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Coastal palaeoenvironmental record of the last 7 kyr BP in NW France: Sub-millennial climatic and anthropic Holocene signals

Assia Fernane,1,2 Emmanuel Gandouin,2 Aurélie Penaud,1 Brigitte Van Vliet-Lanoë,2 Jérôme Goslin,1 Muriel Vidal1 and Christophe Delacourt1

Abstract
To date, Holocene palaeoecological signatues on the Northwestern coast of France have not been thoroughly investigated. In this study, environmental changes related to both climate processes and human disturbances were reconstructed over the last 7000 cal yr BP, based on pollen and chironomid assemblages from four coastal cores retrieved in Western Brittany (Porsmolin beach, NW France). Pollen and chironomid records show an environmental response to both millennial- and centennial-scale climate changes. During the mid-Holocene (until around 4200 cal yr BP), when human impact was low, Porsmolin’s landscape was characterized by a mixed oak forest dominated by Corylus trees. Comparison of our data with other palynological data from the NW French Atlantic coast shows that this dominance may be related to more humid conditions prevailing at Porsmolin. Furthermore, over the mid-Holocene, Corylus and Quercus patterns appear anti-correlated, with Quercus declines appearing synchronous with major storms revealed by data recently compiled for the Holocene in Brittany, suggesting that cold conditions probably caused the Quercus contractions, which in turn favoured the rise of the heliophile taxon Corylus. Regarding chironomids over this period, our results show variations in lotic-lentic taxa related to hydrological conditions, more specifically changes in river run-off strength. Over the Late Holocene, the Porsmolin palynological record shows progressive vegetation changes mainly related to human settlement and development of agricultural activities.

Keywords
Brittany, chironomids, North Atlantic Oscillation, peat, pollen, storm events

Introduction
Climate in the Holocene has not been uniformly warm, but was instead influenced at an orbital scale by decreasing summer insolation in the Northern Hemisphere (Berger and Loutre, 1991), and punctuated by abrupt climatic events at millennial timescales (Mayewski et al., 2004). At sub-millennial timescales, as well as over the past two millennia, numerous climate reconstructions suggest various forcing mechanisms and climate feedbacks for Holocene climate fluctuations, generally involving North Atlantic Oscillation (NAO) modes, solar activity, ocean circulation or volcanic eruptions (Bond et al., 2001; Fletcher et al., 2013; Van Loon and Meehl, 2014; Wanner et al., 2008). Consequently, multidisciplinary palaeoclimatic studies from different ecosystems are necessary to understand how these forcing mechanisms interact to define Holocene climate at a regional/global scale. At a local scale, multiproxy investigations are also crucial since some ecosystems, such as coastal ones, do not respond homogeneously to global changes, as suggested by Najjar et al. (2000).

In the northern part of Brittany, few palaeoenvironmental studies have been conducted, and have often applied a monodisciplinary palynological approach (e.g. Giot and Monnier, 1972; Morzadec-Kerfourn, 1974; Van Zeist, 1963). However, southern Brittany is well documented with continental (e.g. Barbier and Visset, 1997; Bernard, 1995; Gaudin, 2004; Joly and Visset, 2009; Visset et al., 1994) and marine (Naughton et al., 2007) palaeoenvironmental investigations. In this study, we present a multidisciplinary Holocene climate study based on pollen and chironomid (Insect, Diptera) analyses, as well as radiocarbon and lithological data from a Northwestern Brittany coastal area (Porsmolin (PM); NW France; 48°21'N). Studies associating chironomid and aquatic botanical proxies, both considered to be fast migrating biotic proxies, have increasingly been used in palaeoecological analysis of past environmental changes. Chironomids have emerged as promising tools for palaeoecological investigations. Their modern characteristic assemblages give diversified information such

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as water trophic status (Walker and Mathewes, 1987), dissolved oxygen content (Francis, 2001), water-level change (Massaferto and Brooks, 2002), types of alluvial habitats (Gandouin et al., 2006) and water (Rossaro, 1991) and air temperatures (Luoto, 2009). Their use has also enabled the successful detection of past human presence and its impact on water quality (Heiri and Lotter, 2003; Langdon et al., 2004; Ruiz et al., 2006; Taylor et al., 2013). Furthermore, pollen analysis can reveal climate variations and agro-pastoral activities over adjacent watersheds.

In this paper, we will attempt to decipher the respective influence of local versus regional/global factors over the last 7 kyr BP at PM, based on a cross-correlated pollen–chironomid analysis. At a regional scale, the present study constitutes the first coastal palaeoenvironmental investigations based on chironomids in Brittany. Furthermore, it should help improve palaeoecological knowledge of river chironomids.

Environmental setting and studied material

Sedimentary context

The PM beach is located in Northwestern Brittany within the Bay of Bertheaume and is surrounded by the Iroise Sea (Figure 1a). More precisely, our studied site is located at the outlet of a short valley and consists today of a sandy pocket beach surrounded by cliffs and limited landward by a high sandy coastal bar (Figure 1b). This geomorphological configuration encloses a marshy depression located at +6.5 m NGF (‘Nivellement Général de la France’ which is a French reference system for sea level elevation), crossed by a small stream that flows seaward. On the lower part of the beach, extensive thick layers of peat are exposed at low tide (Figure 1b), resulting from retrogressive erosion of sandy barriers.

The PM beach was cored with a 5 cm diameter percussion corer (Eikjelkamp). Twin cores were extracted in the marsh (PM3 and PM4; 48°21′24.88″N, 4°40′44.41″W; Figure 1b) and in the intertidal zone (PM1 and PM5; 48°21′18.67″N, 4°40′45.10″W; Figure 1b). The accessibility of the PM marsh did not allow us to drill intermediate cores to better constrain the geometry of the marsh infill.

Present-day climatic context and associated vegetation cover

The climate in Brittany is characterized by temperate oceanic conditions (because of the North Atlantic Ocean influence) with generally cool, rainy winters and fresh summers. Mean annual precipitation is close to 800 mm, and the mean annual temperature is around 11°C with weak thermic seasonal amplitudes. Meteorological conditions over Brittany mainly result from the NAO pattern (Hurrell, 1995). This natural pattern of atmospheric variability directly influences temperatures and hydrological conditions (precipitation, river discharges, and water table levels), which in turn influence vegetation cover (Gouveia et al., 2008).

Regarding current vegetation, the flanks of the PM valley are partly covered by a forest composed of Quercus, Corylus avellana and Carpinus betulus. The bottom of the valley is surrounded by woodlands dominated by Salix fragilis and several other herbaceous species such as Cyperus longus, Juncus effusus, Limonium dodartii, Mentha suaveolens, Polygonum aviculare, Potentilla anserina and Rumex obtusifolius.

Anthropogenic context

Brittany’s coasts are escarped, with intertidal zones that provide both shelter and fishery facilities which have favoured human settlement and economic development (Daire et al., 2011). Nevertheless, the human settlement chronology in Brittany is still obscure because of the scarcity of well-dated and continuous palynological records. In Southeastern Brittany, Cerealia pollen grains are sporadically recorded at the end of the Mesolithic and at the beginning of the Neolithic, indicating probable early human activity at that time (Ougueret and Visset, 2001; Visset et al., 2002). In Northern Brittany, the earliest agricultural activities are marked by the presence of Cerealia and clearing events at the beginning of the Sub-Boreal period, that is, at the end of the Bronze Age (Barbier and Visset, 1997; Marguerie, 1992; Morzade-Kerfourn, 1974). The 1st millennium BC (i.e. from 3 kyr BP), covering the period of the Late Bronze Age, the Iron Age and the first decades of the Gallo-Roman period, constitutes a key period concerning land use for agro-pastoral activities. Based on archaeological studies, human colonization of Northern Brittany at the beginning of the Bronze Age (4 kyr BP) appears conditioned by the presence of dense forest, with the first
clear events related to cultural activities occurring during the mid-Bronze Age (e.g. Daire et al., 2011). At the end of the Iron Age, from 450 to 50 BC (2.5-2 kyr BP), clearing events are intensified in parallel with a demographic expansion recorded in Western France (Giot et al., 1995). This period is marked by the development of enclosures and human settlement increases on coastal areas, which favoured marine exploitation (salt production, fishing) and maritime exchanges (e.g. Daire et al., 2011).

Nowadays, anthropic infrastructures such as the Atlantic Wall (relic of the Second World War), boarding dock, wastewater pipe, car park and access stairs artificialize PM beach. Over the last decades, agricultural activities have strongly declined in favour of tourism infrastructures such as hiking trails and camping areas, which contribute strongly to deterioration of the environment, especially coastal dune vegetation which is very sensitive to human trampling.

**Methodology**

*Sedimentological analysis and chrono-stratigraphy*

The lithology of the cores is mainly characterized by dense peats deposited in freshwater conditions and is interrupted by sand layers (Figure 2). A detailed sedimentological and stratigraphical analysis of the coastal sedimentation of PM was recently undertaken by Goslin (2014). According to this author, major discontinuities represented by sandy layers appear related to
Table 1. List of all ¹⁴C dating obtained on Porsmulin (PM) cores.

<table>
<thead>
<tr>
<th>Depth sample in PM cores (cm)</th>
<th>Altitude NGF (m)</th>
<th>Laboratory code</th>
<th>Dated material</th>
<th>Date (¹⁴C BP)</th>
<th>Age min–max (cal. BP)</th>
<th>Mean age (cal. BP)</th>
</tr>
</thead>
<tbody>
<tr>
<td>PM4/MARSH/46 m NGF</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>110</td>
<td>4.90</td>
<td>Poz-48775</td>
<td>Peat</td>
<td>865 ± 25</td>
<td>726–799</td>
<td>762</td>
</tr>
<tr>
<td>180</td>
<td>4.20</td>
<td>Poz-49814</td>
<td>Wood fragment</td>
<td>1600 ± 30</td>
<td>1411–1544</td>
<td>1477</td>
</tr>
<tr>
<td>210</td>
<td>3.90</td>
<td>Poz-48774</td>
<td>Peat</td>
<td>1870 ± 35</td>
<td>1719–1880</td>
<td>1800</td>
</tr>
<tr>
<td>PM3/MARSH/46.5 m NGF</td>
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<td></td>
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<td></td>
<td></td>
</tr>
<tr>
<td>109</td>
<td>5.41</td>
<td>SacA 26803</td>
<td>Peat</td>
<td>870 ± 40</td>
<td>676–804</td>
<td>750</td>
</tr>
<tr>
<td>190</td>
<td>4.40</td>
<td>Poz-43078</td>
<td>Wood fragment</td>
<td>1760 ± 30</td>
<td>1567–1739</td>
<td>1653</td>
</tr>
<tr>
<td>230</td>
<td>4.40</td>
<td>SacA 26804</td>
<td>Organic silt</td>
<td>1950 ± 45</td>
<td>1814–2000</td>
<td>1907</td>
</tr>
<tr>
<td>280</td>
<td>3.70</td>
<td>Poz-43080</td>
<td>Wood fragment</td>
<td>3470 ± 35</td>
<td>3680–3834</td>
<td>3757</td>
</tr>
<tr>
<td>330</td>
<td>3.20</td>
<td>SacA 26805</td>
<td>Wood fragment</td>
<td>3840 ± 45</td>
<td>4147–4412</td>
<td>4279</td>
</tr>
<tr>
<td>370</td>
<td>2.80</td>
<td>SacA 26866</td>
<td>Wood fragment</td>
<td>4460 ± 30</td>
<td>5158–5285</td>
<td>5221</td>
</tr>
<tr>
<td>PM5/INTERTIDAL ZONE/–0.5 m NGF</td>
<td></td>
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</tr>
<tr>
<td>35</td>
<td>–0.85</td>
<td>SacA 26807</td>
<td>Peat</td>
<td>4410 ± 45</td>
<td>4865–5060</td>
<td>4962</td>
</tr>
<tr>
<td>40</td>
<td>–0.90</td>
<td>Poz-48777</td>
<td>Peat</td>
<td>4480 ± 30</td>
<td>5032–5292</td>
<td>5164</td>
</tr>
<tr>
<td>100</td>
<td>–1.50</td>
<td>Poz-48776</td>
<td>Peat</td>
<td>5010 ± 40</td>
<td>5654–5778</td>
<td>5714</td>
</tr>
<tr>
<td>185</td>
<td>–2.35</td>
<td>SacA 26808</td>
<td>Organic silt</td>
<td>5445 ± 45</td>
<td>6178–6317</td>
<td>6247</td>
</tr>
<tr>
<td>195</td>
<td>–2.45</td>
<td>SacA 26809+</td>
<td>Wood</td>
<td>5890 ± 50</td>
<td>6602–6803</td>
<td>6702</td>
</tr>
<tr>
<td>220</td>
<td>–2.70</td>
<td>SacA 26810</td>
<td>Peat</td>
<td>5725 ± 35</td>
<td>6434–6636</td>
<td>6535</td>
</tr>
<tr>
<td>275</td>
<td>–3.25</td>
<td>SacA 26811</td>
<td>Peat</td>
<td>6015 ± 50</td>
<td>6743–6967</td>
<td>6855</td>
</tr>
<tr>
<td>PM1/INTERTIDAL ZONE/–0.5 m NGF</td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
<td></td>
</tr>
<tr>
<td>0</td>
<td>–0.50</td>
<td>Poz-42857</td>
<td>Wood fragment</td>
<td>4189 ± 40</td>
<td>4610–4767</td>
<td>4688</td>
</tr>
<tr>
<td>100</td>
<td>–1.50</td>
<td>Poz-42858</td>
<td>Wood fragment</td>
<td>5000 ± 35</td>
<td>5646–5769</td>
<td>5707</td>
</tr>
<tr>
<td>175</td>
<td>–2.25</td>
<td>Poz-42860</td>
<td>Wood fragment</td>
<td>5365 ± 35</td>
<td>6171–6223</td>
<td>6197</td>
</tr>
<tr>
<td>210</td>
<td>–2.60</td>
<td>SacA 26802+</td>
<td>Peat</td>
<td>5350 ± 50</td>
<td>5997–6221</td>
<td>6109</td>
</tr>
<tr>
<td>280</td>
<td>–3.30</td>
<td>Poz-42861</td>
<td>Charcoal</td>
<td>6150 ± 40</td>
<td>6946–7163</td>
<td>7054</td>
</tr>
</tbody>
</table>

NGF: Niveau Général de la France.
*Rejected dates for PM stratigraphy.

dune breakouts occurring during huge storm events (Goslin et al., 2013).

Twenty-one AMS ¹⁴C dates were obtained on PM cores (PM1, PM3, PM4 and PM5) at the Poznan (Poland) and Saclay (Paris) Radiocarbon Laboratories (Table 1). Calibration of all dates was performed using the Clam program (version 2.2; Blaauw, 2010) associated with the statistical software R (version 3.0.1; R Development Core Team, 2013) and the non-marine (IntCal09) radiocarbon calibration (Reimer et al., 2013).

Because of discontinuities recorded within peat deposits and the occurrence of a sediment hiatus recorded in the PM cores (at 270 cm in PM4, and at 220 cm in PM3; Figure 2), we did not establish a complete age–depth model for PM cores. This hiatus, which corresponds approximately to the interval 3700–2000 cal. yr BP, is recorded all over Brittany and has been recently been reported to relate to a major stormy period (Goslin, 2014). However, for the homogenous and well-dated part of the PM cores, from 6221 to 4688 cal. yr BP (PM1 core), we carried out a linear interpolation between radiocarbon points (Figure 3).

Chironomid preparation

The PM5 and PM4 cores have both been entirely sampled for chironomid analysis (5 cm interval), representing 62 and 75 levels, respectively. Sub-fossil chironomid head capsule extractions followed the laboratory procedures of Rolland and Larocque (2007). First, 60 g of sediments were deflocculated in heated 10% KOH solution for 20 min. The solution was then filtered through a 100 µm mesh sieve, and petroleum flotation was used to condense the chironomid sub-fossils in each sample. The floating fraction was then decanted and removed in a sieve similar to that used previously. The residue was washed with detergent and stored in a beaker containing a solution of 90% alcohol. Chironomid head capsules were sorted out with fine tweezers and mounted on glass slides with glycerine, ventral side up. In this study, head capsules have been identified according to Wiederholm (1983), Klink and Moller Pilot (2003) and Brooks et al. (2007). A minimum of 50 head capsules per sample were counted under a stereomicroscope at 400× magnification. It is worth noting that on the PM5 and PM4 cores, only 22 and 18 samples, respectively, contained head capsules. However, many of these samples contained less than 50 head capsules and were thus merged with subsequent samples for a complete chironomid stratigraphy. A total of 20 samples are

Figure 3. Calendar age–depth model based on ¹⁴C dates (cf. Table 1).
therefore presented in the simplified chironomid diagrams (10 samples for both PM4 and PM5 cores, Figure 4).

Pollen preparation
Rich organic sediments of PM1 (intertidal zone) and PM3 (marsh) were sampled at 9 and 3 cm intervals, respectively. Pollen extraction followed the laboratory procedures of Moore et al. (1991). A minimum of 300 pollen grains per sample were counted under a stereomicroscope at 500x magnification (oil immersion). Pollen frequencies for trees and herbs are expressed as percentages versus the main pollen sum, classically excluding aquatic plants, Cyperaceae and spores. *Alnus* pollen grains were also excluded from the main pollen sum since they are generally over-represented in alluvial plain sediments. A total of 46 samples are presented in the simplified pollen diagrams (20 samples for PM3 core and 26 samples for PM1 core, Figure 5). Selected pollen taxa have been grouped according to ecological affinities in order to facilitate the interpretation of catchment dynamics with respect to land use change (Behre, 1981).

Results
Chironomid assemblages
Chironomid diversity is illustrated by 51 and 39 different taxa within the PM4 and PM5 cores, respectively. Main chironomid taxa changes are presented in Figure 4. Based on differences in the composition of chironomid fauna, four stratigraphic changes (Pmch1-4) have been established. In both PM cores, Chironomini (e.g. *Polypedilum, Microtylodes, Chironomus*) and Tanytarsini (e.g. *Microspectra*) are the most represented tribes recorded in fossil assemblages. Among main chironomid taxa reaching percentages higher than 15%, we observe more specifically *Polypedilum nubeculosum*-type, *Chironomus plumosus*-type, *Microtylodes pedellus*-type, *Brillia modesta*, Linnothephys and *Microspectra* spp. (Figure 4). Differences between both cores are mainly attributed to individual relative abundance values and especially the *C. plumosus, P. nubeculosum*-type, *M. pedellus*-type and *B. modesta* percentages.

The lower part (Pmch1) of the intertidal core PM5 is characterized by a low number of head capsules (<50), and consequently some levels have been merged. This part of the core is marked by high percentages of *Microspectra* spp. and *B. modesta*. The second ecozone (Pmch2) is marked by an increase of total head capsules (>50) accompanied by an increase in abundances of ubiquitous and lentic taxa such as *P. nubeculosum*-type, *C. plumosus* (up to 50%), *Linnothephys* and *M. pedellus*, while percentages of the stream taxa *B. modesta* decrease. In the eozone Pmch3, the faunal assemblage is characterized by a substantial increase of *Linnothephys* synchronous with the appearance of *Heterotrissocladius grimmshawi, Gymnotricnemus, Pseudortrocladius, Metrognemus* and *Corynoneura edwardsi*. In the last eozone (Pmch4), major changes consist of the disappearance of several species, especially *B. modesta*, synchronously with the exclusive dominance of the ubiquitous taxa *P. nubeculosum*-type, with abundances reaching 80%.

The high chironomid diversity observed in PM coastal cores, with the presence of lentic, ubiquitous, loic and some terrestrial taxa, is consistent with the spatial heterogeneity and hydrological complexity which prevail in this coastal area.

Pollen assemblages
The distribution of main pollen taxa observed in both cores is displayed in Figure 5, with a diversity of 67 and 64 different taxa for the PM1 and PM3 cores, respectively. Globally, the PM1 core is relatively homogeneous in terms of vegetation, while the PM3 core is characterized by significant successive changes in terms of woodland and anthropogenic pollen grain percentages. PM vegetation dynamics between 7000 and 700 cal. yr BP (Figure 5) can be summarized as a general progression from a forested landscape dominated by arboreal taxa (Neolithic to Bronze Age), to a disturbed and cultivated landscape (Gallo-Roman period and Middle Ages).

Based on the evolution of arboreal pollen percentages, three main palynozones have been established (Pmp1-3). Based on the variation of herbaceous taxa, additional sub-palynozones have been proposed (Figure 5). Palynozone Pmp1a is dominated by high percentages of arboreal taxa (AP: 98%) such as *Corylus, Quercus* and *Ulmus*. The riverine forest mainly composed of *Alnus* is also highly represented in pollen assemblages. The following palynozone Pmp1b is marked by the appearance of some herbaceous taxa such as Apiaceae and Primulaceae, suggesting a slight opening of the forest. Palynozone Pmp2a, at about 300 cm (start of the Gallo-Roman period), shows a significant decrease of arboreal forest (50%) revealed by a sharp decrease of the dominant arboreal taxa *Corylus*. This is synchronously associated with the presence of Apiaceae, *Artemisia* and Cyperaceae. Palynozone Pmp2b is characterized by a diversification of the herbaceous taxa with the presence of Asteraceae, *Plantago*, *Rumex* and aquatic taxa such as *Potamogeton*. A clear increase of anthropogenic taxa such as *Rumex* and *Plantago* is then observed within Pmp2c, suggesting increased pastoral/agricultural activities at the studied site. The last palynozone Pmp3 is characterized, from the base to the top, by a persistent increase of *Poaceae, Galium, Cerealia* and *Cannabis*, synchronously with a collapse of tree populations, indicating an opening of the forest directly related to human activities.

Discussion
Vegetation change: General evolution from the mid-Holocene to present
The PM pollen records reflect vegetation changes between 7000 and 700 cal. yr BP with a clear gap from c. 3750 to 1910 cal. yr BP (Figure 2), probably related to a major storm period reported in northern Europe (Sørbel et al., 2012) and Western Brittany (Van Vliet-Lanoe et al., 2014b). Our data suggest that the surrounding landscape was dominated by woodlands between 7000 and approximately 4200 cal. yr BP (Figures 6 and 7), with high percentages of *Quercus* and especially *Corylus* — the latter probably occupying the drained slopes of the watershed, while *Alnus* developed on the alluvial plain and surrounding marsh. Conversely, the late Holocene from 4200 cal. yr BP to present shows a decreasing abundance of arboreal trees (Figure 5), mainly linked to agropastoral practices at the onset of the Bronze Age. These results are in agreement with archaeological and palaeobotanical observations carried out in Northern Brittany, suggesting human installation in littoral arboreal areas at the Bronze Age (Daire et al., 2011; Marguerie, 1992). Furthermore, previous studies carried out on littoral peats revealed that the whole Iron Age occurred in a context of open landscape (Morzadec-Kerfourn, 1974) in agreement with our observations from 2000 cal. yr BP (Figure 5).

At PM, the vegetation recorded during the mid-Holocene, between 7000 and 4200 cal. yr BP, is characterized by a dominance of *Corylus* (Figures 5 and 6). However, in Brittany, it is widely recognized that this period is marked by the dominance of *Quercus over Corylus* (Barbier and Visset, 1997; Joly and Visset, 2009; Morzadec-Kerfourn, 1974). *Corylus* is a thermophilous and heliophilic arboreal taxon and has a wider ecological tolerance as compared with the other thermophilous taxa (Huntley, 1993). Several palynological studies carried out over Europe showed that the *Corylus* expansion at the onset of the Holocene was related to a high seasonality signal (Naughton et al., 2007;
Figure 4. Percentages of chironomid species identified in PM cores. The bands represent the breakdown of the sandy dune.

PM: Porsmilin.
Figure 5. Percentages of pollen species identified in PM cores. The bands represent the breakdown of the sandy dune.
PM: Porsmilin.
Figure 6. Palaeoenvironmental synthesis. (a) Summer insolation values at 65°N (Berger and Loutre, 1991); (b) δ¹⁸O record from Greenland ice core (NGRIP, GICC05 timescale); (c) hematite-stained grain abundances recorded over the North Atlantic (Bond et al., 2001); (d) total solar irradiance (Steinhilber et al., 2009) (grey bands represent Bond events); (e) seasonality quantification (°C) from a pollen analysis carried out in the Bay of Biscay, France (Naughton et al., 2007); (f) Brittany storms events (Van Vliet-Lanoë et al., 2014b); (g) Northern Europe storms (Sorrel et al., 2012); (h) NAO (Olsen et al., 2012).
Tallantire, 2002) directly caused by the orbital configuration (Berger and Loutre, 1991) (Figure 5). Pollen-based temperature reconstructions carried out on a marine core from the northern Bay of Biscay, close to our studied area, showed that cold winter and warm summer conditions prevailed during the early Holocene, implying a high seasonality signal especially from 8700 to 8000 cal. yr BP, that probably favoured the expansion of Corylus (Naughton et al., 2007). After the well-known cooling event recorded at 8200 cal. yr BP, a collapse of Corylus is observed (Naughton et al., 2007). The same observation is reported from Swiss Lake (Tinner and Lotter, 2001), and the authors attribute this observation to the irreversible change in climatic conditions, consisting of a significant decrease in drought seasonal stress, which probably allowed the establishment of the other thermophilous taxa, confirming the role played by climate in vegetation changes. The persistent dominance of Corylus from 7000 to 4200 cal. yr BP in our study suggests that seasonality, which varied over this period, is not the main driver explaining arboreal vegetation recorded at PM (Figure 5). Throughout the mid-Holocene, contrary to our observations at PM, the palynological record from the Bay of Biscay (Naughton et al., 2007) displays a persistent dominance of Quercus over Corylus. This difference may be related to the fact that the Bay of Biscay marine core recorded a wider continuous regional signal throughout the Holocene while PM is a coastal core recording a more local signal of vegetation. The marine site is influenced by the huge and more meridional rivers situated along the Atlantic coast, including the Loire, Gironde and other French rivers (Naughton et al., 2007). Conversely, high percentages of Corylus recorded during the Sub-Boreal at PM may be related to local environmental conditions. Furthermore, a regional Holocene palynological study for Brittany based on a compilation of coastal pollenic spectra evidences the dominance of Corylus in the arboreal pollenic spectra during the Sub-Boreal (Gaudin et al., 2008). The spatial analysis of Corylus distribution furthermore indicates that the highest percentages of this tree characterize sites located in the most humid areas (Gaudin et al., 2008). The same result is reported from Germany where the spread of Corylus during the early Holocene mainly occurred in sites characterized by wetter conditions (Theuerkauf et al., 2014). In addition, Corylus is more adapted to an oceanic climate compared with Quercus, which is considered as being more continental, as reported from a Holocene study carried out in Scandinavia (Giesecke et al., 2008). Hence, at an orbital scale at PM, the dominance of Corylus, synchronously with significant percentages of other taxa that require moist conditions (Ulmus, Alnus; Figure 5), is probably related to prevailing humid conditions during the mid-Holocene at our studied site.

**Climate variability over the mid-Holocene: Multi-centennial to multi-decadal timescale**

It is worth noting that the whole mid-Holocene is characterized at PM by recurrent short-term contractions of Quercus coeval with a short-term increase of Corylus, superimposed on the long-term vegetation pattern described above, still dominated by Corylus trees (Figure 6). These Quercus drops are recorded around approximately 4000, 4500, 5400 and 6000 cal. yr BP and possibly 6500 cal. yr BP according to the radiocarbon pointers we have. The same anti-correlation pattern between these two taxa has been reported in the Armorican domain over the mid-Holocene and has been related to human impact (Joly and Vissot, 2009). At PM, the absence of anthropogenic taxa through the PM1 core leads us to consider an alternative hypothesis and to examine more deeply the possibility of climate forcing. On the other hand, it is well known that several cold millennial-scale events, or Bond events, have been recorded in the North Atlantic throughout the Holocene and are
attributed to variability in solar activity and/or internal ocean-atmosphere mechanisms (Bond et al., 2001; Giradou et al., 2010, Mayewski et al., 2004; Moros et al., 2004; Nesse et al., 2004; Wanner et al., 2008). These major periods of drift ice measured in deep-sea sediment cores (Bond et al., 2001) are also the basis of significant reduction of North Atlantic Deep Water production (Bond et al., 2001; Mayewski et al., 2004). Furthermore, palaeohydrological conditions reconstructed from peat bogs in Britain (Langdon et al., 2003) suggest that Holocene Bond events were synchronous with humidity increases and associated humification in the peat record. As for investigated areas in Southern Europe, repeated Mediterranean Forest collapses during the Holocene have been observed at multi-centennial timescales in the Iberian margin (Chabaud et al., 2014) and the Alboran Sea (Fletcher et al., 2010), reflecting repeated dryness and/or cold conditions at multi-centennial timescales related to Bond events. They were interpreted as decreases in the North Atlantic subtropical gyre sea surface temperature (SST) driven by freshwater inputs and thermohaline circulation slowdown. We therefore also observe in Western Britain the impact of the Holocene North Atlantic SST on the coastal arboreal vegetation through our observed Quercus drops at PM and increases in humidity. Furthermore, we also find that short-term Quercus contractions–Corylus expansions coincide with major Holocene storm events reported for Northern Europe (Sorrel et al., 2012) and Western Britain at the Holocene scale (Van Vliet-Lanoë et al., 2014a) (Figure 6). According to Sorrel et al. (2012), storm events at that time may have been induced by a global reorganization of North Atlantic oceanic and atmospheric configurations, resulting in a southward shift of the subtropical gyre and in a major change in the track and strength of the westerlies. This would also be consistent with our observed vegetation and faunal changes as mentioned above.

The time resolution of our palynological record during the period between 6200 and 5000 cal. yr BP (70 years on average) allows us to describe more precisely the two specific shifts centred on 6000 and 5400 cal. yr BP (Figure 7). We then identify contractions of Quercus synchronously with SST decreases and sea ice cover increases North of Iceland (Solignac et al., 2006). We can therefore argue that the significant and recurrent contractions of Quercus during the millennial climatic cold event 6–5 kyr BP would be attributed to more humid, colder conditions and more frequent storms at that time (Figure 7). Our data would therefore suggest that Quercus–Corylus trends recorded at PM over the mid-Holocene are not an anthropic signal (Joly and Vissset, 2009) but instead result from climate shifts that prevailed over the Northern Hemisphere, implying atmospheric and oceanic reorganizations.

Land use over the Late Holocene (Bronze Age to Middle Ages)

Major vegetation changes in PM can be observed at the transition between mid- and Late Holocene (PMP1-PMP2 limit around 300 cm, also equivalent to the Sub-Boreal-Sub-Atlantic limit, Figures 5 and 6). This change is characterized by a shift from a wooded landscape dominated by Corylus–Quercus to a cleared landscape (PMP2, Gallo-Roman period; Figures 5 and 6). Until the end of PMP2, any signal of cereals stays relatively weak as suggested by low percentages of Cerealia and Humulus/Cannabis. The good representation of Plantago lanceolata during PMP2-3 (Figures 5 and 6) probably marks both compaction and nitrogen enrichment of soil by increasing livestock activities. The persistence of a well-developed forest landscape at the beginning of the Bronze Age (4500 cal. yr BP) at PM (end of PMP1 palynozone, Figure 5) is in good agreement with previous studies carried out in the Armorican domain, suggesting the late apparition of agricultural activities (Gaudin, 2004), as we observe just before 750 cal. yr BP (cf. PMP3c; Figures 5 and 6).

Then, a regional decrease of agro-pastoral activities is observed from about 1800 to 1653 cal. yr BP at PM, together with a slight but significant forest recovery (Corylus–Quercus) accompanied by alluvial trees (Alnus). This partial abandonment of agricultural practices has previously been observed in the Armorican domain (Barbier and Visset, 1997, Visset, 1994) and has been related to an economic decline all over Europe during the Roman Empire (beginning of the 3rd century AD). Climatically, the 3rd century is also reported as being colder and drier in Western Europe (McCormick et al., 2012) as also marked by the onset of historical dune development in Western Brittany (Van Vliet-Lanoë et al., 2014a). These climate variations disrupted food production over decades and arboreal vegetation was restored during these dark periods.

After this agricultural regression observed at PM, a major new clearing event is observed after 1653 cal. yr BP (start of PMP3 palynozone, Figure 5) on the watershed, with a significant decrease of Corylus–Quercus marking the start of the Middle Age (Figure 5). This major forest drop is followed by the decrease of Alnus around 200 years later and the anthropization reached its maximum before 750 cal. yr BP (i.e. around the 12th–13th century AD) with a maximum of Cerealia and Cannabis recorded at PM (PMP3c; Figures 5 and 6).

Impact of clearance events on the insect fauna

Regarding the Sub-Atlantic period, the decrease of Alnus at PM is mainly related to clearing events associated with human settlement, including (1) direct cutting in order to develop farming activities on alluvial plains and (2) increases in water supply caused by deforestation and agricultural use of the surrounding dry land areas (Barthelmes et al., 2010). Several studies around Europe have previously shown the relationship between Alnus declines and human settlement during the late Sub-Atlantic period in England (Brown, 1988), Finland (Sarmaja-Korjonen, 2003) and in the eastern Baltic sea (Saarse et al., 2009).

Simultaneous faunal declines associated with waterlogged forests, such as B. modesta (Cranson et al., 1983), are evidence that the floodplain deforestation by cultivators, and its transformation into humid meadow by drainage, lead to faunal habitat loss. Disappearances of insects related to vegetation changes have been observed several times, allowing for the detection of human impact (Andrieu et al., 1997; Ponel et al., 2007). A previous insect-based study carried out in Northern France on peat bogs evidenced a disappearance of tree-dependent insects from 4000 cal. yr BP (Gandouin and Ponel, 2010). Major ecological changes related to large-scale forest clearance have also been attributed to the spread of cultivation and pastoralism during the Neolithic period. At Saint-Omer (Northern France), pollen data also show evidence of deforestation post 4000 cal. yr BP (Gandouin et al., 2009) while, before 4000 cal. yr BP, the coleopteran record did not contain any species associated with cultivated grounds. On the contrary, several coleopteran species typical of pristine, undisturbed primeval forests were identified (Ponel et al., 2007). These results suggest that insects can be used as clearing event indicators. In the case of PM, we suggest that B. modesta, which is reported to prosper in woods, can also be used as an indicative species for the development of undisturbed alluvial forests.

Faunal changes: General evolution from the mid-Holocene to present

Chironomids are sensitive to several factors such as temperature, oxygen, trophic level and precipitation. At PM, chironomid
changes are not directly related to temperature, since quantitative reconstructions show that there is no significant correlation between chironomid distribution and temperature. In floodplain areas, according to Gandouin et al. (2005), some chironomids can be qualitatively assigned to different ecological conditions related to current speed (lentic versus lotic conditions), and hence allow for reconstruction of hydrodynamic conditions. At PM, variations of lotic- and lentic-based taxa show that the level of connection of the studied site to the river has probably impacted faunal assemblages. During the mid-Holocene, from about 6900 to 6600 cal. yr BP, high percentages of lotic taxa such as B. modesta, Chaetocladius and Micropsectra spp. (Figure 6), synchronously with low percentages of Limnophyes which presently occurs in very shallow water (Hofmann, 1998) characterized by low precipitations (Hofmann, 1998; Massafero and Brooks, 2002), indicate that the site was humid and that river run off was high at that time. Later, from about 6000 to 5700 cal. yr BP, higher percentages of lentic taxa such as C. plumosus-type and M. pedellus-type, associated with a slight decrease of B. modesta percentages may suggest a slight decrease of river run off, probably related to lower precipitation as suggested by the slight increase of Limnophyes. This hypothesis is confirmed by the decrease of Corylus (Figures 5 and 6). Many of the changes observed in chironomid assemblages can be also related to eutrophic conditions, as reported for C. plumosus-type and P. nubeculosum-type. Between 6000 and 5700 cal. yr BP, lentic taxa increases such as C. plumosus-type and ubiquitous taxa such as P. nubeculosum-type could then be related to eutrophic changes. However, changes in vegetation observed at that time, illustrated by a decrease of the humid taxa Corylus, exclude the trophic level as a trigger factor.

Regarding the upper core, the period from 1800 to 1000 cal. yr BP is marked by a collapse of lotic taxa, especially B. modesta (Figures 4 and 6), associated with an increase of lentic taxa such as Limnophyes and ubiquitous taxa such as P. nubeculosum-type. This suggests a probable decrease of river run off related to agricultural activity, as suggested by the increase of anthropogenic taxa such as Plantago, Rumex and Cannabis (Figures 5 and 6). Between 1000 and 750 cal. yr BP, the increase of P. nubeculosum-type associated with the disappearance of several chironomid taxa (Figure 4) is contemporaneous with the expansion of agricultural activity in the region as suggested by the increase of Ceratia and Cannabis (Figures 5 and 6). The chironomid signal detected throughout PM4 then indicates that human disturbance is present and has certainly impacted faunal assemblages. Changes of current speed observed during this period may then directly be related to the development of agricultural activities; natural areas have been embanked and drained, inducing a loss of chironomid habitats.

Conclusion
This present work represents the first study using combined Holocene chironomid and pollen analysis in a coastal marsh environment in Brittany (PM studied site, Northwestern France). We demonstrate the great potential of this sedimentary context to record widespread climatic variation over the mid-Holocene (7-4.2 cal. kyr BP) as well as anthropization forcing on the Late Holocene (from 4.2 kyr BP to 750 cal. yr BP). Dense arboreal mixed forests, composed of thermophilous taxa (Corylus—Quercus), clearly prevailed over the mid-Holocene with a clear dominance of Corylus, related to the humid oceanic conditions prevailing at PM. Furthermore, superimposed on the long Holocene climate trend, the well-dated and resolved pollen results between 6 and 5 kyr BP reveal a short-term anti-correlation between slight Corylus expansions and Quercus contractions at 6, 5.8 and 5.4 cal. kyr BP, with a probable signature of oceanic and atmospheric reorganizations at that time. Concerning chironomid records, our data show variations attributable to river flooding and temperature fluctuations, probably related to climate changes over the mid-Holocene. Regarding the late Holocene (Sub-Atlantic), starting with the late Bronze Age, vegetation and chironomid changes appear to be related to human activities as suggested by synchronous high abundances of pollinic anthropogenic taxa and decline in chironomids associated with alluvial forest.

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