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A turning-point in the evolution of the Variscan orogen: the ca. 325 Ma regional partial-melting event of the coastal South Armorican domain (South Brittany and Vendée, France)

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Keywords. – Partial-melting, Strain localisation, Gravitational collapse, U-Th/Pb on monazite, South Armorican domain

Abstract. – By drastically reducing the bulk strength of crustal materials, partial-melting is one of the main parameter controlling the rheological behaviour of the continental crust. With more than ca. 50% of the outcropping surface characterised by migmatites and granites, the coastal South Armorican domain, offers an opportunity to study deep-orogenic processes and more particularly, to understand the role of partial-melting for the late-evolution of the Variscan belt. To date, time-constraints are scarce hindering the understanding of this crucial stage in the Variscan belt evolution. This paper provides 29 new U-Th/Pb chemical ages on monazite collected over five sampling areas consisting in migmatite domes and late regional classic plutons. Based on structural, textural and chemical criteria, three main U-Th/Pb age-groups are distinguished. The first group, settled at ca. 335-330 Ma concerns samples of restites and core-domains of the monazite crystals for most of the granite massifs. Its significance is ascribed to inherited crystallisation ages probably recording the crossing of prograde monazite forming reactions (i.e. metamorphic isograds) during increasing P-T conditions in an overall nappe-stacking context. The second group that clusters at ca. 325-320 Ma corresponds to newly formed monazite grains that crystallised from juvenile silicate melts. Ages of this group are interpreted as crystallisation ages of leucosomes after a major partial-melting event that affected the whole domain. The last ca. 320 Ma group corresponds to rim-domains of monazite crystals. It is interpreted as the emplacement age of most of the large-scale granite massifs and therefore fixes the end of the partial-melting event.

The inception and drastic generalisation of partial-melting at peak-P conditions therefore coincides with a major change in the tectonic regime recorded at regional-scale. In the lights of these results, this implies that (1) either continuous stacking of continental crustal units, rich in radiogenic elements, led to an increase of temperature within the orogenic wedge provoking partial-melting, the resulting drop in the crustal strength inducing collapse and lateral expansion of the belt, or (2) a drastic change of the boundary conditions has induced hot asthenospheric upwelling which in turn led to coeval extension and partial-melting. At a more local scale, strain benefited of the low-strength of the magmatic bodies prior to complete crystallisation promoting intense strain localisation within the South Armorican domain large-scale laccoliths often referred to as synkinematic plutons.

Un tournant dans l'évolution de l'orogène Varisque : l'événement régional de fusion partielle à 325 Ma du domaine Sud Armoricaïn côtier (Bretagne Sud et Vendée, France)

Mots-clés. – Fusion partielle, Localisation de la déformation, Effondrement gravitaire, U-Th/Pb sur monazite, Domaine Sud Armoricaïn

Résumé. – En réduisant considérablement la résistance des matériaux crustaux, la fusion partielle est un paramètre majeur contrôlant le comportement rhéologique de la croûte continentale. Avec plus de 50% de surface d'affleurement caractérisée par des migmatites et des granites, la partie côtière du domaine Sud Armoricaïn offre l'opportunité d'étudier les processus orogéniques profonds et plus particulièrement de cerner le rôle de la fusion partielle dans l'évolution tardive de la chaîne Varisque. Pour l'heure, très peu de contraintes géochronologiques sont disponibles entravant la compréhension de cette étape cruciale dans l'évolution de la chaîne. Cette étude apporte un total de 29 nouvelles datations chimiques U-Th/Pb sur monazite provenant de cinq zones d'échantillonnage composées de dômes migmatitiques et de granites associés ou tardifs. Sur des critères structuraux, texturaux et chimiques, trois principaux groupes d'âges U-Th/Pb ont été reconnus.

Le premier groupe, principalement centré vers 335-330 Ma a été mis en évidence pour des échantillons de restites et pour les domaines de cœur des cristaux de monazite de la plupart de massifs de granite. Sa signification est attribuée à des âges de cristallisation hérités enregistrant probablement un incrément prograde lors de l'augmentation des conditions P-T dans un contexte de mise en place de nappes. Le second groupe d'âges, vers 325-320 Ma correspond aux cristaux de monazite ayant cristallisés à partir de liquides silicatés néoformés. Il est interprété comme l'âge de cristallisation des leucosomes des roches partiellement fondues suite à un événement régional de fusion partielle. Le dernier groupe d'âges, vers 320 Ma, correspondant aux domaines de bordure des cristaux de monazite des massifs de granite "tardifs". Il est interprété comme l'âge de mise en place de la plupart des grands massifs granitiques et clos ainsi l'épisode de fusion partielle.

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L'introduction et la généralisation de la fusion partielle initiée dans les conditions du pic métamorphique coïncident donc avec un changement majeur du régime tectonique à l'échelle régionale. A la lumière de ces résultats, cela implique (1) soit que l'empilement d'unités crustales, riches en radioéléments a conduit à une augmentation de la température au sein du prisme orogénique provoquant la fusion partielle, la chute de résistance de la croûte entraînant alors l'effondrement et l'échappement latéral de la chaîne, (2) soit qu'un changement important des conditions aux limites ait induit la remontée d'asthénosphère chaude ayant conduit à l'apparition synchrone d'extension en surface et de fusion partielle en profondeur. A une échelle plus locale, la déformation semble avoir profité de la faible résistance des corps magmatiques avant leur cristallisation complète favorisant la localisation de la déformation dans les grands laccolithes du domaine Sud Armoricaïn souvent désignés comme des plutons syncinématiques.

INTRODUCTION

Partially molten rocks or migmatites are a common feature of mature orogens [i.e. Sawyer, 1994, 1999; Brown, 1994, 2001; Vanderhaeghe and Teyssier, 2001; Vanderhaeghe, 2004]. Together with strain rates and thermal conditions, partial-melting controls the rheological behaviour of crustal materials by drastically reducing the bulk strength of the crust [i.e. Arzi, 1978; Vigneresse *et al.*, 1999; Rosenberg and Handy, 2005]. Along with a decrease of tectonic boundary forces, a decrease in the strength of a thickened crust is associated with gravitational collapse of an orogenic belt and the transition from crustal thickening to crustal thinning [i.e. Rey *et al.*, 2001; Vanderhaeghe and Teyssier, 2001; Tirel *et al.*, 2008; Rey *et al.*, 2009a; 2009b]. Inception of partial-melting in the core of an orogenic belt may thus promote drastic changes in the balance of forces and thus represents a major turning-point in its evolution. Partial-melting in the thickened crust can either be an ultimate consequence of the stacking of radiogenic continental crustal slices or is due to a drastic change in heat supply from underneath as in the case of convective removal or slab retreat [England and Thompson, 1986; Houseman, 1988; Vanderhaeghe *et al.*, 2003].

Besides, intrusion of magmas in higher structural levels or even in the upper brittle crust can locally promote strain localisation as shown by the formation of detachments at the roof of plutons [i.e. Daniel and Jolivet, 1995; Grasemann and Petrakakis, 2007; Jolivet *et al.*, 2010] or even along volcanic sills or feeders [Famin and Michon, 2009]. The low strength of a magmatic body must have, at least temporarily, drastic consequences on strain localisation [Hutton, 1982; Neves *et al.*, 1996]. It is thus crucial to recognise the timing of both the bulk loss of crustal strength during large-scale partial-melting events and localising events during magmatic intrusions at more local-scale. In both cases, accurate dating of the crystallisation of melts either trapped at depth in migmatite complexes or after upward migration as plutons or dykes may put crucial age-constraints on important changes on mountain belts dynamics.

The case of the West-European segment of the Variscan belt is quite exemplary of this topic. Indeed, one of the most striking features of the belt is the abundance of both migmatitic and plutonic rocks, partly obliterating the initial architecture of the orogen. In the coastal South Armoricaïn massif, an almost continuous belt in which migmatites and

of typical "migmatite-cored metamorphic core complexes" [Brown, 2005; Turrillot *et al.*, 2011a]. Moreover, most of the major shear zones are underlined by syntectonic granite laccoliths, strengthening the link between strain localisation and presence of melt at the scale of the pluton [Gapais *et al.*, 1993; Cagnard *et al.*, 2004; Turrillot, 2010]. However, to date, time-constraints are scarce and partly conflicting, hindering the understanding of this crucial stage in the Variscan belt evolution. This paper provides 29 new, and sometimes pioneer, U-Th/Pb chemical ages on monazite from the main occurrences of migmatites and regional classic plutons exposed along the coast of the South Armoricaïn domain. Together with an accurate separation of inheritance from neof ormation, statistics on the dataset ensures internal consistency of the time-constraints at regional-scale, shedding light on the dynamics of the late-orogenic stages of the West-European Variscanides.

GEOLOGICAL SETTING

The French Armoricaïn massif (fig. 1) is one of the largest pieces of the Variscan belt that extends throughout western and central Europe [i.e. Matte, 1991, 2001]. Together with the Alleghanian belt, further to the West, the Variscan belt

resulted from the convergence of Laurussia and Gondwana that participated in the formation of the supercontinent Pangaea during the Late Paleozoic [e.g. Ballèvre, 2007; Ballèvre *et al.*, 2009; Nance *et al.*, 2010]. The structural pattern of the Armoricaïn massif is currently dominated by Carboniferous W-E to NW-SE striking, dextral strike-slip faults, namely the North Armoricaïn shear zone (NASZ), and the South Armoricaïn shear zone (SASZ) (fig. 1B [e.g. Cogné, 1974; Watts and Williams, 1979; Jégouzo, 1980; Matte, 1986; Le Corre *et al.*, 1991; Tartèse *et al.*, 2011a]). Furthermore, the SASZ is subdivided into an E-W, and a NW-SE striking northern and southern branches, respectively. From the paleogeographic point of view, the Armoricaïn massif is divided into three main distinct parts, namely the Léon, North-Central, and South-Armoricaïn domains bounded by major tectonic contacts whose importance is compared to suture zones: the Conquet-Penzé (LCPS), and Nort-sur-Erdre (NSES) sutures (fig. 1) [e.g. Cogné, 1974; Jégouzo, 1980; Le Corre *et al.*, 1991; Jégouzo and Rossello, 1988; Cartier *et al.*, 2001; Cartier and Faure, 2004; Gumiaux *et al.*, 2004; Faure *et al.*, 2005]. These domains drastically differ by their metamorphic grade and structural

granites currently represent ca. 50% of the outcropping surface has long been described [e.g. Cogné, 1974]. Besides, a close association between crustal partial-melting, magmas and the development of extensional shear zones (i.e. detachments) has been recently reemphasised with the description style [e.g. Le Corre *et al.*, 1991; Ballèvre *et al.*, 2009; Faure *et al.*, 2008a]. The North-Central Armorican domains experienced only moderate shortening and shearing between two main suture zones and mainly behaved as rigid, upper-crustal continental domains [i.e. Le Corre, 1977; Le

Corre *et al.*, 1991]. Conversely, the Léon and the South Armorican domains experienced a succession of tectono-metamorphic events related to subduction, collision and finally gravitational collapse phases [Gapais *et al.*, 1993; Ballèvre *et al.*, 2009]. The South Armorican domain consists of two parts. Between the Nort-sur-Erdre eo-Variscan suture and the southern branch of the SASZ, the Ligerian domain is an area characterised by Early Paleozoic ophiolites, high-pressure mafic and felsic rocks, and Devonian migmatites recently dated at ca. 387 ± 8 Ma (fig. 1) [Marchand, 1981; Ballèvre *et al.*, 1994; Bosse *et al.*, 2000; Cocherie *et al.*, 2005; Faure *et al.*, 2008b]. To the south of the southern branch of the SASZ, the coastal South Armorican domain exhibits metamorphic and granitic rocks [Cogné, 1974; Berthé *et al.*, 1979; Jégouzo, 1980].

In the following, only the geology of the coastal South Armorican domain is detailed (fig. 1a). Readers unfamiliar with the geology of the Armorican massif are referred to recent publications [Gumiaux *et al.*, 2004; Faure *et al.*, 2005, 2008b; Ballèvre *et al.*, 2009]. Regionally, the coastal South Armorican domain can be divided in three main groups of units, from top to bottom (fig. 1b) [e.g., Cogné, 1949, 1974; Le Corre *et al.*, 1991; Gapais *et al.*, 1993; Cagnard *et al.*, 2004; Turrillot *et al.*, 2011a]:

1) *the Upper units*, formed by metamorphic rocks that have been equilibrated along a ca. $10^\circ\text{C}/\text{km}$ HP/LT subduction gradient. The uppermost units, exposed in the Bois de Céné and Ile de Groix klippe are made of chloritoid-micaschists, lawsonite-blueschists, garnet-blueschists and eclogites derived from metasediments, metabasites and rare felsic layers [Bosse *et al.*, 2000; El Korh, 2010]. The presence of both tholeiitic and alkaline signatures of the metabasites and Mn-rich metasediments argues for an initial oceanic setting or an intracontinental rift [Bernard-Griffiths *et al.*, 1986; Audren and Plaine, 1986; Audren *et al.*, 1997; El Korh *et al.*, 2012]. Metamorphic peak-conditions reached around 18-20 kbar and $450\text{-}500^\circ\text{C}$ [Bosse *et al.*, 2002; Ballèvre *et al.*, 2003] and were followed by a limited retrogression into the greenschist facies conditions [Bosse *et al.*, 2002]. The lower ones (i.e. Belle-Ile group) consist of monotonous series of metamorphosed volcanics intercalated with black-shales (i.e. the Vendée and Belle-Ile-en-Mer porphyroids) and graphite-rich micaschists (i.e. Saint-Gilles micaschists). Age of the series was recently settled at ca. 480 Ma in Vendée by dating the emplacement age of the porphyroids [Ballèvre *et al.*, 2012]. P-T estimates for the Belle-Ile group rocks yielded HP/LT metamorphic peak-conditions of 7-9 kbar and $350\text{-}400^\circ\text{C}$ [Le Hébel *et al.*, 2002a];

2) *the Intermediate units* (i.e. the Vilaine Estuary and Sables d'Olonne units) made of metasediments in association with rare metavolcanics were equilibrated under a MP-MT gradient [Brown and Dallmeyer, 1996] while relics of HP/LT metamorphism have not been recognised. Metamorphism increases downward from albite greenschist-facies to garnet-staurolite-biotite amphibolite-facies conditions [Bossière, 1988; Triboulet and Audren, 1988; Brown and Dallmeyer, 1996] and locally to partial-melting conditions to the southern Vendée [Goujou, 1992; Goujou *et al.*, 1994]. Estimates of the metamorphic peak-conditions yielded 7-9 kbar for 620°C in the Vilaine Estuary area [Triboulet and Audren, 1985, 1988; Brown and Dallmeyer, 1996] and 9 kbar for $650\text{-}700^\circ\text{C}$ to the north of the Sables d'Olonne area, where localised partial-melting has been widely

described [e.g. Ters, 1972; Brillanceau, 1978; Goujou, 1992; Cagnard *et al.*, 2004]:

3) *the Lower unit* consists of high-grade rocks, migmatites and large volumes of granite recognised as the South Brittany Migmatite belt that extends to the south in Vendée [Marchildon and Brown, 2003]. Migmatites include restitic rocks with granulites, and kinzigites derived from pre-Carboniferous metasediments and Ordovician orthogneiss protoliths [Audren and Le Métour, 1976; Peucat *et al.*, 1978, 1979; Brown, 1983; Peucat, 1983; Audren, 1987; Jones and Brown, 1990; Marchildon and Brown, 2003]. P-T estimates yielded metamorphic peak-conditions at 8-10 kbar and $750\text{-}800^\circ\text{C}$ accompanied and followed by distinctive melt extraction stages down to 4 kbar and 750°C [Brown, 1983; Jones and Brown, 1989, 1990; Jones, 1991; Goujou, 1992; Audren and Triboulet, 1993; Brown and Dallmeyer, 1996; Marchildon and Brown, 2003; Johnson and Brown, 2004].

Until now, the knowledge of the tectonic evolution of the coastal South Armorican domain remains incomplete. However, an overall two-stepped tectonometamorphic evolution of the domain can be drawn including an initial episode of subduction and crustal-thickening tectonics followed by pervasive extensional tectonics [Brun and Burg, 1982; Burg *et al.*, 1987, 1994; Gapais *et al.*, 1993]. Evidence for the subduction stages are quite well preserved in the upper units, particularly in the Ile de Groix and the Bois de Céné klippe where typical subduction-related blueschists and eclogites occur in different tectonic slices [e.g. Barrois, 1883; Carpenter, 1976; Quinquis *et al.*, 1978; Bosse *et al.*, 2002; Ballèvre *et al.*, 2003]. Architecture of the Ile de Groix has even been interpreted as a remnant of a south-east-verging accretionary wedge related to a northward dipping subduction [Ballèvre *et al.*, 2007]. Metamorphic phengite from HP assemblages yielded Late Devonian-Early Carboniferous Rb/Sr and $^{40}\text{Ar}/^{39}\text{Ar}$ ages at 370-360 Ma, while retrogression and final cooling under greenschist-facies conditions occurred at around 355-345 Ma [Le Hébel, 2002; Bosse *et al.*, 2005]. These uppermost units are systematically found at the top of the Variscan edifice, overthrust on lower pressure continental units [Bosse *et al.*, 2005]. Subsequent accretion of continental slices by thrusting in the orogenic wedge is also locally preserved as exemplified by the tectonic superimposition of the porphyroids HP/LT units over low-grade metasediments in Vendée [Ters, 1972; Iglesias and Brun, 1976; Burg, 1981; Maillet, 1984; Colchen and Poncet, 1987; Vauchez *et al.*, 1987; Goujou, 1992; Geoffroy, 1993]. While N-S directions and a preferred top-to-the South sense of shear were proposed [Cannat and Bouchez, 1986], kinematics of thrusting at a regional-scale remain currently not documented [Le Hébel, 2002]. Similarly, the age of final nappe-stacking is only locally settled to the Tournaisian (i.e. 359-347 Ma) as recorded by syn-thrusting sedimentation [Maillet, 1984; Colchen and Poncet, 1987; Rolin and Colchen, 2001].

The tectono-metamorphic record of the crustal thickening stages were erased by pervasive partial-melting and extensional tectonics during a subsequent late-orogenic stage [Burg, 1981; Brun and Burg, 1982; Cannat and Bouchez, 1986; Gapais *et al.*, 1993; Marchildon and Brown, 2003; Cagnard *et al.*, 2004; Brown, 2005; Turrillot *et al.*, 2009, 2011a]. During this stage, the Lower unit was exhumed as a series of domes locally referred to as

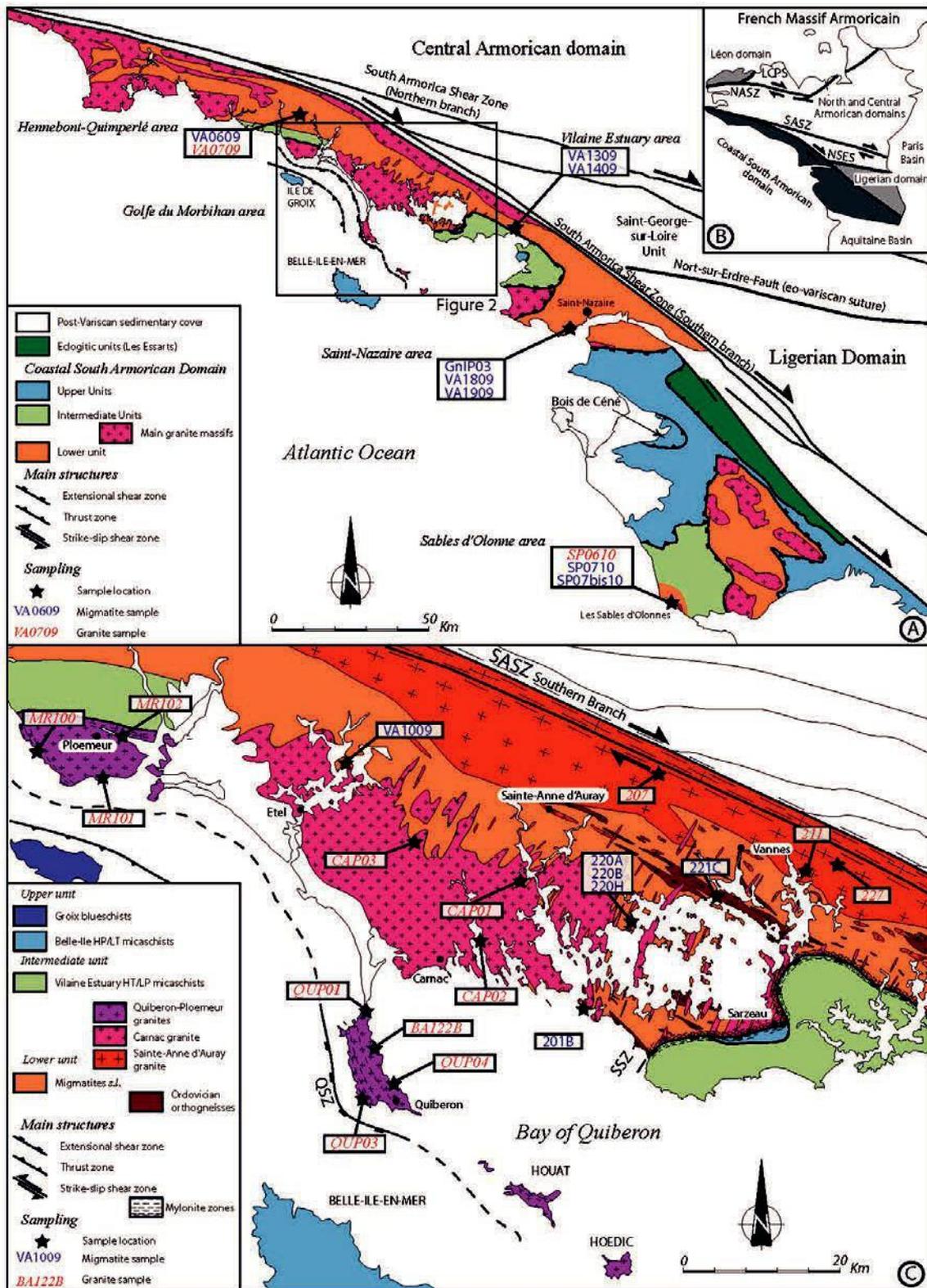


FIG. 1. – Geological maps of the studied area.

(a) Simplified geological map of the coastal South Armoricain domain [modified after Gapais *et al.*, 2009 and Bitri *et al.*, 2010]. Indicated is the location of the samples used for U-Th/Pb chemical datings. (b) Inset is a sketch map of the French Armoricain massif and its main subdivisions. The black rectangle is the location of Golfe du Morbihan area presented in figure 1c. (c) Simplified geological map of the Golfe du Morbihan area [Turrillot, 2010]. The location of samples that provided ages for Sainte-Anne d'Auray and Sarzeau granites is also indicated [Turrillot *et al.*, 2009]. Only the units corresponding to the coastal South Armoricain domain have been represented. Note the close association between the migmatite fabrics and the anatexis bodies and particularly the Sainte-Anne d'Auray anatexis granite massif. Conversely, the Carnac, Ploemeur and Quiberon granite intrusions clearly cut across the migmatite fabrics and the Sainte-Anne d'Auray granite. These plutons locally emplaced within the upper units.

migmatite-cored metamorphic core complexes [Brown, 2005; Turrillot *et al.*, 2011a] surrounded by the overlying intermediate and Upper units (fig. 1) [Gapais *et al.*, 1993; Burg *et al.*, 1994; Le Hébel, 2002]. At regional-scale, deformation associated with the exhumation of the Lower Unit is marked by a regional flat-lying planar fabric, and a dominantly W-E to WNW-ESE striking stretching lineation with the exception of the Saint-Nazaire area where the stretching direction strikes almost ca. N-S [Bouchez *et al.*, 1981; Gapais *et al.*, 1993; Turrillot, 2010]. Besides, despite striking structural similarities, two contrasted modes of crustal extension were opposed within the coastal South Armorican domain [Cagnard *et al.*, 2004], and one can distinguish from southeast to northwest: 1) the southeastern end of the domain, in the Sables d'Olonne area, where extensional tectonics is mainly accommodated by pervasive vertical thinning of the entire metamorphic series [Iglesias and Brun, 1976; Goujou, 1992] and 2) the Golfe du Morbihan area and to some extent, the Saint-Nazaire area toward the north-west of the domain, where, conversely, most of extension appear concentrated along major extensional shear zones [e.g. Bouchez *et al.*, 1981; Guillet *et al.*, 1983; Gapais *et al.*, 1993; Turrillot *et al.*, 2009]. Strikingly, most of these extensional shear zones are marked by large tabular two-mica plutons emplaced during shearing within the base of overlying units [Vignerresse, 1978; Gapais *et al.*, 1993; Gumiaux, 2004; Turrillot *et al.*, 2011b]. Most of these plutons remain undated.

Initially attributed to a Proterozoic basement [Ters, 1979; Wyns and Le Métour, 1983], migmatites of the Intermediate and the Lower units yielded two age groups: a Devonian group including ages at 376 ± 19 Ma and 372 ± 24 Ma in the Golfe du Morbihan area [Peucat, 1983] and at 388 ± 8 Ma in the Sables d'Olonne area [Vidal, 1980] by whole-rock Rb/Sr, and zircon U/Pb methods and a younger, ca. 320 Ma age group (i.e. $^{207}\text{Pb}/^{206}\text{Pb}$ 323 ± 4 Ma and 315 ± 9 Ma ages on monazite) [Peucat, 1983]. However, these Devonian ages were recently questioned in the Golfe du Morbihan area by monazite *in situ* U-Th/Pb chemical dating from a foliation-parallel granitic stock closely associated to the migmatites that yielded a 322 ± 3 Ma age [Turrillot *et al.*, 2009]. This Late Carboniferous age complies with the clustering of many isotopic systems around ca. 315-300 Ma both in the Lower unit and the base of the Intermediate unit, suggesting a common final cooling [Carpena *et al.*, 1979; Peucat, 1983; Goujou, 1992; Gapais *et al.*, 1993; Brown and Dallmeyer, 1996; Turrillot *et al.*, 2011a]. Besides, geochronological data of synkinematic crystallisations along the Quiberon and the Sarzeau extensional shear zones yielded similar 305-300 Ma ages [Brown and Dallmeyer, 1996; Ruffet *in Le Hébel*, 2002; Turrillot *et al.*, 2011a; Tartèse *et al.*, 2011b].

SAMPLING, METHOD AND PROCEDURE

Sampling strategy and description of the outcrop conditions

Fourteen samples of migmatite were collected from five zones distributed along the coastal South Armorican domain (fig. 1a), namely, from the NW to the SE, the Hennebont-Quimperlé, the Golfe du Morbihan (fig. 1c), the Vilaine Estuary, the Saint-Nazaire and the Sables d'Olonne

areas. Samples were all collected from outcrops showing unambiguous fabrics. Besides, samples sometimes carry evidence of weathering but are systematically devoid of macroscopic alteration. In order to avoid mixing between age-inheritance and melt-related crystallization ages, a particular attention has been paid to separate the juvenile melt fractions, represented by leucosomes either parallel or discordant to the foliation, from cumulative restites (i.e. melanosome) or from remnants of the initial gneiss (i.e. paleosome or mesosome) (fig. 2) [*sensu* Mehnert, 1968; Brown, 1973; Vanderhaeghe, 2009]. To gather independent age-constraints, this set of samples was completed by 15 granite samples. Among the collection, 5 samples correspond to the diffuse-boundary granites often referred to as “anatectic” granites, closely associated with the migmatites [Carron *et al.*, 1994]. The 10 other samples correspond to “late”, sharp-boundary granites cutting across the migmatite fabrics either at the scale of the map (fig. 1c) or the scale of the outcrop (fig. 2, 3, 4). A particular attention was paid on the Golfe du Morbihan area, which is, to date the most studied segment of the coastal South Armorican domain as it offers different generations of intrusive rocks of contrasted chemistry and mineralogy (fig. 1c) [Augier *et al.*, 2011]. Accordingly, a set of 19 samples was specifically picked from migmatites and different granite generations (i.e. 6 and 13 samples, respectively). Location of all samples is re-ported in figure 1.

The Golfe du Morbihan area

The Golfe du Morbihan area appears as the best exposed occurrence of migmatite of the Lower unit roofed by two major extensional shear zones characterised by opposite kinematics [e.g., Gapais *et al.*, 1993; Turrillot *et al.*, 2009]. In the core of the Lower unit, migmatites present a N110-120°E striking subvertical foliation [Audren, 1987; Turrillot *et al.*, 2009, 2011a; Augier *et al.*, 2011]. Conversely, foliation adopts a gently dipping attitude in the vicinity of shear zones. There,

a conspicuous N115°E stretching lineation is associated with non-coaxial kinematic indicators [e.g. Gapais *et al.*, 1993; Turrillot *et al.*, 2009]. Several generations of granite were distinguished including anatectic granite bodies that are diffuse within the migmatite and “late” granite plutons and dykes clearly cutting across the migmatite fabrics [e.g. Carron *et al.*, 1994; Augier *et al.*, 2011]. Anatectic granites occur in close association with migmatites as various scale bodies and develops North of Vannes as a 70 km long elongated massif, regionally known as the Sainte-Anne d'Auray magmatic complex [Cogné, 1974; Carron *et al.*, 1994; Augier *et al.*, 2011]. Among the “late” intrusive granites, the Carnac magmatic complex is composed of a main pluton and a pervasive dyke array [Carron *et al.*, 1994; Augier *et al.*, 2011]. Finally, to the southwest a series of two-mica leucogranite plutons including the Ploemeur, Quiberon and Houat-Hoedic plutons are emplaced within the Intermediate and the Upper units. Intensively sheared during their emplacement, these plutons may have triggered, at least locally, strain localisation [Gapais *et al.*, 1993; Turrillot *et al.*, 2009; Turrillot, 2010].

In the field, depending on the bulk protolith chemistry, on the mechanisms of melt extraction or on subtle differences of peak-metamorphic conditions, migmatites display a wide range of textures and facies (e.g. compare fig. 3a to

fig. 3b, fig. 3c and fig. 3d) [Brown, 1983; Audren, 1987; Jones and Brown, 1990; Audren and Triboulet, 1993; Brown and Dallmeyer, 1996; Marchildon and Brown, 2003]. However, high-grade gneisses, metatexites, diatexites and anatectic granite stocks are unequally distributed and an overall northward increase in the amount of both the *in situ* melt and the melt collected from the center of the Golfe du Morbihan is classically described [i.e. Audren, 1987; Audren and Triboulet, 1993]. 19 samples were selected throughout the area. 6 samples of migmatite were picked from the northern part of the Ria d'Étel, to the west, to the southern part of Golfe du Morbihan, to the east (fig. 1c). Underlined by the alternation of quartz-feldspar irregular leucosomes and dark biotite (-sillimanite-garnet) melanosomes, the foliation is dominantly oriented N110-130°E and locally isoclinally folded (fig. 3a). Samples 201, 220A and 220B were sampled from leucosome-rich metatexites (fig. 3b). Conversely, samples 220H and 221C were selected from biotite-sillimanite-garnet restitic layers (fig. 3c). In order to evaluate the effects of local inheritance on the U-Th/Pb system, samples 220A, 220B and 220H were selected on the same outcrop 10 centimetres apart (fig. 3b). Sample VA1009, located to the north of the Ria d'Étel, is a diatexite that presents a well-developed subhorizontal foliation and a

conspicuous N115°E stretching lineation (fig. 3d). To strengthen the probable genetic link between the migmatites and the anatectic massifs, 3 samples were collected in different petrographic facies throughout the Sainte-Anne d'Auray massif (i.e. samples 207, 211 and 227). Samples 207 and 211 were selected from a porphyritic biotite granite and coarse-grained granite facies (e.g. fig. 3e), while sample 227 was selected from a two-mica fine-grained granite [Augier *et al.*, 2011].

In addition, 13 samples are representative of the "late" intrusive granite massifs (fig. 1c). 3 samples were collected from the Carnac massif to illustrate the petrological variability within the pluton from a biotite (-cordierite) facies (i.e. sample CAP01) to the west, to two-mica granite facies (i.e. samples CAP02 and CAP03). In the field, Carnac granite dykes clearly cut across the migmatite fabrics and the Sainte-Anne d'Auray granite (figs 2 and 3f). 7 samples were collected from the Ploemeur and the Quiberon massifs (i.e. 3 and 4 samples, respectively). The dense sampling in the Ploemeur and the Quiberon two-mica leucogranites reflects the commitment to explore the possible effects of the mylonitic deformation and/or associated fluid circulations on the U-Th/Pb system in monazite. Sample BA122b, collected

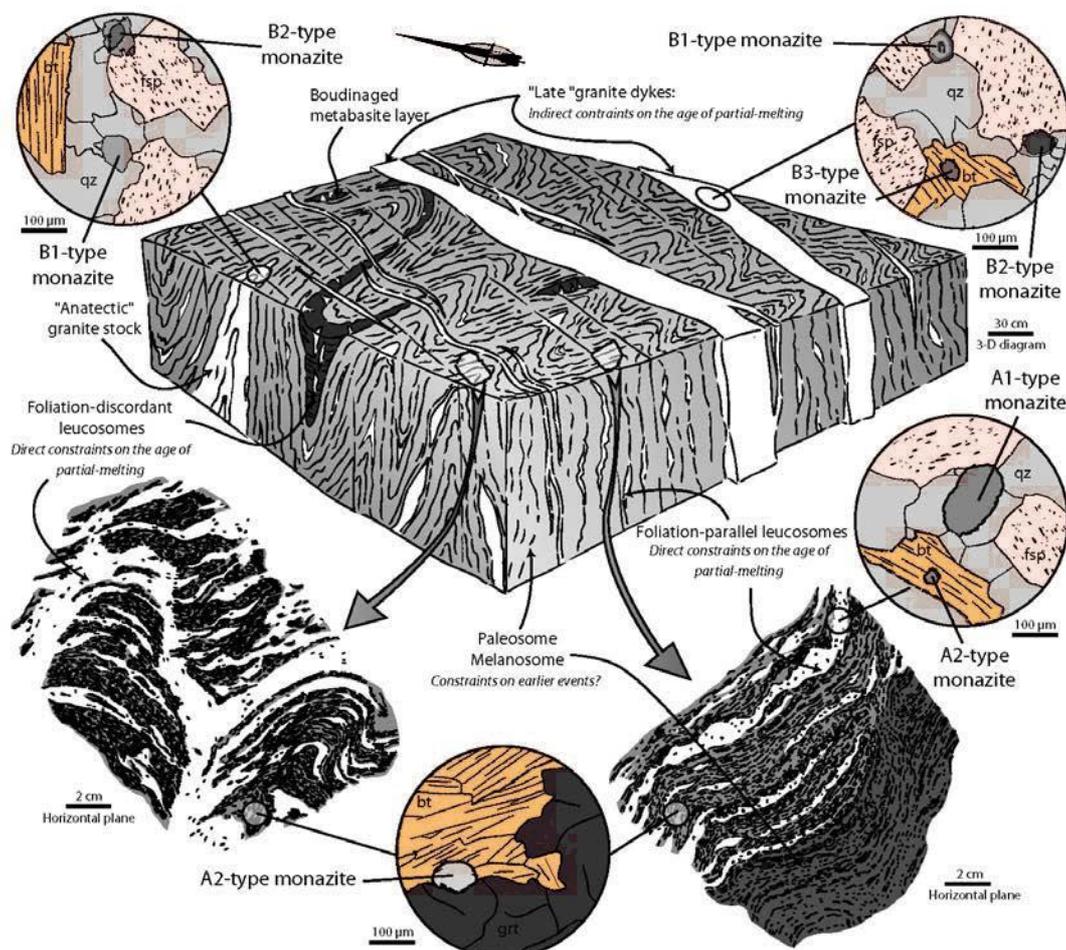


FIG. 2. – Sampling strategy scheme.

For each zone, melt-bearing migmatite samples were picked to gather the opportunity to sample newly-formed material and ultimately newly-formed monazite grains. When possible, associated anatectic, generally foliation-parallel granite bodies were also sampled. Conversely, melanosome layers were collected to explore metamorphic record inheritance. Late to post-migmatite granite intrusions that remains poorly dated were also sampled to ensure an independent minimum age for the partial-melting event. Mineral abbreviations are after Whitney and Evans [2010].

in the eastern part, of the Quiberon peninsula is a weakly deformed granite sample. Samples MR100 and MR101 from the Ploemeur granite and samples QUP01 and QUP04

from the Quiberon granite are deformed granite samples (fig. 3g). Representative outcrops are characterised by heterogeneously developed planar-linear fabrics and locally

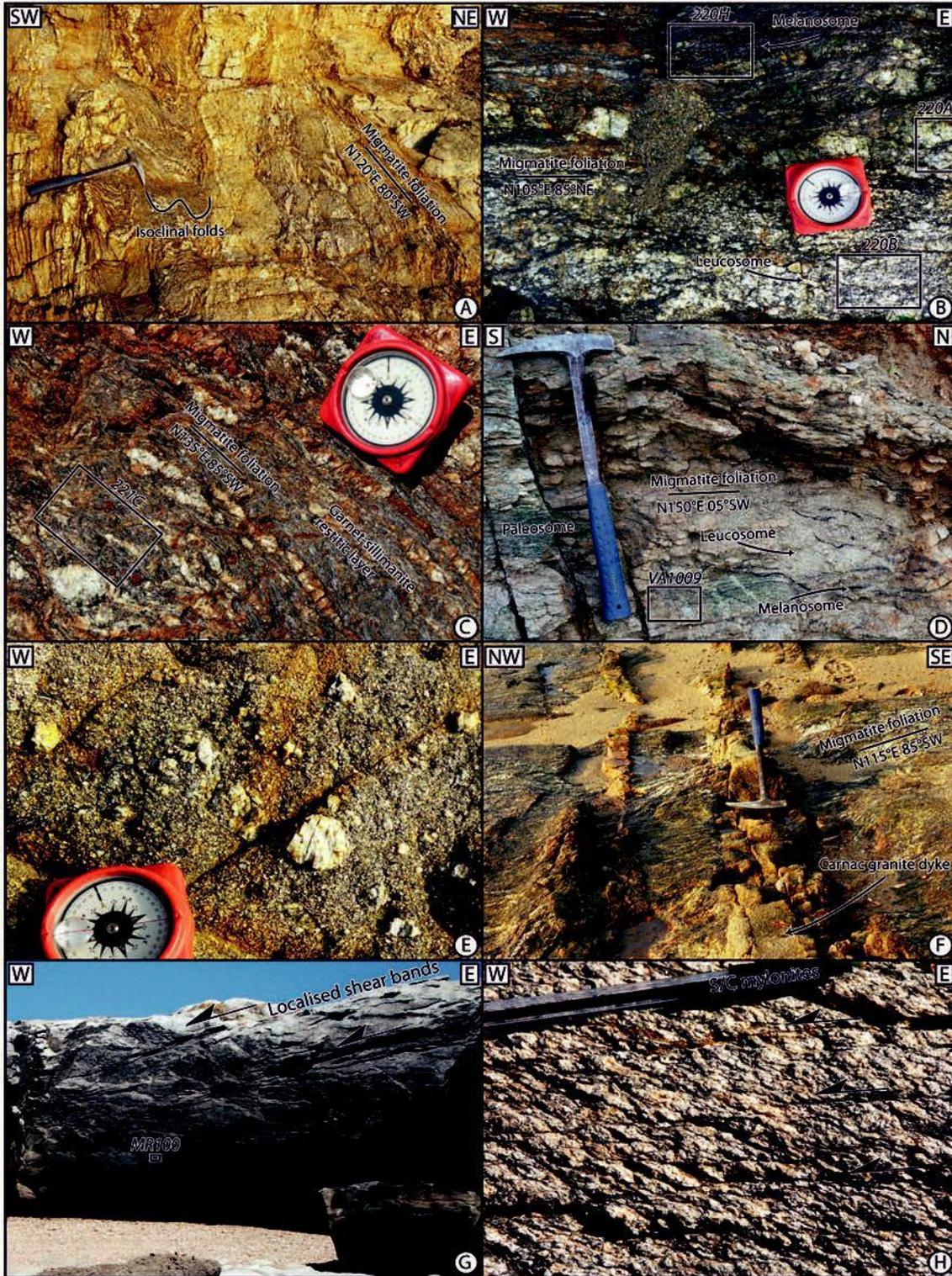


FIG. 3. – Representative pictures of the outcrop conditions of the migmatites and the granites in the Golfe du Morbihan area. (a, b and c) Different outcrops of the Golfe du Morbihan showing the general aspects of the migmatites. Most outcrops are characterized by metatexites (a and b) while some consists in restitic layers (c). (d) Typical outcrop of the Etel area where migmatites present a subhorizontal attitude. (e) Detailed view of the heterogeneous porphyritic biotite granite facies of the Sainte-Anne d’Auray granite, few tens of meters north of the contact with the migmatites. (f) Cross-cutting relationships between the Carnac granite and the migmatites. Note that granite dykes clearly cut across the migmatite fabrics. (g) Deformed facies of the Ploemeur granite. Foliation is heterogeneously developed and most of the deformation concentrates along low-angle top-to-the-W extensional shear zones. (h) S/C mylonites developed at the expense of the Quiberon granite within the Quiberon extensional shear zone.

displays evidence of meter-scale top-to-the-West shear zones (fig. 4g) [Turrillot, 2010]. Finally, sample MR102 from the Ploemeur granite and sample QUP03 from the Quiberon granite are typical S-C granite mylonites (fig. 4h) collected along the most deformed zones.

The Hennebont-Quimperlé area

The Hennebont-Quimperlé area is characterised by high-grade paragneisses and migmatites closely associated with anatectic granite bodies (fig. 1). Migmatites present N110-120°E striking subvertical foliation [Cogné *et al.*, 1986]. To the south, the migmatites present a rather gradual transition to the overlying micaschists units without evidence of important detachment zones [Cogné *et al.*, 1986; Béchenec *et al.*, 2012].

In the field, metatexites exhibit a rhythmic alternation of biotite (sillimanite melanosomes and scarce well-developed quartz-feldspar leucosomes or small-scale anatectic granite stocks (fig. 4a). Laterally, the migmatites evolve to larger-scale lensoid stocks of anatectic granite bodies with 10-100 m typical dimension. Sample VA0609 was picked from a metatexite (fig. 4a). Sample VA0709 was sampled 5 km to the SE along the strike of the foliation from a ca. 2 km thick anatectic granite stock [Cogné *et al.*, 1986; Béchenec *et al.*, 2012].

The Vilaine Estuary area

The Vilaine Estuary area is characterised by a 2 km large band exposing migmatites, limited to the north by the Sainte-Anne d'Auray anatectic granite through a diffuse transition [Plaine *et al.*, 1984]. To the south, the migmatites are separated from the Intermediate units by a ca. 100-150 m thick gently-dipping mylonite zone belonging to the Sarzeau extensional shear zone [Plaine *et al.*, 1984; Turrillot *et al.*, 2009, 2011a; Augier *et al.*, 2011]. While the N120°E strike of foliation remains consistent regionally, the dip of the metamorphic layers evolves from subvertical in the vicinity of the Sainte-Anne d'Auray granite to gentle (20-40°) toward the south and the Sarzeau shear zone [Plaine *et al.*, 1984; Augier *et al.*, 2011].

At the outcrop-scale, most rocks present conspicuous evidence of partial-melting and the increase in the amount of the *in situ* melt toward the north, where a ca. 500 m large diffuse zone marks the southern boundary of the Sainte-Anne d'Auray granite massif (fig. 1) [Plaine *et al.*, 1984]. Two samples, (VA1309 and VA1409) were picked from leucosome-rich metatexites at ca. 500-700 m to the south of the Sainte-Anne d'Auray granite massif (fig. 4b).

The Saint-Nazaire area

The Saint-Nazaire area appears as a large outcrop of the Lower unit (fig. 1). The overall structure of the Lower unit defines a structural dome, namely the St-Nazaire dome [Cogné, 1974], with a dominant gently north-dipping foliation to the northwest (i.e. La Baule area) and a gently south-dipping foliation to the south (i.e. Saint Brévin area) [Ters *et al.*, 1978; Ters, 1979]. On both flanks, the foliation bears a N-S striking lineation [Bouchez *et al.*, 1981; Guillet *et al.*, 1983]. To the south, the dome is roofed by an extensional shear zone [e.g. Gapais *et al.*, 1993; Turrillot, 2010; Bitri *et al.*, 2010]. To the north, the contact with the

Intermediate unit is underlined by a large-scale syn-kinematic gently north-dipping laccolith called the Guérande pluton [Vignerresse *et al.*, 1978; Guillet *et al.*, 1983].

In the field (mostly along the coast), the migmatites consist of coarse-grained alternations of quartz and feldspar or even pegmatitic layers and of dark biotite, sillimanite and garnet layers. Locally, diatexites and anatectic granite stocks crop out in alternation with massive restitic layers or boudins embedded in a heterogeneous granite (fig. 4c). Sample GnIP03 corresponds to a migmatitic gneiss. Sample VA1909 was collected from a large leucosome stock within typical metatexites (fig. 4d), while sample VA1809 was sampled from a large restitic body (fig. 4e).

The Sables d'Olonne area

The regional-classic coastal section of the Sables d'Olonne area shows a complete Barrovian metamorphic sequence, from low-grade metasediments to the north, to migmatites to the south in the Sable d'Olonne bay [Ters, 1972; Iglesias and Brun, 1976; Goujou, 1992; Goujou *et al.*, 1994] without evidence of major detachment zones [Cagnard *et al.*, 2004]. E-W-striking and N-dipping in the northern part of the area, and WNW-ESE-striking and NE-dipping in the eastern part (fig. 4f), the overall attitude of the foliation is consistent with the dome-shaped structure [Goujou, 1992]. Along with metamorphic grade, the intensity of the foliation strongly increases toward the north [Goujou, 1992; Goujou *et al.*, 1994; Rolin and Colchen, 2001]. The foliation carries a clear E-W to WNW-ESE striking stretching lineation [Iglesias and Brun, 1976], while a dominant top-to-the-west sense of shear is observed [Brun and Burg, 1982; Cannat and Bouchez, 1986; Goujou, 1992; Cagnard *et al.*, 2004; Turrillot, 2010].

In the field, conspicuous traces of partial-melting are only restricted to the southernmost parts of the section (fig. 1a). There, the N-dipping foliation is locally outlined by cm to dm-sized leucosomes. A dense array of variably deformed dykes cuts across the migmatite fabrics (figs. 4f and 4g). Most of the dykes display conspicuous internal (i.e. foliation and lineation) or external strain (e.g. folds, boudins; fig. 4g) consistent with the strain axes defined from the migmatites [Cagnard *et al.*, 2004]. Interestingly, some leucosome layers appear connected to the dykes indicating a clear synkinematic melt extraction. Samples SP0710 and SP07bis10 were sampled few meters apart from leucosome-poor metatexites of probably contrasted meta-sedimentary protoliths. Rich in biotite and sillimanite, sample SP0710 rather corresponds to a metapelitic protolith while SP07bis10, rich in quartz to an impure mica-rich meta-quartzite. Conversely, sample SP0610 was sampled within a granite dyke transposed in the regional foliation (fig. 4g).

Sample preparation, analytical procedure and data processing

Depending on the nature of the rock, preparation followed two different strategies. Because of the intrinsic complexity of migmatite in which neocrystallisation and inheritance are common, migmatite samples were studied on polished thin-sections to perform *in situ* chemical datings on monazite (fig. 2). *In situ* analysis allows for combining geo-chronology and rock-texture (i.e. phases equilibria and ultimately P-T conditions) to make inferences on the crystallization conditions of monazite [e.g. Pyle and Spear,

2003; Goncalves *et al.*, 2004], particularly relevant for dating metamorphic rocks and migmatites. To ensure an accurate textural control, thin sections were cut in the XZ

structural plane. Conversely, granites were mostly studied by mineral separation that provides the largest monazite grains likely to contain useful internal information on the

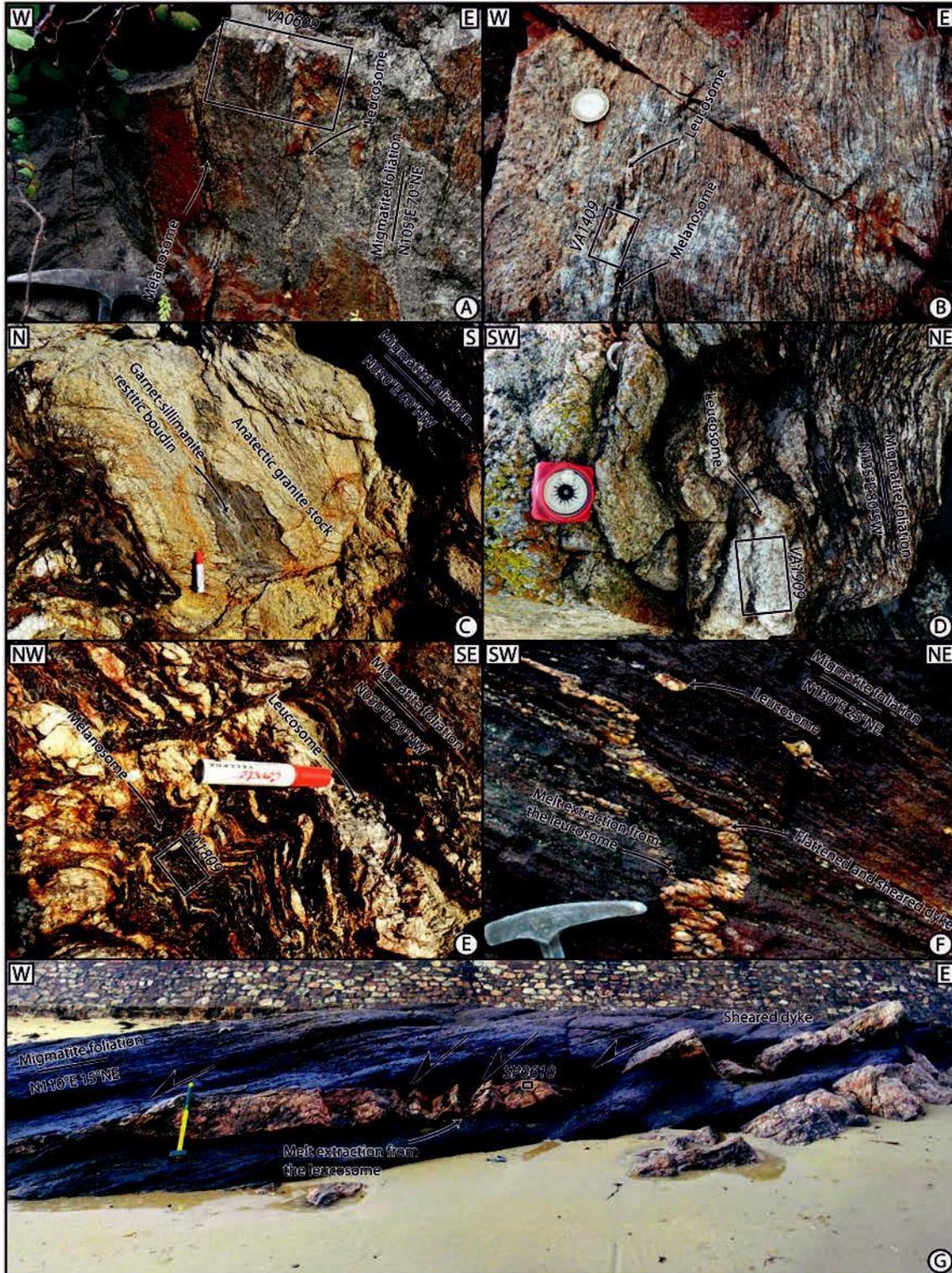


FIG. 4. – Representative pictures of the outcrop conditions of the migmatites and the granites in the other sampling areas. (a) Interlayering of paleosome and finely alternating leucosome and melanosome in the Hennebont-Quimperlé area. (b) Leucosome-rich migmatites exposed in the Vilaine Estuary area. Note that the paleosome is still locally preserved. (c, d and e) Representative pictures of the migmatites exposed on the coastal outcrop on the Saint Nazaire area. Picture (c) and (d) present different sizes of melt accumulation as foliation-concordant large leucosome (c) or as genuine small-scale anatectic granite stock (d). (f and g) General aspects of the migmatites and the granite dyke array in the Sable d’Olonne area. Note that leucosomes appear sometimes connected to the dykes indicating that melt extraction is synkinematic.

growth and dissolution history of the grains [i.e. Kelsey *et al.*, 2008]. Rocks were first finely crushed using disk grinders. Obtained powders were sieved and only the fraction size below 0.5 mm was retained. Then, mineral density separation using heavy liquids, followed by magnetic separation using a Frantz separator were applied to sort the monazite from other accessory heavy minerals. A final optical control based on the morphological characteristics of phosphates was made. The separated monazite grains were hand-picked and mounted in epoxy as plots, with further fine polishing. Mineral separation has been made at BRGM (French Geological Survey). Reader unfamiliar with the detailed procedure is referred to recent publications [Be Mezeme, 2005; Turrillot *et al.*, 2009].

The same analytical procedure was applied to all samples. Monazite grains have been observed under JSM-6400 JEOL Scanning Electron Microscope (SEM) hosted at Institut des Sciences de la Terre d'Orléans (ISTO). Thin sections and polished plots were first coated by a thin carbon layer. Acceleration voltage and beam-current intensity were 20 kV and 10 nA, respectively. For thin-sections, SEM observations provided for the distinction of monazite from other accessory minerals, using Energy-Dispersive X-ray spectrometer (EDS) and IdFix Software package for data processing. Then, back-scattered electrons (BSE) imaging mode was used for evaluating intra-grain composition variations within monazite, whereas the textures were examined with secondary electrons (SE) imaging mode. Only BSE images are here presented. Monazite grains were further analysed with the Cameca SX50 electron probe micro-analyser (EPMA) cooperated by BRGM and ISTO, equipped with four Wavelength Dispersion Spectrometers (WDS). For this study, we have followed the procedure described by Cocherie *et al.* [1998]. The accelerating-voltage and beam-current intensity conditions were routinely fixed to 20 kV and 100 nA providing absolute error of 150 ppm for U, Th and Pb [Cocherie and Albarède, 2001]. Beam current intensity was increased to 200 nA, for samples SP0710 and SP07bis10, providing lower absolute errors for U, Th and Pb of 105 ppm, 130 ppm and 110 ppm, respectively. As the respective amount of U and Th may significantly vary in natural monazite [Linthout, 2007], relative errors for U, Th and Pb can be very low especially when large amount of U and Th are incorporated in the monazite structure, and this may hide the errors due to instrumentation and background measurements. Therefore a minimum value of 2% is assumed for the relative errors. Natural and synthetic silicates, phosphates, oxides, and sulfides were used as standards. Total duration of background and peak counts is 240 s for Pb, 200 s for U, and 40 s for other elements (Ca, Ce, Gd, La, Nd, P, Pr, Si, Sm, Th, Y). The EPMA resolution of 2 µm, combined with the high-resolution SEM images allows for avoiding the altered domains of the grains with possible significant Pb loss. Structural formulae of monazite (A₄B₄O₁₆) were also calculated to reveal the composition variations due to chemical substitutions between and within samples. Raw or treated data and chemical tables that cannot be presented here for space limitations are available upon request from the first author. Only the substitution between monazite and cheralite (Ln³⁺ + Si⁴⁺ ↔ Ca²⁺ + P⁵⁺) is presented here.

Data sorting was further performed using the "EPMA dating" Excel® package [Pommier *et al.*, 2002], to calculate

individual U-Th-Pb age. Several filters (15% < Ce < 35%; 9% < P < 17%; Pb > 0,015%; 97% < Total 103%) were then applied to reject anomalous punctual results. Using ISOPLOT Excel(r) package [Ludwig, 2003], the mean age of each data population was calculated according to two methods:

- weighted average (mean) age calculation based on the regression line that intercepts the individual ages, which are weighted by the inverse of their related error. The aberrant values are thus easily rejected and the scattering of individual ages with respect to the average age can be evaluated by calculating the Mean Square Weighted Deviation (MSWD);

- isochron U-Th/Pb age following the procedure of Cocherie and Albarède [2001]. As monazite may present a strong variability of the Th/U ratio, Th/Pb versus U/Pb plots can be used to represent the individual data and their respective error (represented by ellipse of confidence). The isochron U-Th/Pb age of the sample was computed from different graphical parameters such as the slope with error of the best-fit line, the value of the Th/Pb and U/Pb intercept ages, and the coordinates of the centroid of the best-fit line.

RESULTS

All U-Th-Pb age-results calculated either as isochron U-Th/Pb age or as weighted average age, are presented in figures 5 to 10. A representative BSE image of the internal texture of monazite is given for each sample. For the sake of clarity, results are displayed according to sampling zones discriminated in figure 1.

The Golfe du Morbihan area

Migmatite samples

Sample VA1009, collected in the Ria d'Étel (western part of the Golfe du Morbihan) contains monazites of typical size below 50 µm. Monazite occurs as numerous sub-euhedral grains (fig. 5a) at the boundary between quartz and feldspar and frequently presents patchy zoning due to sectorial variations of U and Th contents. Individual data are scattered along a line in the Th/Pb versus U/Pb plot and a robust best-fit line can be drawn. The theoretical isochron is not strictly parallel to the best-fit line but both are similar within error and an U-Th/Pb isochron age at 325 ± 6 Ma has been calculated (fig. 5a).

Sample 201B contains sub-rounded monazite grains of small size (below 50 µm) set at the boundary between quartz and feldspar or as inclusions in biotite. No compositional zoning has been observed by SEM imaging and the narrow range of Th/Pb and U/Pb ratios do not provide a well-constrained best-fit line. Nevertheless, the best-fit line is similar within error to the theoretical isochron and an isochron U-Th/Pb age at 326 ± 5 Ma was calculated for this migmatite leucosome (fig. 5b).

Samples 220A, 220B and 220H, collected from the same outcrop include monazite of size ranging from 40 µm to 100 µm. In the leucosome (samples 220A and 220B), monazite occurs as sub-euhedral grains at the boundary between quartz and feldspar and rarely as inclusion within biotite. Conversely, in the melanosome (sample 220H),

monazite occurs as slightly corroded euhedral grains enclosed in biotite or garnet. In these samples, no textural and chemical zoning has been observed or measured. For sample 220A, the best-fit line presents a large error envelop due to close values of the Th/Pb and U/Pb ratios. But the best-fit line is similar within error to the theoretical isochron, providing an isochron U-Th/Pb age at 323 ± 6 Ma (fig. 5c). For

sample 220B, the wide range of Th/Pb and U/Pb ratios gives a good constraint for the best-fit line, which coincides within error with the theoretical isochron to yield an isochron U-Th/Pb age at 320 ± 5 Ma for this leucosome (fig. 5d). For sample 220H, the values of the Th/Pb and U/Pb ratios slightly vary and the best-fit line displays a large error envelope that incorporates the theoretical

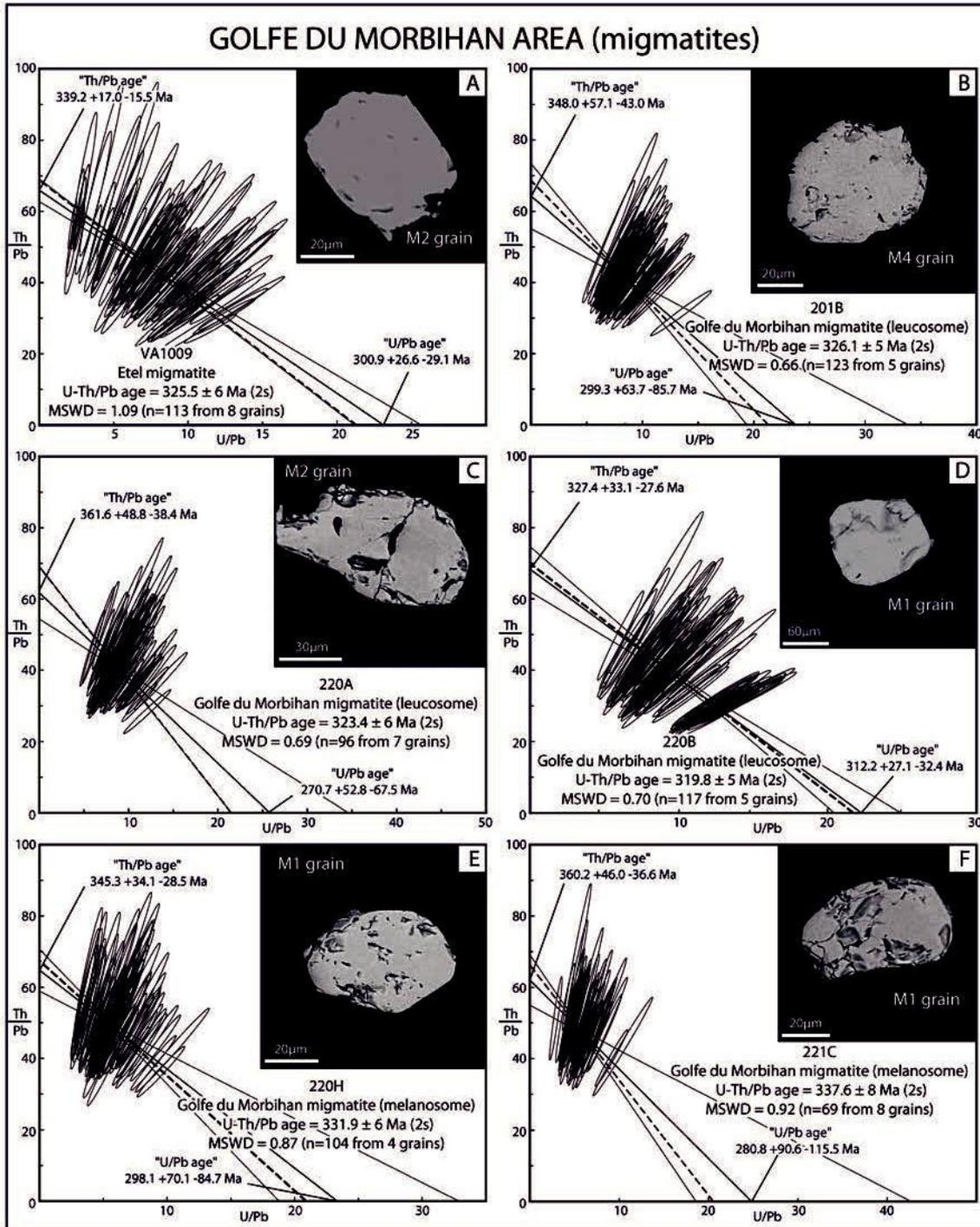


FIG. 5. – Geochronological results from migmatite samples collected in the Golfe du Morbihan area. For Th/Pb vs. U/Pb isochron diagrams, all standard deviations are quoted at 95% confidence level and error ellipses are plotted at 2. Thick (thin) solid line corresponds to best-fit line (envelop error); dotted line is the theoretical isochron for the calculated isochron age and should be comprised in the error envelope. SEM images were obtained using back-scattered electrons (BSE) imaging mode. Chemical transect presents the spatial variations of Ca + P and LnL + Si, according to monazite-cherilite substitution. To illustrate internal composition zoning, average UO₂ and ThO₂ contents are also presented for each grain domain. Mineral abbreviations are after Whitney and Evans [2010]. T refers to as transects across monazite grains.

isochron. The isochron U-Th/Pb age at 332 ± 6 Ma has been calculated for the melanosome (fig. 5e).

Sample 221C (melanosome) contains small monazite grains ranging from 20 μm to 50 μm . The grains are sub-rounded and occur as inclusions in the biotite and in the garnet. SEM imaging does not reveal significant compositional zoning. The narrow range of Th/Pb and U/Pb ratios do not provide a well-constrained best-fit line, but the error envelope includes the theoretical isochron. The isochron U-Th/Pb age at 338 ± 8 Ma has been calculated (fig. 5f).

Anatectic granite samples (Sainte-Anne d'Auray massif)

Sample 207 contains very large sub-euhedral crystals of monazite (up to 200 μm) set at the boundary between quartz and feldspar or enclosed in biotite. Due to variable U and Th contents, a slight compositional zoning was observed (fig. 6a) and a darker core may be distinguished from a lighter rim. This mantle-core texture was additionally re-vealed by high-Ca and P content in the rim and high-REE, Y and Si content in the core (fig. 6b). However, for the analyses from the grain core, the range of Th/Pb and U/Pb ratios is very narrow; this leads to extremely large error envelope for the best fit line and high value of MSDW above validity criteria [Wendt and Carl, 1991]. The Th/Pb versus U/Pb plots cannot be drawn and no isochron U-Th/Pb age has been calculated for this sample. Weighted average (or Mean) ages have been determined instead, yielding ages of 318 ± 3 Ma and 327 ± 9 Ma for the rim and the core, respectively (figs. 6c and 6d).

Sample 211 also presents very large monazite grains (up to 200 μm), which are associated with zircon and apatite. Grains are euhedral (fig. 6e) or rounded and they do not present any internal variation of composition. The Th/Pb and U/Pb ratios strongly vary and give a robust constraint to the best-fit line that is parallel to the theoretical isochron. An isochron U-Th/Pb age at 317 ± 6 Ma was thus calculated (fig. 6e).

Sample 227 contains large sub-euhedral crystals of monazite (ca. 100 μm) set at the boundary between quartz, feldspar and biotite or enclosed in biotite. SEM imaging does not reveal any particular internal texture of the grains, as they present homogeneous content in U and Th (fig. 6f). However, based on other elements incorporated in the monazite composition, a rim rich in Ca and P can be distinguished from a core rich in REE, Y and Si (fig. 6g). The individual analyses that correspond to the rim present weak variations of the Th/Pb and U/Pb ratios, inducing a significant error for the best-fit line. This latter one is similar within error to the theoretical isochron, though, yielding an isochron U-Th/Pb age at 317 ± 3 Ma (fig. 6h). Conversely, individual data from the core well constrain the best-fit line that parallels the theoretical isochron and an isochron U-Th/Pb age at 335 ± 6 Ma has been calculated (fig. 6i).

Carnac massif samples

Sample CAP01 contains monazite, whose size ranges between 50 μm and 200 μm . The grains are often euhedral and SEM observation has revealed various types of internal zoning, such as patchy zoning, mantle-core zoning or combination of both (fig. 7a). Zoning is caused by significant variations of the U and Th contents in the monazite.

Variations of the other elements are also observed and two populations, namely a core and a rim can be distinguished. The two calculated isochron U-Th/Pb ages are similar within error (320 ± 4 Ma for the rim and 323 ± 6 Ma for the core) and an average isochron U-Th/Pb age has been computed. The Th/Pb and U/Pb ratios are well distributed in the diagram providing very good constraints for the best-fit line that parallels the theoretical isochron. Hence, an isochron U-Th/Pb age at 321 ± 3 Ma was calculated (fig. 7a).

Sample CAP02 presents very large euhedral monazite grains (up to 300 μm) showing growth and patchy zoning. This is due to differences of the U and Th contents, and variations of other elements are also observed. Nevertheless, it is difficult to discriminate two populations, as observed growth zoning reveals successive chemical stages during crystallization. Despite significant variations of the U and Th contents, the range of the Th/Pb and U/Pb ratios is narrow and individual data are grouped, providing poor constraints for the best-fit line. This latter one is similar to the theoretical isochron, though, yielding an isochron U-Th/Pb age at 319 ± 5 Ma (fig. 7b).

In sample CAP03, small to medium-sized (below 100 μm) euhedral grains of monazite do not display evidence of zoning. The large variations of the Th/Pb and U/Pb ratios give robust constraints for the best-fit line that parallels the theoretical isochron. An isochron U-Th/Pb age has been calculated at 323 ± 5 Ma (fig. 7c).

Ploemeur-Quiberon leucogranite samples

Sample MR100 from the Ploemeur massif contains medium-sized (80 μm to 150 μm), generally sub-euhedral monazite grains that do not present evidence of zoning. The range of the Th/Pb and U/Pb ratios is wide and well constrains the best-fit line, which parallels the theoretical isochron. An isochron U-Th/Pb age at 335 ± 6 Ma has been calculated for sample MR100 (fig. 7d). Sample MR101, presents medium size (80 μm to 150 μm) sub-rounded monazite grains. SEM observations did not reveal internal zoning (fig. 7e). Although there are no significant variations of the U and Th contents, the composition of the monazite is heterogeneous. A horseshoe-shaped core with low Ca and P contents and high REE, Y and Si contents can be distinguished from a rim with high Ca and P contents and low REE, Y and Si contents (fig. 7f). Hence two diachronous populations have been distinguished. For the rim, the distribution of the Th/Pb and U/Pb ratios provides a good constraint for the best-fit line and its error envelope that encompasses the theoretical isochron. Analyses from the rim yielded an isochron U-Th/Pb age at 320 ± 8 Ma (fig. 7g). For the core, although the Th/Pb and U/Pb ratios are distributed in two separate groups, the best-fit line is similar within error to the theoretical isochron, yielding an isochron U-Th/Pb age at 333 ± 8 Ma (fig. 7h). Sample MR102 presents large, up to 200 μm euhedral monazite grains. They do not present significant element variations within grains, and no internal texture has been observed under SEM. Despite a distribution of the Th/Pb and U/Pb ratios into two separate groups, the best-fit line presents a narrow error envelope and parallels the theoretical isochron. An isochron U-Th/Pb age at 316 ± 4 Ma has been calculated for sample MR102 (fig. 7i).

Sample BA122b from the Quiberon massif contains medium-sized (30 μm to 80 μm) sub-euhedral monazite grains, set between quartz and feldspar. The grains do not display any textural or compositional zoning, but the range of the Th/Pb and U/Pb ratios is wide, constraining the position of

the best-fit line accurately. This latter one is parallel to the theoretical isochron, and an isochron U-Th/Pb age at 320 ± 3 Ma has been calculated for sample BA122b (fig. 8a). Sample QUP01 generally presents large (up to 300 μm) euhedral monazite grains. SEM observations have revealed

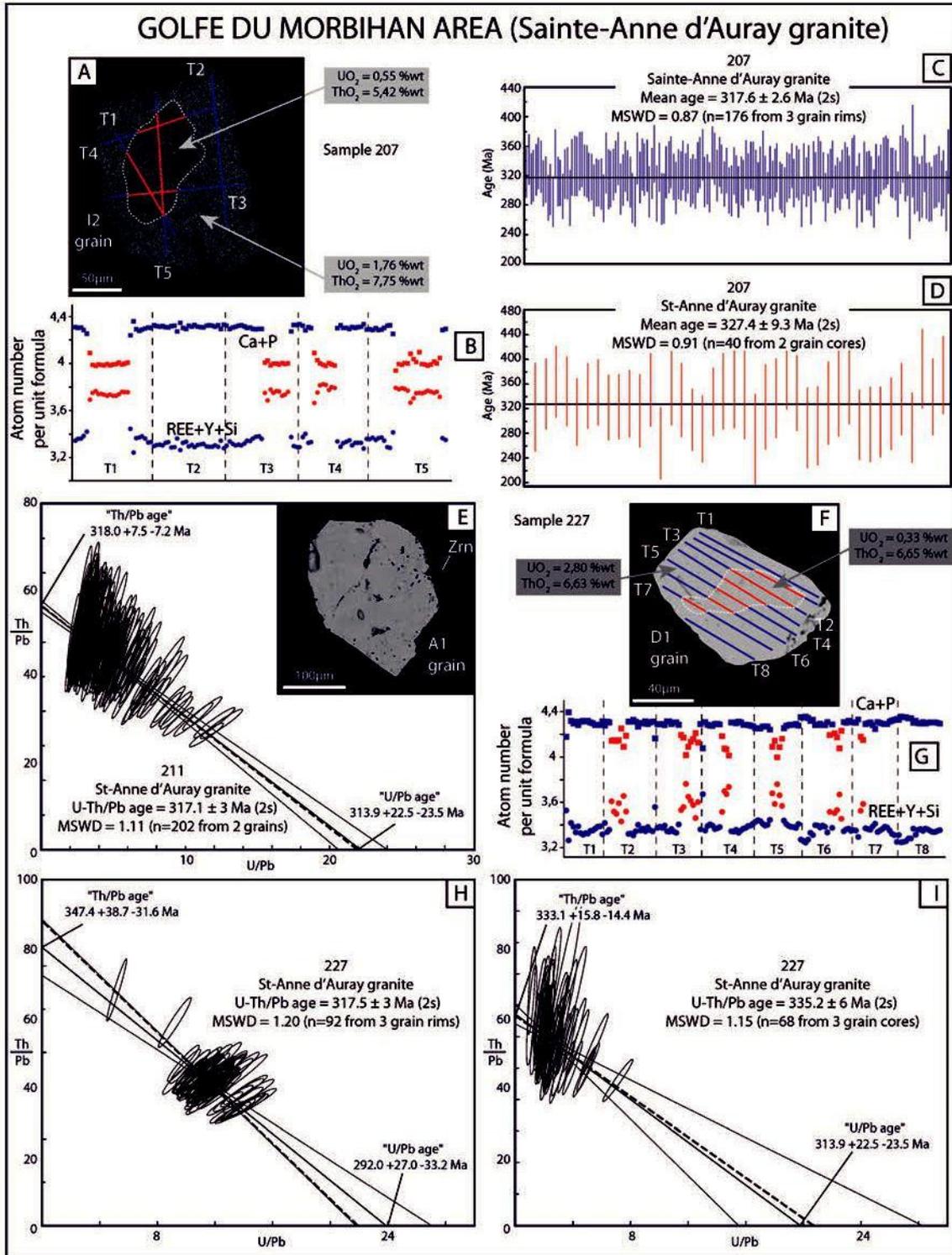


FIG. 6. – Geochronological results from Sainte-Anne d'Auray granite samples collected in the Golfe du Morbihan area. Details may be found in the caption of figure 5. For sample 207, weighted average ages are presented for the rim (C) and the core (D), as the weak chemical variations preclude isochron age calculation. Thick line corresponds to the mean age. Error bars are plotted at 2s. Mineral abbreviations are after Whitney and Evans [2010].

mantle-core texture in accordance with the significant differences of Th content between the core and the rim (fig. 8b). In addition, the core presents low Ca and P contents and high REE, Y and Si contents, whereas the rim is relatively rich in Ca and P and relatively poor in REE, Y and Si (fig. 8c). Hence, two Th/Pb versus U/Pb plots have been produced. For the rim, the Th/Pb and U/Pb ratios are grouped, providing a significant error for the best-fit line. This latter one is parallel to the theoretical isochron, though, and we have calculated an isochron U-Th/Pb age at 323 ± 6 Ma (fig. 8d). Conversely, individual Th/Pb and U/Pb ratios from the core display a linear dispersion, which provides robust constraints for the best-fit line. The best-fit line parallels the theoretical isochron, yielding an isochron U-Th/Pb age at 330 ± 5 Ma (fig. 8e). Sample QUP03 also contains large euhedral monazite grains (up to $300 \mu\text{m}$), but no textural or composition zoning has been found. Although the Th/Pb and U/Pb ratios are divided into two different groups, the best-fit line presents a relatively narrow error envelope and parallels the theoretical isochron. Hence an isochron U-Th/Pb age at 335 ± 6 Ma has been calculated (fig. 8f). Sample QUP04 contains large monazite grains (up to $300 \mu\text{m}$) that are often euhedral and present mantle-core texture superimposed to growth zoning (fig. 8g). The rim is slightly lighter than the core because of higher U content. In addition, other elements incorporated in the monazite composition argue for a two-stage grain crystallization, with a rim rich in Ca and P and a core rich in REE, Y and Si (fig. 8h). For the rim, the Th/Pb and U/Pb ratios are separated into two different groups, but the best-fit line is well constrained. This latter one is similar within error to the theoretical isochron and we have calculated an isochron U-Th/Pb age at 320 ± 4 Ma (fig. 8i). Conversely, the range of the Th/Pb and U/Pb ratios is wide and a larger error for the best-fit line is observed. The best-fit line parallels the theoretical isochron, though. Hence, an isochron U-Th/Pb age at 328 ± 6 Ma has been calculated (fig. 9j).

The Hennebont-Quimperlé area

Sample VA0609 contains monazite grains ranging from $60 \mu\text{m}$ to $100 \mu\text{m}$. The grains are included in biotite flakes or set at the boundary between quartz and feldspar. They display an apparent rim (darker in SEM images), which may correspond to post-crystallization alteration due to iron (hydro-) oxides mixtures (fig. 9a). Individual microprobe analyses from these rim domains have been excluded from the isochron U-Th/Pb age calculation. The Th/Pb and U/Pb ratios present important variations giving a good constraint for the best-fit line that is very close to the theoretical isochron. An isochron U-Th/Pb age has been calculated at 319 ± 5 Ma (fig. 9a). Conversely, sample VA0709 contains small monazite grains (i.e. less than $60 \mu\text{m}$) that were repeatedly observed as biotite inclusions. SEM imaging does not reveal internal zoning, as U and Th contents are relatively constant (fig. 9b). But variations of other elements are significant and a Ca and P-poor core was distinguished from a Y and Si-rich rim (fig. 9c). The separation of the two inferred populations provides two different isochron U-Th/Pb ages, which meets the validity criteria despite a narrow range for Th/Pb and U/Pb ratios. For the rim, an isochron U-Th/Pb age has been calculated at 318 ± 4 Ma (fig. 9d), whereas the isochron U-Th/Pb age of the core is 328 ± 6 Ma (fig. 9e).

Migmatite samples from the Vilaine Estuary area

Sample VA1409 presents monazite, whose size ranges between $50 \mu\text{m}$ and $100 \mu\text{m}$. The grains are often interstitial and set between chloritised biotite, quartz and feldspar. Growth, patch and mantle-core zoning (fig. 9f) have been observed, but no significant variations of the element contents were observed. Conversely, the range of the Th/Pb and U/Pb ratios is wide and the well-constrained best-fit line parallels the theoretical isochron, yielding an isochron U-Th/Pb age at 318 ± 2 Ma (fig. 9f). Sample VA1309 contains large monazite crystals (up to $150 \mu\text{m}$). Monazite occurs as sub-rounded grains, without prominent internal zoning and mostly included in biotite, but also as set between quartz and feldspar. The relatively large range of the Th/Pb and U/Pb ratios provides a good constraint for the best-fit line that is parallel to the theoretical isochron. An isochron U-Th/Pb age has been calculated at 326 ± 5 Ma (fig. 9g).

Migmatite samples from the St-Nazaire area

Sample GnIP03 contains various-sized sub-euhedral monazite grains ($30 \mu\text{m}$ to $200 \mu\text{m}$) set between quartz and feldspar or included in biotite. It is worth to notice that each grain presents evidence of incipient alteration (corrosion and dissolution) along crystal cleavage, possibly due to late weathering. Monazite is associated with apatite and zircon (fig. 10a). In Th/Pb versus U/Pb plot, individual data are well scattered and constrain a best-fit line parallel to the theoretical isochron. We have calculated an isochron U-Th/Pb age at 328 ± 4 Ma (fig. 10a). Sample VA1909 presents two morphological types of monazite. The first type is characterised by very large crystals of sub-rounded monazite (up to $300 \mu\text{m}$) sometimes with dissolution evidence at the quartz, feldspar and biotite grain boundaries (fig. 10b). Conversely, the second type is characterised by rare, medium-sized euhedral monazite ($80 \mu\text{m}$) included in biotite is rarely observed (fig. 10c). Beyond morphological differences, each population presents distinct composition.

The large monazite grains present significant internal element variations within a grain, with low contents in Ca and P and high contents in LREE, Y and Si. Conversely, the small monazite grains present regular element compositions characterized by high Ca and P contents and low LREE, Y and Si contents (fig. 10d). The population composed by the large monazite present weak variations of the Th/Pb and U/Pb ratios, but the number of individual data (122) well constrains the position of the best-fit line and its error. Hence, the best-fit line is similar within error to the theoretical isochron, yielding an isochron U-Th/Pb age at 331 ± 3 Ma (fig. 10e). For the population of small monazite grains, Th/Pb and U/Pb ratios slightly vary and the best-fit line presents a large envelope error, which encompasses the theoretical isochron. An isochron U-Th/Pb age has been calculated at 319 ± 4 Ma (fig. 10f). Sample VA1809 is a restite that contains medium-sized sub-euhedral monazite grains (ca. $100 \mu\text{m}$) set at the boundary between quartz and feldspar or included in quartz, feldspar and chloritised biotite. The range of the Th/Pb and U/Pb ratios is wide and the best-fit line is well constrained. This latter one is similar within error to the theoretical isochron, providing an isochron U-Th/Pb age at 352 ± 10 Ma (fig. 10g).

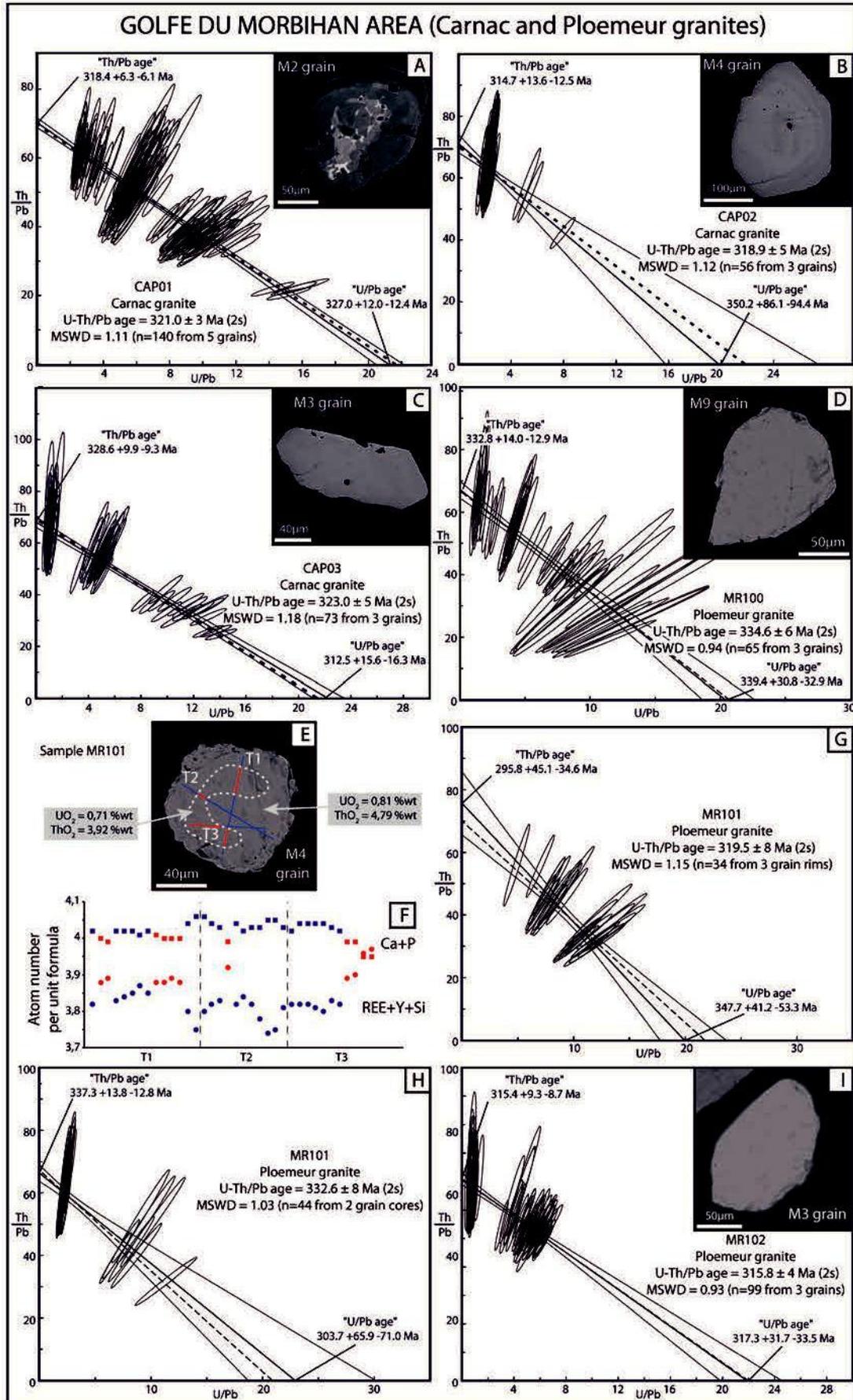


FIG. 7. – U-Th/Pb geochronological results from Carnac (A to C) and Ploemeur (D to I) granites collected in the Golfe du Morbihan area. Details are given in the caption of figure 5.

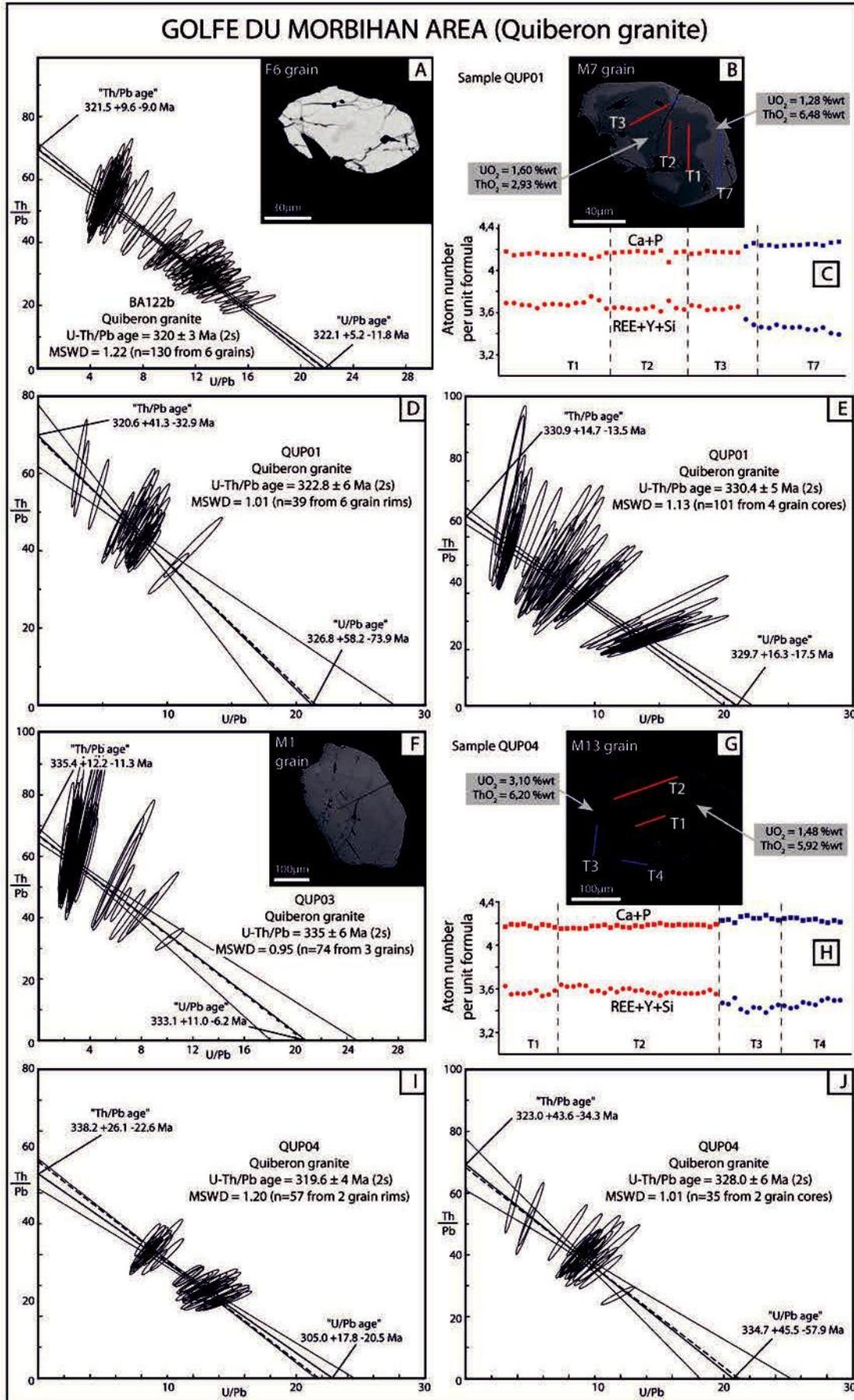


Fig. 8. – U-Th/Pb geochronological results from Quiberon granite samples collected in the Golfe du Morbihan area. Details are given in the caption of figure 5.

The Sables d'Olonne area

Samples SP0710 and SP07bis10 contain euhedral to sub-rounded monazite grains, with variable size (up to 200 μm in sample SP0710 and less than 60 μm in sample SP07bis10). In sample SP0710, the grains are set between quartz, feldspar and biotite, whereas in sample SP07bis10, they may also be located as inclusions in biotite. Although both samples present different ranges and distributions of the Th/Pb and U/Pb ratios, the best-fit lines are robust and parallel the respective theoretical isochron. Hence we have calculated an isochron U-Th/Pb age at 320 ± 5 Ma for sample SP0710 (fig. 10h) and an isochron U-Th/Pb age at 323 ± 6 Ma for sample SP07bis10 (fig. 10i). The spatially (and probably genetically) associated granite sample (sample SP0610) contains small sub-rounded monazite grains (less than 50 μm), which often occur at the boundary between quartz, feldspar and muscovite, or rarely trapped as inclusion in the muscovite. The Th/Pb and U/Pb ratios are well distributed and provide excellent constraints for the best-fit line. This latter one is similar within error to the theoretical isochron, yielding an isochron U-Th/Pb age at 328 ± 4 Ma for sample SP0610 (fig. 10j).

INTERPRETATIONS : RELIABILITY AND SIGNIFICANCE OF THE AGE DATA

EPMA U-Th/Pb chemical dating of monazite has become an increasingly popular, non-destructive and *in situ* method providing high-resolution texturally-controlled ages. This method was successfully applied to various types of magmatic and metamorphic rocks, including migmatites [e.g. Cocherie *et al.*, 2005; Be Mezeme 2005; Be Mezeme *et al.*, 2006; Faure *et al.*, 2008a, 2010; Gagné *et al.*, 2009; Majka *et al.*, 2012]. While new geochronological constraints were accumulated and data reduction methods were refined, the significance of the ages was questioned, as conditions of monazite crystallization remain not clearly understood, especially for metamorphic monazite [e.g. Vavra *et al.*, 1996; Forster *et al.*, 2000, 2002; Forster and Parrish, 2003; Pyle and Spear, 2003; Goncalves *et al.*, 2004; Pyle, 2006; Kelsey *et al.*, 2008; Janots *et al.*, 2009]. In the following, the reliability of the calculated ages and their significance are discussed in terms of geological events, in the light of the present knowledge of monazite chemical system.

Reliability of ages

Reliability of the U-Th/Pb chemical ages is based on two main assumptions related to the intrinsic mineral structure of monazite [Suzuki and Adachi, 1991] including that 1) common Pb is negligible compared to the amount of thorogenic and uraniumogenic Pb [Parrish, 1990; Cocherie *et al.*, 1998] and 2) no radiogenic Pb loss has occurred since the closure of the system [e.g. Smith and Giletti, 1997; Cherniak *et al.*, 2004; Seydoux-Guillaume *et al.*, 2002; Gardés *et al.*, 2006]. These hypotheses were verified on natural samples by numerous studies, where results obtained by the chemical method were compared to isotopic techniques like TIMS, SIMS and LA-ICPMS [i.e. Cocherie *et al.*, 1998; 2005; Montel *et al.*, 2002].

The isochron method [Cocherie and Albarède, 2001; Cocherie *et al.*, 2005] was used to represent the individual

ages and to compute an isochron age for each sample. This method requires a variation of the Th/U ratio to provide a large data span in the Th/Pb vs. U/Pb diagram, constraining the error envelope of the best-fit line. According to [Cocherie *et al.*, 2005] the computed regression must satisfy 3 criteria: 1) the theoretical isochron age must be comprised in the error envelope of the best-fit line and must crosscut each individual data within the ellipse error; 2) the calculated MSWD must be lower than the statistical parameter [Wendt and Carl, 1991] and 3) the two intercept ages (i.e. Th/Pb and U/Pb ages) must be similar within error. These conditions were satisfied for all samples except for sample 207 that presents very low variations of the Th/U ratio. In this case, the weighted mean age method was used as an alternative age calculation procedure. Similarly to the isochron method, the calculated age is statistically acceptable, since, for a given population, the calculated MSWD value falls below the deviation parameter (table I) [Wendt and Carl, 1991]. This method assumes that age errors are symmetric, producing ages, which are usually younger (ca. 3 to 5 Ma) than the age obtained from other methods [Be Mezeme, 2005].

Classification of the monazite grains from the dataset

Validation of all criteria of the U-Th/Pb chemical dating method allows interpreting each calculated age (weighted mean and isochron age) as related to a single geological event. However, monazite grains present morphological and chemical differences and provide a variety of ages for monazite (re-) crystallisation (table I). The occurrence of growth zoning among numerous analyzed grains suggests a more intricate crystallisation evolution for monazite crystals, which cannot be portrayed with the time resolution of present geochronological tools. Depending on crystal habits and composition together with phase-relationships, five types of monazite (A1 and A2 types for migmatite, B1, B2 and B3 types for granite) have been recognized as following (fig. 2).

– **A1-type** monazite grains of various sizes that occur either included in biotite flakes or set at the boundary between quartz and feldspar in the leucosome (migmatite samples VA0609, VA1009, 201B, 220A, 220B, SP0710, SP07bis10, VA1409, VA1309 and GnIP03). Although growth, patchy and mantle-core zoning have been sometimes observed among the samples, only one population has been distinguished and a single age for each sample has been calculated. The ages range from 326 ± 5 Ma to 318 ± 6 Ma (table I). For sample GnIP03 (migmatitic gneiss), we cannot ascertain the significance of the age (328 ± 4 Ma), as both melanosome and leucosome parts were collected together. Despite a partial overlap of ages within error brackets, this rather old age could show mixing between two age populations of monazite, without significant composition variations.

– **A2-type** contains small to medium-sized monazites that occur either at grain boundary or enclosed in biotite and garnet in the restite (samples 220H, 221C and VA1809). No internal zoning has been revealed. The calculated ages range between 352 ± 10 and 332 ± 6 Ma and (table I) that is significantly older than the ages provided by the monazite of the leucosomes.

– Sample VA1909 presents two populations of monazite (table I). Euhedral small grains of medium-sized monazite sharing similarities with **A1-type** monazite are included in biotite, whereas large crystals of sub-rounded monazite similar to **A2-type** monazite are set at grain boundaries and present evidence of corrosion. Each population presents distinct contents in Ca, P, LREE, Y and Si. The age of the population marked by the small grains (319 ± 4 Ma) is consistent with

the age of other leucosome samples. The old age (331 ± 3 Ma) is very similar to the age of restite samples.

– Sub-euhedral to euhedral **B1-type** monazite up to $300 \mu\text{m}$ was identified in granite samples BA122b, MR102, CAP01, CAP02, CAP03, and 211. The grains that display growth zoning occur at the boundary between quartz, feldspar and muscovite. The calculated ages range from 316 ± 4 Ma to 323 ± 5 Ma (table I).

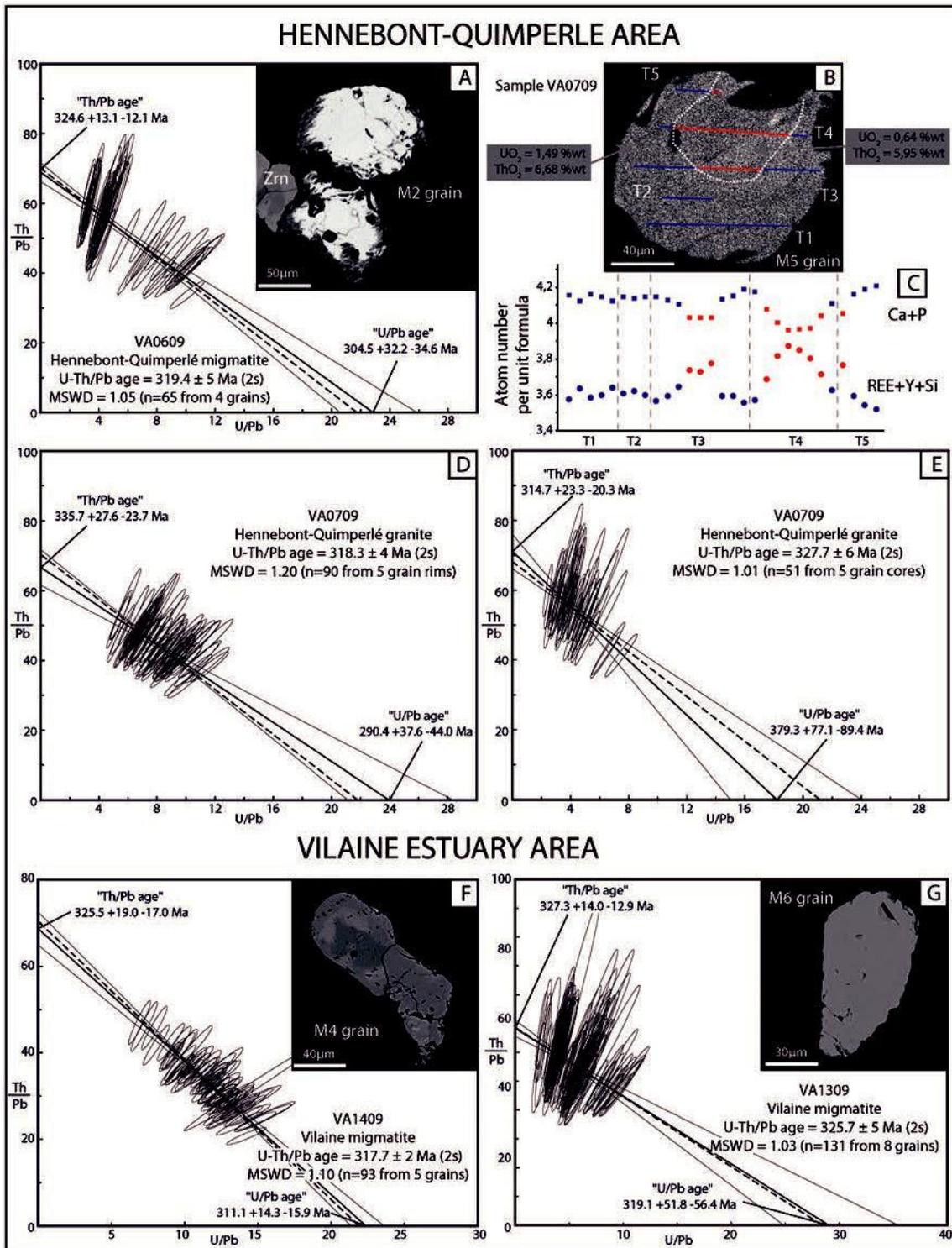


FIG. 9. – U-Th/Pb geochronological results from migmatite and granite samples collected in the Hennebont-Quimperlé (A to E) and the Vilaine Estuary areas. Details are given in the caption of figure 5.

– **B2-type** monazite, with a magmatic rim and an inherited core was identified in granite samples QUP01, QUP04, MR101, 207, 227 and VA0709. The grains are usually euhedral, medium to large in size. They occur as inclusions in biotite or at the grain boundary between quartz, feldspar and mica. Due to repeatedly observed low variations of U

and Th contents, mantle-core textures were not systematically observed on BSE images. However, obvious variations of the Ca, P, Si, REE and Y contents and significant differences in the degree of chemical substitution help to distinguish a core from a rim. The core presents a sub-rounded shape, showing embayment-like textures. It is always low in Ca and P and rich in REE, Y and Si. Conversely, the rim is marked by euhedral shape and internal growth zoning has been observed for some grains. Rim and core populations often present different ranges of value for the U/Th ratio. The age of the rim vary between 323 ± 6 Ma and 317 ± 3 Ma, whereas the age of the core is bracketed between 335 ± 6 Ma and 327 ± 9 Ma (table I).

– **B3-type** monazite was identified in granite samples QUP03, MR100, and SP0610. The grains are generally small (except for sample QUP03) and often occur included in biotite or rarely at grain boundary between quartz and feldspar. The grains do not present internal zoning, such as

growth zoning. The isochron U-Th/Pb ages are bracketed between 338 ± 9 Ma and 328 ± 4 Ma (table I).

Geological significance of the calculated ages

Monazite has been proposed to resist dissolution at temperatures up to 900°C [Braun *et al.*, 1998; Cherniak *et al.*, 2004], ensuring, at least partially its preservation in the studied rocks [Brown, 1983; Jones and Brown, 1989, 1990; Goujou, 1992; Brown and Dallmeyer, 1996; Marchildon and Brown, 2003; Johnson and Brown, 2004]. **A2-type** monazites from restite present ages ranging between 352 Ma and 332 Ma. This age group may represent inherited pre-partial-melting stages, either during the prograde evolution of the rocks, or during an early metamorphic event. Melt extraction and formation of the leucosomes may also provoke encapsulating of these old monazite grains. In addition, monazite solubility remains generally limited in peraluminous melts [Montel, 1986; 1987; 1993] allowing the persistence of round-shaped old monazite grains in the leucosome. In this case of partial dissolution, new grains of monazite may crystallise, while melt cools to form the leucosome [Hawkins and Bowring, 1999; Rubatto *et al.*, 2001; Kohn *et al.*, 2005]. The 326-318 Ma ages of **A1-type**

TABLE I. – Results of U-Th/Pb geochronology on monazite. MSWD is the calculated mean square weighted deviation; DP is theoretical deviation parameter [Wendt and Carl, 1991]. All ages are isochron ages, except for sample 207 (spotted by *) for which a weighted average age has been calculated.

Area	Type	Massif	X Long (°)	Y Long (°)	Sample	Age (Texture/Structure)	MSWD/DP
Golfe du Morbihan	Granite	Sainte-Aune d'Auray	-2.76687	47.69901	207	318 ± 3 Ma (Rim) 327 ± 9 Ma (Core)	0.87/1.21 0.91/1.45
			-2.64583	47.62890	211	317 ± 6 Ma	1.11/1.20
			-2.62172	47.63511	227	317 ± 3 Ma (Rim) 335 ± 6 Ma (Core)	1.20/1.30 1.15/1.35
		Carnac	-2.99442	47.64120	CAP01	321 ± 3 Ma	1.11/1.24
			-3.02187	47.59121	CAP02	319 ± 5 Ma	1.12/1.38
			-3.10194	47.66289	CAP03	323 ± 5 Ma	1.18/1.33
		Ploemeur	-3.47818	47.71515	MR100	335 ± 7 Ma	0.94/1.35
			-3.41500	47.70031	MR101	320 ± 8 Ma 333 ± 8 Ma	1.15/1.49 1.03/1.43
			-3.40333	47.73311	MR102	316 ± 4 Ma	0.93/1.29
		Quiberon	-3.12964	47.52340	BA122b	320 ± 3 Ma	1.20/1.25
			-3.13143	47.53810	QUP01	323 ± 6 Ma 330 ± 5 Ma	1.01/1.46 1.13/1.28
			-3.12867	47.52501	QUP03	335 ± 6 Ma	0.95/1.33
	-3.11828		47.49780	QUP04	320 ± 4 Ma 328 ± 6 Ma	1.20/1.38 1.01/1.49	
	Migmatite	-3.14948	47.71601	VA1009	325 ± 6 Ma (leucosome)	1.09/1.27	
		-2.91310	47.54286	201B	325 ± 6 Ma (leucosome)	0.66/1.26	
-2.86416		47.61073	220A	323 ± 6 Ma (leucosome)	0.69/1.29		
-2.86417		47.61074	220B	320 ± 5 Ma (leucosome)	0.70/1.26		
-2.86418		47.61075	220H	332 ± 6 Ma (restite)	0.87/1.28		
Hennebont-Quimperlé	Granite	-3.39141	47.83899	VA0709	318 ± 4 Ma (Rim) 328 ± 6 Ma (Core)	1.20/1.30 1.01/1.40	
		-3.39511	47.83585	VA0609	319 ± 5 Ma (leucosome)	1.05/1.35	
Vilaine River Mouth	Migmatite	-2.45181	47.54941	VA1409	318 ± 2 Ma (leucosome)	1.10/1.29	
		-2.48122	47.55001	VA1309	326 ± 5 Ma (leucosome)	1.03/1.25	
St-Nazaire Dome	Migmatite	-2.46636	47.27041	GnIP03	328 ± 4 Ma (leucosome)	1.20/1.37	
		-2.30674	47.23741	VA1909	331 ± 3 Ma (large grains) 319 ± 4 Ma (small grains)	1.02/1.26 0.91/1.42	

monazite, collected from leucosomes are interpreted as melt crystallisation ages. As monazite corrosion by the melt is limited to the grain boundary [Braun *et al.*, 1998], mixture of **A1-type** and **A2-type** monazites may occur in the leucosome as for sample VA1909.

The occurrence of monazite in granitic rocks depends mainly on the availability of REE and actinides content of the original magma [Förster, 1998]. Because of their composition, peraluminous granites usually host a significant fraction of monazite, which often crystallised at near-solidus conditions. **B1-type** monazite grains occur at grain boundary and repeatedly display growth zoning that reveals successive chemical stages during crystallisation, and interaction between monazite growth and the melt composition evolution. Their ages (323-316 Ma) can be therefore interpreted as the crystallization age. As the emplacement of granite pluton is characterized by a cooling period shorter than the error bars of the calculated age, it is generally assumed that the age of monazite crystallisation represents the age of granite intrusion [Parrish, 1990]. For several granite intrusions, **B2-type** monazite grains showing mantle-core texture and the obvious differences of chemical composition between the rim and the core support crystallisation during two separate stages. Age of the rim (323-317 Ma) is consistent with the age of **B1-type** monazite, whereas the ages of grain core are significantly older (335-327 Ma). While core ages may represent an inherited magmatic or metamorphic event, rim ages are interpreted as the late crystallisation episode, coeval with melt cooling and pluton emplacement. Pre-anatectic monazite may therefore be preserved and coexists with grains formed during late magmatic events. If the melt contains sufficient P and REE, a new generation of monazite can crystallise through nucleation on inherited grains. If the melt is poor in P and REE, no monazite will form during magma crystallisation and only inherited monazite such as **B3-type** monazite will be found. Their age (338 ± 9 Ma to 328 ± 4 Ma) is very similar to the core age of **B2-type** monazite grains; hence it may correspond to inherited, magmatic or metamorphic events.

In the lights of these results, it is proposed that sub-solidus ductile deformation and the possible associated fluid circulations did not significantly affect the U-Th/Pb system in monazite, as variously deformed samples of Ploemeur and Quiberon granites present similar ages. This observation is in agreement with geochronological data from other peraluminous plutons associated to ductile shear zone that provided same U-Th/Pb isochron ages on monazite for the undeformed granite and for the mylonite [Choulet *et al.*, 2012]. Conversely, the lack of age reset due to deformation does not comply with other U-Th/Pb isotopic ages from monazite from southern Brittany granites [Tartèse *et al.*, 2011b], whose younger ages can be explained rather by pervasive fluid circulations along permeable conduits than by methodological bias.

AGE OF PARTIAL-MELTING AND TECTONIC IMPLICATIONS

Among the main results of this study, textural/chemical analyses allow recognising five monazite types. However, conspicuous similarities between different types and results

of the U-Th/Pb datings permit to distinguish only two genuinely different generations of monazite in both migmatites and granite rocks: i) inherited grains occur in mesosomes and melanosomes in migmatite (**A2 type**), as inclusions in biotite (**B3 type**) or in the core of younger crystals in the granite rocks (**B2 type**) that yielded ca. 352-330 Ma ages and ii) newly formed grains are found in leucosomes (**A1 type**) and in (**B1 type**) crystals and rim around old grains (**B2 type**) crystals settled at ca. 325-315 Ma. Implications of these ages will be now discussed in the regional and the geodynamical context.

Age-constraints for the evolution of the Lower unit

New U-Th/Pb chemical age-results provide absolute time-constraints on the evolution of the Lower unit for the whole coastal South Armorican domain (fig. 11). In the Golfe du Morbihan area, ca. 326-318 Ma ages were obtained for A1 monazites from small-scale leucosomes within 4 migmatite samples (fig. 11), interpreted as monazite crystallisation ages during melt cooling. These results are consistent with the parallel dating at ca. 318 ± 3 Ma of B1 monazites from the closely associated large-scale anatectic granite massif (i.e. Sainte-Anne d'Auray). Besides, samples from 4 other migmatite sampling zones yielded crystallization ages for A1 monazites from leucosomes, clustering at ca. 326-318 Ma. Furthermore, these ages are in agreement with the scarce previous time constraints [e.g. Turrillot *et al.*, 2009] and support a strikingly synchronous partial-melting event over the whole coastal part of South Armorican domain. Interestingly, the solidus was intersected during burial along a typical ca. 30°C/Km Barrovian gradient (fig. 12) [Brown, 1983; Jones and Brown, 1989, 1990; Goujou, 1992; Brown and Dallmeyer, 1996; Marchildon and Brown, 2003; Johnson and Brown, 2004]. Effective melt generation was subsequently initiated at peak-P metamorphic conditions at ca. 8-9 kbar for 750-800°C at variance with the examples of the Canadian Cordillera and the Aegean Sea, where partial-melting occurs during decompression some 20 Ma after peak-P conditions [Vanderhaeghe *et al.*, 1999; Hinchey *et al.*, 2006; Duchêne *et al.*, 2006; Jolivet and Brun, 2010]. Ages of late granite intrusions, such as Carnac, Quiberon and Ploemeur granites also cluster at ca. 320 Ma, at the end of the partial-melting event at conditions of ca. 4-5 kbar for 700-750°C [Johnson and Brown, 2004]. Despite clear cross-cutting relationships at map-scale (fig. 1) and limited single-age error brackets the ability to geochronologically distinguish the different late intrusive granite complexes has not been possible.

Melanosome samples (i.e. A2 monazite type) mostly yielded a subordinate age cluster at ca. 352-332 Ma, which is similar to the inherited ages, repeatedly observed in the B2 monazites from intrusive granites. Apart the ca. 352 Ma age that may have recorded an earlier event, this older age group that significantly differs from the age of partial-melting is relatively new for the South-Armorican domain. Based on microtextural and compositional criteria, this age group is believed to reflect a regional metamorphic event, related to the crossing of monazite producing reactions (fig. 12). Despite a complex behaviour of monazite during the metamorphic evolution, peak-crystallisation conditions mainly occur during increasing P-T conditions from the upper-green schist facies and in the amphibolite facies through

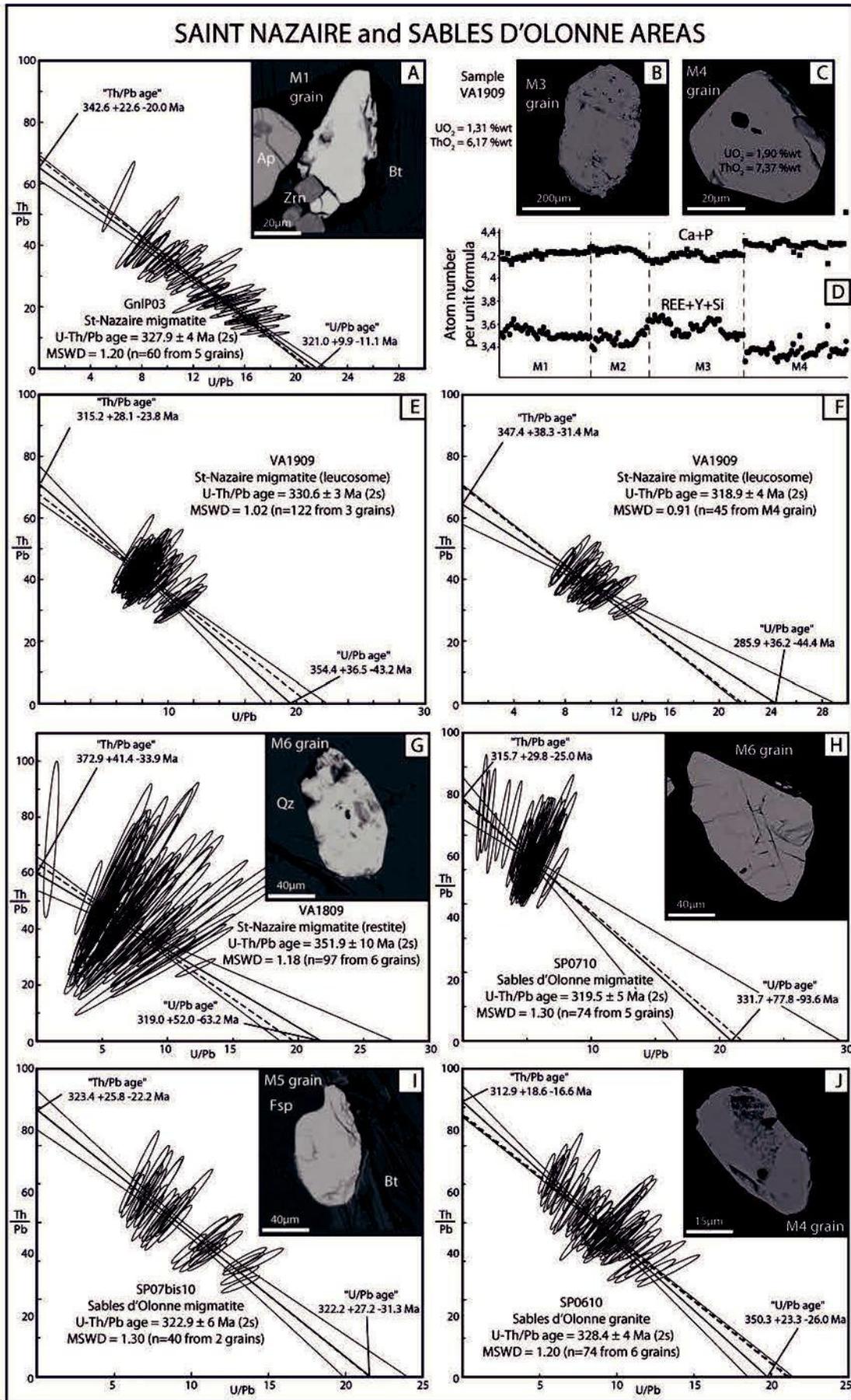


FIG. 10. – U-Th/Pb geochronological results from migmatite samples collected in the Saint-Nazaire (A to G) and in the Sables d'Olonne areas (H to J). Details are given in the caption of figure 5.

monazite-producing reactions in the range of ca. 400 to 650°C [e.g. Smith and Barreiro, 1990; Pyle *et al.*, 2001, Pyle, 2006; Forster *et al.*, 2002; Janots *et al.*, 2009; Spear, 2010]. This assumption is strengthened by a ca. 338 ± 7 Ma metamorphic age recently retrieved from a non-melted micaschist sample from the overlying Vilaine Estuary unit [Augier *et al.*, 2011]. The ca. 335-330 Ma ages therefore probably record monazite crystallisation ages related to the crossing of a monazite producing reaction during increasing P-T conditions in an overall nappe-stacking context [i.e. Goujou, 1992; Johnson and Brown, 2004; Gloaguen, 2006; Janots *et al.*, 2009; Spear, 2010].

The ca. 325-320 Ma turning-point in the evolution of the coastal South Armorican domain

Gravitational collapse corresponds to the decay of lateral contrast in the gravitational potential energy stored within orogenic belts by gravity-driven flow and ultimately triggers crustal thinning of the orogenic crust [e.g. Rey *et al.*, 2001; Vanderhaeghe and Teyssier, 2001]. Gravitational collapse appears controlled by three main mechanisms including 1) a decrease or even an inversion in the tectonic forces applied at the boundary of the orogenic system, 2) an increase of the potential energy caused by convective removal,

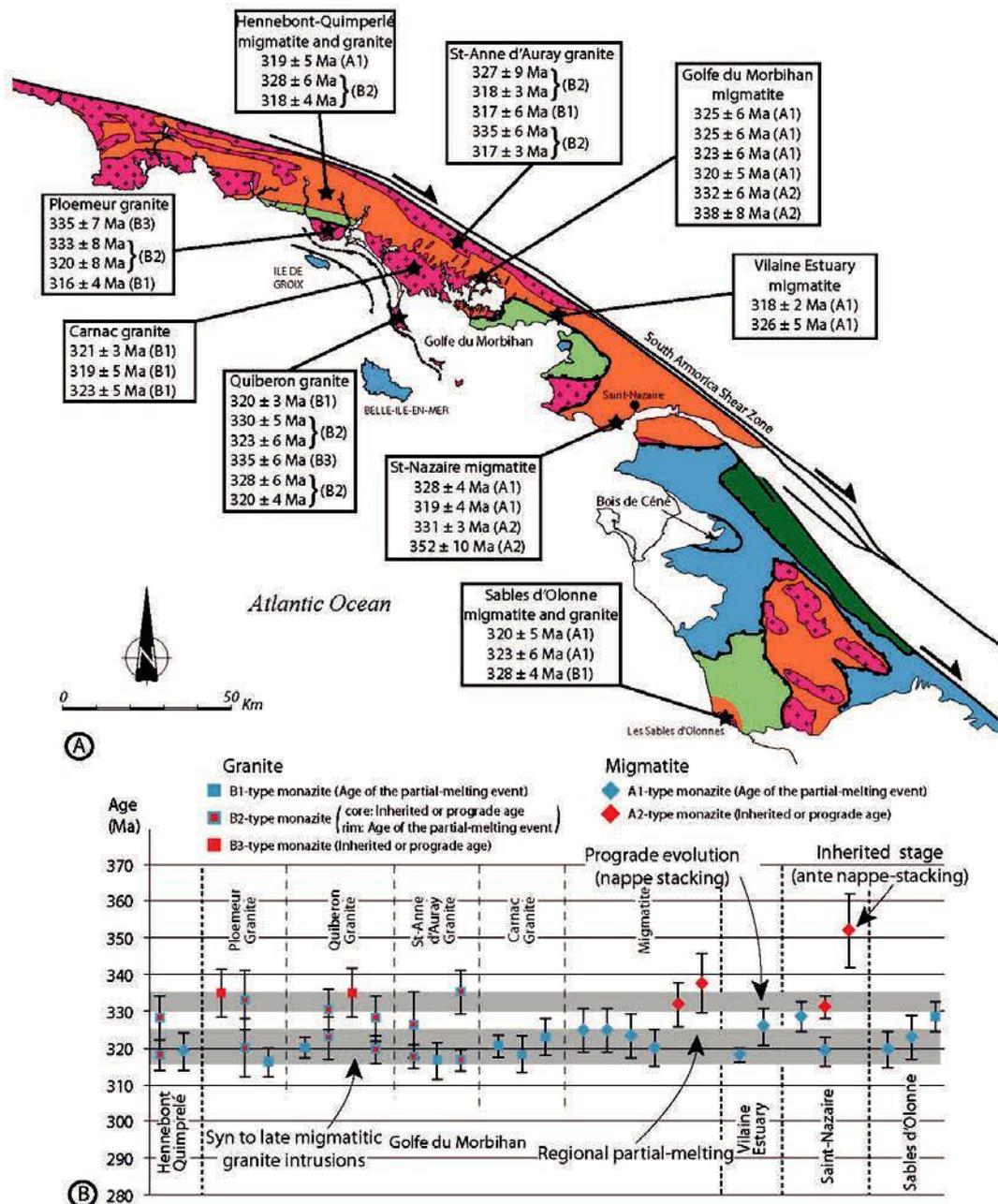


FIG. 11. – Synoptic chart of the U-Th/Pb geochronological results across the coastal South Armorican domain. Despite local partial age overlapping, 3 different stages can be distinguished. Old inherited ages remain rare (e.g. the 352 Ma age), whereas a pre-anatectic, prograde episode probably associated to nappe-stacking is recorded throughout the study area at ca. 335-330 Ma. Partial-melting that probably starts at ca. 325 Ma, and became widespread during the 325-320 Ma period, accompanied by the accumulation of large volume of anatectic melts as anatectic granites (e.g. Sainte-Anne d'Auray granite). The ca. 320-318 Ma age is ascribed to late or even post-migmatitic granites (e.g. Carnac, Ploemeur and Quiberon granites).

delamination of the lithospheric mantle root or slab break-off or 3) a decrease in the strength of the thickened orogenic crust due to thermal weakening thereby emphasising the role of partial-melting [e.g. Vanderhaeghe and Teyssier, 2001].

At variance with the examples of the Canadian Cordillera or the Aegean domain where 1) precise geodetic or geophysical-constraints on plate-kinematics, 2) the behaviour of the mantle or 3) the geometry of the continents involved in the collision are reasonably well-constrained, the case of the West-European Variscan orogenic segment is quite complex as none of these parameters is accurately constrained. Conversely, one of the most important advantages of the coastal South Armorican domain is to offer an opportunity to study deep-orogenic processes over series of crustal-scale sections and more particularly, to understand the role of partial-melting for the late-evolution of the Variscan orogenic belt.

An unexpected major result of this study is the recognition of a common, regional-scale probable signature of prograde metamorphism dated at ca. 335-330 Ma and ascribed to crystallisation ages during monazite-forming reaction(s) before partial-melting [e.g. Gloaguen, 2006; Janots *et al.*, 2009]. A single, common ca. 326-318 Ma age-cluster was also retrieved for the melt crystallisation age following a major partial-melting event throughout the coastal South Armorican domain. Partial-melting thus occurred at regional-scale some ca. 35-40 m.y. after the youngest age of subduction-related HP metamorphism recognised within the uppermost units [Le Hébel *et al.*, 2002a, 2002b; Bosse *et al.*, 2005] and ca. 20-25 m.y. after the integration of the first continental-derived units in the orogenic wedge during collision [Maillet, 1984; Colchen and Poncet, 1987; Rolin and Colchen, 2001]. Traces of initial stacking-structures are however extremely rare and heavily reworked by subsequent extensional deformation except at the top of the structural pile [e.g. Iglesias and Brun, 1976; Maillet, 1984; Goujou, 1992; Geoffroy, 1993; Le Hébel *et al.*, 2002a, 2002b; Bosse *et al.*, 2002, 2005]. The currently outcropping architecture of the coastal South Armorican Domain thus displays large-scale portions of Variscan lower crust heavily affected by partial-melting and pluton intrusions that were exhumed in several migmatite-cored metamorphic domes roofed by extensional shear zones [Brown, 2005; Turrillot *et al.*, 2009; 2011a].

As inferred from existing P-T constraints of the studied migmatite belt [Brown, 1983; Jones and Brown, 1989, 1990; Goujou, 1992; Brown and Dallmeyer, 1996; Marchildon and Brown, 2003; Johnson and Brown, 2004], the solidus was intersected during burial along a typical Barrovian collisional gradient (fig. 13) [e.g. Gardien *et al.*, 1997]. Accordingly, inception of partial-melting at peak-P conditions therefore coincides with a major change in the tectonic regime recorded at regional-scale. The inception and drastic generalisation of partial-melting within the coastal South Armorican orogenic segment was coeval with a fast transition from net crustal thickening to net crustal thinning. Accordingly, partial-melting that must have lowered the overall strength of the orogenic crust would be an important factor facilitating the late-orogenic gravitational collapse of the Armorican orogenic segment or the Variscan belt. However, transition from compression to extension may also imply a contemporaneous change of kinematic boundary conditions. Two scenarios can then be envisaged: (1) continuous stacking of

continental crustal units, rich in radiogenic elements, led to an increase of temperature within the orogenic wedge provoking partial-melting, the resulting drop in the crustal strength inducing lateral flow of a low-viscosity partially molten layer, like in the recent evolution of the Tibetan plateau [Nelson *et al.*, 1996], or (2) a drastic change of the boundary conditions such as convective removal of the lithospheric mantle, or inception of slab retreat, slab detachment or tearing has induced hot asthenospheric upwelling which in turn led to coeval extension and partial-melting.

At a more local scale, interrelations between partial-melting or magmatism in a broad sense, and deformation were observed from the scale of large-scale plutons (fig. 11) to the scale of the outcrop (fig. 4g) [Cagnard *et al.*, 2004; Turrillot *et al.*, 2009]. Strain thus benefited of the low-strength of the magma prior to complete crystallisation promoting intense shearing of large-scale laccoliths often referred to as synkinematic plutons [Jégouzo, 1980; Gapais *et al.*, 1993; Tartèse *et al.*, 2011a]. The Ploemeur-Quiberon leucogranite set of plutons is particularly exemplary of that discussion as it concentrated large amount of strain over the km-thick top-to-the-WNW mylonitic Quiberon extensional shear zone [Gapais *et al.*, 1993]. By accurately constraining the emplacement-age of these plutons within non-melted overriding units, this study retrieved an early increment of the extensional deformation at ca. 321 ± 3 Ma. Cooling and further strain localisation at the roof of leucogranite laccoliths was subsequently recorded by ca. 310-300 Ma Ar/Ar datings (fig. 12) [Ruffet *in* Le Hébel, 2002; Turrillot, 2010]. By comparison with the French Massif Central, the Brême leucogranite laccolith [Gébelin *et al.*, 2006, 2009] bounded by the top-to-the-WNW Nantiat mylonitic extensional shear zone [Mollier and Bouchez, 1982; Faure, 1989, 1995; Faure and Pons, 1991], was emplaced syntectonically at ca. 324 ± 4 Ma [Holliger *et al.*, 1986]. Active exhumation, strain localization and cooling occur from ca. 320 to 305 Ma including a rapid exhumation stage (1.5 mm/yr) at ca. 305 Ma [Scaillet *et al.*, 1996a]. These results reinforce the recognition of a large-scale WNW-ESE strain continuum without significant kinematic change during the exhumation of metamorphic domes from the end of the partial-melting event in the upper amphibolite facies at ca. 320 Ma and continued to near surface brittle faulting until ca. 302-298 Ma controlling the sedimentation in intramontane coal-bearing basins [e.g. Pin and Vielzeuf, 1983; Gébelin *et al.*, 2007]. Similarly, strain localisation over series of major extensional shear zones seem greatly influenced by the emplacement of thin leucogranite laccoliths from ca. 320 Ma [Gapais *et al.*, 1993; Gébelin *et al.*, 2006, 2007, 2009; Turrillot *et al.*, 2011].

Geometry, and in some extent, kinematics and timing of exhumation of migmatite-cored metamorphic domes of the coastal South Armorican domain or even of the western parts French Massif Central are therefore quite comparable [Caen-Vachette *et al.*, 1984; Gapais *et al.*, 1993; Burg and Vanderhaeghe, 1993; Burg *et al.*, 1994; Scaillet *et al.*, 1996b; Ledru *et al.*, 2001; Vanderhaeghe *et al.*, 2001; Cagnard *et al.*, 2004; Gébelin *et al.*, 2006, 2007, 2009; Turrillot, 2010; Turrillot *et al.*, 2011a]. Besides, as constrained by emplacement ages of synkinematic plutons and both cooling and deformation ages (see synthesis on figure 12) [Turrillot, 2010; Augier *et al.*, 2011; Tartèse *et al.*, 2011a; 2011b, 2012], dextral displacement along the South Armorican shear zone

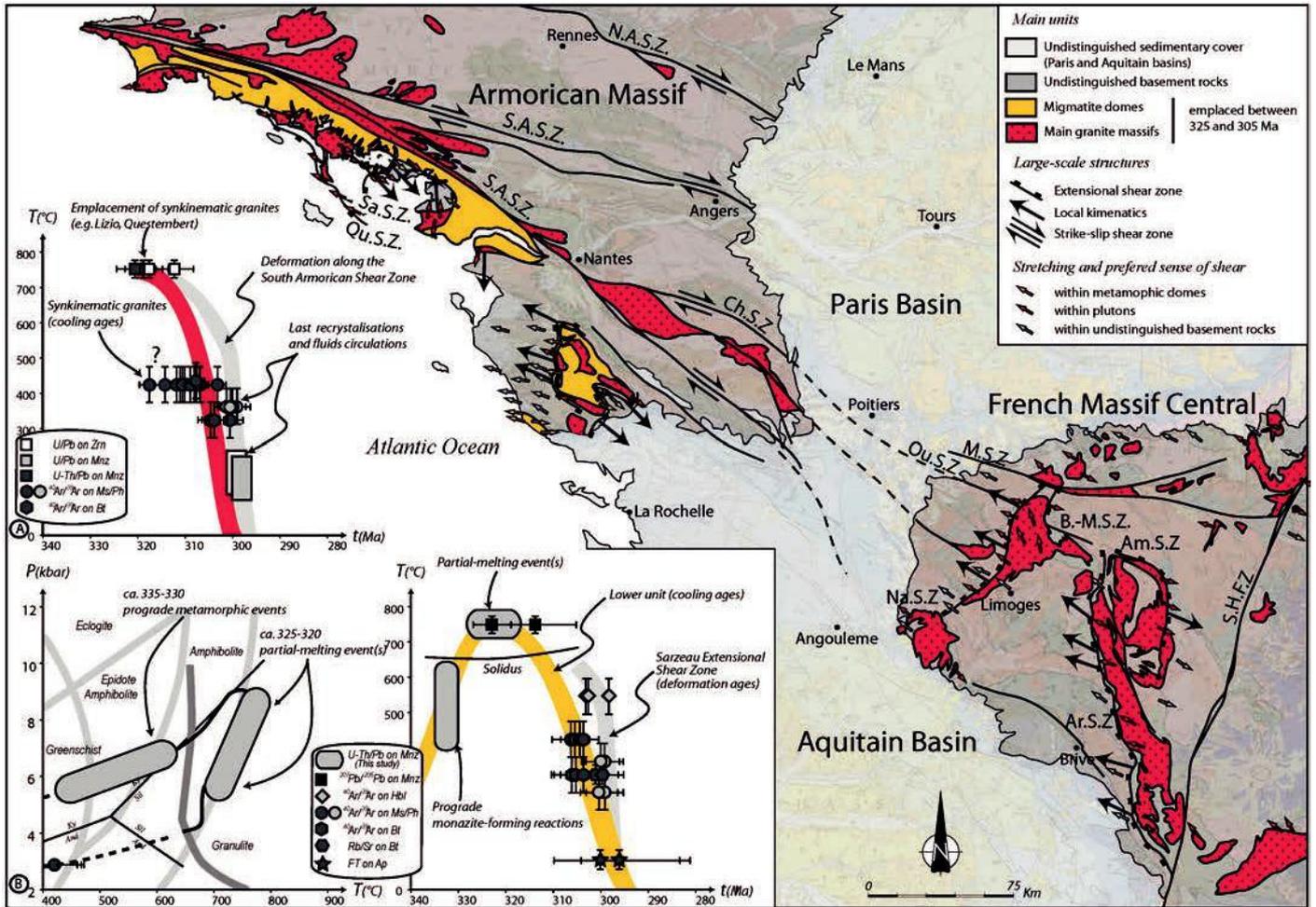


FIG. 12. – Correlations between Armorican massif and the French Massif Central. Particularly highlighted are correlations of wrench and extensional shear zones, (leuco-) granite massifs and metamorphic domes (inspired by Gébelin *et al.* [2007]; data are from Autran and Lameyre [1980]; Faure [1989, 1995]; Faure and Pons [1991]; Goujou [1992]; Cagnard *et al.* [2004]; Gébelin *et al.* [2006, 2007, 2009]; Turrillot *et al.* [2011a, 2011b] and unpublished data from our team). Qu.S.Z.: Quiberon shear zone; Sa.S.Z.: Sarzeau shear zone; N.A.S.Z.: North Armorican shear zone; S.A.S.Z.: South Armorican shear zone; Ch.S.Z.: Cholet shear zone; B.-M.S.Z.: Bussière-Madélène shear zone; Ou.S.Z.: Ouzilly shear zone; M.S.Z.: Marche shear zone; Am.S.Z.: Ambrugeat shear zone; S.H.F.Z.: Sillon Houillier fault zone. (b) Deformation ages and inferred T-t evolution of the South Armorican shear zone and the associated synkinematic granites. U/Pb ages on zircon (zrn) and monazite (Mnz) are from Tartèse *et al.* [2011a, 2011b, 2012]; $^{40}\text{Ar}/^{39}\text{Ar}$ time-constraints on muscovite, phengite (Ph) and biotite (Bt) from Ruffet *in* Le Hébel [2002], Tartèse *et al.* [2011a, 2011b], Turrillot [2010] and Augier *et al.* [2011]. (c) P-T-t evolution and inferred T-t evolution of the migmatites of the Golfe du Morbihan area from prograde to final cooling stages. Quantitative P-T path where obtained by MnNCKFMASH pseudo-section modelling [Johnson and Brown, 2004]. Metamorphic facies are after Spear [1993]. $^{207}\text{Pb}/^{206}\text{Pb}$ ages on monazite (Mnz) are from Peucat [1983]; $^{40}\text{Ar}/^{39}\text{Ar}$ time-constraints on hornblende (Hbl), phengite (Ph) and biotite (Bt) from Brown and Dallmeyer [1996] and Turrillot *et al.* [2011a]; Rb/Sr on biotite (Bt) data from Peucat [1983]; apatite (Ap) fission-track (FT) ages from Carpena *et al.* [1979]. Available geochronological time-constraints on the Sarzeau extensional shear zone mylonites are given for comparison (see details in Turrillot *et al.* [2011a]). Mineral abbreviations are after Whitney and Evans [2010].

appears coeval with extensional tectonics. The common WNW-ESE stretching and the dominant top-to-the-WNW kinematics of the extensional deformation thus imply an overall post-thickening northwestward non-rigid motion of the internal zones with respects to the non-thickened Central Armorican domain controlled by the South Armorican shear zone. It is therefore tempting to propose that the dextral motion over the South Armorican shear zone would be as a major orogeny-parallel transfer zone as initially proposed by Gapais *et al.* [1993]. Furthermore, non-rigid tectonic escape does imply a free-boundary to the northwest of the orogenic system that may have played a major role in the formation, or at least to the amplification of the Ibero-Armorican arc. Constraints on the full spatial extent of ca. 320-300 Ma extensional tectonics as well as better, understanding of the orocline formation, still remain to be

achieved and requiring integration at the scale of the Variscan belt.

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