



Variability of summer precipitation over eastern China during the last millennium

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Variability of summer precipitation over eastern China

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Variability of summer precipitation over eastern China during the last millennium

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We use measurements of recent decades, 1500-yr proxy data, and millennium model simulations with a variety of climate facings to study the temporal and spatial variability of summer precipitation over eastern China. Spectral analysis of the proxy data 5 using multi-taper method reveals three statistically significant bidecadal (15–35-yr), pendadecadal (40–60-yr), and centennial (65–170-yr) oscillation bands. The results of wavelet filtering show that the amplitudes of these bands vary substantially through time depending on the temperature regimes. Weak centennial oscillation and strong pentadecadal oscillation occur in warm conditions, whereas the oscillations are both 10 strong in cold conditions. A model/data intercomparison suggests that the centennial oscillation might be linked to the fluctuation of solar forcing (Gleissberg cycle), and the pentadecadal and bidecadal oscillations could be associated with internal variability of the climate system. It is also found that the increased frequency of drought-in- 15 north/flood-in-south spatial pattern over eastern China during the last two decades is unusual in the past five centuries.

1 Introduction

The Earth's climate has significantly warmed in the last few decades, most likely due to anthropogenic forcing (IPCC, 2001; 2007; Jones et al., 2001). In order to assess the impacts of climate change under global warming on the human societies and natural ecosystems and to develop suitable adaptation and mitigation policies, a better 20 understanding of climate variability under "natural" and "anthropogenic" conditions and accurate climate change predictions are needed at the global and, more importantly, the regional and local scales (Bradley et al., 2003; Jones and Mann, 2004; Giorgi, 2005). Changes in globally average climate are primarily determined by the radiative budget of the coupled atmosphere-ocean-land system and modulated by internal non- 25 linear feedbacks (IPCC, 2001; Giorgi, 2005). A positive global radiative forcing such

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as greenhouse effect generally causes global warming and an increase of global precipitation as a consequence of increased evaporation from the warmer land and ocean surfaces (Allen and Ingram, 2003). However, climates, especially precipitation, are primarily controlled by circulation regimes and internal modes of variability at the regional scale. For example, the Asian monsoon and its interactions with the El Nino/Southern Oscillation (ENSO) and extratropical circulations are the major sources of interannual variability of summer precipitation over Asia (Wang, 2006). Therefore, the study of variability in seasonal precipitation at the regional scales is vital to providing detailed picture we need in our research for an unambiguous “fingerprint” of the climate response to increasing greenhouse gas emissions (Barnett et al., 1999; Jones et al., 2001).

Eastern China is located in the monsoonal Asia, which is vulnerable to small changes in the timing and intensity of monsoon precipitation. The vulnerability is likely to increase in the future with continued population growth and intensified land-use (Feddema et al., 2005; Zhao et al., 2006). The monsoon system is highly variable over a broad range of timescales from days to millennia. We have come to know a great deal about Asian monsoon dynamics on seasonal to interannual time scales due to an increased wealth of new data from satellite observations and field experiments, and the advances in computing power and mathematical representations of coupled climate system (Wang, 2006). For example, It has been known that the ENSO event and Quasi-biennial Oscillation (QBO) have been the primary source of the precipitation over East Asia on seasonal to interannual time scales (Ding, 1991; Waliser, 2006; Yang and Luo, 2006). It is also known that astronomical variations in seasonal insolation have been the primary pacemaker of the Asian monsoon on glacial-interglacial time scales (e.g. Clemens et al., 2003), and a 1500-yr cycle in Holocene monsoon dynamics has been driven by solar activities (e.g. Gupta et al., 2005). However, the patterns and causes of decadal-centennial variability are not well understood due to the limits in the length of observational records and relatively coarse resolution of proxy data. Although work to date has revealed many strong hints of decadal to quasi-centennial oscillations in the

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summer precipitation over eastern China using relatively short observational data and proxy data (e.g. Zhu and Wang, 2002; Ding et al., 2007), the behavior of decadal to centennial variability in time and space is still unclear in the context of millennia. In this study, we analyze the 1500-yr proxy data of summer precipitation to examine long-term variability with a focus on decadal to centennial oscillations.

Some studies using observational and modeled data have indicated an increase in the frequency of drought-in-north/flood-in-south over eastern China during the last few decades (Xu, 2001; Ding et al., 2007). It has been ascribed to the response of summer precipitation to anthropologic forcing (Xu, 2001; Menon et al., 2002; Fu, 2003). However, question is how unusual this variability of summer precipitation in the last few decades in the millennium context of climate. The last millennium is an epoch including Medieval Warm Period (MWP) with climate as warm as the last century. It also encompasses the period before large-scale contamination of the global atmosphere by human activities and global-scale changes in land-surface condition (Bradley et al., 2003; Jones and Mann, 2004). Therefore, the study of the variability in summer precipitation over eastern China during the time interval longer than observational record is essential for assessing the temporal and spatial variability of summer precipitation under “natural” and ‘anthropogenic’ condition and for understanding the role of anthropogenic forcing in future precipitation variations. In this study, we analyze observational, proxy, and modeled data to examine the occurrence probability of different spatial patterns over eastern China during the instrumental and historical times.

Theoretical models of the climate system driven with external forcing can provide important insights into the factors governing climate changes in the past (Jones and Mann, 2004). Simulations of climate changes over the last millennium have been performed using energy balance models (EBMs, e.g. Crowley, 2000), two-dimensional climate models (e.g. Bauer et al., 2003), and general circulation models (GCMs, e.g. Ammann et al., 2007). However, studies of model/data intercomparison mainly focused on temperature. There has been little attempt to test climate models with precipitation proxy data or interpret such proxy data using model simulations (Yoshimori et al., 2006).

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In this study, the millennium simulations driven with natural and anthropogenic forcing and fixed forcing (control run) using CCSM2.0 are used to analyze the variability of summer precipitation over eastern China during the last millennium. The comparisons of modeled, observed, and proxy data are conducted to assess the performance of CCSM2.0 in simulating the spatial patterns and temporal variability of summer precipitation over eastern China, and to provide some potential insights into the factors influencing temporal and spatial variability of summer precipitation over eastern China during the last millennium.

2 Data and methods

- The data used in this study include observational, modeled, and proxy data of summer precipitation over eastern China (Table 1). Observational precipitation data we analyze are from the Chinese National Climate Center (CNCC) and the Climate Research Unit of the University of East Anglia. The former is a data set of precipitation from 1951 to 2000 at 160 stations over China, and the latter is a data set of land-only precipitation in a $3.75^\circ \times 2.5^\circ$ resolution covering a period of 1901–2000. Proxy data used in this study is a dataset of the dryness/wetness index (DWI), a proxy data of summer rainfall, over eastern China from 1470 to 2000, derived from Chinese historical documents and instrumental measurements (CNMA, 1981; Zhang et al., 2003). Additionally, two 1500-yr time series of regional DWI covering two regions, i.e. North China (NC, $34\text{--}41^\circ \text{N}$, $107\text{--}120^\circ \text{E}$) and the middle-lower Yangtze River Valley (MLYRV, $26\text{--}34^\circ \text{N}$, $109\text{--}122^\circ \text{E}$) over eastern China (Zheng et al., 2006) are also analyzed. Model data are three millennium simulations of NCAR_CCSM2.0, including a control run, a run with orbital and solar forcing, and another run with full forcing. Forcing time series for solar activity, volcanic activity, anthropogenic greenhouse gases and aerosols used in the model runs of millennium simulation of climate are from the EBMs and GCMs studies (Crowley, 2000; Crowley et al., 2003; Ammann et al., 2007).

In this study, principal components analysis (PCA) is employed to reveal the spatial

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patterns of summer precipitation over eastern China. Multi-taper method (MTM) (Ghil et al., 2002) and wavelet analysis (Torrence and Campo, 1998) are applied to detect decadal to centennial oscillations of summer precipitation and their temporal evolution during the past 1500 yr over eastern China.

5 3 Results and discussion

3.1 Variability of summer precipitation in observational data

Eastern China is located in the East Asian monsoon region. The summer monsoon brings about 70% of the total annual precipitation to this region. On average, the summer monsoon onset occurs in early to mid-May when heavy convective rainfall develops over southern China along the pre-Meyu front. It is then followed by abrupt northward shifts of the summer monsoon, causing an elongated rain belt referred to as Meyu or “Plum Rains”, a regional rainy season in June and July over the Yangtze River Valley in eastern China. The summer monsoon then shifts to northern China in July, triggering the rainy seasons of northern China. At the end of August, the summer monsoon begins to withdraw southward (Ding, 1991; Samel et al., 1999; Chang et al., 2000). Following this summer monsoon northward penetration, the summer precipitation over eastern China exhibits a significant gradient from south to north.

Chinese National Climate Center has developed a dataset of monthly temperature and precipitation from 1951 to the present. It consists of 160 stations, covering the most land areas of China. Here we selected a subset of summer precipitation consisting 37 stations over eastern China (Fig. 1) to conduct PCA. The result of PCA shows that the first two PCs explain 19.4% and 15.2% of the total variance in this dataset. PC1 indicates a three-belt mode of spatial pattern, in which NC and southeastern coast (SEC) share the same precipitation conditions (either increase or decrease), and the MLYRV experiences the opposite conditions. PC2 reveals a dipole mode of spatial pattern. When NC experiences a decrease in summer precipitation, the MLYRV and

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SEC witness an increase in summer precipitation. It is evident that the result of PCA shows a distinct difference of summer precipitation variability in NC and the MLYRV during the past 50 yr. Therefore, we choose NC and the MLYRV as our study regions to explore the tempo-spatial variability of summer precipitation during the instrumental and historical times over eastern China. Another reason why we choose them as our study regions is that 1500-yr proxy data of summer precipitation is available only in these two regions (Zheng et al., 2006).

During the last century, large interannual variability in summer precipitation is observed in both regions (Fig. 2a, b). In the MLYRV, average monthly summer precipitation varied from 110 mm (1900) to 230 mm (more than 2 standard deviation, 1958), while it varied from 55 mm (1900) to 113 mm (1964) in NC. Spectral analysis of these two time series reveals some significant cycles. In the MLYRV, 2-yr cycle is significant at the 95% confidence level, whereas 5-yr, 7-yr, and 26–27yr cycles are significant at the 90% confidence level (Fig. 2c). In NC, 3-yr, 5-yr, and 7-yr cycles are significant at the 95% or 90% confidence levels. A number of studies have indicated that the interannual variability in summer precipitation over the Asian monsoon region is associated with the tropospheric biennial oscillation (TBO, Shen and Lau, 1995; Webster et al., 1998; Yang and Lau, 2006), ENSO events (Wang et al., 2000; Lau and Wu, 2001), and snow cover (Wu and Qian, 2003; Yang and Lau, 2006). 2–3-yr and 5–7-yr cycles revealed by summer precipitation in the MLYRV and NC are typical TBO and ENSO signals. Decadal variability in summer precipitation in NC is not as significant as in the MLYRV, where a bidecadal oscillation exhibits at the 90% confidence level. Wavelet filtering of the time series shows that decadal-centennial band oscillations only explain 19.4% and 18.3% variance of summer precipitation in the MLYRV and NC, respectively.

3.2 Variability of summer precipitation in proxy data

Proxy data of summer precipitation used here are DWI derived from the Chinese historical documents. This proxy data covers the past 1500 yr. Due to that 10-yr moving average was involved in developing regional DFI for the MLYRV and NC, high frequency

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the last two millennia in China occurred around AD 1230. It seems reasonable to speculate that this transition in the phase relationship of centennial oscillation between NC and the MLYRV could have been caused by a major shift of climate over China in 12th or 13th century. A comparison of reconstructed temperature from multi-proxy data in China (Yang et al., 2004) with our results further supports our speculation. The warm period (so-called Medieval Warm Period) ended around AD 1100 in China. This comparison also reveals an interesting fact that dominant long-term oscillation modes varied during the last 1500 yr under different temperature conditions. During the warm periods such as the Medieval Warm Period (AD 800–1100) and 20th century, the centennial oscillation was weak and the pentadecadal oscillation was strong. During the cold periods such as AD 500–800 and AD 1400–1700, both the centennial and pentadecadal oscillations were strong, although the pentadecadal oscillation in the MLYRV was not as strong as that in NC during the former cold period. During the normal conditions such as periods of AD 1100–1400 and AD 1700–1900, the centennial oscillation was relatively strong whereas pentadecadal oscillation was very weak. The correlation of temperature conditions with the bidecadal oscillation is not as clear as that with the centennial and pentadecadal oscillations, although strong signals of the bidecadal oscillation more frequently occurred in cold or normal conditions.

3.3 Variability of summer precipitation in modeled data

Here we present the results of millennium simulations using the version of NCAR-CCSM2.0. Three experiments including an AD 1000 control run, a run with orbital forcing and solar forcing, and another run with full forcing, were carried out in this study. In the AD 1000 control run, the model was first run fully coupled for 300 yr in a spin up mode with constant 1990 conditions. The run was then continued from year 300 without any forcing changes (fixed at AD 1000 forcing conditions, i.e. solar forcing is 1365.0 W.m^{-2} , the concentration of CO_2 , CH_4 , N_2O , CFC-11, and CFC-12 are 280, 680, 266, 0, and 0 ppbv, respectively) and was run fully coupled for 1000 yr. The forced simulations were started from fully coupled 300 yr. The orbital parameter is almost

constant during the last 1000 yr; it is thus neglected in our interpretations of model data.

Figure 5 shows the spatial patterns of summer precipitation (May–September) over East Asia ($10\text{--}50^\circ\text{N}$, $90\text{--}130^\circ\text{E}$) in observed and simulated data from 1950 to 1995. 5 Generally, CCSM2.0 underestimates the summer precipitation in the regions south to 30°N , whereas it overestimates in the regions north to 30°N , especially the Sichuan Basin. This might be an indication that CCSM2.0's performance in simulating the effect of the Tibetan Plateau on precipitation needs to be improved. However, the spatial pattern of summer precipitation in the regions beyond the Tibetan Plateau, such as 10 eastern China, is better simulated in the three experiments, although the summer precipitation is obviously underestimated in the southeastern costal region. Nevertheless, the summer precipitation is well simulated in our study regions, the MLYRV and NC, where a zonal gradient of summer precipitation similar to observed pattern exhibits significantly in the three experiments. Figure 6 shows a comparison of distribution 15 statistics (full range, 25th, 50th, 75th quartiles and the mean) in observed and modeled summer precipitation from 1951 to 1995 in the MLYRV and NC. The simulated means and medians of summer precipitation in three experiments are rather stable and they are close to the observed ones in the MLYRV. The dispersion ranges of simulated summer precipitation for this period in the MLYRV are also comparable to that from 20 observation. The simulated means and medians in NC are about 20–30 mm higher than observed ones. The dispersion range for NC in the experiment with full forcing is smaller than other two experiments and it is close to the observed one. Although the difference in modeled summer precipitation between the two regions is not as distinct 25 as observed one, one-sample T-test indicates that the means of summer precipitation in the two regions are significantly different at the 99.99% confidence level, suggesting that these two regions apparently belong to different regimes of summer precipitation as indicated by observed data. Therefore, it seems reasonable to use these modeled data in these two regions to determine the variability of summer precipitation in millennium simulations and then to further compare with proxy data.

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icant centennial band (65–170-yr) oscillation shows up in the control run. In the solar forcing run, the dominant oscillations with the 95% confidence level in the MLYRV are 16–17-yr, 22-yr, 27–35, and 105–170-yr periodicities, whereas they are 17-yr, 25–29-yr, 30-yr, 37–41-yr, 68–90-yr, and 170–260-yr cycles in NC. Additionally, there are two
5 cycles at 50–60-yr and 65–80-yr significant at the 90% confidence level in the MLYRV. In the full forcing run, the major oscillation modes for the MLYRV are 17–21-yr, 22–25-yr, and 37–41-yr variations, which are significant at the 95% confidence level. A 50–72-yr cycle with the 90% confidence level is also seen in the MLYRV. In NC, significant oscillation modes include 19–21-yr, 23–27-yr, 29-yr, and 56–80-yr periodicities.
10 Theses modes are significant at the 95% confidence level.

As mentioned above, three dominant cyclic modes revealed by proxy data are bidecadal band (15–35-yr), pentadecadal band (40–60-yr), and centennial band (65–170-yr) oscillations. Centennial band oscillation is not found in the control run, indicating that it is not an internal variability. This oscillation is clearly visible in the solar-forcing and full forcing runs, especially in the solar-forcing run. This suggests that in broad terms the model correctly captures the impact of solar forcing on long-term variation in summer precipitation over eastern China, although peaks in the centennial band oscillation revealed by proxy do not match those in the two simulations with solar forcing and full forcing. Since there are chaotic components of internal variability in models
15 and some uncertainties in the forcing reconstructions, it is not realistic to anticipate that models reproduce the exact variability registered in proxy data. Therefore, this centennial oscillation could be linked to external forcing. The most potential pacemaker of this centennial oscillation in summer precipitation is the century-type periodicity known as Gleissberg cycle in solar forcing, as suggested by some previous studies (e.g. Zhu
20 and Wang, 2002). The change in the length of this variation in proxy data (from 68 to about 170 yr) may reflect the respective changes in the length of the Gleissberg cycle, which has been considered not constant and varies from about 65 to more than 130 yr (70–80-yr and 100–120-yr modes dominate) (Ogurtsov et al., 2002). The behavior of
25 the amplitude of this centennial oscillation shows significant episode changes in proxy

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coupled ocean-atmosphere mode of oscillation. These findings support our interpretation of this pentadecadal oscillation as an internal variability of the ocean-atmosphere system.

The bidecadal band oscillation (15–35-yr) in proxy data consists of a series of oscillation modes. Dominant modes include 15–19-yr, 21–29-yr and 31–35-yr periodicities. These oscillation modes can be found in both control run and forced runs. Thus it is difficult to determine whether it is an internal or external variability. Previous studies suggested that the decadal-interdecadal oscillations within bidecadal band found in the Pacific sea-surface temperature, ENSO events, and PDO are largely internal, resulting from the ocean-atmosphere interaction and the tropical-extratropical interaction (Zhang et al., 1997; Nakamura et al., 1997; Minobe, 1999; Evans et al., 2001; Liu et al., 2002; Deser et al., 2004; Schneider and Cornuelle, 2005; Nonaka et al., 2006). Therefore, the bidecadal band oscillation in our proxy and model data is likely an internal variability.

15 3.4 Variability in spatial pattern of summer precipitation in proxy and modeled data

As mentioned in the introduction, some studies using relatively short observed data (several decades to one century) indicated an increase in the frequency of drought-in-north/flood-in-south over eastern China during the last few decades (Xu, 2001; Ding et al., 2007). It is evident that records longer than multi-decade are necessary to examine 20 whether this conclusion is solid in the millennium or multi-century context of climate. Here we use both proxy and model data to examine the dominant spatial patterns of anomalous summer precipitation events over eastern China and to determine how unusual this spatial pattern of summer precipitation recorded in the last few decades. The 25 annually-resolved proxy data of summer precipitation used here is the area-weighted regional DWI developed from the network of DWI over eastern China (Shen et al., 2007). It covers the last 530 yr. Model data is from the full forcing run. Both data are standardized using the means and standard deviations of 1951–1995.

Figure 9 is 4-quarter plot of proxy and model data showing the spatial patterns of

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summer precipitation in the MLYRV and NC. We use one standard deviation as the criterion to define an anomalous summer precipitation event. When the standardized DWI values in two regions are within the range of one standard deviation, this year witnesses a normal summer in precipitation. As the standardized DWI value in any region is beyond the range of one standard deviation, this year is defined as an anomalous summer precipitation event. Its spatial pattern is then assigned through comparing the standardized DWI values in the two regions. Two modes, coherence and dipole consisting four spatial patterns, are then defined. Coherence mode includes coherent drought and coherent flood, whereas the dipole mode includes drought-in-north/flood-in-south and flood-in-north/drought-in-south. After this procedure, we use a moving 20-yr window to calculate the occurrence probability of four spatial patterns of anomalous summer precipitation events and normal conditions.

The distribution statistics of their occurrence probability in 530-yr proxy and model data is shown in the lower panel of Fig. 9. As expected, normal conditions have the highest occurrence probability in both proxy and model data. During the past 530 yr, the most frequent anomalous summer precipitation event was coherent flood with a mean of 0.18. The second one was coherent drought with a mean of 0.1. Two spatial patterns of dipole mode occurred less frequently than the spatial patterns of coherence mode. There are some extreme values in the occurrence probability of coherent drought, and two dipole patterns. These extreme values indicate some time intervals with extremely high occurrence probability of these spatial patterns. The periods with extremely high probability (0.35 or 0.30) of coherence drought were AD 1481–1490 and 1927–1942. The high frequency of drought-in-north/flood-in-south occurred in AD 1555–1565, 1601–1610, and 1980–2000 (with probability of more than 0.25). The flood-in-north/drought-in-south was the most frequent anomalous summer precipitation event in AD 1645–1654 and 1955–1964. For the last two decades, the occurrence probabilities of coherent drought, coherent flood, and flood-in-north/drought-in-south are close to their means of the last 530 yr. The probability of normal conditions is markedly lower than its means, close to a statistically significant level, indicating

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that anomalous summer precipitation events happened more frequently in recent two decades. The probability value of drought-in-north/flood-in-south occurs as an outlier, i.e. this value is significantly different from its mean of the last 530 yr at the 99% confidence level. It means that this variability in spatial pattern of summer precipitation during the last two decades is unusual even in the five-century context. Model data seems to support observed result, although there is somewhat difference in the probability distribution of normal conditions and four spatial patterns of anomalous summer precipitation events. For the last two decades, a decrease in the occurrence probability of normal conditions and an increase in the probability of drought-in-north/flood-in-south are visibly apparent, although statistically they are not significantly different from their means.

4 Concluding remarks

This study analyzes observational, proxy, and model data to investigate the variability of summer precipitation over eastern China during the last millennium with a focus on the middle and lower Yangtze River Valley and North China. On the interannual scale, 2–3-yr and 5–7-yr cycles, typical TBO and ENSO signals are well presented in observational data over the MLYRV and NC. Oscillations on the decadal scale are not as distinct as these on the interannual scale in the instrumental times. Spectral analysis of regional proxy data of summer precipitation reveals three statistically significant oscillation bands, i.e. bidecadal (15–35-yr), pendadecadal (40–60-yr), and centennial (65–170-yr) band during the past 1500 yr. The time evolutions of these bands revealed by the wavelet filtering indicate that the amplitudes of these oscillation bands are not constant through time. They vary substantially from one period to another over the past 1500 yr. A comparison of the behaviors of these oscillation bands with temperature conditions in China shows that these periods distinguished by the temporal evolutions of oscillations corresponds those episodes with different temperature conditions very well. Weak centennial oscillation and strong pentadecadal oscillation occurred in warm

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conditions during the MWP and 20th century. During the cold periods of AD 500–800 and AD 1400–1700, both the centennial and pentadecadal oscillations were strong. When the temperature was in relatively normal conditions, pentadecadal oscillation was very weak while the centennial oscillation was relatively strong. Our analyses also show that the phase relationship in centennial oscillation between NC and the MLYRV was not constant during the past 1500 yr. It changed from “in-phase” to “out-of-phase” at about AD 1100. It may suggest a major shift of climate in China around AD 1100.

A comparison of observational data with modeled data from millennium simulations using NCAR-CCSM2.0 shows that the model correctly reproduces the tempo-spatial patterns of summer precipitation over East Asia in broad terms. Analyses of model results indicate that the model also captures the impact of climate forcing on long-term variation in summer precipitation over eastern China. A model/data intercomparison suggests that the centennial oscillation might be linked to the fluctuation of solar forcing (Gleissberg cycle), and the pentadecadal and bidecadal oscillations be associated with internal variability of climate system. However, further analyses on the model results such as ocean and land temperature, atmospheric circulations, and moisture transportation are needed to separate the factors governing these oscillations and better understand underlying mechanisms.

In this study, we also analyze the occurrence probability of spatial patterns of anomalous summer precipitation events using proxy and model data. Our analyses on both model and proxy data demonstrate that an increase in the frequency of drought-in-north/flood-in-south over eastern China during the last two decades is unusual in the past five centuries. The study of Menon et al. (2002) suggested that this trend might be related to increased black carbon aerosols.

Acknowledgements. This work is supported by a grant from the Office of Biological and Environmental Sciences, Office of Sciences, US Department of Energy (DOE) to SUNY-Albany; and the China Meteorological Administration, Chinese Academy of Sciences, the National Natural Science Foundation of China. The millennium simulations of climate were conducted by Y. Peng as a part of her Ph.D dissertation.

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Table 1. Observed, modeled, and proxy data of summer precipitation over eastern China.

Data Sets	Spatial resolution	Temporal resolution	Length	Reference
CNCC_dataset	point	Monthly	1951–2000	This study
UK CRU Land Surface	3.75×2.5	Monthly	1850–2000	Hulme, 1992, 1994
D/W index	Point	Seasonal	1570–2000	CNMA, 1987
Regional D/W index	Area	Seasonal	505–1995	Zheng et al., 2006
NCAR-CCSM 2.0	T31L26	monthly	1000–1999	This study

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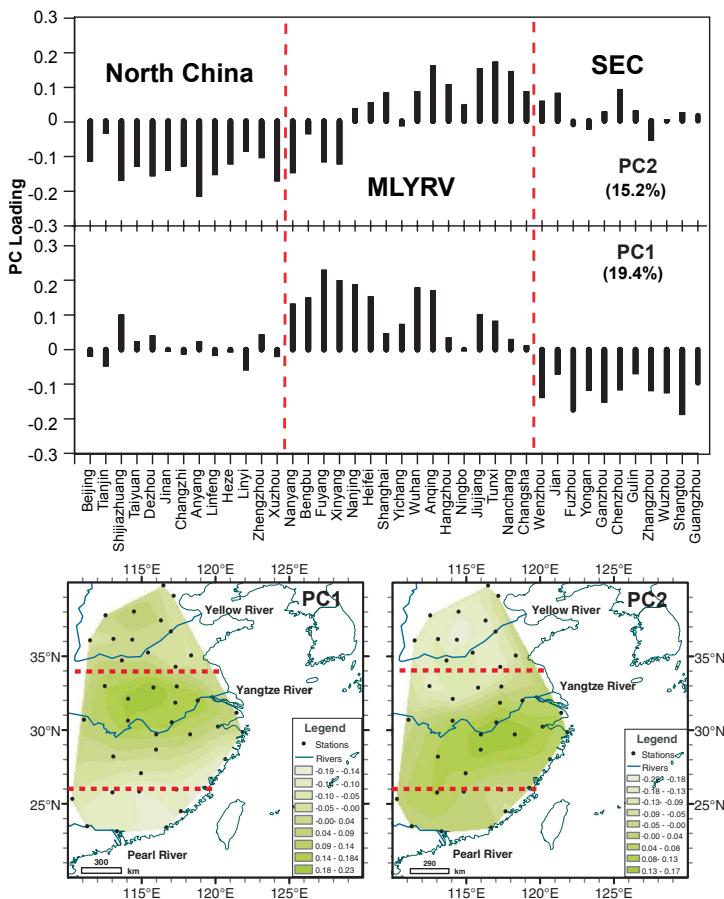


Fig. 1. Result of PCA on the observed data from 1951 to 2000 consisting of 37 stations over eastern China. Location of stations and two study regions are also shown.

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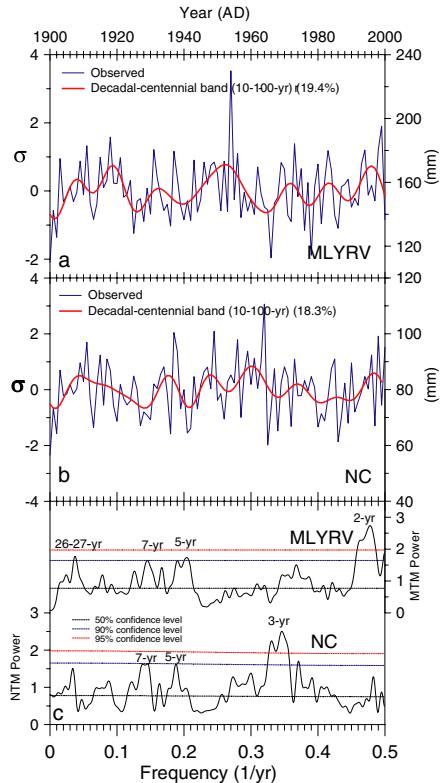


Fig. 2. (a) Observed summer precipitation in the MLYRV and its Wavelet filtered ($T_{\text{pass}}=10\text{--}100\text{-yr}$, paul basis) time series. (b) Observed summer precipitation in NC and its Wavelet filtered ($T_{\text{pass}}=10\text{--}100\text{-yr}$, paul basis) time series. (c) MTM results of the MLYRV and NC summer precipitation time series. Resolution and taper number used in the MTM analysis are 2 and 3, respectively.

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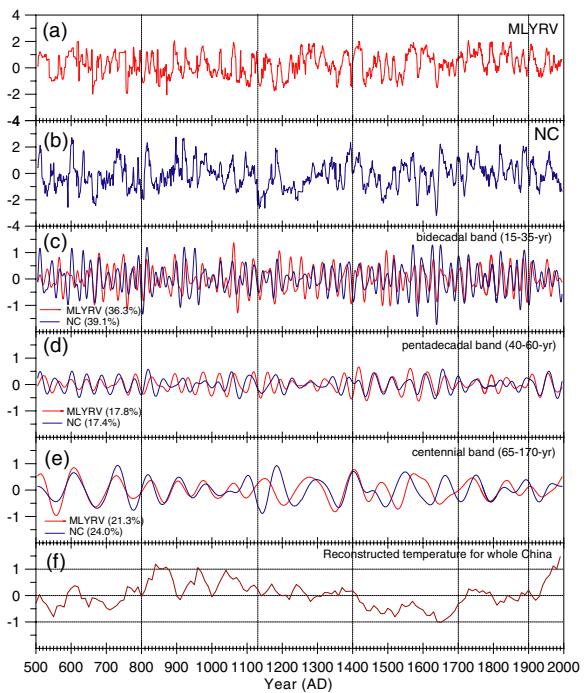


Fig. 3. The normalized time series of 10-yr moving averages of DWI for the MLYRV **(a)** and NC **(b)**. Wavelet filtered DWI time series for the MLYRV and NC in bidecadal band ($T_{\text{pass}}=15\text{--}35$ -yr, paul basis) **(c)**, pentadecadal band ($T_{\text{pass}}=40\text{--}60$ -yr, paul basis) **(d)**, and centennial band ($T_{\text{pass}}=65\text{--}170$ -yr, paul basis). A curve of reconstructed temperature for whole China (Yang et al., 2004) is also shown for a comparison of summer precipitation variability over eastern China with temperature variation in China.

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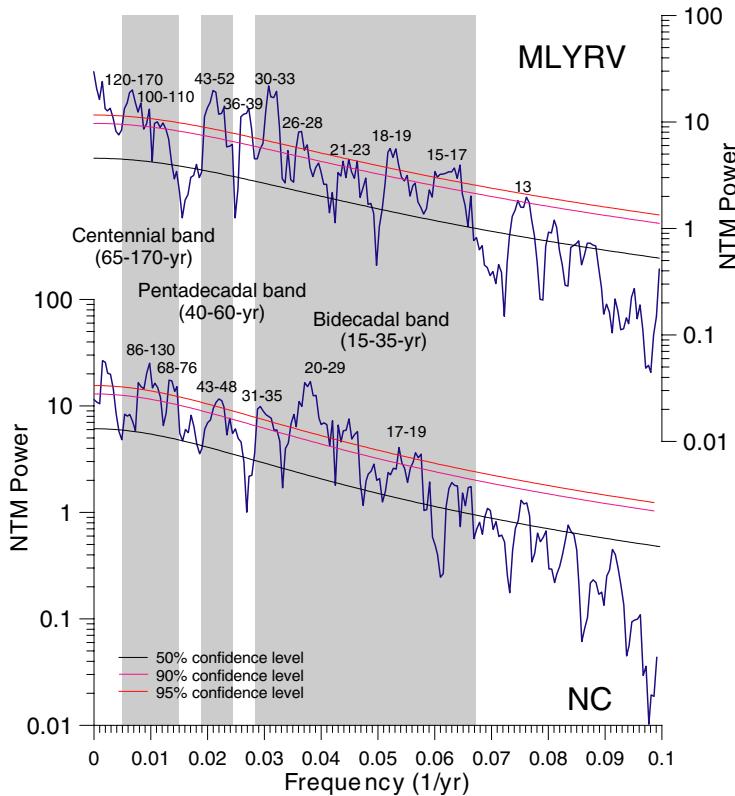


Fig. 4. MTM results of the MLYRV and NC DWI time series. Resolution and taper number used in the MTM analysis are 2 and 3, respectively.

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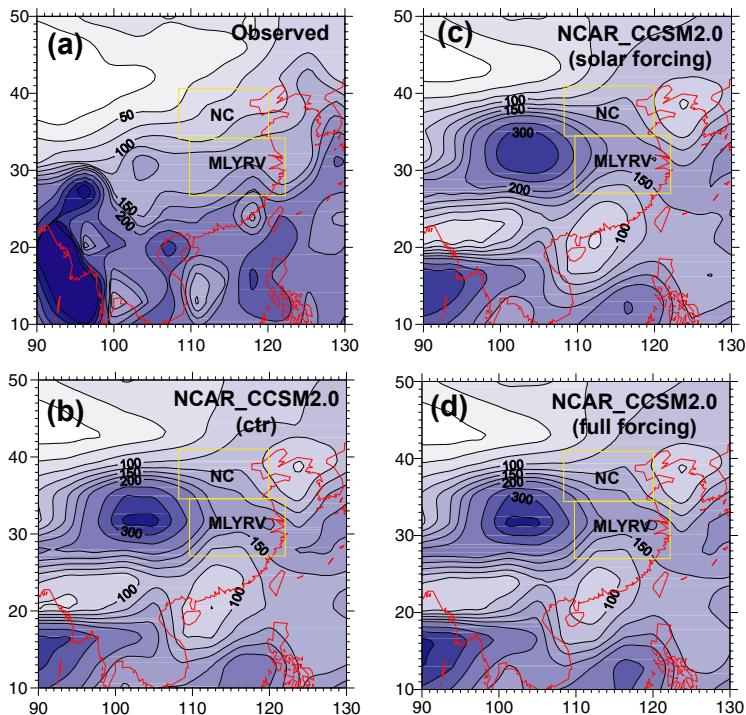


Fig. 5. Spatial patterns of summer precipitation (unit: mm/month) in observed data **(a)**, and modeled data of three experiments, i.e. control run **(b)**, orbital and solar forcing driven run **(c)**, and full forcing driven run **(d)**. All data are 45-yr averages of 1951–1995.

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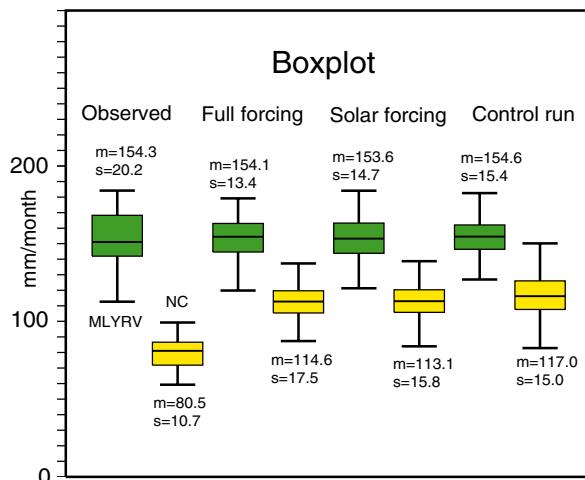


Fig. 6. Boxplots summarizing the distribution statistics of observed and modeled summer precipitation (May–September) from 1951 to 1995. In each boxplot the box extends from the lower (first) to the upper (third) quartile, with the line inside the box denoting the median. Lines indicating the tails extend outward from the quartiles to the minimum/maximum value. The mean values and standard deviations are also shown.

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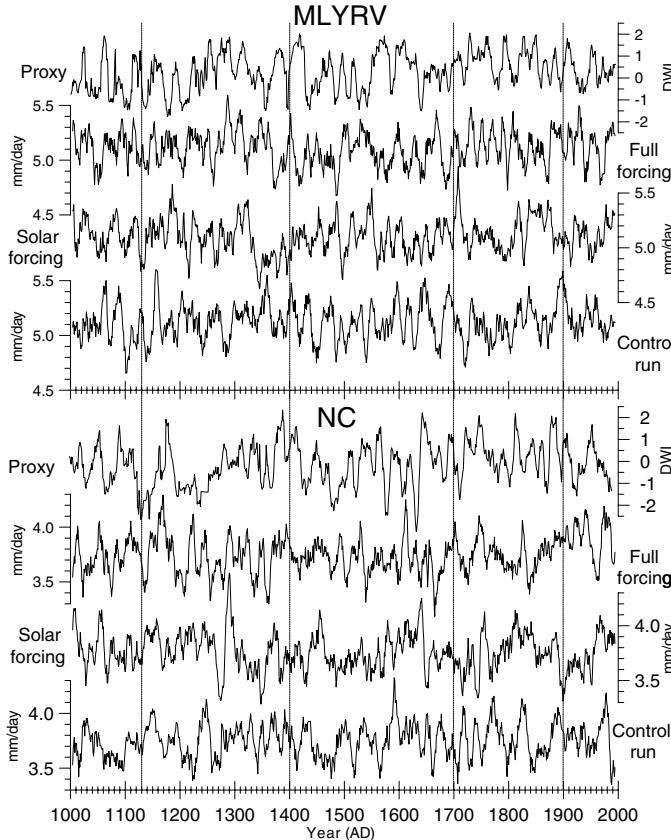


Fig. 7. A comparison of summer precipitation from proxy data and modeled data in the MLYRV and NC. All data are 10-yr moving averages. The dot lines represent the period boundaries recognized by temporal evolutions of decadal to centennial oscillations in Sect. 3.2.

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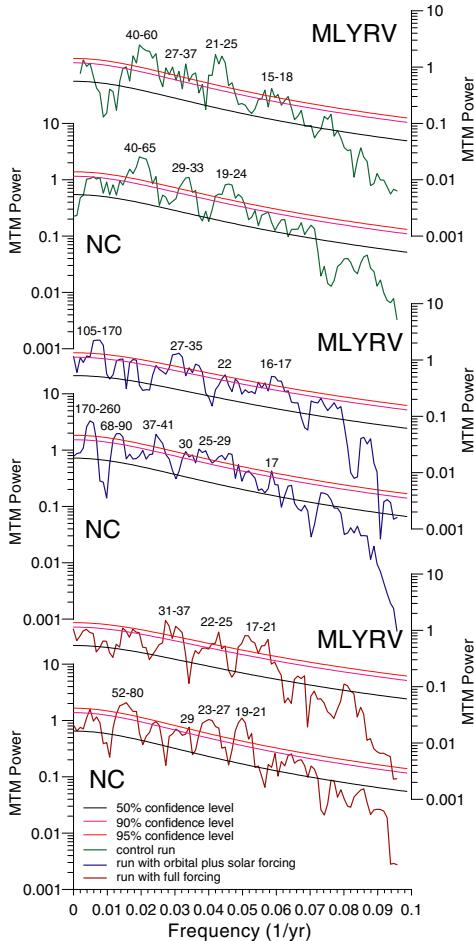


Fig. 8. MTM results of summer precipitation in three experiments of millennium simulation. Resolution and taper number used in the MTM analysis are 2 and 3, respectively.

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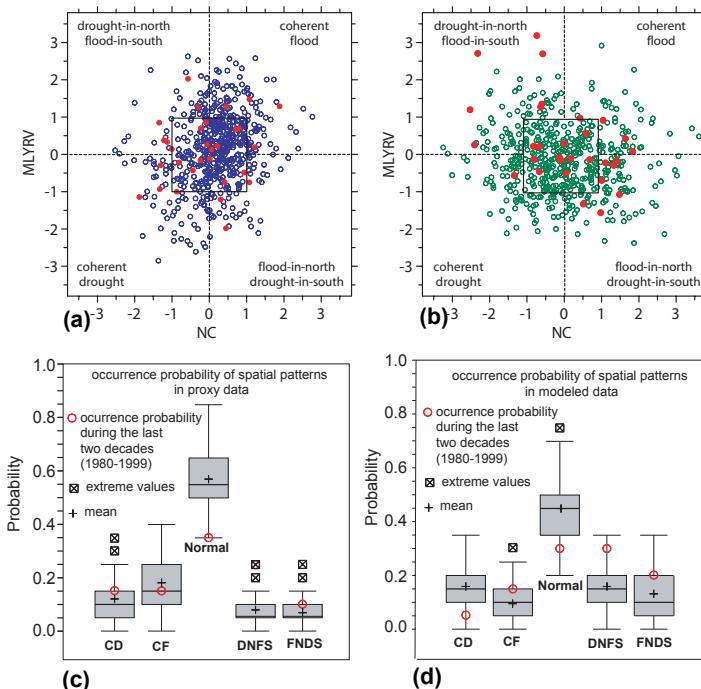


Fig. 9. Four-quarter plots of standardized proxy data (DWI) **(a)** and modeled data (experiment with full forcing) **(b)** of summer precipitation during AD 1470–1999 over the MLYRV and NC showing the spatial pattern of anomalous summer precipitation events. The last 30 yr (1970–1999) are presented by red dots. The proxy and modeled data are standardized using their means and standard deviations of 1951–1995. Boxplots show the probability distributions of different spatial patterns of anomalous summer precipitation events revealed by proxy **(c)** and modeled **(d)** data. The shaded box around the median represents the inner quartile range (IQR), which identifies the region between the 25th and 75th percentiles. The whiskers extending from the IQR represent the largest/smallest value that is still within the median ± 1.5 (IQR). Any value beyond the whiskers is subjectively considered to be an extreme value or an outlier.