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Research paper

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A B S T R A C T

The Cambrian–lower Ordovician volcanic units of the South Armorican and Occitan domains are analysed in a tectonostratigraphic survey of the French Variscan Belt. The South Armorican lavas consist of continental tholeiites in middle Cambrian–Furongian sequences related to continental break-up. A significant volcanic activity occurred in the Tremadocian, dominated by crustal melted rhyolitic lavas and initial rifting tholeiites. The Occitan lavas are distributed into five volcanic phases: (1) basal Cambrian rhyolites, (2) upper lower Cambrian Mg-rich tholeiites close to N-MORBs but crustal contaminated, (3) upper lower–middle Cambrian continental tholeiites, (4) Tremadocian rhyolites, and (5) upper lower Ordovician initial rift tholeiites. A rifting event linked to asthenosphere upwelling took place in the late early Cambrian but did not evolve. It renewed in the Tremadocian with abundant crustal melting due to underplating of mixed asthenospheric and lithospheric magmas. This main tectono-magmatic continental rift is termed the “Tremadocian Tectonic Belt” underlined by a chain of rhyolitic volcanoes from Occitan and South Armorican domains to Central Iberia. It evolved with the setting of syn-rift coarse siliciclastic deposits overlain by post-rift deep water shales in a suite of sedimentary basins that forecasted the South Armorican–Medio-European Ocean as a part of the Palaeotethys Ocean.

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1. Introduction

In SW Europe, the Variscan Ibero-Armorican Arc of the Gondwana margin comprises two branches (Fig. 1). The NE branch includes (1) the Armorican Massif, (2) the northern Massif Central, (3) the South Armorican Domain (southwestern Bretagne and Vendée), (4) the Occitan Domain (Albigeois, Montagne Noire, Mouthoumet, and Cévennes parts of the southern Massif-Central), and its lateral prolongation into Corsica and Sardinia, and (5) the Pyrenean Domain. The SW branch is completed with the Variscan zones of the Iberian Peninsula (Quesada, 1991; Martínez Catalán et al., 2009).

A common Cambrian–early Ordovician geodynamic framework is recognized throughout both branches (Balletre et al., 2014). After the post-Cadomian Ediacaran tectonic and magmatic events, a Cambrian transgression was linked to a long-term continental extensional process that led to the rifting of the Rheic Ocean in the mid Cambrian and its drifting in the early Ordovician, with the separation of Neoproterozoic arc terranes from the northern continental margin of Gondwana, such as Avalonia, Ganderia, Carolina and Meguma (Nance et al., 2008, 2010). Separation
occurred along the line of former Neoproterozoic sutures following the onset of subduction in the outboard Iapetus Ocean (Murphy et al., 2006). Opening of the Rheic Ocean propagated from southwest to northeast, in relationships with intracontinental rifting announcing the opening of the Palaeotethys oceanic Domain (Linnemann et al., 2008; Stampfl et al., 2013; Von Raumer et al., 2013, 2015). This Cambrian—early Ordovician tectono-magmatic history is incompletely understood in the South Armorican and Occitan domains, hereafter named SAOD, due to partial studies of disconnected volcanic exposures, the lack of well-constrained geochemical data, including incompatible elements useful for geodynamic interpretations, and the necessity of thorough stratigraphic revisions particularly of the volcanic units and their host formations.

The aim of this paper is to review the tectonostratigraphic pattern of the SAOD and to decipher the magmatic activities contemporaneous with the Cambrian—lower Ordovician sedimentary succession. Our targets are to understand the magmatic signatures of this activity and the tectonic significances of magma products in the geodynamic history of the northwestern Gondwanan margin. For this purpose, petrographical and geochemical investigations have been performed, using a new set of geochemical analyses and selected data from the literature.

2. Geological background and stratigraphic framework

The stratigraphic successions of the SAOD display numerous episodes of effusive and explosive events that are not necessarily correlatable throughout neighbouring Variscan tectonostratigraphic units bearing Cambrian—lower Ordovician exposures. We investigated the Vendeian area and the southern Massif Central from Albigeois to Mouthoumet regions. The lithostratigraphic nomenclature for the Vendée (southern Armorican Domain) has been revised and a new structural scheme is proposed. For the Montagne Noire and Albigeois Mountains (Occitan Domain), we followed the recent nomenclature of their tectonostratigraphic units by Alvaro et al. (2014a). Published available lithostratigraphic patterns are retained for the other areas: southern Cévennes (Verraes, 1979; Alabouvette, 1984; Ortenzi, 1986), and Mouthoumet Massif (Bessière et al., 1989; Berger et al., 1997).

2.1. South Armorican Domain

The South Armorican Domain including the Vendée area (Fig. 2) is limited with the Central Armorican Domain by the Nort-sur-Erde Fault (NEF), a former suture zone of a basin between the Gondwanan margin and the Middle Armorican terrane (Faure et al., 1997, 2008). Based on structural arguments (Cartier et al., 2001; Cartier and Faure, 2004; Faure et al., 2005, 2008; Ballèvre et al., 2009), it is suspected that an oceanic crust subducted to the north beneath Armorica. After ocean closure, collision of the margins evolved to dextral wrenching and northward backthrusting of the South Armorican northern area above Central Brittany between about 335 and 305 Ma (Gumiaux et al., 2004a).

South of the so-called suture zone area, we divide by the South Armorican Domain into four structural areas or zones bounded by major NW–SE trending faults: (1) the Northern Area between the NEF system and a southern branch of the South Armorican shear-zone that is the Montaigu-Secondigny Fault Zone in Vendée (MSZ); (2) the Central Structural Zone between the MSZ and the Chantonnay Fault (ChaF) continued with the Mervent Shear Zone (MSZ); (3) the Intermediate Structural Zone between the ChaF-MSZ, the Sainte-Pazanne-Mervant tectonic line (SPTL) and the Saint-Martin-des-Noyers tectonic line (SMTL); and (4) the Western or Littoral Structural Zone between SMTL and the Atlantic coast.

The structural nomenclature of South-Armorican pre-Variscan units is defined according to the autochthonous-allochthonous scheme of Ballèvre et al. (2014), with modifications based on the present study in Table 1. A synthetic cross section (Fig. 2b) illustrates the relationships between the allochthonous units or complexes in the Vendean area.

2.1.1. Northern structural area

The Ligerian region or domain located between the northern and southern branches of the South Armorican Shear-Zone is divided into two subdomains by the NEF: the northern subdomain including the Saint-Georges-sur-Loire Basin Unit and the southern subdomain that is defined as the northern structural area of the South Armorican Domain. This area consists of the Champtocoeux Complex thrust over the Neoproterozoic Maues paraautochthonous basement, the allochthonous Neoproterozoic Mauges Unit and its
Figure 2. Structural sketch of the Southern Armorican Domain in the Vendean area. (a) Geological map after removing the Quaternary deposits. ChaF: Chantonnay Fault; MCF: Mortagne-Cholet Fault; MSZ: Mervent Shear Zone; MSSZ: Montaigu-Secondigny Shear Zone; SMLT: St Martin-des-Noyers Tectonic Line; SPTL: Sainte-Pazanne-Mervent tectonic Line. Sites: Br: Brétignolles; Ch: Chantonnay; Co: Cholet; GL: Grand-Lieu Lake; Le: Les Essarts; Lf: Le Fenouiller; Lh: Les Herbiers; LR: La Roche-sur-Yon; Ls: Les Sables-d’Olonne; MAR: Mareuil-sur-Lay; Mo: Montaigu; Mt: La Meilleraye; Nt: Nantes; Pz: Pouzauges; Sg: Saint-Gilles. AB, cross section. (b) A–B, synthetic cross-section of the Vendean area.
limited Palaeozoic cover, the allochthonous Palaeozoic Cholet and Haut-Bocage units, and the Ancenis Basin (Cartier and Faure, 2004; Ballèvre et al., 2009; Ducassou et al., 2011).

The Northern Structural Area is limited to the south by the left-lateral Montaigu-Cholet Fault (MCF) along which Variscan granites and gabbros were intruded and mylonitized. Between the MCF and the MSSZ, a large granitic arch was emplaced during late Devonian to late mid-Carboniferous times, represented by the Les Herbiers migmatites (368 ± 7 Ma; Cocherie and Albarède, 2001), the Pouzauges monzogranite massif (347 ± 4 Ma; Poncet and Bouton, 2010), and the Mortagne leucogranitic batholith (313 ± 15 Ma; Guineberteau, 1986).

The lower Palaeozoic volcanic rock-bearing formations are distributed into two units, the Cholet and the Haut-Bocage units. The Cholet Unit includes a basal detrital formation that unconformably overlies the Neoproterozoic Les Herbiers migmatites (368 ± 7 Ma; Cocherie and Albarède, 2001), the Pouzauges monzogranite massif (347 ± 4 Ma; Poncet and Bouton, 2010), and the Mortagne leucogranitic batholith (313 ± 15 Ma; Guineberteau, 1986). The overlying volcanic series is divided into two units by an intermediate detrital formation of conglomerates and sandstones, and shales dated by trilobites as mid Cambrian (Cavet et al., 1979), a zircon dating of the Pouzauges granite, a shaly sequence transformed to schists of the MLAC from Pouzauges to northwest of Montaigu. The rocks are foliated NW–SE and dip to the southwest. The contact of the SPP and the Chantonnay Basin is faulted. On both fault sides, sedimentary features, structural patterns and metamorphism are similar. Thus, we consider that the Saint-Paul-en-Pareds Unit, the lower Palaeozoic Chantonnay Basin, and the Carboniferous coal-bearing basin.

The Montaigu Gneiss Unit is a migmatized and orthogneissified metagranite extended along the southwest of the MSSZ in the Montaigu area. It is dated at 488 ± 12 Ma (Godard et al., 2010). The planar fabric trending NW–SE and dipping 30° to 50° to the southwest. The tectonic contact with the orthogneiss is underlined by an elongated slice of serpentinitized peridotite indicative of a crustal-scale main fault. This fault is underlined with a suite of SE–NW elongated and brecciated bodies of peridotites parallel to the MSSZ (Fig. 2). This salient tectonic feature is termed the “MSSZ Peridotite Line”. The MLAC was thrust to the northeast above the Montaigu orthogneiss and wrenched along the MSSZ. It is assigned to the lower Allochthon (Table 1). Based on the lack of dating, Godard et al. (2010) allotted the MLAC to the basement of the Chantonnay Basin, and assumed that its magmatic activity was related to the Cambro–Ordovician history.

The Saint-Paul-en-Pareds Unit (SPP), made of shales and sandstones, extends along the southwestern margin of the MSSZ and the MLAC from Pouzauges to northwest of Montaigu. The rocks are foliated NW–SE with a 50 to 80° dip to the southwest. The contact with the MLAC is characterized by a tectonic mélangé of metabasement, amphibolite and leptynite slices. In addition, lenses of metasediments are enclosed in the complex, implying that the MLAC was thrust by the SPP in a number of slices. To the southwest, the contact of the SPP and the Chantonnay Basin is faulted. On both fault sides, sedimentary features, structural patterns and metamorphism are similar. Thus, we consider that the Saint-Paul-en-Pareds Unit belongs to the eastern lower part of the Chantonnay Basin.

### Table 1

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<th>Northern structural zone</th>
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<th>Intermediate structural zone</th>
<th>Littoral structural zone</th>
<th>Tectonic setting of the Gondwana multi-rifted platform</th>
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<tr>
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<td>Northeastern marginal basins</td>
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<td>Lower Allochthon</td>
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<td>Montaigu LAC</td>
<td>Les Essarts Complex eclogite lense</td>
<td>Mid-Carboniferous</td>
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<td>Les Essarts Complex orthogneiss</td>
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Figure 3. Geologic map of the Chantonnay Basin. (a) Map; (b) cross sections. Ant: Antigny; CF: Cheffois; CH: Chantonnay; LC: La Châtaigneraie; MC: Moncoutant; Mp: Mouilleron-en-Pareds; Mt: Montournais; PZ: Pouzauges; SG: Sigournais; SM: St-Pierre-du-Chemin; MSSZ: Montaigu-Secondigny Shear Zone. A-B, C-D, cross sections. U-Pb age dating: 486 Ma, orthogneiss of Mervent (Diot et al., 2007); 485 Ma, metarhyolite (blaviérite) of St-Pierre-du-Chemin, La Châtaigneraie Volcanic Formation (Bouton and Branger, 2007); 347 Ma, granite of Pouzauges (Poncet and Bouton, 2010). Ant, Smn1, Smn2: analysed samples.
The Chantonnay Unit is a basinal piling of Cambrian to middle–upper Devonian sedimentary formations including several interbedded mafic and acidic volcanic lavas. The basin is elongated NW–SE from the east of the Grand-Lieu Lake to Saint-Maixent-l’École, with a lenticular shape, an overall length of 140 km and a maximum width of 25 km. It is limited to the northeast by the Montaigu-Leptyno-Amphibolite Complex, the Saint-Paul-en-Pareds Unit, and the Montaigu-Secondigny Shear Zone (MSSZ) that obliquely cuts the sedimentary formations, and to the southwest by the Chantonnay Fault, the Mervent Shear Zone, and the Jurassic limestones that partly cover the eastern tectonic basin boundary until the Parthenay shear zone (Rolin and Colchen, 2001a,b). The southern area is crosscut by the southeastern extension of the Carboniferous coal basin. Considering that the Saint-Paul-en-Pareds Unit constitutes the base of the Chantonnay Basin, this basin is thrust above the Montaigu-Leptyno-Amphibolite Complex and is assigned to the Upper Allochthon (Table 1). The geological map and structural pattern are illustrated in Fig. 3 and the lithostratigraphy summarized in Table 2. The lowermost formations of lower to middle Cambrian age are dominated by basinal turbiditic sequences (Ménardière, Roc-Cervelle). The overlying middle Cambrian formations consist, in the lower part, of black shales (Gerbaudières) and interbedded sandstones that predominate in the southeastern side of the basin (Marillet-Puy Hardy), and in the upper part, of turbiditic sequences that hosted sills of metabasites and rhyolites (Bourgneuf). Sedimentation is interrupted by the flowing of a rhyolitic activity, the La Châtaigneraie Volcanic Formation, dated to Tremadocian by two samples: 470 ± 11 and 485 ± 14 Ma (Fig. 3a) (Cocherie in Bouton and Branger, 2007; Poncet and Bouton, 2010). The sedimentation renewed with a coarse detrital deposition of conglomerates, sandstones and slithstones that attest for a major tectonic event related to a syn-rift stage and erosion of rift shoulders (Cheffois, Sigournais). Then, a major change occurs with a deep-water post-rift deposition of black shales interbedded with Silurian phanitites in the lower part and Devonian limestones in the upper part (Réaumur). The basin history ended with the flowing of sub-water basaltic lavas (La Meilleraie).

The last tectonostratigraphic unit, the Carboniferous Coal Basin, is superimposed at the southwestern edge of the Central Structural Zone, along the Chantonnay Fault. It consists of an elongated NW–SE narrow trough extended from the Grand-Lieu Lake to the southeast of Chantonnay, and of two pull-apart troughs emplaced in the southern part of the Chantonnay Basin suggesting an N–S extensional strain. These basins are filled with conglomerates, sandstones, shales, and coal lenses. Their palaeoflora indicates a Namurian to Stephanian age.

![Figure 3](http://dx.doi.org/10.1016/j.gsf.2016.03.002)
2.1.3. Intermediate Structural Zone

The Intermediate Structural Zone is made of the Les Essarts Complex, the Saint-Martin-des-Noyers Unit, and the Mervent orthogneiss. These units are considered as NW–SE tectonic slices (Godard, 2001a).

The Les Essarts Complex (LEC) is an 80 km-long slice from the Grand-Lieu Lake to Chantonnay town. It is limited to the northeast by the fault system of the Carboniferous coal basin, and to the southwest, by the SPTL. It is made of polycyclic gneisses and amphibolites including metre- to kilometre-sized lenses of eclogites and peridotites stretched and boudinaged within the gneisses (Godard, 1981, 1983, 1988, 2001a,b, 2010). The gneisses are both ortho- and paragneisses showing relics of two consecutive high-T and high-P parageneses (Godard, 2001a). An orthogneiss is dated at 483 ± 4 Ma (U-Pb zircon; Lahondère et al., 2009). The eclogites, derived from metagabbros, recorded a high-pressure metamorphic event and then retrograded in the amphibolite facies. The LEC is supposed to be a tectonic mélangé of a partly subducted oceanic crust (eclogites) and a pre-Variscan continental crust (paragneisses) intruded by granite plutons (orthogneissess) (Godard, 2001a, 2010). In its northwestern and main part, the LEC is thrust to the southwest above the Saint-Martin-des-Noyers Unit and is thrust by the Chantonnay Unit (Godard, 2001a). It is therefore assigned to the middle Allochthon (Table 1).

The Saint-Martin-des-Noyers Unit (SMN) is elongated from the Grand-Lieu Lake to the south of Mervent with a total length of 100 km. It is thrust against the Les Essarts Complex and the Mervent Orthogneiss along the SPTL and suspected to belong to the lower Allochthon. It is limited to the southwest by a shear-zone fringed by a mylonitized linear pluton, the L’Aigle metagranite that metamorphosed the basin sediments of the Western Zone but not the SMN (Lahondère et al., 2009). The metamorphosed sediments of the Western Zone are dated to Ordovician and Silurian. It means that the metagranite postdating these sediments is Variscan and that the SMN unaffected by the metagranite was tectonically jointed with the pluton. Consequently, we locate the western boundary of the Intermediate Zone at the southwestern margin of the SMN along a newly defined tectonic line, the SMTL. The SMN consists of amphibolites and amphibole-rich gneisses belonging to the epidote-amphibolite facies of medium-grade metamorphism. The syn-metamorphic foliation is transposed along a dextral shear foliation trending NNW–SSE and highly dipping toward the southwest. The amphibolites display few relics of doleritic textures and derived from volcanic rocks (Thiébemont et al., 1987b). The gneisses range from gneissic amphibolites to amphibole-bearing gneisses. They may be either evolved magmatic rocks or volcanic-sedimentary formations, in the lack of magmatic or sedimentary remnant features.

The Mervent Orthogneiss is a 60 km stretched body from west of Chantonnay to southeast of Mervent, between the MSZ and the SPTL. Mylonitization planes are vertical with a subhorizontal stretching lineation and a right-lateral motion. The granite crystallization is dated at 486 ± 11 Ma (Diot et al., 2007).

2.1.4. Western Structural Zone

The Western Structural Zone is made of four lithological groups, in chronological order: (1) lower Cambrian igneous bodies of orthogneissified metagranodiorites constituting the Yeu island and the Les Sables-d’Olonne coast; (2) Furongian to middle Carboniferous metasedimentary and metavolcanic formations including five units, from base to top (structurally): the La Roche-sur-Yon, La Chapelle-Hermier, Brétignolles Basin, Saint-Gilles and Mareuyl-sur-Lay, and Bois-de-Céné Units; (3) the middle Carboniferous migmatic formations of Saint-Nazaire, also occurring in the metagranodiorite of Les Sables-d’Olonne; and (4) the Carboniferous granites (not studied here).

The Cambrian granitic massif constitutes the Yeu Island. It is dated to 530 ± 8 Ma (U-Pb zircon; Diot et al., 2015). The massif represents a tectonic pile of increasing metamorphic conditions from protomylonitic granite to kyanite- and sillimanite-bearing orthogneisses. Its foliation dips gently to the north and bears an N–S stretching lineation. Kinematic features attest a top-to-the south thrust (Sassier et al., 2006; Diot et al., 2015). Though less severely deformed, the metagranite north of the Noirmoutier Island resembles that of the Yeu island. It is intruded by late granitic apophyses (Fig. 2).

At the seashore, from northwest to southeast of Les Sables-d’Olonne, a comparable granodioritic orthogneiss is exposed. Its N–S stretching lineation is clearly expressed, but it is overprinted by an E–W stretching lineation (Iglesias and Brun, 1976; Cannat and Bouchez, 1986; Goujou et al., 1994; Cagnard et al., 2004). This orthogneiss body was mylonitized and thrust into the basal sequences of the La Roche-sur-Yon Basin as illustrated in Fig. 4, and exhibits a high-grade metamorphic foliation trending WSW–ENE and dipping 35° to the north. It can be assumed that the crystallization age of the original granite is contemporaneous with that of

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Table 2

<table>
<thead>
<tr>
<th>Chantonnay basin</th>
<th>La Roche-sur-Yon basin</th>
<th>Age</th>
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<td><strong>Upper part</strong></td>
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<td>La Meilleraye</td>
<td>Lava flows of basalts</td>
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<tr>
<td>Réaumur</td>
<td>Black shales, phanitines and rare limestones</td>
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<tr>
<td>Sigournais</td>
<td>Conglomerates, sandstones and siltstones</td>
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<tr>
<td>Chefois</td>
<td>Conglomerates and sandstones</td>
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<td>La Châtaigneraie</td>
<td>Lava and pyroclastite flows of rhyolite</td>
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<td>Bourgneuf</td>
<td>Turbidites, sills of metabasites and metahyalites</td>
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<td><strong>Lower part</strong></td>
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<tr>
<td>Formations</td>
<td>Lithology</td>
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</tr>
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<td>Turbidites</td>
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<tr>
<td>Roc-Cervelle</td>
<td>Turbidites</td>
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<tr>
<td>Payré</td>
<td>Sandstones</td>
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<table>
<thead>
<tr>
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<td>La Vierge</td>
<td>Siltstones</td>
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</tr>
<tr>
<td>Née-le-Dolent</td>
<td>Black shales and phanitines</td>
<td></td>
</tr>
<tr>
<td>Le Grouard</td>
<td>Shales</td>
<td>Mid Ordovician</td>
</tr>
<tr>
<td>Grosbreuil</td>
<td>Siltstones and sandstones</td>
<td>Floian</td>
</tr>
<tr>
<td>Les Sables-d’Olonne</td>
<td>Shales, rare dolostones, sills of metabasites and rhyolites</td>
<td>Floian, Tremadocian</td>
</tr>
<tr>
<td></td>
<td></td>
<td>Furongian</td>
</tr>
<tr>
<td></td>
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</tr>
<tr>
<td></td>
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<td>Mid Cambrian</td>
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the Yeu Island, i.e. 530 Ma. The migmatites are dated between 328 and 320 Ma (monazite U-Th/Pb chemical dating; Turillot, 2010; Turillot et al., 2011; Augier et al., 2015). A late granitic pluton intruded both the La Roche-sur-Yon Basin sediments and the migmatized orthogneiss, and provided numerous apophyses and dykes of pegmatite and aplite, inducing thermal metamorphism (Fig. 4). This pluton could be contemporaneous with the neighbouring Avrillé granite, 25 km east of Les Sables-d’Olonne, dated to 313 ± 3 Ma (Cocherie, 2008 in Béchennec et al., 2010).

The La Roche-sur-Yon Unit emplaced in the largest basin of the Vendean area and extends from the Saint-Martin-des-Noyers tectonic line (SMTL) to the Atlantic coast (Fig. 2). The sedimentary infill...
resembles that of the Chantonnay Basin with a piling of Cambrian to Silurian detrital formations and interbedded mafic and acidic volcanic lavas. The lithostratigraphy is summarized in Table 2. The geological map of the westernmost and lower part of the basin is illustrated in Fig. 4. The structural pattern of the basin is controlled by the anatectic domes of Les Sables-d'Olonne and La Roche-sur-Yon. The lower and oldest sediments crop out at the southern sea shore and are dominated by sandstones (Payré). The overlying formations consist of shales (Les Sables d'Olonne) interbedded with sandstones and rare dolostones and intruded by sills of basalts and rhyolites, a sequence of shales with interbedded sandstones and arkoses (Grosbreuil), and a shale-dominated sequence (Le Girouard). These last formations are suspected to date to lower and middle Ordovician owing to rare palaeontological remnants (Deflandre and Ters, 1966) and assuming a correlation of the rhyolitic activity with the Tremadocian one of the Chantonnay Basin. The overlying pile is made of black shales rich in phtanite beds dated to Silurian by fossils (Deflandre and Ters, 1970; Ters, 1970; Ters and Viaud, 1987) (Nieul-le-Dolent), and of undated upper siltstones and shales (La Vierge, Chassenon).

The La Chapelle-Hermier Unit is located to the northwestern side of the La Roche-sur-Yon Basin. It consists of slices of rhyolitic lava and pyroclastic flows southward thrust above the Ordovician–Silurian basin formations, along a reverse fault system dipping to the north. The northwestern limit of the unit is the basal thrust plane of the Saint-Gilles-sur-Vie nappe, and, in its west side, the Dinantian Brétignolles Basin. The rhyolite flows are characterised by a magmatic planar fabric 30° dipping to the north and a magmatic prismatic normal to the planar fabric. The slices resulted from basal delocalization of thick and massive lava flows in the La Roche-sur-Yon sedimentary basin during a north-to-south compressive event. The La Chapelle-Hermier Unit rhyolites are dated to Tremadocian by Béchennec et al. (2008) (U-Pb zircon): 486 ± 4 Ma (Chie-loup), 483 ± 10 Ma (Coex), and 478 ± 14 Ma (La
The Brétignolles Basin Unit consists of lower Carboniferous sediments filling a tectonic trough located east of Brétignolles, E-W trending and averaging 20 km in length (Comble et al., 1985; Colchen and Poncet, 1989). This basin is thrust at its northern side by the metarhyolite of the Saint-Gilles Unit nappe and is thrust to the south above the southernmost La Chapelle-Hermier slices. The whole basin sediments are folded in E-W trending and northward dipping reverse folds.

The Saint-Gilles-sur-Vie and Mareuil-sur-Lay Units (reduced to "Saint-Gilles" and "Mareuil") are located to northwest and southeast of the La Roche-sur-Yon Basin, respectively (Fig. 2). They were related to a former single nappe overlying the La Roche-sur-Yon Basin. The tectonic pattern of the nappe is evidenced by subhorizontal mylonitic foliation and shear planes gently dipping to the east and bearing an E–W stretching lineation indicating a westward transport. The basal thrust is reworked by NNW-dipping reverse faults along the northwestern edge of the La Roche-sur-Yon Basin. The nappe extends to the northwest until the St-Gildas cape, south of the Loire estuary, and is correlatable with the La Vilaine and Belle-île nappe (Audren et al., 1975). To the northeast, the nappe is limited by a metagranite that intruded and metamorphosed the nappe sediments (Lahondère et al., 2009), indicating that the granitic intrusion post-dated the nappe setting. This northeastern border is parallel to the SPTL and recorded the strong dextral shearing with highly dipping to the southwest foliation and subhorizontal stretching lineation.

The base of the nappe is a stack of rhyolitic lavas and pyroclastics dated to Tremadocian by five U-Pb zircon ages from Saint-Gilles.
(491 ± 12, 481 ± 14, and 477 ± 7 Ma) and Mareuil (479 ± 4 and 478 ± 2 Ma), and two U-Th/Pb monazite ages from Mareuil (486 ± 6 and 483 ± 4 Ma) (Béchennec et al., 2008, 2010; Ballèvre et al., 2012). This stack averages 500 m in thickness and is termed the “La Sauzaie Volcanic Formation”.

The overlying sedimentary deposition is made of a lower transition part made of rhyolitic pyroclastites and volcanosediments, and an upper part with fining upward sandstones and interbedded shales (“Saint-Gilles Formation”). In this sedimentary pile, a second volcanic sequence is recorded and termed the “Le Fenouiller Member”. It consists of two kilometre-scale rhyolitic lenses overlain by metabasites and located east of Saint-Gilles-sur-Vie (Lf, Fig. 2).

The Bois-de-Céné Unit is a 12 km-scale klippe emplaced above the Saint-Gilles Formation in the middle area of the Saint-Gilles nappe. The rocks consist of shales and interbedded sandstones transformed to micaceous with abundant metre- to hectometre-sized lenses of metabasites and serpentinites, strongly foliated and folded with axial planes trending NW–ESE. The paragenesis indicates a high-pressure metamorphism in the blueschist facies (Ters and Viaud, 1983; Guiraud et al., 1987). There is a symmetamorphic shear contact separating the Bois-de-Céné and the underlying Saint-Gilles formations. These were folded together during the deformation events that postdated the nappe thrusting and are only distinguished according to their different metamorphic patterns.

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### Chemical analyses of magmatic rocks: ICP and ICP-MS methods at ACME-LABS in Canada and at the "Service d'Analyse des Roches et des Minéraux du Centre de Recherches en Pétrographie et Géochimie" of Nancy, France: – no analytical data.
2.2. Limousin and Occitan domains

To the southeast of the South Armorican Domain, and beyond a Mesozoic cover (the so-called “Seuil du Poitou”) Cambro–Ordovician sedimentary and volcanic formations are distributed along the western to southern Massif Central. There, a major Variscan structural contact allows differentiation between the northern and southern Massif Central domains. The former domain is dominated by the lower and upper Gneiss Allochthonous Unites, but includes some low-grade sedimentary basins in the Limousin area. The latter domain, here named Occitan Domain, comprises the Variscan tectonostratigraphic units of the Thiviers-Payzac, Rouergue-Albigeois, Montagne Noire, Southern Cévennes and Mouthoumet massifs (Fig. 1). The Occitan Domain represents a proximal early Palaeozoic Gondwana margin that linked to the NW with the South Armorican Domain, as shown by the seuil du Poitou structural map of Rolin and Colchen (2001a), and to the SE with Corsica and Sardinia. The domain is limited to the south by the North Pyrenean frontal thrust.

2.2.1. Limousin

Lower Palaeozoic low-grade sedimentary series, the Bas-Limousin Group, are preserved in the western and southwestern margins of the high-grade metamorphic Limousin Massif. From north to south, these series are the La Gartempe, Saint-Salvador and Mazerolles Units (Santailier and Floch, 1989; Floch et al., 1993), the Thiviers-Payzac and Génis Units (Guillot et al., 1977; Roig et al., 1996), and the Leyme Unit (Guillot et al., 1989, 1992). The more complete sequence is recorded in the Thiviers-Payzac and Genis area and shows: (1) lower Cambrian sandstones, (2) a metadolerite and metagabbro stack, (3) middle Cambrian–Furongian shales and sandstones, (4) a lower Ordovician metahyolitic pile, (5) lower Ordovician sandstones, and (6) Ordovician to upper Silurian shales and amphibites. All these formations were metamorphosed in the gneissic facies, folded and displaced along a shear corridor.

2.2.2. Albigeois and northern Montagne Noire

Thick lower Palaeozoic sedimentary and volcanic formations occur in the Albigeois nappes, namely the Saint-Sernin-sur-Rance and Saint-Salvi-de-Carcavès nappes. These nappes southward overthrust the Lacaune Mountains (northern Montagne Noire), which consist of a succession of imbricated slices named Lacaune, Barre-Peux, Merdelou, Brusque, and Mélagues. All these units are closely related to the Avène-Mendic parautochthon and its eastern prolongation, the Lodevois inlier (Guérangé-Lozes and Burg, 1990; Álvaro et al., 2014a). The geological map of this region and location of the study area are illustrated in Fig. 5.

To the south, the northern Montagne Noire is separated by the metamorphic Axial Zone forming an elongated dome of migmatized orthogneiss (Somail orthogneiss) and micacchists tectonically overlain by lower Palaeozoic schists and marbles. Dating of orthogneiss has been revised to Ordovician (460–450 Ma; Roger et al., 2004). The southern Montagne Noire is made of south-facing nappes including a complete and fossiliferous Cambrian–Carboniferous succession (Vizzalino and Álvaro, 2001, 2003).

The lithostratigraphy of the northern Montagne Noire and Albigeois has been updated by Álvaro et al. (2014a) based on recent discovery and re-study of significant fossiliferous sites and on stratigraphic re-setting of volcanosedimentary units (Álvaro et al., 1998; Álvaro and Vizzalino, 1999; Devaere et al., 2013). This region recorded a basinal stratigraphy with an uppermost Ediacaran–lower Ordovician continuous sedimentation. A noticeable volcanic activity is widespread in many of the sedimentary sequences. The lithostratigraphic succession followed in this study, focused on latest Ediacaran—early Ordovician times, is given in Fig. 6. From base to top, eleven units are distinguished: exposures of (1) the Grandmont Formation and (2) its overlying Rivernous Rhylotic Complex are limited to the southeastern parautochthon sites; they are dated to latest Ediacaran, and the Ediacaran—Cambrian boundary interval, assuming a correlation with the “Schistes X” of the Axial Zone (Fournier-Vinas and Débat, 1970) dated at 545 ± 15 Ma by Pb isotope single zircon method (Lescuyer and Cocherie, 1992); they are overlain by the volcanosedimentary Layrac Formation; (3) the carbonate- and shale-bearing Marcou Formation has been recently dated to Terreneuvian (Devaere et al., 2013); (4) the siliciclastic Marcory Formation is widespread in the northern and southern Montagne Noire where it is dated at the Tommotian–Atabdanian interval (Álvaro et al., 1998; Álvaro and Vizzalino, 1999), and includes thick basaltic flows known as the Ensèges Volcanic Complex; (5) the Pardailhan, Série Noire, and Lastours formations consist of carbonates and black shales dated from Atabdanian to Toyonian (Thoral, 1935; Courtoiselle, 1973; Donnot and Guérangé, 1978; Debenre and Courjault-Radé, 1986); in the lower part, abundant volcanic units of mafic and acidic pyroclastic flows and deposits are distributed into two areas of the Brusque slice, namely the Moulergues–Lacan and Rocozels–Soubra volcanics; in the upper part, upper Ordovician basaltic activity is embedded mainly in the Saint-Méen Volcanic Complex; (6) the siliciclastic Coulouma and Ferrals formations are dated to lower middle Cambrian (Bogdanoff et al., 1984; Torno, 2002, 2003); (7) the thick pile of shales of the La Gardie Formation deposited during mid Cambrian times (Thoral, 1935; Torno, 2003; Cohen and Torno, 2006); (8) the Larroque Volcanic Formation is a thick (500–1000 m) and widespread package of porphyroclastic metarhyolites exposed from Albigeois to Rouergue (allochthonous in the former and parautochthonous in the latter; Fig. 5); this volcanic activity emplaced above the Furongian and the so-called “série chisto-gréseuse verte” (sensu Guérangé-Lozes et al., 1996; Guérangé-Lozes and Alabouvette, 1999) in the uppermost part of La Gardie Formation; (9) the Mendic megagranite intruded the Grandmont Formation and is suspected to be contemporaneous with the Larroque Volcanic Formation; (10) the La Dentelle Formation is a 15 m-thick quartz-sandstone stratigraphic marker; and (11) the Saint-Chinian and Albigeois formations are 500–1000 m thick black shales dated to Tremadocian–Floian (Bogdanoff et al., 1984; Guéranjé-Lozes and Guérangé, 1991; Vizzalino and Álvaro, 2001, 2003); the lower sedimentary pile of the latter includes numerous basaltic flows and sills referred to the Peyrebrune Volcanic Complex.

2.2.3. Southern Cévennes

In the Cévennes region, the lower Palaeozoic low-grade sedimentary series are stratigraphically equivalent to those of the Albigeois. They are considered as a single tectonostratigraphic entity, separated by Mesozoic and Cenozoic basins (Ledru et al., 1994), and named the Albigeois–Cévennes Unit. As in the northern Montagne Noire, the Albigeois–Cévennes Unit overthrusts to the south of the southern Cévennes area, equivalent of the Lacaune mountains (Marignac et al., 1980; Guérangé-Lozes, 1987). Various thrust slices have been described involving Cambrian units (Alabouvette, 1988). In the parautochthon, the lower sequence of sandstones and shales is a lateral equivalent of the Marcory Formation (Álvaro et al., 2010). It is overlain by shales and archaeocyath-bearing limestones, Botoman in age and correlated with the Pardailhan Formation (Debenre et al., 1976), and then by a thick stack of dolostones: equivalent of the Lastours Formation and dated to the Botoman–Toyonian (late early Cambrian). The lower part of this stack contains interbedded acidic to mafic volcanosedimentary deposits and basaltic lava and pyroclastic flows: the Saint-Bresson Volcanic Complex (Verraes, 1979; Alabouvette, 1988).
of the analysed samples in Table 3 and chemical analyses in the literature are retained. A petrographic study was done to the residual magmatic pyroxenes for microprobe analyses. Particular attention was paid to the early Ordovician volcanics for microprobe analyses.

3. Magmatic composition of the Cambrian—early Ordovician volcanism

The volcanic products of the above-reported domains dated around the Cambrian and early Ordovician have been sampled for petrographical and geochemical studies. Particular attention was done to the residual magmatic pyroxenes for microprobe analyses. Some accurate chemical analyses from the literature are retained. A petrographic study was done to the residual magmatic pyroxenes for microprobe analyses.

Table 5: Selected chemical analyses of pyroxenes. Microprobe analyses using Cameca SX 100 at the Camparis Analytical Laboratory of the University of Paris VI, France.

<table>
<thead>
<tr>
<th>Sample site</th>
<th>Volcanic formation</th>
<th>Sample#</th>
<th>Rock type</th>
<th>Major elements (wt.%)</th>
<th>Trace elements (ppm)</th>
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<tr>
<td>St-Méen sill</td>
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<td></td>
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<td>TiO₂ 0.54</td>
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<td></td>
<td></td>
<td></td>
<td></td>
<td>MnO 0.25</td>
<td>MgO 17.53</td>
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<td>Cr₂O₃ 0.30</td>
<td>MnO 0.25</td>
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<td></td>
<td></td>
<td></td>
<td>Total 100.22</td>
<td>MgO 17.53</td>
</tr>
</tbody>
</table>

3.1. South Armorican Domain

Magmatic products of interest are located in the following areas:

- **Cholet Basin** (Thiéblemont et al., 2001), these volcanics postdate a sedimentary sequence also palaeontologically dated as mid Cambrian (Cavet et al., 1996), so they may be younger.
- **Western Structural Zone** (La Roche-sur-Yon, La Chapelle-Hermier, Saint-Gilles and Bois-de-Céné units).

2.2.4. Mouthoumet massif (Corbières)

Finally, the Mouthoumet massif lies south of Montagne Noire and north of the North Pyrenean frontal thrust. It contains four tectonostratigraphic units, from east to west, the Serre de Quintillan, Félines-Palairac and Roc de Nitable thrust slices, and an unnamed stratigraphic unit dated from lower Ordovician to Carboniferous.

As in the case of the Pyrénées and Montagne Noire, the middle Ordovician is absent and its gap allows differentiation between a lower Ordovician sedimentary sequence and a para- unconformable overlying upper Ordovician—Silurian sedimentary sequence (Alváro et al., 2016). This package includes black shales that resemble those of the Albigeois and Vendean basins.

The most significant lower Ordovician volcanics consist of (1) a thick flow (ca. 100 m) of porphyritic metarhyolites in the sandstones of the parautochthon and dated to Tremadocian by acritarchs (Cornet, 1980; Bessière and Schulze, 1984; Bessière and Baudelot, 1988; Bessière et al., 1989; Berger et al., 1997). Mafic and acidic volcanic episodes are widespread in the various stratigraphic units dated from lower Ordovician to Carboniferous.

As in the case of the Pyrénées and Montagne Noire, the middle Ordovician is absent and its gap allows differentiation between a lower Ordovician sedimentary sequence and a para- unconformable overlying upper Ordovician—Silurian sedimentary sequence (Alváro et al., 2016). This package includes black shales that resemble those of the Albigeois and Vendean basins.

The most significant lower Ordovician volcanics consist of (1) a thick flow (ca. 100 m) of porphyritic metarhyolites in the sandstones of the parautochthon and dated to Tremadocian by acritarchs (Cocchia, 1982), and (2) flows or sills of metarhyolites overlain by flows of metabasalts in the lower Ordovician shales and sandstones of the lower part of the Serre de Quintillan slice, in its western side and in the Davejan tectonic window. These lower Ordovician volcanics are referred to the Davejan Volcanic Group.

Although the volcano-plutonic rocks of the Cholet Basin are broadly dated as mid Cambrian (Thiéblemont et al., 2001), these volcanics postdate a sedimentary sequence also palaeontologically dated as mid Cambrian (Cavet et al., 1996), so they may be younger.

Major and some trace elements of magmatic interest were presented by Thiéblemont (1988) for 14 metagabbros and meta-dolerites, 22 metarhyolites and microgranites, and 6 intermediate rocks of the upper magmatic unit. No new analyses were done.

The mafic rocks are metamorphosed in the greenschist and amphibolite facies. Their composition is subalkaline basaltic. Most of these rocks are saturated olivine-basalts (2.7 < Òl < 16.3), one being slightly under-saturated (Ne = 1.86), and a few others over-saturated tholeiites (0.1 < Qtz < 4.3). According to the Mg number, ranging from 70.3 to 45.2, the basaltic rocks are primitive to fairly evolved, fitting with the light rare earth low to moderate values (8.0 ppm < La < 21.8 ppm). Only La to Tb rare earth elements were analysed except for one sample where Yb was determined. The chondrite normalized La/Tb ratios range from 1.1 to 3.0. Taking into account the profile trends and the Yb analysed sample, the chondrite normalized La/Yb ratios range from 1.4 to 3.7. The complete analysed sample displays a chondrite normalized (La/Yb)NC ratio of 3.0. These values indicate a weak rare earth element fractionation. The lithophile moderate enrichment is illustrated by averaged values of N-MORB-normalized ratio of (Ta/La)NM = 1.1, (Th/La)NM = 2.6, and (Rb/La)NM = 8.3. In spite of the lack of some
key elements (Nb and heavy rare earth elements), the N-MORB-normalized element patterns of the Cholet metabasites are similar to those of the Montaigu Leptyno-Amphibolite Complex (see the following section) (Fig. 7a). Both magmatic series are characterized by moderate lithophile element enrichment, no enrichment or slight Ta-Nb negative anomalies, and moderate Ti negative anomalies limited to the more evolved rocks. These features are characteristic of continental tholeiites shown by an averaged profile of Holm (1985) in Fig. 7. The continental tholeiite signature is illustrated in the Th-Tb-Ta ternary diagram of Cabanis and Thiéblemont (1988) (Fig. 8a).

The acidic rocks gather microgranular porphyritic, subaphyric and granophytic facies. They are foliated and partly recrystallized in a felsic assemblage of quartz, albite, microcline, muscovite and biotite. All these rocks display similar silica- and alkaline-rich compositions. They are moderately enriched in incompatible elements (La = 3.1 ± 0.2 ppm; Yb = 7.2 ± 2.3 ppm; La/Yb = 4.5 ± 1.2; Th = 13.7 ± 2.9 ppm; Hf = 10.4 ± 4.3 ppm; Ta = 1.3 ± 0.3 ppm), in the Q-Ab-Or diagram (Tuttle and Bowen, 1958 in Thiéblemont, 1988), the rhyolites and microgranites plot close to the thermal eutectic for $P_{H2O}$ of 0.5 to 1 kbar. According to the petrographical features, the microgranitic texture of microgranites and the perthitic texture of all the alkaline feldspars, and the chemical compositions consistent with anhydrous granitic eutectic, it can be concluded with Thiéblemont (1988) that the rhyolites and microgranites are crustal melts derived from low-pressure eutectic liquids. This genetic model is supported by the abundance of rhyolitic and granitic rocks compared with the low volume of mafic rocks that discards a genesis of the acidic magmas from evolving mafic magmas.

The intermediate volcano-plutonic rocks consist of foliated felsic rocks enriched in amphibole and biotite. They display SiO$_2$ values ranging from 54.2 to 66.2 wt.% and MgO from 7.6 to 1.2 wt.%. Similarly, La varies from 12.9 to 26.0 ppm, Rb from 29.2 to 106.0 ppm, Th from 6.7 to 9.9 ppm, Zr from 67 to 193 ppm, and Hf from 1.8 to 5.4 ppm, while Sr decreases from 645 to 92 ppm. Petrographical features suggest that these rocks are mixtures of mafic and acidic rocks, rather more than products of variously evolved magmas (Thiéblemont, 1988). This mixing process is illustrated in the Th-Tb-Ta diagram by the Th enrichment trend from microgabbros to rhyolites (Cabanis and Thiéblemont, 1988) (Fig. 8a). It is noticeable that the rhyolites display the same Th-Tb-Ta ratio than the averaged upper continental crust (UCC; Rudnick and Gao, 2004).

### 3.1.2. Montaigu Leptyno-Amphibolite Complex

The Montaigu Leptyno-Amphibolite Complex (MLAC) comprises a lower thick amphibolite pile and upper alternating amphibolites and leucocratic to mesocratic fine-grained and banded gneisses commonly named “leptynites”. A broad Cambro–Ordovician age has been assumed. In the notice of the Montaigu geological map,
Figure 8. (a) Th-Tb-Ta diagram of the Cholet Basin and Montaigu Leptyno-Amphibolite Complex igneous rocks (diagram after Cahans and Thiéblemont, 1988). Mixing process between continental tholeiitic magma and upper continental crust material. CT, continental tholeiite UCC, upper continental crust after Rudnick and Gao (2004); (b) Th vs. Rb/Ba diagram of the gneisses from the Montaigu Leptyno-Amphibolite Complex (LAC). UCC, upper continental crust after Rudnick and Gao (2004). LC, lower continental crust after Weaver and Tarney (1984). This diagram discriminates the metabasites and the meta-acidic rocks by the Th enrichment. Compared to the composition of lower Palaeozoic rocks of the Variscan realm (Saint-Gilles nappe and Larroque Volcanic Complex of the South Armorican and Occitan domains in this work); compositional area of the Ollo de Sapo and Urra Volcanic formations in the Iberian Domain after Solà et al. (2008) and Diéz Montes et al. (2010), the LAC gneisses show a lower Rb/Ba ratio due to low Rb and high Ba.

Godard et al. (2010) analysed 9 amphibolites and 5 gneisses. These data being accurate, no new analyses were done.

The amphibolites are foliated and metamorphosed in the epidote-amphibolite facies. Any residual magmatic textures or minerals are lacking. However, according to their setting in thick elongated massifs, the amphibolites could be former gabbros or basalts. These rocks are subalkaline basaltic in composition. In the normative nomenclature, they range from tholeiites (1.1 < Qtz < 8.0) and some others saturated olivine basalts (9.6 < Ol < 17.6). The Mg numbers ranging from 58.9 to 46.2 indicate a fairly evolved mafic magma. The rare earth elements display a weak enrichment with a chondrite normalized (La/Yb)NC ratio of 1.7 ± 0.7 for a La value of 5.7 to 18.5 ppm. The N-MORB normalized incompatible element patterns (Fig. 7a) show a lithophile element enrichment, except Sr and K, with (Th/La)NM = 4.6 ± 2.5, (Rb/La)NM = 11.8 ± 10.3, and (Ba/La)NM = 4.9 ± 2.4. However, Ta is not enriched, (Ta/La)NM = 1.1 ± 0.6 and Nb is relatively depleted, (Nb/La)NM = 0.6 ± 0.2. In addition, a moderate negative anomaly concerns the titanium (Ti/Ti’ = 0.7 ± 0.1). These geochemical features (lithophile element enrichment, lack of Ta enrichment, and moderate Nb and Ti negative anomalies) are common in continental tholeiites (Holm, 1985). Orogentic calcalkaline magmas have a similar lithophile enrichment, but with less Rb-enrichment, more important Ta-Nb negative anomalies (andesites) or less rare earth element abundances (arc tholeiites) (Holm, 1985). In the Th-Tb-Ta diagram (Fig. 8a) amphibolites display a Th enrichment suggesting mixing with gneisses.

The gneisses have silica- and alkaline-rich compositions. They are ferroan, calc-alkaline and peraluminous in the classification of Frost et al. (2001). They are moderately enriched in incompatible elements, fractionated in light rare earth elements in showing a negative anomaly of Eu, and undepleted in the heavy rare earth elements (La = 43.2 ± 11.3 ppm; Yb = 5.3 ± 1.0 ppm; Th = 16.6 ± 1.0 ppm; Nb = 12.2 ± 1.6 ppm; La/Yb = 8.2 ± 2.2; Eu/Eu* = 0.3–0.4). Compared to the average composition of the upper continental crust (Rudnick and Gao, 2004), the gneisses display depletions in Ti, V, Fe, and Eu that may be explained by magmatic fractionation of oxides and plagioclase (Fig. 7b). Profiles are plotted with those of the neighbouring South Armorican rhyolites, e.g. the Saint-Gilles nappe. In the Th-Tb-Ta diagram (Fig. 8a), the gneisses are close to the acidic magmatic rocks of the Cholet Basin. But, compared to the composition of lower Palaeozoic rocks of the Variscan realm (Saint-Gilles nappe and Larroque Volcanic Formation of the South Armorican and Occitan domains in this work), and Ollo de Sapo and Urra volcanic formations in the Iberian Domain (Solà et al., 2008; Diéz Montes et al., 2010), the gneisses show a lower Rb/Ba ratio due to low Rb and high Ba.

3.1.3. Chantonnay Basin

A suspected synsedimentary magmatic activity is recorded in the lower Cambrian formations. It consists of scarce sills of metabasites and volcanosedimentary beds including mafic and acidic debris (Wyns et al., 1988; Bouton and Branger, 2007). No analytical data are available, not even from volcanic products of the neighbouring similar formations of the Haut-Bocage Basin (Wyns et al., 1988).

More significant volcanic activity is widespread in the Furongian formations with numerous sills and flow-sills of metabasites. Most of these sills are metre-sized in thickness, but there are some decametre-sized laccoliths (Fig. 3). Thick sills of rhyolites are set in the upper part of the formation. They are related to the underlying Tremadocian flows of the La Châtaigneraie Volcanic Formation (Table 2). Chemical analyses of major elements and a few minor elements are available from Thiéblemont (1988), namely eight metabasites and one rhyolite. Boyer (1974) analysed twelve rhyolites for the major elements. New and complete analyses were done for two samples of the Antigny mafic laccolith in the upper sequence of the Bourgneuf Formation (Ant., Fig. 3; samples Ant 1-3, Tables 3 and 4).

The metabasites are foliated and metamorphosed in the greenschist facies. Relics of doleritic texture are discernible in the centre of the thicker sills, with intersertal ordering of pseudomorphed plagioclase and actinolite with residual magmatic clinopyroxene in the core of phenoblastic amphibole. Chemical compositions of the new analyses are subalkaline to fairly alkaline. In the normative nomenclature, they range from tholeiites (0.1 < Qtz < 3.5) to olivine-basalts (5.0 < Ol < 15.4), one being slightly under-saturated (Ne = 1.1). According to the Mg number ranging from 67.2 to 49.9, the basaltic rocks are primitive to fairly evolved, in good agreement with the light rare earth low to moderate contents (3.2 < La < 11.9). Only La, Eu and Tb rare earth elements were analysed in the eight data from the literature. All the rare earth and incompatible trace elements are provided in the new analyses. The chondrite normalized La/Tb ratios range from 1.1 to 2.4 for the whole samples. The more complete analysed samples display a chondrite normalized La/Yb ratio of 1.1–1.7. Taking into account the profile trends for the Yb estimated values and the Yb analysed...
values, the chondrite normalized La/Yb ratios range from 1.1 to 1.7 for the whole samples. These ratios indicate a lack of rare earth element fractionation. The lithophile moderate enrichment is expressed with averaged values of N-MORB-normalized ratio, (Th/La)\textsubscript{NMB} = 2.9, (Rb/La)\textsubscript{NMB} = 27.1, and (Ba/La)\textsubscript{NMB} = 13.3. Ta and Nb display a moderate negative anomaly, (Ta/La)\textsubscript{NMB} = 0.8 and (Nb/La)\textsubscript{NMB} = 0.7. The N-MORB-normalized element patterns of the Chantonnay metabasites are similar to those of the amphibolites in the Montaigu Leptyno-Amphibolite Complex (Fig. 9a). Again, the geochemical fingerprints agree with a continental tholeiitic magma.

The amphibolites of the La Châtaigneraie Volcanic Formation show a felsitic groundmass including 15% to 30% of quartz and microcline phenocrysts. The magmatic texture is devitrified and fine-grained porphyritic or eutaxitic corresponding to thick lava flows and to ignimbritic flows. The metamorphism led to recrystallization of the groundmass in sodo-potassic feldspar, quartz and sericite. Close to St-Pierre-du-Chemin (Fig. 3), a felsitic body is elongated in a shear zone, severely mylonitized and finely recrystallized in quartz, sericite, chloritoid and pyrophyllite. The rock termed “pierre des plocères” has been worked in quarries for using in lime kiln because of its refractory property. We determined this rock as a “blaviérite” in the sense of Munier-Chalmas (1881-82).

The amphibolites are either massive or banded with a granular texture the chondrite-normalized diagram (\((\text{La} / \text{Yb})_{\text{NC}} = 0.6-0.3\)) indicating a depleted source and N-MORB signature. Some other amphibolites are enriched in Mg, Cr and Ni, but depleted in the whole rare earth elements with a Eu positive anomaly featuring a feldspar-olivine-pyroxene cumulative process. Some other amphibolites are enriched in the light rare earth elements and moderately enriched in the lithophile elements (Furongian metasediments).

For comparison, both types of eclogites show a tholeiitic composition with a light rare earth depletion in the chondrite-normalized diagram (\((\text{La} / \text{Yb})_{\text{NC}} = 4.5\)) suggesting a different and enriched source ((\text{La} / \text{Yb})_{\text{NC}} = 0.6-0.3)).

The amphibolites are either massive or banded with a granoblastic texture of hornblende, plagioclase, epidote, magnetite and garnet. Only few amphibolites display remains of doleritic texture of former magmatic plagioclase and clinopyroxene transformed to amphibole. All these mafic rocks can be termed metabasites having a chemical composition of subalkaline to fairly alkaline olivine basalts. The Mg number ranging from 64.7 to 38.9 shows that the rocks are primitive to evolved. The amphibole-bearing gneisses are intermediate in composition and may be considered as either evolved lavas or mixed volcano-sediments. The gneisses are silicic and may have derived from sediments or acidic volcano-sediments.

Figure 9. (a) N-MORB normalized incompatible element diagram of Furongian metabasites from the Chantonnay Basin. Normalized values after Sun and McDonough (1989). Partial analyses of Thiéblemont (1988) and complete analyses of this work (Ant 1 and 3). CT, average composition of continental tholeiites after Holm (1985). Comparison with the amphibolites of the Montaigu Leptyno-Amphibolite Complex (LAC, analyses of Godard et al., 2010) displays close similarities except a higher enrichment of Th and U in the LAC. (b) N-MORB normalized incompatible element diagram of metabasites of Saint-Martin-des-Noyers Unit. Magmatic groups of Thiéblemont et al. (1987b); Group 1, amphibolites of Pont-Charron (south of Chantonnay); Group 2, amphibolites of Petit-Lay (south-west of Chantonnay); Group 3, metavolcanics of Pont-Charron interbedded intrusion. New analyses have been done in the Pont-Charron site: Smn 2 belonging to Group 1, and to south-west Chantonnay: Smn 1 belonging to Group 2 and indicating that this site includes amphibolites of both groups 1 and 2. The Group 1 rocks are pure basaltic lavas depleted in the light rare earth elements and moderately enriched in the lithophile elements. The Group 2 rocks are basaltic lavas contaminated or mixed with silicic crustal material as shown by the UCC profile. The Group 3 metaamphibolites have different profiles with REE fractionation and the lack of Ta and Nb negative anomalies.

The amphibolites are either massive or banded with a granoblastic texture of hornblende, plagioclase, epidote, magnetite and garnet. Only few amphibolites display remains of doleritic texture of former magmatic plagioclase and clinopyroxene transformed to amphibole. All these mafic rocks can be termed metabasites having a chemical composition of subalkaline to fairly alkaline olivine basalts. The Mg number ranging from 64.7 to 38.9 shows that the rocks are primitive to evolved. The amphibole-bearing gneisses are intermediate in composition and may be considered as either evolved lavas or mixed volcano-sediments. The gneisses are silicic and may have derived from sediments or acidic volcano-sediments.
light rare earth elements and Ta, and enriched in Rb and Ba. This group is located southwest and south of Chantonnay and includes the new analysed sample Smn2. Group 2 is very fairly enriched in light rare earth elements. The Ta depletion is moderate, but Th and U are greatly enriched. This group gathers the banded amphibolites located to the west of Chantonnay and includes the new sample Smn1. Group 3 is significantly enriched in light rare earth elements, Ta, Ba and Rb. It is limited to the doleritic metabasite forming a thick interbedded body in the amphibolites of SW-Chantonnay, and is interpreted as a lately intruded sill.

The magmatic chemical features are interpreted in the Ta/Tb vs. Th/Tb diagram (Fig. 10b). Group 1 takes place in the mantle array close to the depleted MORB mantle and derived from a depleted source. However, the lithophile enrichment suggests either a former and limited enrichment of this source, or a crustal contamination of the magma. This second explanation is retained for the Group 2 magma because of the high Th values placing the lavas in a trend from Group 1 composition to the upper continental crust composition (UCC). Mixing of the depleted magma with crustal material would explain the overall element enrichment from Group 1 to Group 2 rocks as shown by the UCC profile drawn in Fig. 9b. The simplest explanation would be that the banded amphibolites were formerly a volcanosedimentary material. Group 3 is enriched in all the elements and derived from a slightly enriched source very different to the depleted source of Group 1. Owing to the suspected intrusive setting of the doleritic metabasites, Group 3 lavas are related to a second and distinct magmatic activity that resembles the continental tholeiite activity commonly present in the neighbouring lithostructural units. Concurrently, the Group 1 depleted source may be close to the depleted source of the eclogites though this latter source is only known by the rare earth element analyses in the eclogites.

The trace element contents and ratios of the felsic gneisses do not fit well with composition of acidic magmatic rocks. These gneisses are more probably of sedimentary origin and derived from clayey sandstones and shales with variable contribution of mafic volcanoclastic material as shown by the Ta/Tb vs. Th/Tb diagram (Fig. 10b).

3.1.6. La Roche-sur-Yon Unit

The volcanic products of the La Roche-sur-Yon Unit are limited to the Les Sables-d’Olonne Formation in the lower sequence of the basin dated to Tremadocian (Fig. 4, Table 2). They consist of metabasalts and metarhyolites interbedded in shales and averaging some ten metres in thickness and some hundred metres in width, except a few km-sized rhyolitic masses.

The metabasalts are moderately foliated and recrystallized in the greenschist facies. Relics of interseptal and intergranular doleritic textures are preserved with laths of plagioclase and subidiomorphic pyroxene transformed to amphibole. We selected samples in a 20 m-thick and 150 m-wide sill close to Talmont-Saint-Hilaire and far from the Variscan granite intrusives (TSH, Fig. 4; analyses Tsh 1, 2 and 3, Tables 3 and 4).

The metarhyolitic large bodies appear under two facies, either as porphyritic facies rich in phenocrysts of alkaline feldspars and quartz or as aphyric facies, related to lava flows and to pyroclastic flows from an aerial or a shallow-water volcanic activity. The metamorphic effect led to crystallization of muscovite and quartz and to abundant silica impregnation. Due to the latter effect, no analysis was attempted.

The chemical composition of the metabasalt is subalkaline tholeiitic (Qz = 3). The Mg number averaging 40.5 complies with an evolved magma. The rare earth elements are moderately fractionated (La = 16.7 ppm; (La/Yb)NC = 3.2) and the lithophile elements slightly enriched as shown by the averaged values of N-MORB-normalized ratio, (Th/La)NM = 2.5 and (Rb/La)NM = 2.1

![Figure 10](image-url)
3.1.7. La Chapelle-Hermier, Saint-Gilles-sur-Vie, and Mareuil-sur-Lay units

The La Chapelle-Hermier Unit slices, the lower part of the Saint-Gilles nappe (La Sauzaie Volcanic Formation), and the Mareuil nappe consist of rhyolitic lavas sharing similar petrographical features and ages. Dated and analysed by Béchennec et al. (2008, 2010) and Bâllevre et al. (2012), they consist of subvolcanic plugs and of subaerial lava and pyroclastic flows of porphyritic and aphyric rhyodacites and rhyolites. From the published data, we retained a set of fourteen analyses of major and trace elements. The data being accurate, no new analyses were done. SiO2 ranges from 67.1 to 77.6 wt.%, Na2O from 0.3 to 3.3 wt.% and K2O from 4.6 to 9.8 wt.%. The lowest alkali contents can be due to secondary leaching. In spite of this effect, rhyolites are alkali-calcic in the nomenclature of Frost et al. (2001). The Al2O3/(CaO + Na2O + K2O) mole ratio is peraluminous (1.1–2.0). The rare earth elements are moderately abundant (22 ppm < La < 39 ppm) and fractionated (chondrite normalized ratio: 5.5 < (La/Yb)NMB < 15.4), with increasing values from dacites to rhyolites. The Eu/Eu⁺ negative anomaly varies from 0.6 to 0.2, from rhyodacites to rhyolites. Trace element compositions are close to that of the upper continental crust (UCC) (Rudnick and Gao, 2004), except depletion of Ti, Fe and Mn explained by oxide and Ti-bearing phase fractionation. Negative anomalies of Zr and Eu suggest zircon and feldspar fractionations (Fig. 14a). The result of these fractionations is illustrated in the Zr/HF vs. Nb/Ta diagram according to experimental studies of Linnen and Keppler (2002) concerning the zircon fractionation in peraluminous melts and the data of Pfänder et al. (2007) (Fig. 13b).

Acidic, peraluminous and potassium-rich characters together with the rare earth and trace element patterns are typical of S-type peraluminous granites. These rhyolitic rocks possibly derived from melting of crustal material according to experimental studies of Patiño-Duce and Johnson, 1991; Patiño-Duce and Beard, 1995). A peraluminous and potassic shale or metamorphic equivalent is the best candidate. The crustal origin for the rhyolites is also supported by the abundance of inherited zircons dated around 600 Ma and by the abundance of inherited zircons dated around 600 Ma.

The metabasites capping the rhyolites of the second volcanic sequence are recrystallized in a planar fabric of fine granular actinolite, epidote, chloride, albite and quartz, as a consequence of thrust shearing. One sample of the Le Fenouiller volcanic Member has been analysed (Lf, Fig. 2). The composition is subalkaline (SiO2 = 55.6 wt.%; TiO2 = 1.2 wt.%; MgO = 5.5 wt.%; Na2O = 5.0 wt.%; K2O = 1.1 wt.%) though albitionization is suspected. In the normative nomenclature, the lava is a tholeite (Qtz = 2). The Mg number of 54.6 corresponds to a slightly evolved magma. The rare earth element contents are low and very moderately fractionated (La = 11.1 ppm; (La/Yb)NMB = 1.9) with slight Eu negative anomaly (Eu/Eu⁺ = 0.6). The N-MORB normalized diagram shows a clear enrichment of the lithophile elements: (Th/La)NMB = 5.2, (Rb/La)NMB = 9.2 and (Ba/La)NMB = 12.4 (Fig. 11). Ti, Ta and Nb display moderate negative anomalies: Ti/Ti⁺ = 0.6, (Ta/La)NMB = 0.6 and (Nb/La)NMB = 0.4. The normalized profile mimics that of the averaged continental tholeiite (CT). Taking into account a weak Th-enrichment possibly due to crustal contamination, the Saint-Gilles nappe metabasalt is close to the mafic rocks of the regional tektoson stratigraphic units: Chantonnay and Chelet Basin metabasites, and MLAC amphibolites (Figs. 12b and 13a).
3.1.8. Bois-de-Céné klippe

Interbedded metabasites of the Bois-de-Céné klippe are highly altered and impregnated with calcite. No analysis can be done. It is impossible to decipher the initial petrographical composition.

3.2. Limousin and Occitan Domain

In the Limousin and Occitan Domain of the Massif Central, the magmatic products consist of abundant volcanic rocks distributed in lower Palaeozoic sedimentary basins. The synsedimentary volcanic activities are only dated by their chronostratigraphic position. We can then describe the volcanics following a chronological succession and organizing them in volcanic complexes.

Many petrographical and geochemical studies have been done for these rocks (e.g. Durand and Gagny, 1966; Boyer, 1974; Prian, 1980; Gachet, 1983; Alsac et al., 1987; Guérangé-Lozes, 1987; Béziat et al., 1992; Pin and Marini, 1993). As a general rule, the chemical data were limited to major elements and a short selection of trace elements and many samples are not carefully localized.

Consequently, no accurate data have been retained from the literature. For that reason, an important set of new and complete analyses has been obtained from a thorough sampling (Fig. 5; Tables 3 and 4).

3.2.1. Rivernous Rhyolitic Complex

The Rivernous Rhyolitic Complex was emplaced in the Ediacaran–Cambrian boundary interval between the Grandmont and Layrac formations (Fig. 6). The rhyolites are historically known as “blaviérites” (Bergeron, 1888). This term, coined by Munier-Chalmas (1881-82), is applied to slaty silicic and sericite-rich rocks that are refractory, easy to carve in brick-shape and traditionally used for lime or brick kiln. Bergeron (1888) correctly used the term “bla-viérite” for very fine-grained micaceous talcschists located in fault zones close to rhyolitic bodies that he named “porphyroïdes”. Unfortunately, the following authors wrongly used the term “blaviérite” for the whole rhyolite formations by neglecting or ignoring the right determination. Earliest petrographical examination proved that blaviérites are metarhyolites (de Lapparent, 1909) severely mylonitized and totally recrystallized in quartz, sericite, paragonite...
and pyrophyllite (explaining the refractory property) in a planar fabric. The ultramylonitic features defining the blaviérites were highlighted in the Tremadocian and Tournaisian blaviérites of the Armorican Massif by Klein and Trichet (1968). These rocks occur in shear zone corridors across or along rhythmic bodies. Ordinary Riven-
ernous rhyolites are moderately deformed with a fracture cleavage and do not comply with the original blaviérite mineralological and tectonic definition, except in limited fault zones.

Three rhyolites have been analysed from (1) the Layrac summit, (2) the stratotype (LOD) at the Rivournivu rift, both in the Avené-
Mendic Unit, and (3) the Murat-sur-Vèbre area (Bl, Id, Mu, Fig. 5). The composition is silica-rich (68.7 wt.% < SiO₂ < 77.0 wt.%) and alkali-
potassic (0.6 wt.% < Na₂O < 2.8 wt.%; 4.1 wt.% < K₂O < 6.0 wt.%). The Al₂O₃/(CaO + Na₂O + K₂O) mole ratio (A/CNK) is peraluminous (1.3–2.6). Acidic, peraluminous and potassium-rich elements may be related to S-type peraluminous granitic melt.

In the Layrac and Rivournivu stratotype rhyolites, the rare earth elements are moderately abundant (14.5 ppm < La < 18.9 ppm) and fractionated (chondrite normalized ratio: 4.2 < (La/Yb)NC < 5.8) with Eu negative anomalies (0.24 < Eu/Eu* < 0.40). Compared with the upper continental crust (UCC), the light rare earth elements (LREE) are depleted (0.38 < (La/Ta)UCC < 0.40) (Fig. 1A-B). As commonly recorded in rhyolites, depletion of Ti, V, Fe and Mn are due to oxidation fractionation. Negative anomalies of Zr and Eu suggest zircons and feldspar fractionations, respectively. Depletion of the light rare earth elements (LREE) can be due to apatite and/or allanite fractionation according to their mineral/melt distribution coefficients in acidic rocks (Henderson, 1984). Fractionation is particularly important in the Murat-sur-Vèbre rhyolite with high LREE depletion (La = 4.7 ppm; (La/Yb)NC = 3.2).

Compared with the Orдовикian rhyolites, the Rivournivu Complex rhyolites are discriminated by higher Rb/Ba ratios.

2.2.2. Ensèges and Moulergues Volcanic Complexes

The Ensèges and Moulergues Volcanic Complexes comprise the lower Ensèges mafic lava pile and the upper Moulergues mafic to acidic lava and volcaniclastic counterpart.

The Ensèges lava pile conformably underlies the Atdaban-
ian–Botoman Pardaillan Formation in the northern part of the Brusque slice (En, Fig. 5). It consists of a ca. 300 m thick basaltic under-water pyroclastic flow pile exposed along a 10 km-long outcrop (Gachet, 1983). Lava outpouring occurred on a steep slope that explains the fragmentation of the quenched lava. Gravity driven aggregation of plastic debris produced decimetre-sized subrounded lavas. The highly brecciated rock is a mixture of angular fragments of vitreous to microlitic lava with plagioclase phenocrysts (An 74–80), partly transformed to albite and pyroxene transformed to actinolite. The lava was palagonitized and locally spilitized with recrystallization of epidote, actinolite, albite and calcite. Quartz and feldspar xenoliths are abundant.

Two samples were selected for analyses from the middle part of the flow pile (Ens 1 and 2, Table 4). The chemical composition is subalkaline (44.3 wt.% < SiO₂ < 44.6 wt.%; 0.8 wt.% < TiO₂ < 0.9 wt.%; 12.9 wt.% < MgO < 14.6 wt.%; 1.9 wt.% < Na₂O < 3.5 wt.%; 0.3 wt.% < K₂O < 0.8 wt.%). The high Mg number averaging 75 and the high MgO, Cr and Ni contents (Cr = 240 ppm; Ni = 140 ppm) are features of an olivine-pyroxene cumulate lava. But the rock is greatly altered (LOI = 6%), and no residual olivine can be discerned. In the normative nomenclature, the lava is saturated (Qz = 0), rich in olivine and hypersthene (Ol = 16–26; Hy = 12–30), and thus determined as an olivine-tholeite. The rare earth elements (REE) are poor (La = 1.5 ppm) and depleted in the LREE (La/Yb = 0.69; (La/Yb)NC = 0.47). However, the lithophile elements are enriched as shown by the averaged values of N-MORB-normalized ratio ((Th/La)NM = 2.1; (Rb/La)NM = 31–71; (Ba/La)NM = 12.7) (Fig. 15a). There are no significant Nb and Ta anomalies compared to N-MORB.

The REE pattern is that of the N-MORB and implies a depleted mantle source, as it is also documented in the Ta/Yb vs. Th/Yb diagram (Fig. 16a). The lithophile element enrichment suggests a magmatic calc-alkaline affinity. But, compared with the average pattern of the island arc tholeiites, the N-MORB normalized profiles of the Ensèges tholeiite show a lack of Nb-Ta anomalies and a distinct enrichment of Rb. This latter feature can be properly explained by a crustal contribution as shown by the upper crust continental average composition (Rudnick and Gao, 2004) (UCC, Fig. 15b). Although the lava includes quartz and feldspar xenolithic fragments that are not easy to eliminate in the sample preparation, the major element composition is not really modified by mechanical contamination of felsic material. The crustal contamination only concerns the major lithophile elements that are commonly transported by volatile phases. It is thus inferred that the tholeiitic magma was contami-
nated by volatile transfer during the crustal passage.

The Moulergues Complex overlies the Ensèges flows in two neighbouring areas of the Brusque slice: the Moulergues-Lacan subcomplex to the southwest and the Rocozels-Soubra sub-
complex to the northeast (Lc, Og, Kg, Fig. 5) (Guérangé-Lozes and Guérangé, 1990; Álvaro et al., 2014a). Both subcomplexes are encased in the lower part of the Pardaillan Formation dated to the Atdabanian–Botoman transition (Álvaro et al., 1998).

The basaltic part is made of volcanioclastic deposits of basaltic and rhyolitic composition (20–40 m) followed by an alternation of basaltic lava and pyroclastic flows, basaltic and rhyolitic breccias, detrital volcanosedimentary layers, and dolostone/shale interbeds (100–200 m thick). The upper part is a thick outpouring of quartz- and albite-rich rhyolitic breccias overlain by a felsic tephra layer (ca. 100 m). The mafic component is common in the Moulergues-Lacan sector, but rare in the Rocozels-Soubra sector.

The mafic lava flows are hyalo-microclitic with more or less abundant feldspar phenocrysts. Most of the flows are auto-
breciated and invaded by quartz and sodic feldspar angular frag-
ments of various sizes. The acidic lavas are rich in quartz and albite phenocrysts. All the lavas are highly spilitized and display albite-
phyr or keratophyre facies.

Two basaltic rocks were analysed in lava flows of the Lacan area (Lacan 1 and 3, Table 4). The chemical composition is subalkaline (43.5 wt.% < SiO₂ < 44.6 wt.%; 0.8 wt.% < TiO₂ < 0.9 wt.%; 11.6 wt.% < MgO < 12.9 wt.%; 2.7 wt.% < Na₂O < 3.8 wt.%; 0.6 wt.% < K₂O < 2.6 wt.%). The high Mg number averaging 73 and the high MgO, Cr and Ni contents (Cr = 277 ppm; Ni = 175 ppm) indicate a cumulate facies. The rock is highly altered (LOI = 6.5), and no residual olivine was observed. In the normative nomenclature, the lava is saturated (Qz = 0), rich in olivine and hypersthene (Ol = 24%; Hy = 13%), and thus determined as an olivine-basalt close to an olivine-tholeite. The rare earth elements (REE) are poor (La = 1.2–1.9 ppm) and depleted in the LREE (La/Yb = 0.6–1; (La/Yb)NC = 0.4–0.7). However, the lithophile elements are enriched as shown by the averaged values of N-MORB-normalized ratio ((Th/La)NM = 2.6–3.3; (Rb/La)NM = 64–101; (Ba/La)NM = 105–120) (Fig. 15a). There are no significant Nb and Ta anomalies compared to N-MORB. The whole chemical features are strictly similar to those of the Ensèges lavas assuming the same scenario, that of magma tapping a depleted mantle source (Fig. 16a), enriched in mobile lithophile elements from the crust.

Associated pyroclastic and volcanosedimentary brecciated rocks have been analysed in three samples. Due to the felsic component, silica and alumina are moderately enriched, and magnesia depleted (SiO₂ = 54–62 wt.%; Al₂O₃ = 16–20 wt.%; MgO = 1.1–3.3 wt.%). The spilitization effect caused anomalous enrichment in Na₂O (7–8 wt.%) and leaching of CaO and K₂O. The minor element...
contents recorded the felsic contribution with enrichment in LILE and lithophile elements, but not in Nb and Ti (Fig. 15b). The mafic lavas, volcanosediments and rhyolites are discriminated in the Th vs. Rb/Ba diagram (Fig. 17) due to the positive correlation between Th and Rb enrichments. The volcanosediments properly plot between the mafic lavas and the felsic rocks.

The rhyolite of a brecciated dyke displays a keratophyric facies rich in quartz and albite (SiO$_2$ = 74.5 wt.%; Na$_2$O = 6.7 wt.%). The A/CNK mole ratio is peraluminous (1.1). The trace elements are moderately enriched except Th and U indicating crustal contribution (Fig. 16a, b). The Ti negative anomaly can be explained by oxide fractionation. Owing to the spilitization effect, with leaching of the mobile elements, it is not possible to decipher about the acidic magma source.

3.2.3. Saint-Méen Volcanic Complex

Abundant volcanic products are interbedded in the Série Noire Formation from the central part to the southwesternmost Merdelou slice and in the northeastern corner of the Barre-Peux slice (Cd, Cf, Dg, La, LB, Lb, Mn, Mon, Po, Rf, Sm, Fig. 5) (Gachet, 1983; Guérangé-Lozes and Guérangé, 1990; Alsac, 1991; Béziat et al., 1992; Demange et al., 1995; Guérangé-Lozes and Alabouvette, 1999; Álvaro et al., 2014a). Consequently, renewal of the volcanic activity took place in the latest early Cambrian (Botoman–Toyonian).

The volcanic succession is composed of a basal thick deposit of volcanic and sedimentary breccia mingling mafic and felsic pyroclastites, commonly transformed to albitophyres and keratophyres (ca. 100 m). The following is an alternation of dolostone layers and volcanic material consisting of lava and pyroclastic flows and brecciated tephra deposits of mafic and felsic compositions (200–300 m). In the middle to upper parts of the sequence, numerous mafic sills are conformably interbedded, notwithstanding local tectonic disconformities. They recorded the same deformation and metamorphism than the encasing sediments. Their thickness ranges from 10 to ca. 100 m, and their length from 50 m to 4 km. The rock is massive and doleritic in texture. Some sills
are clearly associated with overlying lava flows. It can be inferred that the setting of these sills was subcontemporaneous with the sedimentation. The Saint-Méen sill is one of the largest subvolcanic bodies and gives the name to the complex.

Few mafic dykes have been encountered, particularly in the Lacaune area. Most of them are not deformed and metamorphosed, so they are related to post-Variscan tectonomagmatic events and not to the synsedimentary volcanism. The question of the feeding dykes of the interbedded sills is discussed in the following section.

Samples have been obtained from four lava flows, three sills, and five pyroclastic and volcanosedimentary deposits.

3.2.3.1. Composition of flows and sills. The lavas are autocrystalline and pyroclastic flows including fragments of hyalo-microlitic and porphyritic lavas and quartz-feldspar xenoliths. The sills are subophitic dolerites with phenocrysts of clinopyroxene, aggregated plagioclase and clinopyroxene, and local poecilitic clinopyroxene. Flows are more or less severely spilitized with high albite content. The low grade metamorphic effect developed a fracture cleavage and recrystallization of the groundmass and phenocryst margins in the greenschist facies with actinolite to Mg-hornblende (Mg/(Mg + Fe) = 0.67–0.81), albite, chlorite, clinozoisite (Cz 65–89), pumpellylite, calcite, muscovite and quartz. Chlorite is pyrochlore and crystallized at 252 ± 21 °C according to the Zang and Pye's (1995) thermometer (36 measurements).

Migmatitic plagioclases have the composition of calcic labradorite (An 61–74) only preserved in the core of phenocrysts of lava flows and sills. Best preserved pyroxenes have been analysed in the sills of Saint-Méen and Raffanel. In the Mg-Ca-Fe/Mn diagram, they show a calcic augite composition (32 wt.% < MgO < 45 wt.%; 14 wt.% < FeO1 + MnO < 24 wt.%; 38 wt.% < CaO < 45 wt.% (Fig. 18). Covariation diagrams attest for low contents of Tschermak moles (Fig. 19). The averaged mole composition is made of diopside (55%), hedenbergite (20%), enstatite (15%), ferrosilite (5%), and acmite + jadeite + Tschermak moles (5%). These features are consistent with pyroxenes of tholeiitic magmas. The composition of the parental magma is estimated by assuming that pyroxenes were in equilibrium in the magma chamber using tests based on Putirka (2008) and gives a crystallizing pressure between 5 and 6 kbar.

The composition of flows and sills is subalkaline (49.3 wt.% < SiO2 < 50.0 wt.%; 0.44 wt.% < TiO2 < 1.02 wt.%; 4.5 wt.% < MgO < 8.8 wt.%; 2.5 wt.% < Na2O < 5.5 wt.%; 0.59 wt.% < K2O < 2.62 wt.%). Spilitized lavas are commonly enriched in Na2O. The high Mg number between 63.8 and 73.1 indicates a partly cumulate facies. However, rocks are altered (LOI = 3.2–7.1) and no residual olivine has been observed. In the normative nomenclature, discarding the contaminated samples, lavas are saturated (Qtz = 0), moderately rich in olivine (Ol = 3.5–6.9%), and defined as olivine-basalts. The rare earth element content (REE) is low (La = 1.9–2.3 ppm) and very weakly fractionated in LREE (La/Yb = 1.9–4.2; (La/Yb)NC = 1.3–2.9). However, the lithophile elements are enriched as shown by the N-MORB-normalized ratios (Fig. 15a). There is a significant negative anomaly for Nb, and moderate for Ta compared to N-MORB ((Nb/La)NM = 0.2–0.5; (Ta/La)NM = 0.4–1.0). The magmatic source may be undepleted or less depleted than that of the Ensèges-Moulergues lavas (Fig. 16a). Magmas of the Ensèges-Moulergues and Saint-Méen complexes are quite different (Fig. 16b) though a similar scenario may be pertained for the magma genesis with mobile lithophile element enrichment from the crust.

3.2.3.2. Compositions of tephra and volcanosediments. Tephra deposits are pyroclastic layers rich in felsic clasts and impregnated with calcite. The quartz-feldspar xenocrysts are mantled with clinopyroxene in a reaction rim indicating their origin from the magma chamber or dyke walls. The volcanosediments are volcanioclastic mixtures of angular fragments of lava and silico-aluminous sediments. Lava fragments are either spilitized or fresh.

Due to the felsic component, silica and alumina are enriched (SiO2 = 59–69 wt.%; Al2O3 = 19 wt.%). The minor element contents recorded the felsic contribution with enrichment in LILE and lithophile elements in the N-MORB normalized diagram (Fig. 15b). Taking into account this enrichment, the patterns mimic those of the lavas with LILE fractionation and Nb and Ta anomalies. The conspicuous feature is the heavy depletion of Ti, the single element that is not enriched. The crustal contribution is illustrated in Figs. 16 and 17.

Figure 17. Th vs. Rb/Ba diagram for the Ensèges-Moulergues, Saint-Méen and Saint-Bresson metabasites, volcanioclastics and volcanosediments, and for the Rivernous, Moulergues, Larroque and Mendic felsic rocks. Compositional area of the Ollo de Sapo and Urra metarhyolites using data from Solá et al. (2008) and Blest Montes et al. (2010).
3.2.3.3. The question of the feeding dykes. A number of doleritic dykes can be encountered when sampling the Merdelou-Lacaune volcanic formations. Many of them are post-tectonic and devoid of schistosity and metamorphism. They commonly display an alkaline signature: ophitic texture with olivine phenocrysts. It has been suggested that these dykes fed the sills of the Saint-Méen Volcanic Complex, and then, these sills did not belong to the Saint-Méen Complex but to a late- or post-Variscan magmatic event, contradicting the interpretation given in this study. In return, there are a few planar intrusions of lamprophyres (spessartites sensu Béziat et al., 1992, 1993) in the Lacaune area, having an ultramafic calc-alkaline to alkaline composition drastically different to that of the metabasite sills. These intrusions are fractured and sheared but not metamorphosed.

To test the feeding dyke hypothesis, a doleritic dyke has been sampled for microprobe and geochemical analyses in the Bayles area, northeastern part of the Merdelou slice. The dyke (Bayles-1) is 1 to 2 m wide, displays a prismatic jointing and crosscuts subvertically the Série Noire and the lower part of the overlying Coulouma Formation, whose lithostratigraphic contact is marked by an interbedded volcanosedimentary level (Bayles-2). The rock is unfolded and devoid of metamorphism unless a late hydrothermal alteration. The rock is an ophitic olivine-rich dolerite dominated by phenocrysts of olivine and clinopyroxene in an aggregated clinopyroxene-plagioclase groundmass with Ti-magnetite (TiO$_2$ = 19–20 wt.%), apatite and secondary actinolite, pumpellylite, sodic feldspar and clinozoisite. The olivine composition is forsterite-rich (Fo 86–88). The olivine-liquid equilibrium temperature is calculated at 1260 °C (Putirka et al., 2007). The pyroxene is a fassaite diopside overlapping the 50 Ca limit of the Mg-Ca-Fe ternary diagram (29.4 wt.% < MgO < 39.2 wt.%; 10.4 wt.% < FeO$^+$ + MnO < 19.7 wt.%; 49.1 wt.% < CaO < 56.0 wt.%). This is due to high Tschermak mole substitutions averaging 25% as shown by the covariations diagrams (Fig. 19). The averaged mole composition is made of diopside (57%), hedenbergite (11%), enstatite (2%), ferrosilite (1%), jadeite + acmite (4%), esseneite (12%), Ti-tschermakite (11%) and Ca-tschermakite (2%). Such a composition is symptomatic of an alkaline magma. Using the pyroxene thermobarometer of Putirka et al. (2003), the crystallizing pressure averages 10 kbar.

The chemical composition is titanian- and magnesian-rich (TiO$_2$ = 2.9 wt.%; MgO = 7.8 wt.%; Mg#$^+$ = 60). The rock has experienced a weak hydrothermal alteration and shows high loss on ignition (8.1 wt.%). Leaching of alkalis probably occurred and the norm calculation has no sense. Owing to the olivine and fassaitic diopside mineral composition, the rock is representative of an undersaturated olivine-basalt magma. Except the most mobile elements (Ba, Rb, K, Sr), the incompatible elements are not too severely modified and their ratios are preserved. The rare earth elements are enriched and highly fractionated (La = 76 ppm; La/Yb = 32; (La/Yb)$_{NOD}$ = 21.5). Enrichments in Ta, Nb and Th can be considered. The N-MORB normalized profile is characteristic of an alkaline basalt (Fig. 15a) and

![Figure 18. Mg-Ca-Fe$^+$-Mn diagram of pyroxenes of Saint-Méen – Raffanel sills and Bayles dyke (a), and of Peyrebrune flow-sills (b). Saint-Méen, Raffanel and Peyrebrune are similar with average composition of calcic augite indicative of a sub-alkaline magma. Bayles is different with a fassaite diopside composition characteristic of an alkaline magma.](image)

![Figure 19. APV-Al$^{III}$, Ti-Al, and Fe$^{II}$-Ti covariation diagrams of per formula unit values of pyroxenes to distinguish the Saint-Méen – Raffanel, Peyrebrune and Bayles chemical compositions. All the pyroxenes are in the phenocrystic area. Compared to Saint-Méen – Raffanel, Peyrebrune pyroxenes are slightly enriched in Ca-Tschermakite, Ti-Tschermakite and Esseneite moles. In contrast, Bayles pyroxenes are highly enriched in the Tschermak moles, explaining the fassaite feature related to alkaline magma composition.](image)
drastically different to the lavas of the Saint-Méen Volcanic Complex and all the regional Cambrian volcanic complexes.

The mineral composition and the salient chemical features of the dyke clearly advocate for a magmatic alkaline signature and an enriched mantle source (Fig. 16). It is concluded that the dyke as well, as all the similar alkaline dykes, are not related to the Cambrian—lower Ordovician sills showing very different composition, but belong to a drastically distinct magmatic phase. Structural features indicate that these dykes were linked to a post-Variscan intracontinental fracturing.

3.2.4. Larroque Volcanic Formation

The most noteworthy facies of the Larroque volcanic rocks is represented by the sheared porphyroclastic rhyolites or microgranites rich in lacunous quartz (rhyolitic quartz) and alkali feldspar largely fragmented phenocrysts giving the name of “porphyroid” commonly used in the whole Variscan realm for these widespread Ordovician metarhyolites. These rocks are also named “augen gneiss” though they do not show a high-grade gneiss paragenesis but, in most cases, a lower grade metamorphic mineralogy. Because of the large bluish quartz phenocrysts, similar rocks are nicknamed “Ollo de sape” in Portugal and Spain. These porphyritic lavas belong to thick rhyolitic flows and to sills. Meanwhile, many other facies are associated such as aplitic lava flows, aphyric and porphyroclastic flows of welded or unwelded ignimbritic varieties, fine to coarse tephra deposits, and volcaniclastic or epiclastic deposits with variable sedimentary contribution.

We analysed four porphyroclastic rocks from the Larroque and Saint-Géraud areas in the Saint-Sernin-sur-Rance nappe, respectively (Am, Sa, Sg, Fig. 5). Composition ranges from potassic-rich dacite to rhyolite (66.7 wt.% Na2O; 0.8 wt.% K2O < SiO2 < 75.0 wt.%; 4.0 wt.% K2O < K2O < 9.3 wt.%), and belongs to the alkalic group in the nomenclature of Frost et al. (2001). The A/CNK ratio is peraluminous (A/CNK = 1.3–2.4). The rare earth elements are enriched and fractionated (33.2 ppm La/Yb e < 45.6 ppm; 11.2 < La/Yb < 14.5). The upper continental crust normalized diagram (Fig. 14b) shows negative anomalies of Ti, V, Cr, Mn and Fe linked to oxide fractionation, of Zr and Hf linked to zircon fractionation, and of Eu linked to plagioclase fractionation. Profiles are similar to those of the Vendéan Saint-Gilles rhyolites. Fractionation of zircon in peraluminous melt is illustrated in Fig. 13b. The Th vs. Rb/Ba features are also similar to those of the Saint-Gilles rhyolites, as well as those of the Iberian Ollo de Sapo and Urra rhyolites (Sola et al., 2008; Díez Montes et al., 2010) (Fig. 17).

3.2.5. Mendic Pluton

Though it is not a volcanic formation, the Mendic Pluton is concerned by this study because it may be related to the same magmatic activity than the Larroque Volcanic Formation. The Mendic metagranite is a biotite-muscovite porphyritic granite showing subhedral potassic feldspar in granular aggregates of micas, sodic plagioclase (An 10–30), alkaline feldspar, quartz and accessory magnetite, apatite and zircon. The initial granular texture is heavily constrained by a protomylonitic planar deformation. Dykes are dominated by microgranites recording the same deformation than the intruded metasediments.

We analysed two granites and one microgranite (Md, Fig. 5; Men 1–2 and Gra2, Tables 3 and 4). The granite is silica rich and sodopotassic (SiO2 = 75 wt.%; Na2O = 5.0–5.8 wt.%; K2O = 1.4–5.0 wt.%). The microgranite is more potassic (K2O = 7.0 wt.%). The A/CNK ratio is peraluminous (1.2). The rare earth elements are enriched and fractionated (La = 35.5–36.9 ppm, La/Yb = 9.5–11.0). The lithophile elements are enriched (Rb = 57–143 ppm; Ba = 113–511 ppm; Th = 19–21 ppm). The microgranite is characterized by depletion of light rare earth elements (LREE) but not of lithophile elements. An apatite fractionation may be assumed for the REE loss. The upper continental crust normalized patterns are similar to those of the Larroque metarhyolites (Fig. 14c). The noteworthy LREE depleted trend of the microgranite is evidence of apatite and/or allanite fractionation based on their mineral/melt distribution coefficients in acidic rocks (Henderson, 1984). The negative Eu anomaly is symptomatic of plagioclase fractionation. The Th vs. Rb/Ba diagram supports a geochemical similarity of the Mendic metagranite with the Larroque metarhyolites (Fig. 17).

3.2.6. Peyrebrune Volcanic Complex

The Peyrebrune Volcanic Complex includes several decametre-thick metadoleritic sills and metre-thick metabasaltic flows and pyroclastites encased in the black shales of the lower Ordovician Albigeois Formation of the Saint-Salvi-de-Carcavès and Saint-Sernin-sur-Rance nappes (Guérangé-Lozes, 1987; Guérangé-Lozes et al., 1996, Guérangé-Lozes and Moulène, 1998; Guérangé-Lozes and Alabouvette, 1999). The complex is named after the notorious mining area of Peyrebrune, western side of the Saint-Salvi-de-Carcavès nappe, where well preserved pillow lava flows were described by Durand and Gagny (1966), Alsac et al. (1987), and Marini (1987). The rocks have been sampled in the Peyrebrune-Teillet area and in the Bonneval-La Roque-Trebias-Mercadal area, located in the western and northern parts of the Saint-Salvi-de-Carcavès nappe, respectively (Py, Te, Bo, Lq, Tx, Fig. 5). The rocks recorded a moderate brittle deformation and recrystallized in the greenschist facies. The metamorphic chlorites are Fe-ripidolite to pycnochlorite and crystallized at 254 ± 18 °C according to the Zang and Pyfe’s (1995) thermometer (17 measurements). Spilitization process has occurled with phenoblastic albite enrichment at the base and top of the flows. Magmatic textures are always visible and residual magmatic clinoxyroxene has been analysed. The lavas are microlitic porphyritic with phenocrysts of plagioclase (An81–69) partly altered to albite and clinoxyroxene partly replaced by secondary actinolite. The inner part of the thickest lava flows (more than 4 m) and the sills are doleritic intergranular and locally porphyritic.

Several pyroxenes have been analysed in flow-sills of Peyrebrune. In the Mg-Ca-Fe+ Mn diagram, they display a calcic augite composition (36 wt.% < MgO < 48 wt.%; 11 wt.% < FeO+ MnO < 25 wt.%; 34 wt.% < CaO < 44 wt.%). These data confirm previous analyses performed by Marini (1987). The Peyrebrune pyroxenes seem to resemble those of the Saint-Méen Volcanic Complex but include higher contents of Tschermak molar (Fig. 19). The averaged molar composition is made of diopside (51%), hedenbergite (17%), enstatite (15%), ferrosilite (5%), acmite + jadeite (3%), Ti-Tsch (4%), esseneite (4%) and Ca-Tsch (1%), for a total of 9% of Tschermak molar. The pyroxene composition is consistent with enriched tholeiitic or transitional magmas. The composition of the parental magma is estimated by assuming that pyroxene was in equilibrium in the magma chamber (Putrika, 1999). The pyroxene thermobarometer of Putrika et al. (2003) gives a crystallizing pressure between 6 and 7 kbar.

The chemical composition of the flows and sills is subalkaline (48 wt.% < SiO2 < 52 wt.%; 2 wt.% < TiO2 < 4 wt.%; 4.3 wt.% < MgO < 7.4 wt.%; 2.6 wt.% < Na2O < 3.9 wt.%; 0.3 wt.% < K2O < 0.6 wt.%), discarding spilitized and silicified contaminated lavas. The Mg number between 40 and 57 corresponds to maﬁc to slightly evolved magmas. In the normative nomenclature, the lavas are oversaturated to saturated with Qtz or O1 and ly (0 < Qtz = 8; 0 < O1 = 10; 14 < H2O < 22), and then deﬁned as tholeiites and olivine-tholeiites. The rare earth element content (REE) is moderate to high (La = 7.9–50.0 ppm) and fractionated in LREE (La/Yb = 2.3–5.8; (La/Yb)NC = 2.2–5.5). The lithophile elements are enriched as shown by the N-MORB-normalized
ratios ($(\text{Th/La})_{\text{NM}} = 1.6–6.9$; $(\text{Rb/La})_{\text{NM}} = 1.0–5.3$; $(\text{Ba/La})_{\text{NM}} = 1.9–6.9$) (Fig. 20a). There are no anomalies for Nb, and Ta compared to N-MORB. The normalized spectra are parallel, indicating that all the lavas were evolved from the same magmatic batch. The more evolved lava shows a Ti negative anomaly. The profiles are compared with the averaged compositions of initial rift tholeiite (IRT) that fits very well with the profiles of mafic terms (neglecting secondary loss of K). The IRT signature is illustrated in the Ti-Nb-Th diagram (Fig. 21a). This conclusion is in good agreement with the previous data of Marini (1988) and Pin and Marini (1993) concerning some lavas of Peyrebrune. The IRT composition is explained by mixing of lithospheric and asthenospheric materials, which occurred during lithosphere thinning and intracratonic rift initiation. In the Ta/Yb vs. Th/Yb diagram (Fig. 21b), the tholeiites plot in the mantle array and may be related to a moderately enriched source. The La/Yb and Sm/Yb values and the lack of heavy rare earth element depletion comply with a mantle source of spinel lherzolite. Batch melting calculations using rare earth elements give a partial melting degree greater than 10%.

3.2.7. Southern Cévennes, the Saint-Bresson Volcanic Complex

The Saint-Bresson Volcanic Complex gathers mafic lavas, pyroclastic and volcanosediments of mixed acidic and mafic components (Vg, Fig. 5). Two basaltic flows were analysed. The rock is a crystal-rich pyroclastite with phenocrysts of pyroxene replaced by actinolite. Lava fragments are microlitic porphyritic and severely altered in secondary actinolite, albite, epidote and chlorite. Rounded xenoliths of quartz and feldspar are locally abundant. The matrix is enriched in sericite and calcite. The chemical composition is shown by the averaged values of N-MORB-normalized ratio (Th/La)NM 4.2; (Rb/La)NM = 14–112; (Ba/La)NM = 3.7–8.7 (Fig. 20b). The highest contents of Rb (and K) are explained by the sericite secondary invasion. There are negative Nb and Ta anomalies ((Nb/La)NM = 0.4–0.5; (Ta/La)NM = 0.4–0.8).

The REE pattern is close to that of the E-MORB and CT. The Nb-Ta negative anomalies and the LILE enrichment are more consistent with the CT pattern, allowing that the lower contents of REE are explained by fractionation process. In the Th/Yb vs. Th/Yb diagram, the mantle source was probably slightly depleted, considering that the moderate Th enrichment resulted from crustal assimilation as suggested by the submolten felsic xenoliths (Fig. 21b).

3.2.8. Mouthoumet Massif, the Davejean Volcanic Complex

A Tremadocian metarhyodacite flow of the parautochthon has been sampled in the Lairière volcanic dome, close to Montjoï village, at the northeastern corner of the Quillan geological map (northwest of Mouthoumet) (Mj, Fig. 5). A metarhyodacite flow interbedded between shales and sandstones of the lower Ordovician sequence of the Serre de Quitain has been sampled in the Quitain-forest area. Metabasaltic flows were investigated in the Davejean window and in the Maisons village, west of the Serre de Quitain slice (Dj, Ms, Qu, Fig. 5). The xenoliths are suspected to belong to the oldest volcanic event. Composition ranges from sodo-potassic dacite to rhyolite (69 wt.% < SiO2 < 76 wt.%; 21 wt.% < Na2O < 37 wt.%; 2.4 wt.% < K2O < 3.8 wt.%), and is alkali-calcic (Frost et al., 2001).

The A/CNK ratio is peraluminous (1.7–2.0). The rare earth elements are enriched and fractionated (59 ppm < La < 63 ppm; 28 < La/Yb < 38). The lithophile elements are enriched (75 ppm < Rb < 75 ppm; 400 ppm < Ba < 657 ppm; 14 < Th = 18). The upper continental crust normalized patterns are similar to those of the Larroque metarhyolites (Fig. 14c), taking into account a more important fractionation of zircon in the more evolved rhyolite. The Th vs. Rb/Ba diagram supports the similarity with the Larroque metarhyolites (Fig. 17).

The basaltic lava flows are microlitic porphyritic and range from mafic composition with abundant phenocrysts of pyroxene and plagioclase to intermediate composition enriched in plagioclase and alkaline feldspar. They were metamorphosed with complete recrystallization of magmatic minerals and matrix in the greenschist facies paragenesis. In addition, many flows were spilitized and enriched in sodium (albite). Weathering alteration is widespread. Consequently, the sampling is hazardous and many analytical results have been discarded. Based on the available analyses, the chemical composition is basaltic to moderately evolved (57 wt.% < SiO2 < 64 wt.%; 0.5 wt.% < TiO2 < 0.9 wt.%; 2.3 wt.% < MgO < 4.6 wt.%; 5.4 wt.% < Na2O < 6.5 wt.%; 0.5 wt.% < K2O < 1.5 wt.%), as shown by the Mg# (54–62). The norm calculation, though poorly constrained,
due to post-magmatic processes, favours an oversaturated tholeiitic magmatic composition (7 < Qtz < 15; 11 < Hky < 20). In return, the incompatible trace element abundances are relatively well preserved. The rare earth elements are enriched and fractionated (28 ppm < La < 42 ppm; 12 < La/Th < 21; 8.8 < (La/Th)NMB < 15.3). The N-MORB normalized diagram shows parallel profiles of the mafic and evolved lavas belonging to the same magmatic batch (Fig. 21c). The lithophile elements are enriched ((Th/La)NMB = 4.4–6.1; (Rb/ La)NMB = 3.1–5.7; (Ba/La)NMB = 3.6–4.9). Nb and Ta negative anomalies are important ((Nb/La)NMB = 0.4–0.6; (Ta/La)NMB = 0.4–0.6). The Ti negative anomaly (Ti*/Ti = 0.3–0.4) may be partly a magmatic feature and the result of oxide fractionation in the differentiated facies. Compared to the CT average composition, the lavas can be defined as continental tholeites, also shown by the Ti-Nb-Th ratios (Fig. 21a). In the Ta/Th vs. Th/Yb diagram, the mantle source cannot be defined because of the Th enrichment resulting from crustal contamination as suggested by the abundant felsic xenoliths (Fig. 21b).

4. Timing of the magmatic activities and geodynamic significance

4.1. South Armorican magmatic activities

In the South Armorican Domain, magmatic activities lasted from mid Cambrian to mid Devonian times according to their dated host sedimentary basin formations. The Cambrian to lower Ordovician volcanic rocks comprise middle Cambrian–Furongian olivine-basalts and tholeiites and Tremadocian rhyolites with a few overlying tholeiites. These mafic rocks share a common signature of continental tholeiites (CT) generated by melting of an enriched source of the continental lithosphere (Fig. 13a). Nevertheless, some Tremadocian tholeiites display an initial rift tholeiite signature (IRT) implying the contribution of a depleted source, and are indicative of an incipient rifting process. The acidic rocks derived from melting of crustal material. Heat for such a melting may have been supplied by the previous mantle-derived mafic magmas that intruded the crust and were underplated at the mantle-crust transition zone. This model is documented by the stratigraphically dated Cambro–Ordovician volcanics of the Chantonnay, La Roche-sur-Yon, and Saint-Gilles Units. It can be advocated that the mafic and acidic rocks of the Montaigu Leptyno-Amphibolite Complex belong to the same magmatic event because they share the same magmatic features.

In contrast, a second and quite different magmatic geochemical composition is illustrated by the amphibolites of the Saint-Martin-des-Noyers and the eclogites of the Les Essarts Complex. These mafic rocks originated from a depleted source (Fig. 10b, and rare earth element data of the literature) of N-MORB composition contaminated by lithophile elements. These mafic rocks are not dated. However, it is worth noting that similar basaltic lavas, originated from mixing of a depleted source and an enriched source, constitute a thick volcanic pile at the top of the Chantonnay Basin sequences named “La Meilleraie Volcanic Formation” (see Section 2.1.2, Table 2) (Thiéblemont, 1988) (Fig. 10b, LMT1). This volcanic stack includes hybridized lavas that resemble some metabasitess of the Saint-Martin-des-Noyers Unit (Fig. 10b, LMT2 and 3). The La Meilleraie Volcanic Formation is middle to upper Devonian in age. Based on identical petrological features, it is straightforward to postulate that the Saint-Martin-des-Noyers Unit was contemporaneous. In that case, the increasing contribution of the depleted source for the La Meilleraie Volcanic Formation, the Saint-Martin-des-Noyers Unit, and Les Essarts Complex is a distinct indication of lithosphere thinning and asthenosphere upwelling during mid to late Devonian times, possibly leading to the opening of a marine basin.

4.2. Occitan magmatic activities

The magmatic activities in the Occitan Domain studied in this work are distributed in five volcanic phases (V1–V5) ranging from earliest Cambrian to early Ordovician times (Fig. 6) (disregarding the youngest activities).

The first phase (V1) is represented by the uppermost Ediacaran to lowermost Cambrian Rivernous Rhyolitic Complex. This complex located in the northern Montagne Noire parautochthon is correlated with the rhyodacitic lavas of the “Schistes X”, an important stratigraphic unit of the Axial Zone dated to the latest Ediacaran time (Leguyer and Cocherie, 1992). It can be also correlated with acidic lavas of the Canaveilles Group from the South Canigou Massif in the eastern Pyrenees. Age datings of these Pyrenean volcanics provide various results (U-Pb SHRIMP method in zircon; Cocherie et al., 2005) due to abundant inherited cores. A previous age of 581 Ma was selected by Cocherie et al. (2005) but was re-investigated by Castiñeiras et al. (2008) who retain the youngest age of 540 Ma. In addition, Castiñeiras et al. (2008) dated to
548.8 ± 8 Ma a metahyolite of the neighbouring Roc de Frausa Massif (= Roc de France) and to 560 ± 11 Ma a similar rock of the Cap de Creus Massif. Nevertheless, new U-Pb geochronological data of Casas et al. (2015) from interbedded acidic rocks of the Canigou and Cap de Creus sequences give ages ranging between 577 and 558 Ma. The amount of inherited zircons is an important problem when interpreting isotopic data. Waiting for more straightforward geochronological analyses, we assume that the V1 volcanic phase took place at the Ediacaran–Cambrian boundary interval, its boundary horizon being dated at 542 Ma (Gradstein et al., 2004).

The significance of this acidic volcanic activity is not clear. Crustal melting should be the evident magmatic process. But, petrological data are too poor to decipher the geotectonic conditions of the magma genesis. A detailed structural and metamorphic analysis (Álvaro et al., 2014a), in order to test the possible onset of a Cadomian tectono-thermal effect, proved that the rhyolitic volcanics were deformed during a single and common deformation stage shared by the overlying Cambrian strata and presumably Variscan in age. In that part of the Montagne Noire, there was no important deformation event across the Ediacaran–Cambrian transition. Consequently, the V1 activity is not related to any late Cadomian or Pan-African orogenic event. It can only be said that a suspected extensional tectonic event occurred around 540 Ma in allowing outpouring of crustal melts possibly residual of a former and undetermined magmatic process.

The second phase (V2) occurred in the late early Cambrian. It began by massive outpouring of magnesian-rich saturated basalts of the Enses-Moulergues Volcanic Complexes in the northern Montagne Noire, and of the Saint-Bresson volcanic Complex in southern Cévennes. The uppermost basaltic lava flows are associated with rhyolitic breccia flows. The mafic lavas are poor in rare earth elements, fractionated in the light rare earth elements, and originated from a depleted mantle source. Meanwhile they are enriched in lithophile elements. This enrichment is attributed to crustal contamination by volatile transfer. Acidic lavas are peraluminous and display a crustal signature. The prevailing contribution of a depleted mantle source for the N-MORB-like basaltic magma implies asthenosphere melting and thinning of the continental lithosphere. However, the following Cambraco–Ordovician sedimentary deposition did not tally with a marine basin. The mafic lavas are limited to high-magnesia cumulate facies having recorded crustal contamination. They include abundant partly melted sialic xenoliths. The associated acidic lavas could have resulted from crustal melting. One may assume that the magma was ponded during a long time inside a crustal reservoir. Any MORB-like fresh magma never reached the surface as a result of which, we conclude to the abortion of the rifting process.

The third latest early Cambrian phase (V3) is recorded in the eastern Lacuana Mountains of the northern Montagne Noire. Abundant olivine-basalt lava flows and sills are interbedded in the Série Noire Formation, a kerogenous black shale and limestone unit deposited in slope-related to deep basinal settings, and constitute the Saint-Méen Volcanic Complex. The basin is saturated, moderately magnesian, REE-poor and weakly fractionated in the LREE, but enriched in lithophile elements. The magma source resembles that of the V2 basalts with a more important contribution of an enriched component of lithospheric mantle. An intra-crustal ponding is also inferred. Owing to the lithosphere contribution, the rifting process did not evolve. On the contrary, the mantle cooled down according to the decreasing involvement of the asthenosphere, the lack of crustal melting and the deepening of the sedimentary basin.

The fourth phase (V4) was the most dramatic volcanic event in the Occitan Domain occurring in the early Tremadocian with rising up of the basin floors and subaerial explosive and effusive abundant rhyolitic activities. This event supplied (1) the Larroque Volcanic Formation, described as parautochthonous units in the Rouergue and allochthonous units in the Albigéois and northern Montagne Noire, (2) the Davejean acidic volcanic part in the Mouthoumet Massif, and (3) the Génis rhyolitic unit of the western Limousin area that is correlated with the Tremadocian rhyolitic activity of the Chantonnay and La Roche-sur-Yon basins, and of the Saint-Gilles nappe. This acidic magmatic output resulted in massive crustal melting requiring a rather important heat supply. Such a great thermal anomaly can only be explained by asthenospheric up-welling leading to lithospheric doming, continental break-up, and a decompressively driven mantle melting. Magmatic products accumulated at the mantle-crust boundary and provided heat transfer for crustal melting (Hupt and Sparks, 1988).

The Tremadocian fifth phase (V5) is the logical complement of the V4 crustal melting event. Emptying of the rhyolitic crustal reservoirs caused collapses and sinking of basin floors. Once the removal of the intra-crustal barrier formed by the acidic magma was achieved, some parts of the underplated basic magma could ascend and reach the surface. This magmatism is depicted by flows and sills of the Peyrebrune Volcanic Complex and the Davejean Volcanic Complex basalts. Basaltic rocks are evolved, REE-rich, and LREE-fractionated. They display a typical initial rift tholeiite (IRT) magmatic signature indicating contributions of both asthenospheric and lithospheric mantle sources.

5. Northern Gondwana volcano-tectonic activity

Correlations are tempted between the volcano-tectonic activities of the South Armorican and Occitan domains and are extended to the Northern Gondwana Variscan Belt in the Iberian Zones of the western Ibero-Armorican Arc branch (Fig. 1).

5.1. From Cambrian extensional pulses to Tremadocian continental break-up

The latest Ediacaran to earliest Cambrian rhyolitic activities are distributed in the northern flank (Rivernous Volcanic Complex, V1 volcanic phase) and the Axial Zone (Schistes X) of the Montagne Noire, and also in the eastern Pyrenees (Canaveilles Group of the southern Canigou Massif). They are dated around 545 and 540 Ma. These rhyolitic activities are younger than those of northern Iberia, dated between 575 and 560 Ma (Rubio-Ordóñez et al., 2015), and attributed to the arc-related final stage of the Cadomian orogeny. They could be contemporaneous with those of the latest stage of the Ediacaran andesitic to rhyolitic Volcanic Chain of the Moroccan Anti-Atlas dated between 578 and 543 Ma (Gasquet et al., 2005; Pouclet et al., 2007), somewhat correlatable with the Bodonal acidic volcanics of the Ossa-Morena Rift (Sánchez-García et al., 2010; Álvaro et al., 2014b). As already reported for the Rivernous Rhyolitic Complex (Álvaro et al., 2014a), the Occitan earliest Cambrian acidic volcanics are not related to any Cadomian or Pan-African tectonic event; until now, they have no pertinent tectonic relationships.

The lower Cambrian mafic and acidic lavas (V2 volcanic phase) are localized in the southeastern tectonic slices of the northern Montagne Noire, Guérangé–Lozès and Burg (1990) suggested that the volcanic activity was initiated along former NE–SW trending fractures that controlled the following décollement of the thrust. Unfortunately, it is not possible to check if the present-day NE–SW elongated setting of the volcanic bodies parallel to the foliation of the metasedimentary formations corresponds or not to the initial volcanic fissural system. The tholeiitic depleted magmatic signature of these earliest lavas fits well with an intracontinental extensional fracturing of incipient rifting. However, the source enrichment of the following lavas implying renewal
contribution of the lithospheric source component (V3 volcanic phase) attests for a decreasing activity of any rifting process.

The lower Cambrian sequences are not well constrained in the South Armorican Domain. The middle Cambrian to Furongian basaltic activities of the Cholet and Chantonnay basins displays a continental tholeiitic magmatic signature have similarities with the Engasten Complex of the Thiviers-Payzac Basin of West-Limousin, but there are no clear equivalents in the Occitan Domain. It is concluded that, except for a limited part of the Occitan Domain where incipient rifting took place (Ensèges Volcanic Complex), the Cambrian volcano-tectonic activity was limited to intracratonic fracturing and outpouring of continental tholeiites and secondary felsic lavas.

The more conspicuous correlative is the widespread early Tremadocian acidic subaerial volcanism. This crustal-originated massive melting was associated with a major continental break-up. From northwest to southeast, the Tremadocian rhylitic activities are distributed in autochthonous and allochthonous terrains originated from various sedimentary basins. Occurrences of these volcanics are located in the Fig. 22, as well as the lower Ordovician orthogneisses.

The more accurate dating is available in the Vendean area. The parautochthonous La Chapelle-Hermier rhylites of the Roche-sur-Yon Basin (472–486 Ma) are contemporaneous with those of the Saint-Gilles and Mareuil nappes (477–486 Ma) and the Chantonnay Basin (470–485 Ma). Structural features attest that the nappes recorded a westward transport (Burg, 1981; Brun and Burg, 1982; Burg et al., 1987; Béchennec et al., 2008; Lahondère et al., 2009). The closest eastern possible position for the thrust material is the Intermediate Structural Zone (Fig. 2), where orthogneisses bodies could be regarded as magmatic chambers of the volcanics, namely Les Essarts orthogneiss dated at 483 ± 4 Ma and the Mervent orthogneiss dated at 486 ± 15–11 Ma (see Section 2.1.3). To the northeast, poorly dated rhyolites of the Cholet Basin may be allocthonous to the same Tremadocian volcanic event. To the northwest, Tremadocian orthogneisses is recognized in the Champtoceaux Complex (485 ± 11 Ma; Paquette et al., 1984; and 481 ± 6–5 Ma; Ballèvre et al., 2002). Thanks to the restoration of the ductile strain of the regional deformation in a simple-shear model, Giumiaux et al. (2004a) demonstrated that the Champtoceaux Complex was the continuation of the Les Essarts Complex and thus, their orthogneisses originated from the same fracture. In the south of Brittany, the Tremadocian magmatic activity is marked out by the Belle-Île—La Vilaine continuation of the rhyolite-bearing nappe of Saint-Gilles (Ballèvre et al., 2012), the orthogneisses of Lanvaux (478 ± 18 Ma; Janjou et al., 1998), Moëlan (486 ± 6 Ma; Calvez, 1976), Loc'h (483 ± 3 Ma; Béchennec et al., 1999), Cosquer and Lanmeur (497 ± 11 Ma and 498 ± 12 Ma; Guerrot et al., 1997), the metamorphites of Merrien (481 ± 11 Ma; Guerrot et al., 1997), and the orthogneiss (480 ± 8 Ma; Paquette et al., 1985) of the eclogite-bearing basic-ultrabasic complex of Baie d'Audierne.

Taking into account the distribution of the rhyolites and related orthogneisses, there are supporting evidences that the rhylitic activity of the South Armorican Domain outfractured from a fracture swarm, though the present-day location of the fracture zone is depending on the thrust displacements and the shear zone motions. However, there are strong evidences that the late Variscan shear zones reworked such a previous fracture system.

To the southeast, the continuation of the structural pattern from South Armorica to Massif Central below the sedimentary cover of the Seuil du Poitou is done in the structural map of Rolin and Colchen (2001a) interpreting geophysical and bore hole data. Using this map, we draw a sketch map of the South Armorican and Occitan domains in Fig. 23. The fracture zone of the Mervent orthogneiss, the southernmost branch of the South Armorican Shear Zone, is prolonged until the NW–SE dextral Parthenay Shear Zone (PSZ). Similarly, the MSSZ joins the PSZ, in setting up the eastern border of the lens-shaped Chantonnay Basin. After a 120 km southeast offset along the PSZ, the fracture system renews along the Estivaux Fault (EsF) bordering the Thiviers-Payzac Basin with orthogneiss elongated bodies dated to the early and mid-Ordovician, and with the Genis Tremadocian rhyolites (Guilhot et al., 1977; Roig et al., 1996; Melleton et al., 2010). Farther to the southeast, the La Bessenoits klippe of eclogite-bearing leptyno-amphibolite complex includes an orthogneiss dated to 481 ± 1 Ma (Paquette et al., 1995). The similar Levezou Complex includes a leptynite dated to 480 Ma (Pin and Lancelot, 1982). Lastly, a thick nappe stack of rhyolite extended from the Rouergue and Albigenses region to the northern Montagne Noire. This is the Larroque Volcanic Complex correlated with the dated early Tremadocian Vendean rhyolites. Fig. 23 clearly illustrates the wide setting of the Tremadocian rhylitic activities and there is no doubt that the Tremadocian rhyolitic fracture system of the Occitan Domain is the continuation of the South Armorican one. Another distinct, though less important, rhyolitic activity was also recorded in the Mouthoumet Massif.

All these volcanoes pertain to a “Tremadocian Volcanic Chain” emplaced along the north Gondwana basins. Most of these basins are parautochthonous. The Occitan basins were moderately thrust to the south or south-west. The Saint-Gilles and cognate basins were thrust to the west. The upper allochthonous position of the northern basins is due to late orogenic northward backthrusting.

5.2. The Iberian link: toward a Tremadocian tectono-magmatic correlation

There is an extensive body of literature on the correlates between the South Armorican-Occitan and Iberian branches of the Variscan Arc based on their structural, lithological and petrological salient features (synthesis in Ballèvre et al., 2014). We already approached the chronostatigraphic and petrological equivalence between Tremadocian South Armorican-Occitan and Iberian metarhyolites (Figs. 8b and 17). The Tremadocian magmatic activity of associated orthogneisses and metarhyolites from the Iberian Domain is illustrated in Fig. 24. The question is: which are the lithostructural relationships between the two branches of the Ibero-Armorican Arc? Discriminating Tremadocian rhyolitic-granitic magmatic complexes could provide appropriate guides for connecting both branches.

In Fig. 22, the suspected location of Iberia is drawn for the Carboniferous–Permian time assuming: (1) a 250 km right lateral offset along the North-Pyrenean fault system (Raymond, 1987); and (2) a 35° clockwise rotation for cancelling the Gulf of Biscay (or Gulf of Gascoing) opening (Perroud and Bonhommet, 1981; Garcia-Mondéjar, 1996; Sibuet et al., 2004). The Iberian lower Ordovician magmatic formations dominated by acidic volcanics are widely distributed into four magmatic complexes or systems at sites #1, 2, 3, and 4. The first magmatic site is located in the Cantabrian Zone and in the easternmost site of the West Asturian-Leonese Zone (#1a and #1b). It is limited to few sills and beds of rhyolitic lavas and volcanioclastics from the upper Barrios Formation at the base of the “Armorican Quartzite” where it is dated to 477.5 ± 0.9 Ma (early Floian), and in the upper part of the lower Tremadocian Borrachón Formation (Gutiérrez-Alonso et al., 2007; Alvaro et al., 2008). The second site is the Olio de Sapo Formation located on eastern side of the Central Iberian Zone and consisting of acidic volcanic and volcanosedimentary rocks interbedded at the base of the lower Ordovician platform-facies sediments (Valverde-Vaquer and Dunning, 2000; Montero et al., 2007, 2009; Díez Montes et al., 2010; Talavera et al., 2013). The formation extends from the Cantabrian coast to the Hiendelaencina region, in the easternmost part of the Central Iberian Zone, and includes orthogneiss bodies of Guadarrama and Miranda do Douro (Bea et al., 2006; Zeck et al.,...
2007) (#2a and #2b). The total length of this belt reaches ca. 570 km, with a large and subcontinuous outcrop of the Sanabria northwestern section of 28 km in width, and a thickness of ca. 2000 m. The formation is folded in a complex anticlinorium structure overlying orthogneissic domes. Reverse faults and recumbent folds developed above the doming core, but the whole complex remains autochthonous and is interpreted as a succession of volcanic domes of dacites and rhyolites. The granitic core is considered as the subvolcanic magmatic reservoir. The third site gathers the Rio Baio, Mora and Saldanha volcanic complexes in the Galicia-Tras os Montes subzone, and associated orthogneisses extended in the Castillan subzone (Valverde-Vaquero et al., 2005; Talavera et al., 2008, 2013; Bea et al., 2009; Andonaegui et al., 2012; Días da Silva et al., 2014) (#3a and #3b). The felsic lavas are interbedded in a thrust sheet between the autochthonous Ollo de Sapo antiform and the Ordenes, Cabo Ortegal, and Morais Massif upper thrusts. They are located in the Schistose Domain of Central Iberian Zone and consist of a lower mafic and acidic suite and of an upper acidic volcaniclastic ensemble. The fourth site includes the Urra Rhyolitic Formation in western margin of the Lusitanian-Alcudian subzone of Central Iberian Zone and the tonalite belt of the Central Extremadure batholith (Solá et al., 2005, 2008; Antunes et al., 2009; Neiva et al., 2009; Rubio-Ordóñez et al., 2012) (#4a and #4b). The Urra Rhyolitic Formation comprises a lower prophyritic unit ca. 200 m thick and an upper volcaniclastic unit ca. 500 m thick. The associated plutons are calc-alkaline granites and related diorites and gabbros. The volcanosedimentary sequence is overlain by the "Armorican Quartzite Formation".

Age datings of the Furongian to early Ordovician magmatic activity limited to the more accurate U-Pb zircon or monazite data are plotted in Fig. 23, also with a distinction between rhyolitic volcanic products and genetically related granitic plutons suspected to belong to magmatic reservoirs of the volcanic counterparts. Ages of rhyolites of the Vendean area (Chantonnay, La Chapelle-Hermier, Saint-Gilles and Mareuil-sur-Lay) sharing the same petrological patterns are not very different to those of the Ollo de Sapo and Guadarrama, taking into account bracketed errors. An earliest Tremadocian age is recorded. Few accurate ages are available for the Mora-Saldanha and Urra volcanics. The Urra site seems to be older and of late Furongian age. But, the Urra volcanic sequence is overlain by the "Armorican Quartzite Formation" that is dated to early Floian (Gutiérrez-Alonso et al., 2007). The associated plutons to volcanics, as demonstrated in the Ollo de Sapo antiform (Díez Montes et al., 2010), are of similar ages. Indeed, ash-fall tuff beds related to the Ollo de Sapo eruptive activity are dated to 477 Ma (4 samples, Gutiérrez-Alonso et al., 2016). One may note that the Pyrenean orthogneisses are clearly younger and belong to a distinct magmatic event.

In the South Armorican Domain, assuming that the Les Essarts Complex orthogneisses are the endogenous sources of Saint-Gilles rhyolitic nappes, there are close likenesses with the Ollo de Sapo magmatic and structural system (#2, Fig. 22) in term of high amounts of magmatic supply and extensive tectonic line. Differences are
explained by much more severe shearing compression and strike slip motion for the Vendean nappe rhyolite leading to décollement of the volcanic cover, while the Ollo de Sapo Formation shows limited thrusting. In the Occitan Domain, owing to its high volume of magmatic flows, the Larroque Volcanic Formation (Figs. 4 and 5), of similar stratigraphic age and petrological significance, is equivalent to the Ollo de Sapo tectono-volcanic system.

It is concluded that a Tremadocian Tectonic Belt extended from South Armorican and Occitan domains to Iberian Zones in a swarm of fractures well developed in the Iberian branch, but compressed in the SAOD due to severe thrust and right-lateral shearing. In the South Armorican Domain, this fracture system was reworked along the southern branches of the South Armorican Shear Zone (Figs. 2 and 23). The other fracture zones with their rhyolitic products are constrained by the onset of the Sardic phase (Álvaro et al., 2016). Consequently, the Tremadocian continental break-up was not immediately followed by a marine basin opening. The IRTs were generated during the Tremadocian asthenosphere upwelling and were stored below rhyolitic reservoirs, the high volume of which constituted a barrier for the ascent of deep mafic magmas. After the emptying of the acidic magma chambers and the collapse of the magmatic domes, residual mafic magmas were able to reach the surface. Indeed, where the volume of rhyolitic magma is low, both acidic and mafic magmas were generated during the same geotectonic event (Fig. 4).

5.3. The question of the propagation of the Tremadocian Tectonic Belt

The Tremadocian Tectonic Belt of the Vendean Intermediate Structural Zone was the source of westward thrust Saint-Gilles – Belle-Île volcanic and sedimentary sequences (Fig. 2). Indeed, eclogites have been exhumed in this structural zone and are suspected to originate from a subducted oceanic crust. Can we infer that the Tremadocian breaking evolved from a continental rift to an ocean basin? In the South Armorican Domain, Tremadocian subaerial rhyolitic lavas are overlain by quartz-rich conglomerates and sandstones belonging to the widespread siliciclastic platform of the “Armorican Quartzite” facies, a common detrital witness of a syn-rift stage, particularly in the Chantonnay Basin (Fig. 3). Into rifted basins, these coarse deposits were fining upward to shales before being disconformably overlain by the uppermost Ordovician and Silurian shales after a marine basin opening. The IRTs were emptied during the Tremadocian asthenosphere upwelling and were stored below rhyolitic reservoirs, the high volume of which constituted a barrier for the ascent of deep mafic magmas. After the emptying of the acidic magma chambers and the collapse of the magmatic domes, residual mafic magmas were able to reach the surface. Indeed, where the volume of rhyolitic magma is low, both acidic and mafic magmas reached the surface together, as in the volcanic arc and the accretion of the Qilian block (Von Raumer and Stampfl, 2008; Von Raumer et al., 2013, 2015), also interpreted as the collision of Gondwana with the Qaidam arc and the accretion of the Rheic Ocean opening patterns (Helbing and Tiepolo, 2005), also interpreted as the collision of Gondwana with the Qaidam arc and the accretion of the Qilian block (Von Raumer and Stampfl, 2008; Von Raumer et al., 2013, 2015), the previous framework of tectonic troughs was reactivated. Local troughs were...
opened and their palaeoreliefs finally sealed by Silurian and Devo-
nian transgressive deposits (Álvaro et al., 2016). The extensional
strain focused along the former Tremadocian Tectonic Belt and an
incipient oceanization was reached, as proved by the eclogite
remnants. Some eclogites show an N-MORB composition and
originated from a depleted mantle, but others display E-MORB, IAT,
IRT, and CT geochemical signatures implying contribution of both
depleted and enriched mantle sources (Bernard-Griffiths and Jahn,
1981; Bernard-Griffiths and Cornichet, 1985; Paquette et al., 1985;
Piboule and Briand, 1985; Berger et al., 2010). These features are
consistent with a continent-ocean transition zone and narrow
oceanic basins. Similarly, the metabasites of the Île-de-Groix nappe
display two contrasting magmatic series mixing transitional tholeiites (E-MORB type) and alkali-basalts originated from unenriched
or moderately enriched mantle sources, though isotopically
depleted (Bernard-Griffiths et al., 1986; El Khor et al., 2013). Such
magmatic features are indicative of both ocean-floor and continen-
tal margin. Ages of the eclogites are poorly constrained owing to
discordant interpretations of the various U-Pb zircon ages as in-
heritance, magmatic crystallization or post-magmatic and meta-
morphic processes. First isotopic data advocated mid Silurian to
early Devonian ages for the high-pressure eclogite facies meta-
morphism (Peucat et al., 1982; Paquette et al., 1985; Berger et al.,
2010). More accurate analyses combining Sm-Nd, Rb-Sr and
40Ar/39Ar methods support a late Devonian age (ca. 390 and 360 Ma)
for two eclogite units of the Champtoceaux Complex (Bosse et al.,
2000). The high-pressure event of the eclogites from the
Montagne-Noire Axial Zone is dated at 358 Ma (Faure et al., 2014).
Nevertheless, the eclogites may have different sources and different
ages. The eclogites of Les Essarts, Champtoceaux and Audierne
complexes are usually allotted to an ocean
floor subducted until to
about 50–60 km in depth (13 kbars) (Godard, 1988, 2001a, 2010;
Lucks et al., 2002). This hypothesis is supported by seismic imaging
showing a suspected slab that subducted northward below the
South Armorican margin (Bitri et al., 2001, 2003, 2010; Judenherc
et al., 2003; Gumiaux et al., 2004b). It is thus admitted that an

Figure 24. Timing of the rhyolitic (a) and granitic (b) activity in the South Armorican, Occitan and Iberian tectono-magmatic lineaments across the Cambrian–Ordovician transition. Age dating from the literature. References in the text.
oceanic crust existed in the Champtoceaux–Les Essarts Siluro–Devonian structural zone and in its prolongation to the western Limousin (Berger et al., 2006, 2010). In the Vendean area, the Chantonnay Basin infilled with under-water flowing of basaltic lavas of the La Meilleria Volcanic Formation above a middle–upper Devonian sedimentary terrigenous sequence (Fig. 3). The volcanic rocks consist of transitional tholeiites originated from two mixed depleted and moderately enriched mantle sources (Thiéblemont, 1988) (see Sections 2.1.2. and 4.1.). Geochemical and unpublished Rb/Sr and Nd/Sm isotopic analyses (Thiéblemont, pers. comm. 2014) proved that the La Meilleria magmatic sources are similar to those of the Les Essarts eclogites that constituted the floor of the neighbouring marine basin drifted from the former Champtoceaux–Les Essarts Ordovician fracture zone. It has also been shown that the western edge of this basin is rimed by the Saint-Martin-des-Noyers metabasites displaying more or less depleted geochemical signatures similar to those of the La Meilleria metabasalts and the eclogites (Section 3.1.5). These metabasites can thus be attributed to the southwestern margin of the basin whilst the eclogites constituted the protorift of the Siluro–Devonian one. Eclogites are located in the Malpica-Tui Unit, and in the Cabo Ortegal, Órdenes, Bragança and Morais upper allochthonous complexes. The Malpica-Tui Unit emplaced in the lower allochthon that represents a distal part of the Gondwana continental margin (Martínez Catalán et al., 1996). The complexes are suspected to derive from the Rheic Ocean (Martínez Catalán et al., 2009). In that case, they are equivalent to the eclogite-bearing Léon terrains of Brittany (Cabanis and Godard, 1987; Faure et al., 2010). This assumption questions the setting of the Rheic Ocean between Iberia and Armorica and in northwestern Iberia, either east (Martínez Catalán et al., 2009) or west (Simancas et al., 2009) of the northwesternmost Schistose Domain autochthon and its Malpica-Tui Unit lower allochthon. The eastern location is supported by the assumed existence of a volcanic arc according to the suspected calc-alkaline composition of mafic Cambrian magmatic rocks (Andonaegui et al., 2002, 2012, 2016). However, the composition of these rocks is consistent with MORB-like olivine-tholeiites with moderate crustal assimilation, but clearly differs from calc-alkaline arc-related basalts. This composition should agree with the pull-apart basin model proposed by Arenas et al. (2014). By contrast, these mafic rocks are similar to Cambrian lavas of the Montagne Noire (i.e. V2 and V3 volcanic phases). Alternatively, the Iberian allochthonous eclogite-bearing formations may have originated from the continuation of the Audierne area that was reactivated during the Siluro–Devonian opening of South Armorican marine basin. The Malpica-Tui eclogites are considered as equivalent to Champtoceaux–Cellier eclogites (Ballèvre et al., 2008, 2014). In this framework, the thrust sheets of the northwest Iberian allochthonous units would be a remnant of a marine basin that extended from the South Armorican Basin, and not of the Rheic Ocean. This last interpretation is in a good agreement with the model of Ribeiro et al. (2007). For these authors, the allochthonous units resulted from obduction of the western Palaeotethys that is the South-Armorican Ocean. This assumption is also proposed as an alternative model by Ballèvre et al. (2014). Indeed, at the time of the Tremadocian continental break-up episode, the Rheic Ocean was diachronously spreading all along the western edge of Iberian and Armorican domains between Gondwana and Avalonia as shown in the Fig. 25. However, the southwestward propagation (in present coordinates) of the suspected South-Armorican Ocean remains questionable. In the Moroccan High Atlas and Anti-Atlas, any large Tremadocian rhyolitic activity is absent. The rift mainly opened in the earliest Cambrian (Pouclet et al., 2008; Alvaro et al., 2014b) and is related to...
the Tremadocian Volcanic Chain, a rhyolitic activity that occurred in the southern part of the Occitan Domain and in the Pyrenean Domain. The lack of pre-Variscan deformational processes recorded in the South Armorican and Occitan domains are revised in order to conduct a petrographical and geochemical study of the Cambrian—Lower Ordovician volcanic formations.

Across the Precambrian—Cambrian transition, a rhyolitic activity is recorded in the southern part of the Occitan Domain and in the Pyrenean Domain. In the lack of pre-Variscan deformational process, this magmatic activity cannot be allotted to the Cadomian or Pyrenean Domain. In the absence of pre-Variscan deformational processes recorded in the southern part of the Occitan Domain and in the Pyrenean Domain, the Tremadocian Volcanic Chain was initiated by tectonic thinning of the lithosphere and upwelling of the asthenosphere. Mantle-derived continental tholeiites having some trace element patterns of N-MORBs. The late early Cambrian, the Tremadocian transition of the South Armorican and Occitan domains built the rift and a changing crustal regime ran out in the continental margin.

6. Conclusions

The geological background and tectonostratigraphic patterns of the South Armorican and Occitan domains are revised in order to conduct a petrographical and geochemical study of the Cambrian—Lower Ordovician volcanic formations.

We postulate that a major continental break-up, the “Tremadocian Tectonic Belt”, was initiated by tectonic thinning of the lithosphere and upwelling of the asthenosphere. Mantle-derived mafic magmas intruded the crust and were underplated at the mantle-crust transition zone. This hot material provided heat for melting the crust which supplied the rhyolitic volcanism of the volcanic chain. After emptying of the rhyolitic crustal reservoirs, underlying mafic magmas could ascend and reach the surface. An attempt is done to continue the Tremadocian Tectonic Belt from South Armorica to Iberia. The extensional process was developed in the Occitan to Iberian continental Cadomian margin while the Rheic Ocean was spreading along this margin in separating the Avalonian and Cadomian terranes. This process was interrupted in the mid-to-late Ordovician times and renewed in the Silurian. Then, the oceanization took place along the inherited fracture zone of the Tremadocian Tectonic Belt that acted as a lithospheric plate margin and focused all the driving forces. Indeed, the Tremadocian fracture network guided the opening of the Siluro—Devonian Massif Central—South Armorican Ocean that evolved to part of the Palaeotethys Ocean.

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