Three years of semicontinuous greenhouse gas measurements at the Puy de Dôme station (central France)

M. Lopez1,a, M. Schmidt1,b, M. Ramonet1, J.-L. Bonne1, A. Colomb2, V. Kazan1, P. Laj3, and J.-M. Pichon2,4

1Laboratoire des Sciences du Climat et de l’Environnement (LSCE – UMR8212), Unité mixte CEA-CNRS-UVSQ, 91191 Gif-sur-Yvette, France
2Laboratoire de Météorologie Physique, Université Blaise Pascal, Clermont-Université, CNRS, UMR 6016, 63171 Aubière, France
3Université Grenoble Alpes, CNRS, Laboratoire de Glaciologie et Géophysique de l’Environnement (LGGE – UMR 5183), 38402 Saint-Martin d’Hères, France
4Observatoire de Physique du Globe de Clermont-Ferrand, Université Blaise Pascal, Clermont Université, CNRS, UMS 833, 63171 Aubière, France

a now at: Environment Canada, Climate Research Division, Toronto, Ontario, Canada
b now at: Institut für Umweltphysik, University of Heidelberg, Heidelberg, Germany

Correspondence to: M. Ramonet (michel.ramonet@lsce.ipsl.fr)

Received: 28 November 2014 – Published in Atmos. Meas. Tech. Discuss.: 20 March 2015
Revised: 19 August 2015 – Accepted: 22 August 2015 – Published: 25 September 2015

Abstract. Three years of greenhouse gas measurements, obtained using a gas chromatograph (GC) system located at the Puy de Dôme station at 1465 m a.s.l. in central France, are presented. The GC system was installed in 2010 at Puy de Dôme and was designed for automatic and accurate semicontinuous measurements of atmospheric carbon dioxide, methane, nitrous oxide and sulfur hexafluoride mole fractions. We present in detail the instrumental setup and the calibration strategy, which together allow the GC to reach repeatabilities of 0.1 µmol mol⁻¹, 1.2 nmol mol⁻¹, 0.3 nmol mol⁻¹ and 0.06 pmol mol⁻¹ for CO₂, CH₄, N₂O and SF₆, respectively. The analysis of the 3-year atmospheric time series revealed how the planetary boundary layer height drives the mole fractions observed at a mountain site such as Puy de Dôme where air masses alternate between the planetary boundary layer and the free troposphere.

Accurate long-lived greenhouse gas measurements collocated with ²²²Rn measurements as an atmospheric tracer allowed us to determine the CO₂, CH₄ and N₂O emissions in the station catchment area were 7.0 ± 4.0 t CH₄ km⁻² yr⁻¹ and 1.8 ± 1.0 t N₂O km⁻² yr⁻¹, respectively. Our derived annual CH₄ flux is in agreement with the national French inventory, whereas our derived N₂O flux is 5 times larger than the same inventory.

1 Introduction

The release of anthropogenic greenhouse gases (GHGs) into the atmosphere leads to a modification of their natural cycles and to a strong increase in atmospheric radiative forcing (Myhre et al., 2013). The Intergovernmental Panel on Climate Change (IPCC) reported that the global average temperature increased by 0.89 °C between 1901 and 2012 (Hartmann et al., 2013) and will continue to increase during the 21st century (Collins et al., 2013). To limit the global temperature rise, most industrialized countries signed the “United Nations Framework Convention on Climate Change” (UNFCCC) treaty in 1992 to stabilize their GHG emissions between 1990 and 2000 and it entered into force in 1994. This convention was enhanced by the Kyoto Protocol, which was signed in 1997 and was ratified by 182 countries.
countries engaged in the Kyoto Protocol aimed to reduce their national emissions of the main long-lived GHGs by 5.2% between 2008 and 2012 compared to the emission levels of 1990. The GHGs in question are carbon dioxide (CO$_2$), methane (CH$_4$), nitrous oxide (N$_2$O), sulfur hexafluoride (SF$_6$), hydrofluorocarbons (HFCs) and perfluorocarbons (PFCs). The European Union (EU) committed itself to a reduction of its GHG emissions by 8% for the same period. In addition, the EU aims to reduce its total GHG emissions by 20% in 2020, relative to emissions in 1990. Despite this commitment, it is extremely difficult to validate the surface GHG fluxes on the country scale using a reliable, transparent method.

Currently, countries report their respective GHG emissions to the UNFCCC on an annual basis. These national emission inventories are based on bottom-up methods and the reliability of these national inventories strongly depends on the uncertainty attributed to each emission factor. To improve and validate the bottom-up methods, it is crucial to better characterize the biogeochemical cycles of the different GHGs, particularly as national inventories report only anthropogenic emissions to the UNFCCC. Therefore, it is important to develop new tools to quantify natural-source emissions and to provide an independent verification of the emission inventories reported to the UNFCCC.

Different methods based on atmospheric measurements can be used to estimate GHG emissions on local to regional scales. Some of these approaches couple atmospheric GHG measurements with measurements of associated atmospheric tracers of air masses, including radon-222 (Biraud et al., 2000; Schmidt et al., 2001), sulfur hexafluoride (Maiss et al., 1996) or isotopes, such as radiocarbon in CO$_2$ (Levin and Karsten, 2007; Lopez et al., 2013). A major advantage of this “multigas” approach is that it avoids the use of complex chemistry-transport models; the tracers that are used are subject to the same atmospheric transport mechanisms as the GHGs. Nevertheless, an accurate assessment of the respective measurement station footprint is required to allocate the estimated surface fluxes to a specific region (Gloor et al., 2001).

The first atmospheric CO$_2$ continuous measurements started in the 1950s at Mauna Loa Observatory using a nondispersive infrared (NDIR) analyzer with a repeatability better than 0.1 µmol mol$^{-1}$ (Keeling et al., 1976). Atmospheric CH$_4$ monitoring began in the late 1970s using gas chromatograph (GC) systems equipped with a flame ionization detector (FID) and a nickel catalyst to enable simultaneous CO$_2$ mole fraction detection (Rasmussen and Khalil, 1981). The repeatabilities were approximately 10 nmol mol$^{-1}$ for CH$_4$ measurements and 0.7 µmol mol$^{-1}$ for CO$_2$ measurements. Coupling an electron capture detector (ECD) to a GC system enabled the detection of N$_2$O and SF$_6$ atmospheric mole fractions with repeatabilities of approximately 1.0 nmol mol$^{-1}$ and 0.1 pmol mol$^{-1}$, respectively (Weiss, 1981; Prinn et al., 1990; Maiss et al., 1996). Subsequently, these two detectors (FID and ECD) have permitted the use of GC to analyze CO$_2$, CH$_4$, N$_2$O and SF$_6$ atmospheric mole fractions simultaneously and on a semi-continuous basis. Since the 1980s, atmospheric monitoring stations that are part of the Global Atmosphere Watch (GAW) have been gradually equipped with GC systems and NDIR analyzers. The GC system technologies described above have continuously evolved to reach repeatabilities better than 0.1 µmol mol$^{-1}$ for CO$_2$, 2.0 nmol mol$^{-1}$ for CH$_4$, 0.3 nmol mol$^{-1}$ for N$_2$O and 0.1 pmol mol$^{-1}$ for SF$_6$, as shown by van der Laan et al. (2009b), Thompson et al. (2009) and Popa et al. (2010).

New types of accurate instruments for CO$_2$, CH$_4$ and N$_2$O atmospheric measurements have recently become commercially available. These instruments are based on optical technologies, including cavity ring-down spectroscopy (CRDS), Fourier transform infrared spectrometry (FTIR) and off-axis integrated cavity output spectroscopy (OA-ICOS). These recent technologies are promising for atmospheric monitoring as they offer high-frequency measurements (on the order of 1 Hz), require low maintenance and achieve equivalent or superior repeatability compared to GC systems (Yver Kwok et al., 2015; Hammer et al., 2013a). Analyzers based on the CRDS technology are generally used for CO$_2$ and CH$_4$ atmospheric measurements, OA-ICOS technology is used for N$_2$O measurements, and FTIR technology is designed to simultaneously measure CO$_2$, CH$_4$ and N$_2$O. These new types of instruments are also more easily transportable, and Hammer et al. (2013b) have demonstrated their feasibility as “traveling” instruments. They could thus be used for comparisons and quality control purposes to ensure data compatibility across a monitoring network.

Regardless of the benefits listed above, these optical technologies cannot yet be used to measure atmospheric SF$_6$ mole fractions; SF$_6$ is an extremely stable GHG, having a global warming potential of 23 900 (Forster et al., 2007). In addition, most of these new technologies need to be continuously flushed, which makes it difficult to analyze flasks, in contrast to the GC systems, which employ discrete samples and can analyze four or five species simultaneously (van der Laan et al., 2009b). Studies of optical technologies are still progressing, and, as noted above, these new technologies are very promising, particularly for a dense monitoring network, such as the European infrastructure ICOS (Integrated Carbon Observation System), which is dedicated to high-precision monitoring of greenhouse gases over Europe. The CRDS technology is slowly replacing the GC systems and NDIR analyzers for CO$_2$ and CH$_4$ monitoring in many stations, but GC is still the reference instrument for N$_2$O and SF$_6$ measurements (see WMO-GAW, 2013). Consequently, we installed a GC system in 2010 at the mountain station of Puy de Dôme (France) to monitor, with a high precision, the long-term atmospheric trend of the main four long-lived greenhouse gases.
This paper focuses on 3 years (2010 to 2013) of ambient air measurements of CO$_2$, CH$_4$, N$_2$O and SF$_6$, obtained using a GC system at the Puy de Dôme station. After a short description of the station, the detailed setup of the GC as well as the calibration strategy are addressed. A paragraph is dedicated to data quality control and atmospheric measurement comparisons, showing to what extent our measurement system are in line with the WMO recommendations (WMO-GAW, 2013). In the last part, we present and analyze our 3-year time series of ambient air measurements. Finally, we demonstrate that these time series can be used to estimate the monthly regional fluxes of CO$_2$, CH$_4$ and N$_2$O in the catchment area of the Puy de Dôme station, using radon-222 as an atmospheric tracer.

2 The Puy de Dôme station

2.1 Site description

The Puy de Dôme station (45$^\circ$46'19"N, 2$^\circ$57'57"E) is located at the top of the Puy de Dôme volcano (1465 m a.s.l.), in Auvergne in the center of France. This station is managed by the Laboratoire de Météorologie Physique (LaMIP) and is part of the Observatoire de Physique du Globe de Clermont-Ferrand (OPGC) located at Clermont-Ferrand, France. According to the French National Institute of Statistics and Economic Studies (INSEE – http://www.insee.fr), the ground cover of Auvergne (26 013 km$^2$) consists mainly of meadows (36.4 %), forests (33.4 %) and arable land (17.6 %), the Puy de Dôme station being located in the center of this region. The major anthropogenic GHG sources are to the east of the station, where the town of Clermont-Ferrand is located: 10 km east of the Puy de Dôme station at an altitude of 396 m a.s.l. Clermont-Ferrand is the largest town in the region, with approximately 150 000 inhabitants.

The CITEPA (Centre d’Etudes de la Pollution Atmosphérique) reports the French national GHG emissions to the UNFCCC but has also provided a regional inventory of Auvergne for the year 2007 (CITEPA, 2010). According to the CITEPA, the anthropogenic CO$_2$ emissions in Auvergne are mainly attributable to road transport and residential and industrial sectors, which represent respectively 45, 25 and 21 % of the total anthropogenic CO$_2$ emissions of the region. The agricultural sector is responsible for 90 % of total anthropogenic CH$_4$ emissions and 97 % of total anthropogenic N$_2$O emissions in the region. Ninety percent of the SF$_6$ emissions are related to the energy transformation sector.

A military base and a telecommunication center are located 20 m northwest of the station, also on the top of the volcano. These facilities consist of a main building (20 m height) and a telecommunication antenna (89 m height). Since 2010, the only access road to the station has been closed to the public and has been replaced by a cog train.

The atmospheric research station hosts different analyzers for long-term atmospheric measurements of GHG, CO, O$_3$, aerosol particles, radon-222, clouds microphysics and radionuclide. The station is part of the European ICOS, AC-TRIS (Aerosol particles, Clouds, and Trace gases Research InfraStructure) and EMEP (European Monitoring and Evaluation Programme) measurement networks and of the global GAW network.

2.2 Atmospheric conditions at the Puy de Dôme station

Meteorological parameters are monitored at the station, including wind speed, wind direction, temperature, relative humidity and atmospheric pressure. A wind shadow area between 300 and 360° is clearly observed in the wind direction due to the building and the telecommunication antenna of the military base, which both induce local turbulences. The planetary boundary layer (PBL) height, wind speed and wind direction were extracted from the European Center for Medium-range Weather Forecasts (ECMWF, 2012; Seidel et al., 2012) at a 3-hour time resolution for the years 2010 to 2012. The grid cell used for the extraction has an area of 15 km x 15 km and is centered at 45°45’N, 3°00’E at an altitude of 575 m a.s.l. In this study, the wind direction from ECMWF was used as the reference because the wind direction provided by the meteorological sensor is influenced by the local turbulences caused by the telecommunication antenna located at the military base. The average difference in wind speed between the meteorological station and the ECMWF data was 3.4 ± 4.3 m s$^{-1}$, the wind speed measured by the sensor was higher because the sensor is located at a higher elevation than the grid cell used for the ECMWF extraction (1465 m a.s.l. compared to 575 m a.s.l.). Therefore, the wind speed from the meteorological sensor was used to correct the ECMWF data. The PBL height and the wind direction from ECMWF were interpolated using a linear regression fit to obtain a 1-hour time resolution.

The Puy de Dôme station is primarily influenced by winds from a southwest direction (48.2 % of the time) with a mean wind speed of 8.4 m s$^{-1}$. The wind blows from the Clermont-Ferrand sector (45°–135°) only 7.7 % of the time, with a mean speed of 4.2 m s$^{-1}$. The PBL height analysis revealed that the Puy de Dôme station is in the free troposphere during more than 70 % of the time and up to 81 % of the time during winter.

Back trajectories were calculated using the Lagrangian dispersion model Flexpart version 8.2.3, based on ECMWF ERA-Interim data at a horizontal resolution of 1° × 1°, with 60 vertical levels and a 3-hour temporal resolution. Eight particles were released every 15 min (96 particles every 3 h) in a 3-D box centered around the Puy de Dôme station (from lower left corner 45.76° N, 2.95° E to upper right corner 45.78° N, 2.97° E; between 1400 and 1500 m a.s.l.) with a lifetime of 3 days. This simulation was performed for particles arriving at the station between 1 January 2010 and 31
December 2012. The footprints were computed on a $1^\circ \times 1^\circ$ horizontal grid, following the method described by Lin et al. (2003) and taking into account the planetary boundary layer height at each particle location. We considered that a particle is influenced by surface emissions from one grid cell when its elevation is under the PBL height and that its influence is inversely proportional to the PBL height. The maps presented in Fig. 1 show the footprints for air masses arriving at the station between 14:00 and 16:00 UTC, when the PBL is usually well developed (Fig. 1a), and between 22:00 and 06:00 UTC, when the PBL is below 1400 m and the station is within the free troposphere (Fig. 1b). The grid cell influence is represented as a relative influence compared to the maximum value (in percent). On both maps, the station is located within the black grid cell. By analyzing back trajectories from HYSPLIT over 437 days at Puy de Dôme (i.e., from Atlantic and continental western Europe areas), Boulon et al. (2011) showed that 87 % of air masses reaching the station are from the west. A statistical analysis of back trajectories over 4 years conducted by Venzac et al. (2009) demonstrated that winter air masses reaching Puy de Dôme travel over longer distances from the west than summer air masses.

3 Instrumental setup

The GHG observations at Puy de Dôme started in 2000 with continuous CO$_2$ measurements using a nondispersive infrared (NDIR) spectrometer. Since 2001, a pair of flasks has been sampled once a week by the LaMP team and analyzed by GC for CO$_2$, CH$_4$, N$_2$O and SF$_6$ mole fractions and by a mass spectrometer for $\delta^{13}$C and $\delta^{18}$O in CO$_2$ at the LSCE in Gif-sur-Yvette, France. In 2010, the GC system for semicontinuous measurements of CO$_2$, CH$_4$, N$_2$O and SF$_6$ was installed at the station. In 2011, the NDIR spectrometer was replaced by a CRDS for continuous CO$_2$ and CH$_4$ measurements. Since 2002, the station has also been equipped with a radon-222 ($^{222}$Rn) analyzer based on the active deposit method. These instruments are housed in a regulated-temperature room, the inlet lines being located on the roof of the station, 10 m a.g.l. This section focuses on the setup of the GC system running at the Puy de Dôme station since July 2010.

3.1 Description of the GC system

The GC system installed at the Puy de Dôme station is a commercial HP-6890N from Agilent that was modified and optimized at the LSCE for automatic and semicontinuous atmospheric measurements of CO$_2$, CH$_4$, N$_2$O and SF$_6$ mole fractions in dry ambient air (Lopez, 2012). Similar instrument configurations are installed at the Gif-sur-Yvette and Trainou stations in northern France (Lopez et al., 2012; Schmidt et al., 2014).

The ambient air is pumped from the roof of the station (pump KNF: PMF 1433-811) through a 10 m long Dekabon tube with an outside diameter of 1/2 in. (~ 1.27 cm). Three filters (140, 40 and 7 µm TF series from Swagelok) are placed in series to protect the pump and the analysis system from dust and aerosol particles. After passing the pump, the ambient air is pressurized and dried in two steps. First, the air passes through a commercial decanting bowl (40 mL volume) placed in a refrigerator set at 5 $^\circ$C for preliminary drying. The water accumulated in the decanting bowl is flushed out every 6 hours for 10 s by opening a solenoid valve. In a second step, the ambient air passes through a glass trap in an ethanol bath that is maintained at $-55$ $^\circ$C by a cryocooler (Thermo Neslab CC-65). The dew point of the air going in the instrument is approximately $-50^\circ$C. The glass trap is

Figure 1. Footprint of the Puy de Dôme station from the Lagrangian dispersion model Flexpart (a) during daytime (14:00 to 16:00 UTC), when the PBL is usually well developed, and (b) during nighttime (22:00 to 06:00 UTC), when the station is usually in the free troposphere.
changed during the weekly maintenance of the station. An electronic box is used to regulate the refrigerator temperature and to open and close the solenoid valve of the decanting bowl. This box also records the temperatures of the fringe, the ethanol bath and the room. In case of power failure, the entire GC system is connected to an uninterruptible power supply (UPS) that allows the system to run for a few hours.

The GC system consists of an injection part, a separation part and a detection part. These three different parts are indicated by different colors in Fig. 2. For analysis, an air sample is first filled into the two sample loops. The sample is then pushed by different carrier gases to the chromatographic columns, where the species are separated. Finally, CO₂ (via a nickel catalyst) and CH₄ are detected by a FID. Between 0 and 0.75 min and between 3.7 and 5.4 min with a 100 mL min⁻¹ flow rate of the carrier gas to eliminate heavy electrophilic molecules from the system to avoid an eventual pollution of the analytical column, which might induce an increase in the ECD baseline. Between 0.75 and 3.7 min, the N₂O and SF₆ molecules are injected first into the pre-column and then into the analytical column, where separation occurs. The analytical column is directly connected to the ECD. The N₂O and SF₆ retention times are 4.3 and 4.8 min, respectively.

The CH₄ and CO₂ molecules are detected by an FID and a Ni-catalyst used to reduce CO₂ to CH₄. Methane molecules elute after 2.7 min and are injected directly into the detector for analysis. Once the CH₄ molecules are released from the analytical column, the Valco valve no. 4 is opened to connect the nickel catalyst, allowing CO₂ molecules to be reduced to CH₄ to enable CO₂ detection by the FID. The retention time of CO₂ is 3.5 min. The FID temperature is controlled at 300 °C, and the flame is fed with hydrogen at a flow rate of a 65 mL min⁻¹ (provided by a NM-H₂ hydrogen generator from F-DBS) and zero air at a flow rate of a 400 mL min⁻¹ (provided by a combination of a compressor from June-Air and a 75–82 air zero generator from Parker-Balston). Hydrogen is also used for CO₂ reduction over the Ni catalyst. The typical efficiency of the catalyst is 97 % in the CO₂ atmospheric mole fraction range.

Figure 3 shows the typical chromatograms obtain by the FID and μECD detectors. The top panel presents the FID’s response in picoampere. The spike observed between the CH₄ peak and the CO₂ peak in the close-up panel is caused by the opening of valve no. 4. The bottom panel of Fig. 3 presents the μECD’s response in Hz. The first large peak observed at approximately 2.7 min is the O₂ peak, which is followed by the N₂O peak and finally by the SF₆ peak.
methods. A typical method is configured to do the following: controlling the GC system parameters through the so-called “methods”. A typical method is configured to do the following:

- control the temperature of the detectors, the catalyst and the oven;
- regulate the flows of the sample (via Aux5), the carrier gases (via Aux3 and Aux4), H₂ and zero air, all via five distinct EPCs;
- schedule the opening and closing of the four six-port two-way Valco valves, controlled via the internal events output GC connector;
- choose the position of the eight-port microelectronic Valco valve, controlled via the external events output GC connector; and
- integrate the results of the analysis (via the chromatograms).

A typical method is presented in Table 1 and corresponds to one analysis of a chosen sample. Table 2 summarizes the GC system setup used between 2010 and 2013. A sequence lasting 3 days is designed by the sequential arrangement of methods which enables the automatic selection of ambient air and calibration gas measurement. The created sequence runs in a loop mode.

The FID and μECD signals (see Fig. 3) are expressed in picoampere and hertz, respectively. The peak integrations (area and height) of the different chromatograms are automatically computed by the Chemstation software at the end of each method and the integration results are stored in “.txt” files. The repeatability of our GC system (see Sect. 3.3) is improved when the peak areas for CO₂, CH₄ and N₂O and the peak heights for SF₆ are used. Once a day, the integration results are transferred to and stored in the LSCE database via ftp, and the mole fractions of the analyzed samples are automatically calculated. Three to five times each week, a trained operator evaluates the performances of the GC through a dedicated graphical application, enabling graphics of the instrument parameters to be drawn (see Sect. 3.3). Based on these graphics, flags are manually assigned to the data.

3.1.3 Calibration strategy

The GC system is calibrated using a two-point calibration strategy. Two working standards containing a known amount of CO₂, CH₄, N₂O and SF₆ in synthetic air (matrix of N₂, O₂ and Ar) are used. The mole fractions of the trace gases in the two working standards are selected to bracket the typical ambient air mole fractions observed at the Puy de Dôme station and are referred to as working standard high (WH) and working standard low (WL). These gas mixtures are used to fill 40 L aluminum cylinders (Luxfer) to 200 bar by Deuste Steininger (Mühlhausen, Germany). All working standards are calibrated at LSCE against the laboratory standard scale of the World Meteorological Organization (WMO scale) provided by the Central Calibration Laboratories (CCL) of the National Oceanic and Atmospheric Administration (NOAA). The calibration scales currently used are WMO-X2007, NOAA-04, NOAA-2006A and NOAA-2006 for CO₂, CH₄, N₂O and SF₆, respectively (Zhao and Tans, 2006; Hall et al., 2007; Dlugokencky et al., 2005).

The response function of μECD for N₂O analysis is non-linear, especially in the range below and above the tropospheric values (see Schmidt et al., 2001; van der Laan et al., 2009b; Lopez et al., 2012). The nonlinearity of our μECD was tested by analyzing five cylinders calibrated by the CCL on the NOAA-2006A scale and with N₂O mole fractions between 302.00 and 338.04 nmol mol⁻¹. In this small mole fraction range, which is important for our measurements, a two-point calibration describes the response function of our μECD sufficiently. It compares very well with an exponential fit through five cylinders, with an average difference of 0.01 ± 0.13 nmol mol⁻¹. This result confirms that a two-point calibration strategy is well adapted to correct for the μECD nonlinearity in atmospheric mole fraction ranges. Similar tests demonstrated that the FID response is linear in the atmospheric mole fraction range for CO₂ and CH₄ measurements.

The two working standards (WH and WL) are analyzed every 30 min to correct for atmospheric (temperature and pressure) changes as well as instrumental drifts, enabling the analysis of five samples between each calibration. The lifetime of our standards is approximately 3 years using this calibration strategy. To limit the risk of drift, each working standard must be replaced before reaching 30 bar pressure.
Table 1. Measurement method used for the GC system at the Puy de Dôme station. The “On” position on the valves presented here corresponds to the dashed lines in Fig. 2 and the “Off” position correspond to the solid lines.

<table>
<thead>
<tr>
<th>Time (min)</th>
<th>Parameter</th>
<th>Value</th>
<th>Comments</th>
</tr>
</thead>
<tbody>
<tr>
<td>0.00</td>
<td>Aux 3</td>
<td>45.0 psi</td>
<td>Carrier gas pressure for N₂</td>
</tr>
<tr>
<td>0.00</td>
<td>Aux 4</td>
<td>20.0 psi</td>
<td>Carrier gas pressure for Ar/CH₄</td>
</tr>
<tr>
<td>0.00</td>
<td>Aux 5</td>
<td>2.5 psi</td>
<td>Sample pressure</td>
</tr>
<tr>
<td>0.00</td>
<td>Valve no. 2</td>
<td>On</td>
<td>Flush of the sample loops</td>
</tr>
<tr>
<td>0.00</td>
<td>Valve no. 3</td>
<td>On</td>
<td>Backflush of the pre-column (N₂O/SF₆)</td>
</tr>
<tr>
<td>0.75</td>
<td>Aux 5</td>
<td>0.0 psi</td>
<td>Sample pressure</td>
</tr>
<tr>
<td>0.75</td>
<td>Valve no. 3</td>
<td>Off</td>
<td>Stop of pre-column backflushing</td>
</tr>
<tr>
<td>1.25</td>
<td>Valve no. 2</td>
<td>Off</td>
<td>Sample injection</td>
</tr>
<tr>
<td>3.10</td>
<td>Valve no. 4</td>
<td>On</td>
<td>Injection of CO₂ into the catalyst</td>
</tr>
<tr>
<td>3.60</td>
<td>Aux 3</td>
<td>0.0 psi</td>
<td>Carrier gas pressure for N₂</td>
</tr>
<tr>
<td>3.70</td>
<td>Valve no. 3</td>
<td>On</td>
<td>Backflush of the pre-column</td>
</tr>
<tr>
<td>5.30</td>
<td>Valve no. 4</td>
<td>Off</td>
<td>Opening of the catalyst valve</td>
</tr>
<tr>
<td>5.40</td>
<td>Aux 3</td>
<td>45.0 psi</td>
<td>Carrier gas pressure for Ar/CH₄</td>
</tr>
</tbody>
</table>

Table 2. GC system equipment and temperature and flow rate settings.

<table>
<thead>
<tr>
<th>Detector FID (CO₂/CH₄)</th>
<th>µECD (N₂O/SF₆)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Carrier gas</td>
<td></td>
</tr>
<tr>
<td>N₂ cylinder (purity &gt; 99.9999 %)</td>
<td>Ar/CH₄ cylinder (ECD quality)</td>
</tr>
<tr>
<td>+ purifier</td>
<td>+ purifier</td>
</tr>
<tr>
<td>Flow rate</td>
<td>100 mL min⁻¹</td>
</tr>
<tr>
<td>Loop sample volume</td>
<td>15 mL</td>
</tr>
<tr>
<td>Oven temperature</td>
<td>80 °C</td>
</tr>
<tr>
<td>Pre-column</td>
<td>HayeSep-Q</td>
</tr>
<tr>
<td>Analytical column</td>
<td>12' × 3/16&quot; SS, 80/100</td>
</tr>
<tr>
<td>Detector temperature</td>
<td>300 °C</td>
</tr>
<tr>
<td>Catalyst temperature</td>
<td>390 °C</td>
</tr>
<tr>
<td>Gas supply</td>
<td>H₂ generator: 60 mL min⁻¹</td>
</tr>
</tbody>
</table>

Table 3. Trace gas mole fractions of GC working standards used during period A (July 2010 to April 2013) and period B (May and June 2013).

<table>
<thead>
<tr>
<th>Species</th>
<th>Period A – 33 months</th>
<th>Period B – 2 months</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td>WH</td>
<td>WL</td>
</tr>
<tr>
<td>CO₂ (µmol mol⁻¹)</td>
<td>425.10</td>
<td>372.45</td>
</tr>
<tr>
<td>CH₄ (nmol mol⁻¹)</td>
<td>2179.90</td>
<td>1732.99</td>
</tr>
<tr>
<td>N₂O (nmol mol⁻¹)</td>
<td>340.90</td>
<td>322.93</td>
</tr>
<tr>
<td>SF₆ (pmol mol⁻¹)</td>
<td>10.05</td>
<td>5.38</td>
</tr>
</tbody>
</table>

3.2 Other instrumentation

3.2.1 Flask sampling unit

A flask sampling unit was installed at the Puy de Dôme station in 2002 for weekly sampling. It consists of a pump that pressurizes two 1 L glass flasks placed in series to 1 bar relative pressure. They are flushed for 15 min with dry ambient...
air prior the pressurization. Ambient air is dried in a distinct cooling trap maintained in the same ethanol bath as the GC trap (see Sect. 3.1). The flasks are then shipped and analyzed at LSCE by a GC system for CO₂, CH₄, N₂O and SF₆ and by a Finnigan MAT-252 isotope mass spectrometer for CO₂ isotopic composition (¹³C and ¹⁸O). The calibrations for trace gas analysis are performed in the same manner as presented in Sect. 3.1.3, and the results are stored in the same database.

### 3.2.2 Radon-222 measurement system

Radon-222 (²²²Rn) is a radioactive noble gas (T₁/₂ = 3.8 days) and is part of the radioactive decay chain of uranium-238. Uranium-238 in the earth’s crust results in the emission of ²²²Rn by the earth’s surface. Atmospheric radon-222 activity has been monitored at the Puy de Dôme station since 2002. The analyzer is based on the active deposit method, which consists of alpha decay counting of ²²²Rn’s solid short-lived daughters: ²¹⁸Po, ²¹⁴Pb and ²¹⁴Bi. The measurement technique has been described in detail by Polian et al. (1986) and Biraud et al. (2000). To avoid the loss of the solid ²²²Rn daughters, the inlet line is a 6 m long straight metal tube with a 31 mm outside diameter. During the first years of measurements, the inlet line was frequently contaminated by room air, and only measurements after October 2006 can be used. Schmidt (1999) estimated a radioactive disequilibrium (see Turner, 1964) at the Schauinsland station (Germany; 47°54’ N, 7°54’ E; 1205 m a.s.l.) of 1.15 ± 0.14. This value was independently confirmed by Xia et al. (2010). The Puy de Dôme station and the Schauinsland station are two medium-elevation mountain stations, having the same geographical environment. They are both frequently above the continental boundary layer, especially in winter. Based on the similarities between the Puy de Dôme station and the Schauinsland station, the measured (²²²Rn) activity at Puy de Dôme has been corrected for the radioactive disequilibrium by using the same value (1.15 ± 0.14).

### 3.2.3 CO₂ continuous measurements by in situ NDIR

Continuous CO₂ measurements with a NDIR gas analyzer at Puy de Dôme began in 2000. The instrument is an integrated system constructed around a LI-COR NDIR (Li-6252 NDIR, LI-COR Inc., Nebraska, USA) analyzer optical bench. The CO₂ measurements are based on the difference in absorption of infrared radiation passing through two cells: the reference cell and the sample cell. Infrared radiation is transmitted through both cell paths, and the analyzer signal is proportional to the difference in absorption between both cells. The measurement frequency is 1 Hz, and the cell flow is typically 20 mL min⁻¹ for the sampling cell and 15 mL min⁻¹ for the reference cell, which is continuously flushed with a reference gas. The calibration strategy is based on four cylinders calibrated on the WMO-X2007 scale. The calibration is performed twice a year by analyzing each calibration cylinder 30 times for 10 min. Data are then corrected using a quadratic fit.

Ambient air is pumped from the roof to the instrument through a 3/8 in. (~ 0.95 cm) outside diameter Dekabon line. The air is dried by passing through a glass trap maintained in a cold ethanol bath (see Sect. 3.1). Ambient air is analyzes for 50 min following the analysis of the reference cylinder for 10 min, the latter passing through both cells at the same time. The NDIR spectrometer was replaced by a CRDS analyzer in April 2011.

### 3.2.4 CO₂ and CH₄ continuous measurements by in situ CRDS

The CRDS analyzer (Picarro G1301) was installed in April 2011. It continuously and simultaneously measures CO₂, CH₄ and H₂O atmospheric mole fractions. We use four calibration cylinders spanning the atmospheric range of 366 to 453 µmol mol⁻¹ for CO₂ and 1722 to 2107 nmol mol⁻¹ for CH₄. The cylinders are calibrated at the LSCE laboratory on the WMO-X2007 and NOAA-04 scales. The instrument calibration is performed automatically every 15 days by injection, according to the following scheme: 4 times for each calibration cylinder for 30 min, beginning with the one with the lowest mole fractions and ending with the one with the highest mole fractions. For each calibration cylinder, the entire first injection and the first 15 min of the subsequent injections are automatically rejected because of equilibration time. The instrument calibration takes 8 h and a linear fit is applied to compute the analyzer response. A target gas is automatically injected into the CRDS every 10 h for 30 min and, again, the first 15 min are automatically rejected. The SDs at 1 sigma of the target gas analysis over 1 year are 0.02 µmol mol⁻¹ for CO₂ and 0.14 nmol mol⁻¹ for CH₄.

Ambient air is injected into the CRDS from the roof (through a 3/8 in. Dekabon line equipped with three filters of 140, 40 and 7 µm) using a pump located after the optical cavity. To avoid the risk of bias caused by the interference of water vapor and trace gases in the CRDS (Chen et al., 2010), the ambient air is dried prior to its injection into the CRDS by the drying system presented in Sect. 3.1. A problem occurred in the cavity in August 2011, and the CRDS measurements were stopped until April 2012. The instrument has been returned to the manufacturer for repair.

### 3.3 Quality control of the GC system and comparisons with different analyzers

A target gas (TGT) is injected into the GC system once an hour for quality control. The target gas cylinder is a 40 L cylinder filled with dry ambient air at Gif-sur-Yvette. After a stabilization time of at least 1 month, the cylinder is analyzed at LSCE using the laboratory primary standards, and CO₂, CH₄, N₂O and SF₆ mole fraction values are assigned to the cylinder (see Table 4).
Table 4. Assigned target gas values with the respective mole fractions measured using the GC system at the Puy de Dôme station over period A and period B. The assigned values were measured by the GC system at the LSCE central lab against WMO calibration gases.

<table>
<thead>
<tr>
<th>Species</th>
<th>Assigned values</th>
<th>Period A</th>
<th>Period B</th>
</tr>
</thead>
<tbody>
<tr>
<td>CO₂ (µmol mol⁻¹)</td>
<td>402.57 ± 0.07</td>
<td>402.42 ± 0.15</td>
<td>402.38 ± 0.46</td>
</tr>
<tr>
<td>CH₄ (nmol mol⁻¹)</td>
<td>1973.87 ± 0.73</td>
<td>1973.81 ± 2.12</td>
<td>1963.72 ± 6.64</td>
</tr>
<tr>
<td>N₂O (nmol mol⁻¹)</td>
<td>325.71 ± 0.23</td>
<td>325.90 ± 0.35</td>
<td>325.76 ± 0.47</td>
</tr>
<tr>
<td>SF₆ (pmol mol⁻¹)</td>
<td>7.23 ± 0.04</td>
<td>7.23 ± 0.06</td>
<td>7.20 ± 0.07</td>
</tr>
</tbody>
</table>

Figure 4 shows the time series of the target gas analysis from July 2010 to June 2013. The target gas cylinder was not changed over the 3 years. The different data gaps observed in CO₂ and CH₄ between March and October 2011 were caused by problems with the hydrogen generator (leaks in the electrolysis cell). The large data gaps in N₂O and SF₆ between April and August 2012 were due to a problem with the power supply of Valco valve no. 2b. The reproducibility (computed here as the SD at 1 sigma of the target analysis over 1 year of measurement) and the typical short-term repeatability (SD of the target analysis over 24 h) are presented in Table 5.

The vertical dark blue lines in Fig. 4 indicate when the working standards were changed and separate the two measurement periods: period A and period B (see Sect. 3.1.3). The average mole fractions of the TGT measurement during period A and period B are presented in Table 4 together with the respective assigned mole fractions. The measured CH₄, N₂O and SF₆ mole fractions agreed well with the assigned mole fractions during period A (considering the uncertainties). The difference between the CO₂ assigned and measured values in period A and period B was 0.15 and 0.19 µmol mol⁻¹, respectively, confirming the consistency between the two scales used in periods A and B. This agreement when using two different calibration scales shows that the problem is probably due to an error in the value attributed to this target gas. During period B, the average CH₄ mole fraction of the TGT gas was lower by approximately 10 nmol mol⁻¹ compared to period A. This decrease was most probably due to a micro-leak in the WH line that affects only the CH₄ mole fractions. Target data as well as ambient air measurements for this period should be recalibrated by applying a one-point calibration to the four trace gases. This micro-leak was detected and fixed in September 2013, when the target cylinder was replaced by a new one.

In addition to internal quality control performed via the target gas analysis, comparisons of in situ ambient air analysis, flask analysis and cylinder analysis performed by different analyzers are relevant. These comparisons enable the validation of the scale consistency between different instruments and the detection of possible leaks or biases introduced by the inlet lines. The WMO-GAW gives recommendations on the scientific level of compatibility for such a comparison in the Northern Hemisphere (WMO-GAW, 2013). These levels are ±0.1 µmol mol⁻¹ for CO₂, ±2.0 nmol mol⁻¹ for CH₄, ±0.1 nmol mol⁻¹ for N₂O and ±0.02 pmol mol⁻¹ for SF₆.

The mean differences between the in situ GC system measurements and weekly flask sampling, in situ NDIR measurements and in situ CRDS measurements are summarized in Table 6. These differences are calculated from the hourly mean measurements. A 2-sigma filter was applied to the differences to flag the eventual outliers. Comparisons between GC and NDIR were based on 9 months of overlapping measurements (July 2010 to April 2011) and revealed a mean CO₂ difference (GC minus NDIR) of −0.14 ± 1.78 µmol mol⁻¹. Comparisons between GC and CRDS were based on 20 months of overlapping measurements, from April 2011 to July 2013 with a break between August 2011 and April 2012. The average differences (GC minus CRDS) were 0.21 ± 0.78 µmol mol⁻¹ for CO₂ and −0.64 ± 5.46 nmol mol⁻¹ for CH₄ over the total overlapping measurements. Because the CRDS instrument was stopped for several months and shipped to the manufacturer in 2011, we compared the results before and after its repair. In the first overlapping measurement period (April to August 2011), the differences were −0.13 ± 0.61 µmol mol⁻¹ and −1.27 ± 3.49 nmol mol⁻¹ for CO₂ and CH₄, respectively.
In the second overlapping measurement period (April 2012 to July 2013) the CH$_4$ difference decreased to $-0.26 \pm 5.02$ nmol mol$^{-1}$, whereas the CO$_2$ difference increased to $0.28 \pm 0.75$ µmol mol$^{-1}$. Over the second comparison period, the observed CO$_2$ difference remained constant with time and did not depend on atmospheric mole fractions. The inlet lines, including the pumps and the dryer systems, were tested for 3 weeks: a common inlet line for ambient air measurements has been used for the GC and CRDS (the GC one, described in Sect. 3.1). During these 3 weeks of testing, the difference between GC and CRDS remained constant and equal to 0.28 µmol mol$^{-1}$, confirming that the inlet lines did not cause the observed bias. Even after changing the GC working standards in late April 2013, the CO$_2$ difference was still observed. This observed difference is stable over time and is not concentration dependent. A second experiment was to analyze the second set of working standards on the CRDS. The results showed a difference between the assigned value (at LSCE) and the CRDS of 0.03 µmol mol$^{-1}$ on the WL and of 0.34 µmol mol$^{-1}$ on the WH.

Comparisons between the GC system and the flask analysis or comparison cylinders provide information on the scale consistency between different laboratories. For the four analyzed long-lived GHGs, the comparisons between GC in situ measurements and flask analyses reached the desirable comparison levels (see Table 6 for more details). The two GC measurements bracketing each sampled flask measurements are linearly interpolated in order to match the time of the flask sampling. The Puy de Dôme station also participates in the “Cucumbers comparison programme” (http://cucumbers.uea.ac.uk/) in the framework of the European Union CarboEurope project (2000–2005), EU IMECC (2007–2011) and InGOS (2011–2015) infrastructure projects. Three cylinders are alternately analyzed on the GC at Puy de Dôme and at the Gif-sur-Yvette, Trainou (France), Kasprowy Wierch (Poland) and Hegyhatsal (Hungary) stations. Table 6 presents the mean differences between the average analysis of the three comparison cylinders at Puy de Dôme and Gif-sur-Yvette (LSCE) between 2011 and 2013.

The different comparison methods presented in this section showed that the GC system installed at the Puy de Dôme station matches the WMO-GAW recommendations for CH$_4$ and SF$_6$ measurements. The CO$_2$ comparison shows different results depending on the method used. The recommendations are met by considering the comparison with cylinders or flasks, whereas they are not if we consider only the comparison between in situ instruments. The WMO-GAW recommendations concerning the N$_2$O measurements are ambitious considering the repeatability obtained with our GC. The N$_2$O measurements are not in line with the WMO-GAW recommendation but we note that the different comparison methods used show differences lower than our instrumental repeatability.

4 Results and discussions

4.1 Three years of ambient air measurements

Figure 5 shows the hourly time series of CO$_2$, CH$_4$, N$_2$O and SF$_6$ ambient air mole fractions together with the $^{222}$Rn activities at Puy de Dôme from July 2010 to the end of June 2013. The different gaps observed in the atmospheric GHG time series are explained in Sect. 3.3. These time series are presented with the respective monthly background values (black lines). The monthly background values were calculated from the monthly average nighttime mole fractions (between 22:00 and 06:00 UTC), when the station is above the PBL (see Fig. 6). The hourly $^{222}$Rn activities are presented in the last panel of Fig. 5 and varied between 0 and 9 Bq m$^{-3}$ over the 3 years of measurements. From February 2012 to the end of April 2012, the computer for $^{222}$Rn data acquisition had hardware and software problems, resulting in the observed gap.

Figure 6 presents the mean diurnal cycles per season of CO$_2$, CH$_4$, N$_2$O, SF$_6$, $^{222}$Rn and for the PBL height (from ECMWF) from June 2010 to June 2013. The GHG diurnal cycles were computed from the detrended hourly time series to the reference of 1 January 2013. The mean yearly increase rate was subtracted from the time series, before computing the seasonal mean diurnal cycles. To represent the thickness of the PBL relative to the ground level, we used the altitude of Clermont-Ferrand (396 m a.s.l.) as the reference altitude to plot the PBL height, Clermont-Ferrand being located at the lowest altitude of the ECMWF extracted grid cell. The horizontal solid black line on the PBL height panel in Fig. 6 gives the altitude of the station above Clermont-Ferrand and enables a quick observation of whether the sta-
Table 6. Results of the comparisons between in situ measurements obtained using the GC system, NDIR, CRDS and flask measurements. The mean differences in the cylinder analysis at the Puy de Dôme and Gif-sur-Yvette stations are presented in the last column. Flasks and cylinders were analyzed at the LSCE central lab at Gif-sur-Yvette.

<table>
<thead>
<tr>
<th>Date</th>
<th>Comparisons</th>
<th>GC in situ – flasks</th>
<th>In situ:</th>
<th>In situ:</th>
<th>Cylinders:</th>
</tr>
</thead>
<tbody>
<tr>
<td></td>
<td></td>
<td>Average</td>
<td>No. of flasks</td>
<td>GC–NDIR</td>
<td>GC–CRDS</td>
</tr>
<tr>
<td>CO₂ (µmol mol⁻¹)</td>
<td>July 2010–July 2013</td>
<td>0.11 ± 1.19</td>
<td>55</td>
<td>−0.14 ± 1.78</td>
<td>0.21 ± 0.78</td>
</tr>
<tr>
<td>CH₄ (nmol mol⁻¹)</td>
<td>July 2010–April 2011</td>
<td>0.04 ± 4.30</td>
<td>57</td>
<td>n/a</td>
<td>−0.64 ± 5.46</td>
</tr>
<tr>
<td>N₂O (nmol mol⁻¹)</td>
<td>April 2011–July 2013</td>
<td>0.12 ± 0.55</td>
<td>47</td>
<td>n/a</td>
<td>n/a</td>
</tr>
<tr>
<td>SF₆ (pmol mol⁻¹)</td>
<td>July 2010–July 2013</td>
<td>−0.01 ± 0.07</td>
<td>53</td>
<td>n/a</td>
<td>n/a</td>
</tr>
</tbody>
</table>

Figure 6. Mean diurnal cycles of CO₂, CH₄, N₂O, SF₆, and ²²²Rn together with the planetary boundary layer height (relative to Clermont-Ferrand altitude – 396 m a.s.l.) at the Puy de Dôme station for each season. Trace gas mole fractions were detrended based on 1 January 2013.

tion is within or above the PBL. The mean diurnal cycles of the PBL height exhibited the same pattern for each season, with an increase in height from 06:00 to 12:00 UTC followed by a stable height until 15:00 UTC. After 15:00 UTC, the PBL height began to decrease, reaching a minimum after 21:00 UTC. On a mean annual scale, the GC sampled the trace gases within the PBL between 09:00 and 18:00 UTC with an enlarged time step in summer and a narrower time step in winter. In winter, the Puy de Dôme station is often above the PBL during several consecutive days.

The mean diurnal cycles of the long-lived GHGs and ²²²Rn observed in this study are typical of mountain sites, as previously described by Schmidt et al. (1996) for the Schauinsland station (1205 m a.s.l.) or Necki et al. (2003) for the Kasprowy Wierch station (Poland, 1987 m a.s.l.). The PBL height is a key atmospheric factor, particularly for mountain sites where measurements alternate between the free troposphere and the PBL. The atmospheric mole fraction variabilities of trace gases are generally larger in the PBL because of the combination between its diurnal variability and the emissions from surface sources. Due to its short radioactive lifetime, ²²²Rn cannot accumulate in the free troposphere, in contrast to other long-lived trace gases. The mean diurnal cycles of ²²²Rn (Fig. 6) exhibited larger variations in summer, when the PBL height is maximal and the inlet line of the station alternates between the PBL during daytime and above during nighttime. In winter, ²²²Rn activities are lower than in summer because the station is usually in the free troposphere.

The mean diurnal cycles for CO₂ and CH₄ exhibit different shapes. As for the ²²²Rn activities, the CH₄ mole fractions at the Puy de Dôme station are higher in the afternoon, when the PBL is well developed, compared with the nighttime mole fractions. On a yearly average, CH₄ afternoon mole fractions
are 3.3 nmol mol\(^{-1}\) higher than the nighttime mole fractions. We observed an opposite diurnal cycle trend for CO\(_2\): the biosphere is a sink for CO\(_2\) during the daytime and counterbalances the atmospheric effects. This is seen clearly during summertime, when the photosynthetic activity is maximal: the amplitude of the CO\(_2\) diurnal cycle is 7.2 µmol mol\(^{-1}\), with a minimum mole fraction of around 16:00 UTC. The CO\(_2\) mole fractions are maximal in winter when the biosphere acts as a CO\(_2\) source, mainly driven by soil respiration.

The amplitudes of the N\(_2\)O and SF\(_6\) diurnal cycles are very small and nearly undetectable, except for the N\(_2\)O in summer which exhibits an amplitude of 0.25 nmol mol\(^{-1}\). The CH\(_4\), N\(_2\)O and SF\(_6\) mole fractions are largest in spring and lowest during summertime because their respective mole fractions are mainly driven by the PBL height and the associated vertical mixing.

### 4.2 The marine boundary layer reference

In this section, the background mole fractions of recorded trace gases at the station (see Sect. 4.1) are compared with the respective marine boundary layer references (MBLRs). Here, the MBLRs are the monthly zonal average trace gas mole fractions for 45.5° N computed from NOAA measurements (Dlugokencky et al., 2013a, b). They were retrieved from the Global Monitoring Division of the NOAA Earth System Research Laboratory. Figure 7 shows the differences between the monthly mean background mole fractions at Puy de Dôme (nighttime values between 22:00 and 06:00 UTC) and the respective monthly MBLRs for CO\(_2\), CH\(_4\), N\(_2\)O and SF\(_6\). These comparisons enable the direct quantification of the influence of sources and sinks on trace gases at the station relative to oceanic air masses. These differences are called continental offsets. The CO\(_2\) continental offset has negative values in spring, indicating the influence of the continental biosphere, which acts as a sink. During summer, autumn and winter, the offsets are positive, revealing the importance of continental fossil fuel and biospheric sources in the Puy de Dôme catchment area. The continental offsets are usually positive for CH\(_4\) and always positive for N\(_2\)O, indicating the strong influence of agricultural sources (see Sect. 2.1) in the Puy de Dôme footprint. Finally, the SF\(_6\) offset varied between −0.10 and +0.12 pmol mol\(^{-1}\), which represents the same order of magnitude as the GC measurement repeatability. In addition, Fig. 7 shows the monthly \(^{222}\)Rn continental offset (for nighttime selection data between 22:00 and 06:00 UTC) at the Puy de Dôme station relative to marine air. The marine air \(^{222}\)Rn reference was computed from the mean activity of 15 years of measurements during maritime background conditions at the European background site of Mace Head (see Bousquet et al., 1996) and is equal to 168 mBq m\(^{-3}\).

### 4.3 The radon tracer method

#### 4.3.1 Method

Once emitted by soils, \(^{222}\)Rn is an excellent tracer of continental air masses due to its physical and chemical properties. Thus the radon tracer method (RTM) has been used in numerous atmospheric studies to estimate trace gas surface emissions on local to regional scales. Detailed descriptions of this method are given in the following studies: Schmidt et al. (2001); Hammer and Levin (2009); Yver et al. (2009); van der Laan et al. (2009a).

The RTM is based on Eq. (1), where \(J_x\) and \(J_{\text{Rn}}\) are the respective fluxes of a trace gas \(x\) and \(^{222}\)Rn. The \(\Delta C_x\) and \(\Delta C_{\text{Rn}}\) terms are the temporal variations in the trace gas \(x\) mole fraction and in the \(^{222}\)Rn activity over a period \(\Delta t\). Finally, \(\lambda_{\text{Rn}}\) is the \(^{222}\)Rn decay constant.

\[
J_x = J_{\text{Rn}} \frac{\Delta C_x}{\Delta C_{\text{Rn}}} \left( 1 - \frac{\lambda_{\text{Rn}} C_{\text{Rn}}}{\Delta t} \right)
\]  

As shown in Fig. 6, the diurnal variations in trace gases at Puy de Dôme are very weak, which makes difficult to correctly assess the \(\Delta C_x\) and \(\Delta C_{\text{Rn}}\) terms on a daily basis. In this study, we apply the RTM approach presented by Schmidt et al. (2003), in which the CO\(_2\) fluxes at the Schauinsland station were calculated using the monthly CO\(_2\) and \(^{222}\)Rn continental offsets (relative to the MBLR). As presented in Sect. 4.2, the continental offsets of trace gases reflect the source and/or sink influence at a continental site relative to a maritime background. In this study, the terms \(\Delta C_x\) and \(\Delta C_{\text{Rn}}\) (see Eq. 1) were calculated as the monthly offsets of trace gases and radon-222, respectively.

The term in brackets in Eq. (1) corresponds to the radioactive decay correction factor which depends on the mean residence time of air masses over the European continent before reaching the station. In Sect. 2.2, it was shown that most of the air masses arriving at the station are from the western part (oceanic air masses) and have an average wind speed of 8.4 m s\(^{-1}\). Based on this, it takes at least 10 h for the oceanic air masses to reach the station from the closest oceanic coast. Considering also that some of the air masses are from other directions, we estimate an average transit time for the air masses arriving at the station of 1 day, leading to a decay correction of 0.91.

The \(^{222}\)Rn emission rate from continental surfaces strongly depends on the type and on the nature of the soils. A study of Karstens et al. (2015) provides a monthly \(^{222}\)Rn emission map at a resolution 0.083° × 0.083° over Europe. The assessment of this map takes into account the soil types and properties, the \(^{238}\)U soil content and the soil moisture evolution over time. According to the mean nighttime footprint at the Puy de Dôme station (see Fig. 1b), we extracted the monthly \(^{222}\)Rn average emission from this map for a 300 km × 300 km region centered on the Puy de Dôme.
Figure 7. Differences between the monthly background at Puy de Dôme and the respective monthly MBLR (Dlugokencky et al., 2013a, b) at 45.5° N latitude for CO₂, CH₄, N₂O and SF₆. The last panel is the ²²²Rn offset relative to marine air.

station (U. Karstens and I. Levin, personal communication, 2014). Over the years 2010 to 2012, the ²²²Rn fluxes range between 75 and 172 Bq m⁻² h⁻¹, with minimums in winter, when the soil is wet or frozen.

4.3.2 Uncertainties

The uncertainties of the radon tracer method presented above result from errors in the ²²²Rn exhalation rate, errors in the ∆Cₓ and ∆Cᵣn terms and error in the decay correction term (see Eq. 1). This section describes how these errors have been assessed to derive a mean relative uncertainty of the flux estimation of each trace gas.

A systematic assessment of the ²²²Rn exhalation rate is quite difficult. The mean ratio of the spatial variability within the extracted area (300 km × 300 km region centered on the Puy de Dôme station) to the mean flux is 30% (Karstens et al., 2015, and U. Karstens and I. Levin, personal communication, 2014). This number is used as a first approximation of the ²²²Rn exhalation rates uncertainties. This estimate does not include systematic errors and therefore is likely an underestimate. Uncertainties in the ∆Cₓ term have been assessed from the MBLRs and from the background mole fraction uncertainties. The monthly MBLR uncertainties for CO₂ and CH₄ were provided by NOAA; mean uncertainties over the measurement period have been taken into account and are equal to 0.6 µmol mol⁻¹ and 5.1 nmol mol⁻¹, respectively. We used a mean N₂O MBLR uncertainty of 0.3 nmol mol⁻¹, which was estimated and provided by E. J. Dlugokencky (personal communication, 2014). The uncertainties regarding the background mole fractions at Puy de Dôme were derived directly from the respective GC repeatabilities (see Table 5). These last two error sources were combined to give the mean absolute continental offset (∆Cₓ) over the entire measurement period. Thus, the mean relative uncertainties in the continental offsets are estimated to be 31, 39 and 42% for ∆C_CO₂, ∆C_CH₄ and ∆C_N₂O, respectively. The ²²²Rn instrument has an absolute error of ±20% for continental measurements (Biraud et al., 2000). Based on the same approach as for the ∆Cₓ term, a constant uncertainty of 28% has been attributed to the ²²²Rn continental offset term. Finally, Schmidt et al. (2003) reported an error of 7% in the decay correction (term in brackets in Eq. 1) estimated at Schauinsland.

These uncertainties were combined using the square root over the quadratic sum. The mean relative flux uncertainties derived for our RTM approach were 52, 57 and 59% for CO₂, CH₄ and N₂O, respectively. The uncertainties estimated here using this continental RTM approach are larger than those found by Biraud et al. (2000), Schmidt et al. (2001), van der Laan et al. (2009a) and Lopez et al. (2012), which are all close to 35% for the CO₂, CH₄ and N₂O flux estimates. The uncertainties presented here are mainly driven by the continental offset uncertainties.

The uncertainties in the SF₆ emissions are up to 300%. Therefore, we do not present any SF₆ emissions in this study.
4.3.3 Estimation of GHG surface fluxes in the Puy de Dôme catchment area

Continental CO₂, CH₄ and N₂O surface fluxes at the Puy de Dôme station were calculated using the radon tracer method. As shown by Gloor et al. (2001), knowledge of the station footprint is an important parameter in interpreting the large time variability in a trace gas mole fraction observed at a measurement station. Figure 1b shows the integrated nighttime footprint (22:00 to 06:00 UTC) of the Puy de Dôme station between 2010 and 2013 (see Sect. 2.2 for more details). The station is mainly influenced by regional air masses, which are well distributed all around the station during nighttime, when the measurements are usually performed in the free troposphere.

The calculated monthly fluxes are presented in Fig. 8 together with the hourly ²²²Rn exhalation rate (U. Karstens and I. Levin, personal communication, 2014) at Puy de Dôme. The units used to express the trace gas fluxes are t km⁻² month⁻¹ for CO₂ and CH₄ and kg km⁻² month⁻¹ for N₂O. Because no ²²²Rn activities were recorded between January and the end of April 2012, no fluxes could be derived from the RTM. The vertical grey lines on each curve are the absolute uncertainties calculated in Sect. 4.3.2.

The CO₂ fluxes integrate the signals from all CO₂ sources and sinks in the nighttime footprint of the station. These are the contributions of the biosphere (emissions and uptakes) and of the anthropogenic emissions (fossil fuel and biofuel). The derived CO₂ fluxes present negative values in spring, emphasizing the net uptake by the plant assimilation with a monthly average value between April and June of −435±226 t CO₂ km⁻² month⁻¹ in the station catchment area. Schmidt et al. (2003) calculated CO₂ fluxes at the Schauinsland station from 1980 to 2000 and observed a long-term monthly mean CO₂ uptake between May and June of 147±107 t CO₂ km⁻², with a maximum uptake of 550±150 t CO₂ km⁻² in the spring of 1989. These values are of the same order of magnitude as the estimations of this study. In summer, fall and winter, the fluxes are positive, which are strongly driven by the small continental offsets between 0.6 and 1.5 nmol mol⁻¹. These differences are also linked to significant uncertainties, which are strongly driven by the small continental offsets between 0.6 and 1.5 nmol mol⁻¹. Despite this difference, the atmospheric approach presented provides an independent estimation of GHG emission over the station footprint as well as new information on flux seasonality.

5 Conclusions

Semicontinuous measurements of four long-lived GHGs at Puy de Dôme started in 2010 with the installation of a GC system. This GC is designed to automatically measure CO₂, CH₄, N₂O and SF₆ atmospheric mole fractions. We described in details three methods which have been used for comparison purposes. They are based on a direct comparison between two in situ analyzers, flask measurements and cylin-
der measurements. For CH\textsubscript{4} and SF\textsubscript{6}, all comparisons show that GC measurements at Puy de Dôme are in agreement with the WMO-GAW compatibility goals. For N\textsubscript{2}O, our measurements do not match the ambitious WMO-GAW compatibility goal. For CO\textsubscript{2}, the comparison based on ambient air flasks and reference cylinders analysis between the GCs operated at Puy de Dôme and at LSCE reaches the desirable comparison level, showing there is no bias in the scale transfer between the two sites. Nevertheless, it does not do so for the in situ comparison with other analyzers (NDIR and CRDS). The comparisons between the GC and the CRDS in situ measurements indicate a constant offset of 0.21 µmol mol\textsuperscript{-1} CO\textsubscript{2} over 20 months of overlapping measurements. Several tests have been performed and are described in the study, but the reason for the observed constant bias is not yet clear. We are continuing to work on this issue and are therefore aware of the order of magnitude of bias that is possible.

At stations that typically run only one analyzer, a bias of 0.25 µmol mol\textsuperscript{-1} might not be detected when the target gas and the comparison cylinders yield good results. For consistency, we thus recommend using different methods based on flask or cylinder comparisons but also based on in situ comparisons to check whether the considered measurements match with the WMO-GAW recommendations.

The diurnal cycles of CO\textsubscript{2} and CH\textsubscript{4} observed at Puy de Dôme are mainly driven by the PLB height, and they present the typical shape of a mountain station, such as the Schauinsland or Kasprowy Wierch stations, while the N\textsubscript{2}O and SF\textsubscript{6} mean diurnal cycles present flat behaviors that are difficult to interpret.
Radon-222 was used in this study as an air mass tracer to estimate the monthly continental fluxes of CO$_2$, CH$_4$ and N$_2$O relative to the maritime background layer references. We derived a yearly net emission of 1310t CO$_2$ km$^{-2}$, 7.0t CH$_4$ km$^{-2}$ and 1.7t N$_2$O km$^{-2}$. The derived CO$_2$ and CH$_4$ fluxes compare well with other European studies or with the national inventory (CITEPA). However, it remains difficult to compare the N$_2$O fluxes with other studies due to large errors. Compared to the GC system presented in this study, the new analysis technique based on CRDS, FTIR or OA-ICOS achieve better precision and require less maintenance. Consequently, the use of these new technologies enables the development of a dense measurement network, such as ICOS, which will further improve uncertainties in the flux estimates.

Acknowledgements. The authors thank the LaMP team for weekly maintenance of the station and weekly flask sampling. We also acknowledge the support provided by the RAMCES group at LSCE for the regular maintenance of the analyzer and for data management. We thank the ECWMF and I. Pison for providing and extracting the meteorological conditions at the Puy de Dôme pixel. We also want to thank I. Levin and U. Karstens for providing and sharing the European 222Rn map emissions. We used the marine boundary layer reference kindly provided by E. Dlugokencky and K. Masarie from NOAA. The authors acknowledge the Regional Council of Auvergne for financial support. This work was partly funded by the InGOS EU project (284274).

Edited by: P. Herckes

References


ECWMF: IFS documentation CY38r1, part IV: physical processes, Tech. rep., ECMWF, available at: https://software.ecmwf.int/wiki/display/IFS/ (last access: 18 March 2015), 2012.


