a small fraction of the total emissions (1.6%). This finding requires further investigation in different sites because of the complexity of measurements involved in indirect N₂O emissions from non-permanent streams.

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Appendix A. Supplementary data

Supplementary data associated with this article can be found, in the online version, at <http://dx.doi.org/10.1016/j.agee.2016.06.015>.

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