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Scientific Research

Spinel-Bearing Lherzolite Xenoliths from Hosséré Garba (Likok, Adamawa-Cameroon): Mineral Compositions and Geothermobarometric Implications

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Abstract

Spinel-bearing lherzolite xenoliths from the Hosséré Garba (1272 m.a.s.l) volcano on the Adamawa Plateau, is located in Likok village, at about 35 km to WSW of Ngaoundéré. These xenoliths (~11 cm size) have been sampled into the host basaltic lava flows from the NE flank of the Hosséré Garba volcano. These xenoliths characterized by porphyroclastic texture consisted of olivine (~55 vol.%), orthopyroxene (~19 vol.%) and clinopyroxene (~21 vol.%) crystals. Spinel crystals (~5 vol.%) are red brown and interstitial between the crystals of olivine and pyroxenes. CaO contents are low (<0.08 wt%) in olivine and similar to those estimated (CaO: 0.05 - 0.1 wt%) for the mantle origin. The values of Al^{VI}/Al^{IV} ratio range between 1.1 and 1.3 for the Cr-diopside crystals from Hosséré Garba xenoliths. The constant value of the volumes V(Cell) and V(M1) for clinopyroxene compositions, indicates the similar pressures. Hosséré Garba represents a residual sequence issued from partial melting of a mantle source. Similar compositions have been recorded in minerals of ultramafic xenoliths from other ultramafic xenoliths domains of the Cameroon Line and the Adamawa Plateau.

Keywords

Xenoliths, lherzolite, Geothermobarometry, Likok, Adamawa

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1. Introduction

Ultramafic xenoliths generally are considered as small parts broken or separated off the lithospheric mantle [1]; or the residues of partial melting of the upper-mantle [2] hosted in basaltic lava during volcanic eruption. These ultramafic xenoliths ejected from volcanoes at the surface, as solid fragmentary material, are important to the knowledge of the textural features, chemical evolutions, temperature and pressure conditions of the upper mantle. Upper-mantle xenoliths were found in basaltic lavas from several localities along the Cameroon Volcanic Line [3]-[5]. However, beneath the volcanoes of the Adamawa Plateau, there are few available data aside from ultramafic xenoliths described from the Dibi volcano [6] [7], Youkou maar [8] and Ngao Voglar volcano [4]. Lherzolite xenoliths described in this paper are discovered in the Hosséré Garba (1272 m.a.s.l) volcano on the Adamawa Plateau, which consist of basaltic lava flows, more or less fragmented into rounded sparse blocks. Hosséré Garba is located in Likok village, at about 35 km to WSW of Ngaoundéré (**Figure 1**).

In this paper, we present the first petrography and mineral chemistry of spinel-bearing lherzolite xenoliths from Hosséré Garba (Likok) in the Adamawa Plateau. Chiefly, the objective is to enhance the knowledge of upper mantle beneath the Adamawa Plateau volcanic domain.

2. Geological Setting and Petrography

The Adamawa Plateau is a tectono-magmatic domain, bounded respectively to the North and South by the Adamawa and Djerem-Mbéré faults, oriented N70˚E [9]. The Adamawa basement rocks consist of Paleo-Proterozoic metamorphic rocks (2.1 Ga; [10] [11]), which were intensively remobilized by the Pan-African orogenesis [12]-[14]. These rocks crossed by the deformed and metamorphosed granitoids with ages of 615 ± 27 Ma to 575 ± 27 Ma [15] are partially covered by Cenozoic basaltic and felsic [16]-[20].

Spinel-bearing lherzolite xenoliths have been sampled into the basaltic lava flows from the NE flank of the Hosséré Garba volcano (Likok sector, Adamawa Plateau). These xenoliths (~11 cm size) characterized by porphyroclastic texture [21] with a cataclastic appearance are greenish yellow and slightly dark. Olivine (~55 vol.%), orthopyroxene (~19 vol.%) and clinopyroxene (~21 vol.%) crystals identified in these xenoliths are more or less equigranular (up to 0.4 - 1 cm). Spinel crystals (~5 vol.%) are red brown and interstitial between

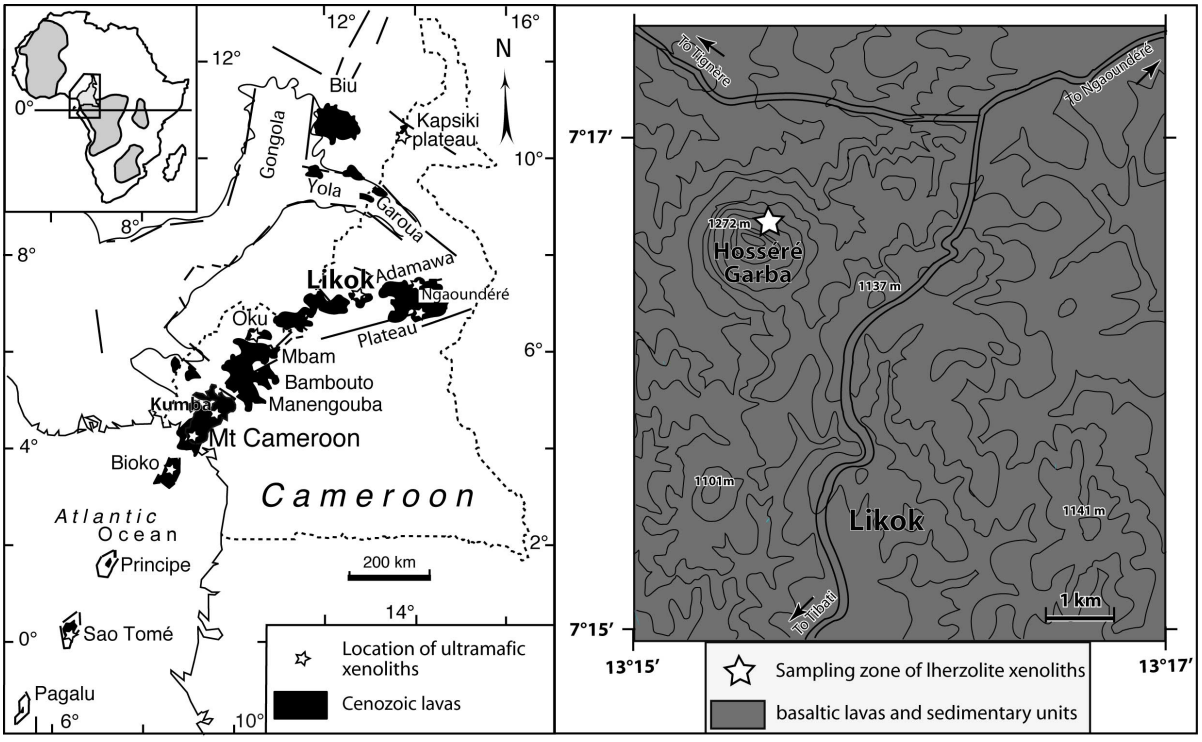


Figure 1. Location of the ultramafic xenoliths from the Hosséré Garga (Likok), and from the other sectors of the Cameroon Volcanic Line and the Adamawa Plateau.

the crystals of olivine and pyroxenes. The "kink-bands" structure characterizes some olivine crystals. Many triple-points involving olivine, pyroxene and spinel show evidence of a particular equilibrium condition between these minerals [22]. The tiny host basalt veins are present in rim sections of the lherzolite xenoliths from Hosséré Garba.

3. Analytical Methods

Chemical analyzes ("Université de Bretagne Occidentale—Brest, France") of major elements were performed using an electron microprobe with beams of: 10 and 40 μm, 15 kV and 15 nA. Errors considered for these analyzes are between 5% and 10% of measured values < 1% and between 1% and 5% of the measured values > 1%. Measurements of trace elements of the different phases have been produced using a technique laser ablation of a mass spectrometer coupled plasma source (LA-ICP-MS). The beam size varies between 40 and 70 μm. Counts where repeated several time for each spot analysis and averaged. The number of repeated analyses and the duration of the counting times depend on the concentration level of the elements studied. To examine homogeneity of constituent mineral, more than seventeen crystals each of clinopyroxène were selected at random from each thin section. Calcium is used as an internal standard, while the NIST 612 is used as an external standard. The accuracy is better than 10%. The whole rock analyzes were performed previously in solution using ICP-AES for major elements and ICP-MS for trace elements.

4. Mineral Chemistry

4.1. Olivine

The forsterite component of olivine range from 89 to 90 (**Table 1**), with the high values of Mg# [Mg# = 100 × atomic Mg/(Mg + Fe^{2+})] (89 - 90).

These values are similar to those calculated for the lherzolite xenoliths from Ngao Voglar. NiO contents reaching 0.41 wt% in the olivine from Hosséré Garba xenoliths. CaO contents are low (<0.08 wt%) and similar to those estimatited (CaO: 0.05 wt% - 0.1 wt%) by [23] for the mantle origin, which characterize equilibrium environment of high pressure (>0.2 GPa; [24]).

Table 1. Olivine compositions (wt% and a.p.f.u. on the basis 4 oygens).

SiO_2 (wt%)	40.96	41.47	41.05	41.13	40.93	40.94
Al_2O_3	0.02	0.02	0.02	0.05	0.01	0.01
FeO	9.77	10.12	9.89	10.11	9.95	9.98
MnO	0.16	0.16	0.12	0.16	0.12	0.16
MgO	49.26	49.53	48.83	49.41	49.24	49.38
CaO	0.07	0.05	0.02	0.06	0.05	0.01
NiO	0.37	0.33	0.34	0.41	0.25	0.36
Sum	100.61	101.68	100.27	101.33	100.55	100.84
Si (a.p.f.u.)	0.998	1.001	1.004	0.997	0.999	0.997
Al	0.000	0.000	0.000	0.001	0.000	0.000
Fe^{2+}	0.199	0.204	0.202	0.205	0.203	0.203
Mn	0.003	0.003	0.003	0.003	0.003	0.003
Mg	1.790	1.782	1.780	1.785	1.791	1.792
Ca	0.002	0.001	0.001	0.001	0.001	0.000
Ni	0.007	0.006	0.007	0.008	0.005	0.007
Fa (%)	10.0	10.3	10.2	10.3	10.2	10.2
Fo	90.0	89.7	89.8	89.7	89.8	89.8

4.2. Pyroxene

Clinopyroxene compositions ($En_{53.1–54.9}Wo_{43.5–44.8}Fs_{3.4–0.3}$) fall within the diopside field, according to Wo-En-Fs diagram **Figure 2** (Morimoto *et al*. 1988).

The TiO_2 contents range from 0.42 to 0.54 wt% and those of Al_2O_3 from 6.5 to 6.9 wt% (**Table 2**).

These diopside crystals are Cr-rich with the values of Cr# [Cr# = 100 × atomic Cr/(Cr + Al)] (6.5 - 7.2) typical of Cr-diopside [25]. The clinopyroxene is characterized by Mg# [Mg# = 100 × atomic Mg/(Mg + Fe^{2+})] values of 94.3 - 97.8, higher than those of coexisting orthopyroxene (see below) and olivine. The values of Al^{VI}/Al^{IV} ratio range between 1.1 and 1.3 for the Cr-diopside crystals from Hosséré Garba xenoliths. The M1 site volume (V(M1)) of these minerals show small variations or are virtually constant (**Figure 3**).

This is due to the Al^{IV} and R^{3+} (Al^{VI}, Fe^{3+}, Cr and Ti^{4+}) relationship. The constant value of the volumes V(Cell) and V(M1) (see **Table 2**) for clinopyroxene compositions, indicates the similar pressures [26] in the spinel-bearing lherzolite xenoliths from Hosséré Garba. V(Cell) and V(M1) relationships are more sensitive than the only chemical composition, to estimate the equilibration pressure of clinopyroxene [27].

Orthopyroxene (**Table 3**) with the compositions of Mg-enstatite (Mg# = 90.6 - 92.4; $En_{88.7–89.4}Wo_{1.2–1.5}Fs_{10.1–9.1}$) have been analyzed in the lherzolite xenoliths from Hosséré Garba. These crystals are characterized by the low TiO_2 (0.05 wt% - 0.14 wt%) and Cr_2O_3 (0.28 wt% - 0.33 wt%) contents.

4.3. Spinel

Spinel crystals display homogenous compositions, with Mg# and Cr# ranging from 81.1 to 81.9 and from 10.6 to 10.8 respectively (**Figure 4**).

TiO_2 contents are low (0.11 wt% - 0.13 wt%) and those of Al_2O_3 are high (up to 57.8 wt%). The Cr_2O_3 contents vary between 10.26 wt% and 10.41 wt% (**Table 4**).

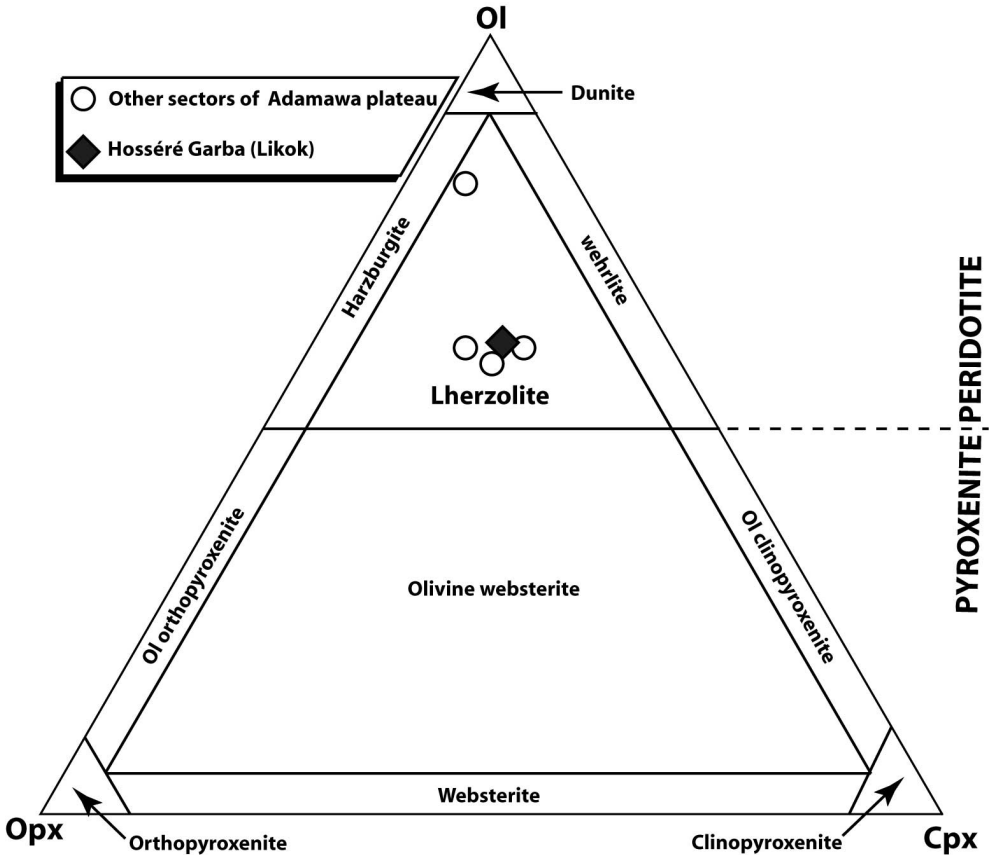


Figure 2. Classification diagram of ultramafic rocks (Streckeisen 1976) on the basis of the modal compositions, used for the nomenclature of Hosséré Garbe ultramafic xenolith types.

Table 2. Clinopyroxene compositions (wt% and a.p.f.u. on the basis of 6 oxygens) and elemental cation partition among T, M1 and M2 sites.

SiO_2 (wt%)	51.88	51.99	52.47	51.76	52.25	51.85	51.97	52.44
TiO_2	0.51	0.46	0.42	0.54	0.52	0.46	0.42	0.52
Al_2O_3	6.70	6.53	6.56	6.71	6.50	6.67	6.65	6.85
Cr_2O_3	0.78	0.73	0.74	0.75	0.67	0.77	0.75	0.72
Fe_2O_3	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
MnO	0.10	0.10	0.14	0.13	0.09	0.10	0.11	0.08
MgO	15.21	15.50	15.54	15.47	15.39	15.45	15.39	15.35
CaO	20.76	20.26	20.55	20.26	20.52	20.51	20.33	20.55
Na_2O	1.64	1.61	1.70	1.72	1.75	1.70	1.57	1.73
Sum	100.45	100.07	100.92	100.29	100.63	100.53	100.01	101.20
Fe_2O_3 (calc.)	1.88	1.47	1.76	2.23	1.93	2.40	1.20	1.58
FeO (calc.)	0.96	1.41	1.04	0.72	1.01	0.61	1.61	1.36
Sum (calc.)	100.45	100.07	100.92	100.29	100.63	100.53	100.01	101.20
T-site								
Si (a.p.f.u.)	1.866	1.875	1.875	1.862	1.874	1.862	1.875	1.871
Al^{iv}	0.134	0.125	0.125	0.138	0.126	0.138	0.125	0.129
Sum	2.000	2.000	2.000	2.000	2.000	2.000	2.000	2.000
M1-site								
Mg	0.737	0.737	0.741	0.738	0.739	0.740	0.733	0.728
Fe^{2+}	0.029	0.039	0.028	0.020	0.027	0.017	0.045	0.037
Al^{vi}	0.150	0.152	0.152	0.147	0.149	0.145	0.158	0.159
Fe^{3+}	0.047	0.039	0.047	0.060	0.052	0.065	0.032	0.043
Ti	0.014	0.012	0.011	0.014	0.014	0.013	0.011	0.014
Cr	0.022	0.021	0.021	0.021	0.019	0.022	0.021	0.020
Sum	1.000	1.000	1.000	1.000	1.000	1.000	1.000	1.000
M2-site								
Ca	0.800	0.783	0.787	0.781	0.788	0.789	0.786	0.785
Na	0.115	0.112	0.118	0.120	0.122	0.119	0.110	0.120
Mg	0.079	0.097	0.087	0.092	0.084	0.087	0.095	0.088
Fe^{2+}	0.003	0.005	0.003	0.002	0.003	0.002	0.006	0.004
Mn	0.003	0.003	0.004	0.004	0.003	0.003	0.003	0.002
Sum	1.000	1.000	1.000	1.000	1.000	1.000	1.000	1.000
V(cell)	437.66	437.48	437.42	437.38	437.48	437.48	437.53	437.48
V(M1)	11.39	11.40	11.39	11.39	11.39	11.39	11.39	11.37
Al^{VI}/Al^{IV}	1.1	1.2	1.2	1.1	1.2	1.0	1.3	1.2
Mg#	96.6	95.2	96.4	97.5	96.5	97.8	94.4	95.3
Cr#	7.2	7.0	7.0	7.0	6.5	7.2	7.0	6.6
Wo (mol.%)	44.8	46.5	44.1	46.6	46.5	44.9	44.7	45.6
En	37.6	38.5	41.0	38.7	39.2	39.0	38.5	38.0
Fs	17.6	15.0	14.8	14.7	14.3	16.1	16.8	16.4

$V(cell) = 11{,}864*Fe^{2+}_{M1} + 9.107*Fe^{3+} - 18.375*Al_{M1} + 11.794*Ti - 1.4925*Cr + 439.97*Ca + 419.68*Na + 431.72*Mg_{M2} + 432.56*^{Fe2+} + 428.03*Mn - 28.652*(Mg_{M2})^2 - 12.741*(Fe^{2+}_{M2})^2$ and $V(M1) = -0.3085*Al_T + 0.8130*Fe^{2+}_{M1} - 0.4173*Fe^{3+} - 2.029*Al_{M1} - 1.0864*Ti - 0.8001*Cr + 11.931*Ca + 11.288*Na + 11.432*Mg_{M2} + 11.885*Fe^{2+}_{M2} + 12.038*Mn + 2.4335*(Mg_{M2})^2 - 1.1661*(Fe^{2+}_{M2})^2$ based on the papers of Nimis (1995) and Nimis and Ulmer (1998).

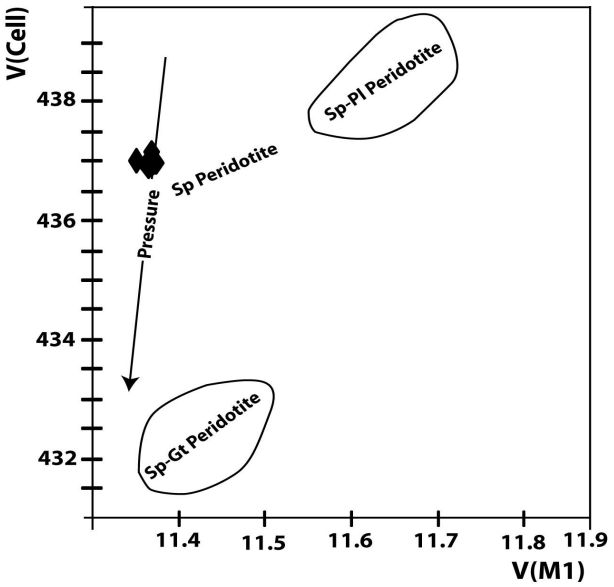


Figure 3. V(M1) vs. V(cell) diagram for clinopyroxene. The fields for clinopyroxene crystals from spinel-garnet (Sp-Gt) peridotite and spinel-plagioclase (Sp-Pl) peridotite are presented (Nimis, 1995).

Table 3. Orthopyroxene compositions (wt% and a.p.f.u. on the basis of oxygens).

SiO_2 (wt%)	54.90	54.96	55.53	55.03	55.61	55.56	55.34	55.32
TiO_2	0.05	0.13	0.19	0.05	0.14	0.14	0.11	0.14
Al_2O_3	4.45	4.44	4.30	4.35	4.32	4.48	4.38	4.51
Cr_2O_3	0.28	0.31	0.28	0.32	0.26	0.32	0.28	0.33
FeOt	6.04	6.40	6.02	6.23	6.12	6.12	6.06	6.25
MnO	0.13	0.13	0.14	0.10	0.13	0.12	0.14	0.19
MgO	33.15	32.99	33.16	33.39	32.97	33.07	33.27	33.53
CaO	0.79	0.74	0.83	0.77	0.71	0.79	0.64	0.76
Na_2O	0.08	0.08	0.12	0.09	0.16	0.11	0.08	0.09
K_2O	0.00	0.01	0.02	0.02	0.01	0.00	0.04	0.00
Sum	99.96	100.24	100.66	100.46	100.46	100.75	100.46	101.27
Fe_2O_3 (calc.)	1.09	0.97	0.42	1.63	0.07	0.18	0.61	1.49
FeO (calc.)	5.06	5.52	5.65	4.76	6.06	5.96	5.51	4.91
Sum (calc.)	99.97	100.27	100.62	100.51	100.42	100.74	100.41	101.28
Si (a.p.f.u.)	1.893	1.893	1.904	1.889	1.910	1.903	1.901	1.885
Al^{iv}	0.107	0.107	0.096	0.111	0.090	0.097	0.099	0.115
Al^{vi}	0.074	0.073	0.078	0.065	0.085	0.084	0.078	0.067
Ti	0.001	0.003	0.005	0.001	0.004	0.004	0.003	0.003
Cr	0.008	0.008	0.008	0.009	0.007	0.009	0.008	0.009
Fe^{3+}	0.028	0.024	0.009	0.040	0.001	0.004	0.013	0.038
Fe^{2+}	0.146	0.160	0.164	0.138	0.175	0.171	0.162	0.140
Mn	0.004	0.004	0.004	0.003	0.004	0.004	0.004	0.006
Mg	1.704	1.694	1.695	1.709	1.688	1.689	1.704	1.704
Ca	0.029	0.027	0.030	0.028	0.026	0.029	0.023	0.028
Na	0.006	0.005	0.008	0.006	0.010	0.007	0.006	0.006
Mg#	92.1	91.4	91.2	92.5	90.6	90.8	91.3	92.4
Wo (mol.%)	1.5	1.4	1.6	1.5	1.4	1.5	1.2	1.5
En	89.2	88.7	89.1	89.1	89.1	89.0	89.4	88.9
Fs	9.3	9.8	9.3	9.5	9.5	9.4	9.4	9.6

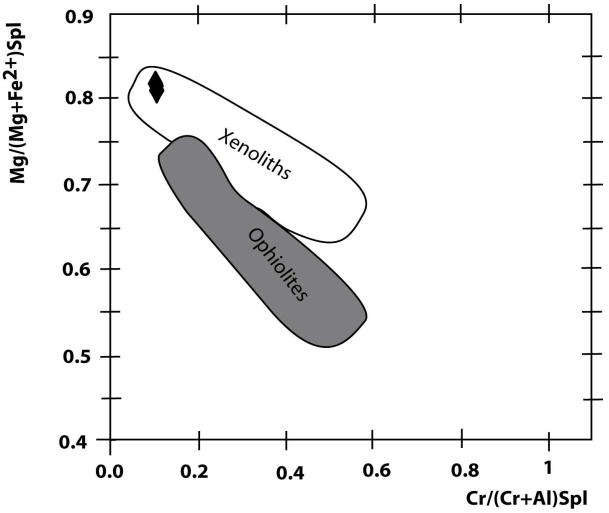


Figure 4. Mg# vs. Cr# diagram for spinels from Hosséré Garba. The domains of xenoliths and ophiolites are established respectively by Conticelli and Peccerillo (1990), and Cabanes and Mercier (1988).

Table 4. Spinel compositions (wt% and a.p.f.u. on the basis of 32 oxygens).

SiO_2 (wt%)	0.05	0.07	0.09	0.07
TiO_2	0.12	0.13	0.13	0.11
Al_2O_3	57.82	57.84	57.50	57.40
Cr_2O_3	10.39	10.26	10.41	10.40
FeOt	10.25	10.17	10.37	10.18
MnO	0.13	0.08	0.09	0.06
MgO	21.20	21.40	21.06	21.02
CaO	0.00	0.00	0.00	0.00
NiO	0.36	0.35	0.29	0.33
V_2O_5	0.00	0.05	0.01	0.00
Sum	100.31	100.35	99.94	99.58
Fe_2O_3 (calc)	1.88	1.91	1.77	1.71
FeO (calc)	8.55	8.45	8.78	8.64
Sum (calc.)	100.50	100.54	100.12	99.75
Si (a.p.f.u.)	0.010	0.014	0.018	0.015
Ti	0.018	0.021	0.019	0.017
Al	13.971	13.954	13.953	13.971
Cr	1.684	1.661	1.694	1.698
Fe^{3+}	0.290	0.295	0.275	0.266
Fe^{2+}	1.466	1.447	1.511	1.492
Mn	0.022	0.014	0.016	0.011
Mg	6.479	6.530	6.464	6.473
Ca	0.001	0.000	0.000	0.001
Na	0.000	0.000	0.000	0.000
Ni	0.059	0.058	0.048	0.055
V_2O_5	0.000	0.007	0.001	0.000
Mg#	81.5	81.9	81.1	81.3
Cr#	10.8	10.6	10.8	10.8

5. Discussion

5.1. Geothermobarometry

Using the orthopyroxene-clinopyroxene geothermometry of [28], equilibrium temperature of ~905˚C has been estimated for the Hosséré Garba spinel-bearing lherzolite xenoliths. Applying the geothermometer of [29], based on the solubility of Al_2O_3 in orthopyroxene and spinel, the temperature expression calibrated at the pressure of 1.5 GPa in a natural spinel lherzolite system, gives the equilibrium temperatures of Hosséré Garba spinel-bearing lherzolite xenoliths of 1215˚C - 1245˚C.

The lack of plagioclase or gamet in Iherzolites and the fact that the lherzolite xenoliths have reached a high degree of internal chemical equilibrium, indicate that equilibration took place at pressures ranging between 0.7 and 1.1 GPa (25 - 36 km). Hosséré Garba xenoliths have probably been sampled by the host-basalt in the upper mantle. At depths of 25 - 36 km, the calculated temperatures are about 200˚C above those of the oceanic geotherm of [30].

These equilibrium temperatures of spinel-bearing lherzolite xenoliths from Hosséré Garba fall more or less in the similar range as equilibrium temperatures of other lherzolite xenoliths along the Cameroon Volcanic Line and the Adamawa Plateau [7]-[23].

A maximum pressure can be calculated from the composition of spinel [31] or established with the clinopyroxene crystal chemistry data to determine the variability of the pressure during the partial melting. Hosséré Garba mantle xenoliths as those from Ngao Voglar [4] are spinel peridotites, containing neither plagioclase nor garnet, and were derived from the spinel stability field in the mantle, which lies in a pressure range of ~0.8 to ~2 GPa [32]-[34]. These estimated pressures can be associated to the depths ranging between 30 and 60 km as reported for the other spinel lherzolite xenoliths [35]-[40].

5.2. Spinel-Bearing lherzolite Xenoliths Origin

The refractory elements as Cr are enriched in the residue, as demonstrated for the Cr-rich clinopyroxene from Hosséré Garbé xenoliths. Thus, the studied spinel-bearing lherzolite xenoliths, characterized by the presence of Cr-diopside, represent probably portions of the upper mantle [41].

However, the compositions of olivine and spinel are strongly dependent on the degree of melting, a feature characterized by a nearly constant values of $(Mg\#)_{Ol}$, $(Mg\#)_{Sp}$, and $(Cr\#)_{Sp}$. Based on these data, the xenoliths from Hosséré Garba represent a residual sequence issued from partial melting of a mantle source. Experiments have shown that the composition of solid solutions, such as olivine, is very sensitive to temperature changes, and olivine could therefore evolve towards refractory compositions during the partial melting. So, an extraction of the melt could leave a residue depleted in highly incompatible elements (e.g. Na, Al, K, and Ti) and enriched in Cr and Ni relative to the most primitive mantle [21]. The range of Ni contents of olivine of lherzolites from Hosséré Garba confirms that they are residues of melting. It appears that high contents of Ni (2500 - 4100 ppm) in olivine suggest that xenoliths from Hosséré Garba could derive from a low degree of melting (~30%) of a relatively homogeneous mantle source. Thus, the model of [41] for portions of the earth's upper mantle appears appropriate for the origin of the Hosséré Garba spinel-bearing lherzolite xenoliths.

6. Conclusion

The spinel-bearing lherzolite xenoliths from Hosséré Garba represent a residual sequence issued from partial melting of a mantle source. The compositions of their silicate minerals are typical of equilibrated Mg-rich olivine, Al-rich enstatite and Cr-diopside in the upper mantle. Similar compositions have been recorded in minerals of ultramafic xenoliths from other ultramafic xenoliths domains of the Cameroon Volcanic Line and the Adamawa Plateau. Thus, ultramafic xenoliths from Hosséré Garba represent peridotites sampled from lithospheric mantle and hosted in the basaltic lava flows.

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